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**A STOCHASTIC DAILY CLIMATE MODEL  
FOR SOUTH AFRICAN CONDITIONS**

BY

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AND

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REPORT DESCRIBING A RESEARCH PROJECT CARRIED OUT BY THE DEPARTMENT OF  
MATHEMATICAL STATISTICS, UNIVERSITY OF CAPE TOWN, UNDER CONTRACT TO THE  
WATER RESEARCH COMMISSION.

DEPARTMENT OF MATHEMATICAL STATISTICS  
UNIVERSITY OF CAPE TOWN  
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## EXECUTIVE SUMMARY

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#### Motivation

Effective water resources management is essential in a country like South Africa which is particularly prone to the adverse effects of drought. This will only become feasible when the risk associated with drought occurrences can be reliably assessed.

Present methods of assessing the risk of adverse weather conditions are based on rainfall and streamflow only and do not take account of the many other climatic factors such as evaporation, humidity, wind run, temperature etc. Such factors play an important role in establishing drought conditions, especially in the agricultural sector.

Methods, such as those based on the Palmer drought index, are purely descriptive and are designed to quantify what has happened in the past rather than what is likely to happen in the future. These methods are therefore of limited use for planning purposes.

This project arose from a need to develop reliable methods to generate artificial climate sequences over any period of the year and thereby enable water resources and agricultural planners to assess the probable consequences of decisions whose outcomes depend on climate factors. For example, sequences generated by a suitable model could be used as the input to plant yield models associated with crops such as maize, wheat and sugar cane, and thereby provide the probability distribution of yield under alternative options regarding, for example, planting date, cultivar and irrigation strategy.

The climate model to be developed in the course of this project was seen as a logical extension of the daily rainfall model which was developed in a previous Water Research Commission project (WRC Report No. 91/1/84 – 91/3/84). The latter model has been used by a number of institutions involved in Forestry, Agriculture, Nature Conservation, as well as by individual researchers at a number of universities and the South African Museum. It is offered as one of the data products available from the Computing Centre

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The steering Committee responsible for the project consisted of the following members:

Dr G C Green	Water Research Commission (Chairman)
Mr P W Weideman	Water Research Commission (Secretary)
Dr P C M Reid	Water Research Commission
Prof T J Stewart	University of Cape Town
Mr S van Biljon	Dept of Water Affairs
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Dr A L du Pisani	Dept of Environmental Affairs
Mr D B Versfeld	FORESTEK, CSIR
Mr S D Lynch	University of Natal

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We wish to express our thanks to the individuals whose advice and cooperation contributed to the completion of this project, especially the staff of the Department of Mathematical Statistics at the University of Cape Town. In particular we wish to thank Professor Theo Stewart for participating in the running of the project during 1989 and 1990.

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of Water Research. The climate model, incorporating several additional variables, would therefore supplement the rainfall model.

### **Objectives**

The objective of this project was to develop a stochastic model for the simultaneous description of climate variables at fixed locations on a daily basis. The variables to be analysed were rainfall, sunshine duration, maximum and minimum temperature, maximum and minimum relative humidity, evaporation and wind run.

Once a suitable model was identified the object was to develop methodology to estimate the relevant parameters from a given historical record and then to develop algorithms to generate artificial daily climate sequences at the given site.

A further objective of the project was that the technicalities of the methodology developed should be transparent to the user, that is, the results should be accessible to users with limited or no knowledge of statistics.

### **Summary of results**

We investigated the properties of the only daily climate model (Model 1) that has been described in the literature. A number of limitations of this model were identified and four alternative models were constructed, Models 2, 3, 4 and 5. (Model 2 was designed as a prototype for the subsequent models and is described in the report for the sake of completeness rather than as a suitable model in its own right.)

The new models, which vary in complexity, are designed to form a compatible family. This allows one to select a model of appropriate complexity for the particular historical record that is available. In general the simpler models outperform more complex models when the historical record is short (as is presently the case at almost all sites in South Africa) whereas the latter can be expected to become increasingly applicable as more data becomes available. Furthermore the compatibility property allows one to model the different climate variables using components from any one of Models 3, 4 and 5 and then to combine these into a single multivariate daily climate model.

To fulfil its purpose a daily climate model must incorporate all the important properties exhibited by climate variables. These include the seasonal cyclical behaviour of climate, its short-term persistence, the interrelationships between the different variables (for example between rainfall and humidity) and the boundedness of some of the variables (for example

the upper and lower limits of maximum and minimum relative humidity). In addition the behaviour of each of the variables on wet days is different to that on dry days. For example, on average, the maximum temperature on dry days is higher than it is on wet days. All these properties have to be preserved, not only qualitatively, but also quantitatively by the climate model.

The results of this project confirm that it is indeed possible to construct models that preserve the above properties. This is in spite of the fact that the historical records which are presently available in South Africa are extremely short for the purpose of modelling a process of the complexity of daily climate. (The length of the records available to us ranged from 6 to 12 years.) An additional factor which reduces the *effective* length of the records for this type of modelling is the average number of *rainy* days which, in many parts of South Africa, is quite small.

The models were calibrated at six sites, namely Elsenburg (South Western Cape), Kakamas (Northern Cape), Middelburg (Eastern Central Cape), Nelspruit (Eastern Transvaal), Cedara (Natal) and Hoopstad (Orange Free State) which, within the constraints of the data available to us, were selected to represent as wide a variety of climate types as possible. Extensive validation tests were carried out and our results show that, on the whole, the models perform remarkably well.

There is no clear-cut answer regarding which model will perform best. As mentioned, one would expect the simpler models to outperform the more complex alternatives when the data records are short. For some sites Model 1 preserves the properties of some of the climate variables better than the more complex alternative models. At other sites the opposite was found to be the case. We therefore recommend that, at new sites, each of the models be applied and tested before a final selection is made.

A major theoretical obstacle that had to be overcome in the course of the project was that of developing methods which could accommodate records with missing observations, with invalid recordings and with outlying observations. Although we had access to some of the best historical records that are available in South Africa, there were considerable gaps and imperfections in these records (amounting to between 1% and 13% of the total record lengths for the records which we examined). As climate variables are both serially correlated and cross-correlated it is not possible to simply ignore missing values. Methods had to be developed to incorporate the estimation of missing values as part of the parameter

estimation procedure.

A substantial portion of the research effort in this project was directed to deriving the mathematical theory for the climate models which were developed. This material, which is rather technical and thus is not accessible to the general user, includes the development of estimation methods both for the individual climate variables as individual time series models and then for the multivariate series which combines these models so as to synchronise the various climate variables.

The second major component of the project was the preparation of computer programs to implement the theory. In order to make the software accessible to as wide a variety of users as possible it was decided at the outset that all programs would be such that they could be implemented on micro-computers. Secondly, it was decided that no use should be made of licensed software packages which may not be available to some potential users. Thus the programs which are listed in this report are self-contained and are coded in ANSI FORTRAN 77, (the HUGE attribute in programs 6 and 8 is an extension to the full ANSI standard but this can be omitted without any problems on a mainframe) a language for which compilers are generally available. This includes the programs to estimate the parameters and to generate the required artificial climate sequences.

An objective of the project was that the results of the project should be accessible to users with limited or no knowledge of statistics. This objective has been mostly met, but with the following qualification. The programs that have been developed to generate the climate sequences are accessible by any user who can operate a micro-computer. Such a person would not have to know anything about programming but merely how to run an existing program. We envisage that most users of the methodology will only be interested in making use of the generating program.

Some training is required to apply the methodology at a new site, that is, to estimate the model parameters from a given historical record. We estimate that, with instruction, it would take a competent programmer between two to three weeks to learn how to make efficient use of the estimation software provided. Most of the training would be concerned with methods for preparing the data for estimation. This aspect of the methodology simply cannot be automated since it requires judgement.

Thus we must distinguish between two types of users; those who wish to calibrate the

model for a new site and those who wish to use the model for a site that has already been calibrated. The former task requires some training but the latter does not. This issue is discussed in the recommendations below.

We believe that the main objectives of the project have been met. We have demonstrated that the models which were investigated in the course of this project meet all the requirements that can be reasonably expected of models for a phenomenon of the complexity of daily climate.

## **Recommendations**

### **Quality of historical records**

The main obstacle to the application of the techniques described in this report on a large scale is the lack of suitable historical records. This refers to both the quantity and the quality of available data. The records which were used for this report represent some of the best available in South Africa. Nevertheless, for the purpose of modelling daily climate, they are barely adequate. Although there is little that can be done to increase the length of records except to wait for more data to be collected, it should be possible to improve the quality of historical records. In particular it would be useful if some measure of the reliability of the observations were also recorded on a regular basis. As we have repeatedly pointed out in the body of the report, one of the problems which we encountered was that of identifying incorrect observations. This task would be considerably simplified if one had some index of reliability associated with (ideally) each recording or set of recordings.

### **Transfer of technology**

For the methods developed in this project to realise their full potential it will be necessary to calibrate the models at many more sites. As was pointed out, no special training is required to use the programs for generating climate sequences once the parameters of the model have been estimated. However, some training is required to use the programs to prepare the data for estimation and to carry out the estimation for a new site.

We recommend that the Computing Centre for Water Research (CCWR) be approached to acquire the expertise to implement the estimation techniques and with the help of users, gradually build up a data base of estimates of the model parameters for as many sites as possible in South Africa. The CCWR already offer a similar data product, namely the parameter estimates of a daily rainfall model for 2550 sites in South Africa. These arose from a previous Water Research Commission project (Zucchini and Adamson (1984)). The

CCWR also offer the artificial rainfall generating program which can be applied to any of these sites. Thus the programs developed in the course of this project constitute a logical extension of a service that the CCWR already offer.

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## CHAPTER 1

### INTRODUCTION

Climate is a critical factor in determining the variety and abundance of vegetation and animal life that can exist in a region. It imposes limits on agricultural and other human activities that are economically feasible. Thus it is not surprising that various aspects of climate, such as precipitation, temperature, solar radiation, humidity, wind speed and others, are recorded on a regular basis throughout the world.

The purpose of measuring these climate variables is to extend our knowledge of the behaviour patterns of climate and thereby, among other things, to identify those activities which are feasible and to determine how these may be most profitably carried out. For this, one has to take account of the fact that both the climate process and human requirements, such as demand for water, are dynamic processes which are stochastic rather than deterministic in character. For example, the annual rainfall in most regions of South Africa varies considerably from year to year and it is obviously inadequate to base water-related decisions solely on the average annual rainfall; the entire distribution of annual rainfall needs to be considered.

Statistical theory provides an ideal framework for expressing our knowledge about the properties of climate. Firstly, it provides a means of quantifying our knowledge in a precise manner. Secondly being designed to describe stochastic phenomena, the theory provides a conceptual framework which accommodates notions such as uncertainty and risk, thereby providing a convenient basis for rational decision making in the face of uncertainty. Thirdly, statistical methodology provides an effective means of synthesising and analysing the information contained in large data sets such as daily climate records. In particular the theory enables one to quantitatively distinguish the systematic patterns in climate (such as the seasonal cycles) from the random fluctuations about these patterns and to express this information in terms of a statistical model.

As well as providing a concise description of the patterns that exist in the different components of climate, a statistical model can be used to generate artificial climate sequences which preserve the properties of real climate sequences, that is, artificial sequences that are indistinguishable from real climate sequences. Among other things, artificial climate sequences are useful as inputs to crop growth models which can then be used to determine

the distribution of yield, the risk of crop failure due to adverse climate, optimal planting dates, the potential profitability of irrigation, and so on. For such purposes artificial climate sequences generated by a good stochastic model are more useful than the original historical record. Firstly, they are free of the typical imperfections which are especially prevalent in historical climate records, for example, incorrect recordings and missing observations. Secondly, the historical records presently available in South Africa are mostly quite short and thus only reflect a small fraction of the different climate sequences that could occur.

It is sometimes argued that artificial sequences generated by a stochastic model constitute no more than complicated extrapolations of the historical record. However, a model contains more than the information that can be extracted from a single historical record. It contains our knowledge (in the form of model assumptions) about the behaviour of climate derived from theory and from observations at other locations. For example, it is reasonable to assume that certain average properties of climate variables vary smoothly with time. Such assumptions give the model a structure which may not be evident in a single short historical record.

The main objective of this project has been to develop a stochastic daily climate model for South African conditions. The time resolution was taken as one day because climate data commonly available are recorded on a daily basis. The variables included in the model are rainfall, maximum and minimum temperature, maximum and minimum relative humidity, evaporation, wind run and sunshine. In fact the models that have been developed can be used to model a subset of the set of variables in cases where some of the above set are not available. Alternatively, it is possible to augment this list if measurements on additional variables are available.

As already mentioned the model needs to preserve the important properties of daily climate sequences. These were identified as being:

- a) **Seasonality.** Each of the climate variables exhibits seasonal behaviour, that is, the recordings fluctuate about a curve which has a cyclical pattern with a period of one year. The shape of the curve is approximately sinusoidal which suggests that it can be parsimoniously approximated by a truncated form of its Fourier representation.
- b) **Wet/dry day effect.** The probability distribution of the climate variables on wet days is different from their distribution on dry days. For example, the maximum relative

humidity is generally higher on wet days than on dry days; and the opposite is true for the number of sunshine hours. Thus it is necessary to treat dry days and wet days differently in the model.

c) **Autocorrelation.** The individual variables exhibit short-term persistence over and above that attributable to seasonality. Generally there is a positive correlation between readings on successive days. This type of persistence needs to be incorporated into the model.

d) **Cross-correlation.** Apart from the wet/dry day effect already mentioned, the variables are cross-correlated. For example, there is a positive correlation between minimum temperature and maximum temperature on the same day. To preserve this property it is not possible to model the climate variables separately — they have to be modelled jointly.

e) **Boundedness.** The values of some of the variables are bounded, for example, relative humidity lies in the range 0% to 100%. Other variables are bounded with respect to others, for example, the minimum temperature on any one day must not be higher than the maximum temperature on the same day. To preserve this type of property the variables have to be transformed.

f) **Non-normality.** The probability distribution of climate variables does not follow the normal distribution. This is problematic because there is practically no other multivariate distribution available that is both sufficiently flexible and mathematically tractable to deal with a phenomenon as complex as climate. It is therefore necessary to transform the variables to achieve normality.

Taken together these properties indicate that we are dealing with a multivariate time series which is non-stationary and which contains a number of variables with special properties. In particular rainfall has the property that it is partly discrete (there is a non-zero probability that it does not rain) and partly continuous (the rainfall depth on rainy days is a continuous random variable). There is no standard statistical model which can be applied directly to such a multivariate time series. A special model has to be constructed for the daily climate process.

Although there is an extensive literature on the modelling of daily rainfall sequences, apart from Richardson (1981), very little work has been reported on models which describe the joint probability distribution of several components of climate. Richardson proposed

separating the observations into two sequences; one for observations which occurred on wet days and one for those which occurred on dry days. For each of these two sequences the seasonal mean and standard deviation of each variable is estimated separately and then a new time series (of residuals) is obtained by deseasonalising the original observations using the appropriate means and standard deviations, depending on whether the observations occurred on wet days or on dry days. The (multivariate) time series computed in this way has mean zero and variance unity and is modelled using a single multivariate autoregressive model.

Richardson's model (Model 1) is probably the simplest structure that has the potential of preserving properties (a) to (d), outlined above, in a sufficiently flexible form. To accommodate (e) and (f) it is necessary to make suitable transformations of the variables at the very start of the modelling procedure. Such transformations are required for all the models which were considered.

At the start of the project Model 1 was fitted to six years of record (1979–1984) at Elsenburg. The model was found to fit some aspects of the historical records quite well but performed poorly on certain other aspects. In particular the annual standard deviation for wind run, maximum and minimum humidity were systematically underestimated. The (lagged) cross-correlations between some of the variables (e.g. maximum temperature and minimum temperature) were not preserved by the model. However the most noticeable deficiency was found to be that the model did not preserve the serial correlation structure of many of the variables. This was attributed to the lack of flexibility of Model 1 in this respect. In particular the model is based on the assumption that the serial correlation function does not depend on the wet/dry status of the days in question. In fact the correlation between variables on two successive days depends on whether the two days are both wet, both dry, wet followed by dry or dry followed by wet. It was therefore decided to develop a model which incorporates additional flexibility in its autocorrelation function, that is, a model which allows for the serial correlations between variables on successive days to depend on their wet/dry status.

In developing a model for a process as complex as daily climate there are two conflicting objectives. On the one hand it is desirable to construct a model that is as flexible as possible so that it can accommodate as many of the special features of the process as possible. On the other hand additional flexibility can only be achieved by increasing the number

of unknown parameters in the model. These parameters have to be estimated from the historical record. Now, for a record of given length, increasing the number of parameters that has to be estimated decreases the precision of the estimates, on average. Put differently, if a model has too many parameters it becomes too specific to the particular historical record that is available and less representative of the population of typical climate sequences that could arise. The appropriate complexity of a model depends on the length of the historical record. In general, simpler models which depend on only a small number of parameters will outperform more complex models if the historical record is short, but the reverse is true if the record is large. A second issue is that some of the climate variables are more appropriately modelled by simpler structures than others. It is therefore not always optimal to use the same model for all the components of climate.

The strategy that we adopted to circumvent the above difficulties was to develop a *family of models* of varying degrees of complexity ranging from the simplest feasible model to more complex alternatives. This allows one to select the particular model from the family which is most appropriate for the historical record that is available. In addition the family which was developed is such that the individual models within the family are compatible in the following sense. One can use different submodels for each of the individual climate variables and then combine these into a multivariate model at the last stage of the modelling procedure. Thus, for example, it is possible to fit a simple model to wind run but a more complex model to minimum temperature. This compatibility feature of the family thus allows for additional flexibility.

Three compatible models were developed which we will refer to as Models 3, 4 and 5 (Model 2 was developed as a prototype to the others and is included in the report for the sake of completeness). Models 3 and 4 are two alternative relatively simple models whereas Model 5 is more general than each of them. Thus one would expect Models 3 and 4 to be suitable for short data records (as are presently available in South Africa) and Model 5 to become preferable as the historical data base increases in length. The method of maximum likelihood was used to estimate the parameters. Since the likelihood equations are extremely complex, it was necessary to develop numerical methods to carry out the estimation. This involved deriving the first and second derivatives of the likelihood function (given in Chapter 3) and the development of procedures to compute initial estimates of the parameters.

One of the major problems which we encountered in applying the estimation procedures was the presence of missing observations and also of outliers, mainly in the form of incorrect readings (e.g. outside the admissible range of values). It is necessary to filter the data in order to remove such outliers before attempting to estimate the parameters — this introduces additional gaps in the record. Thus the estimation procedure that was developed had to be able to cope with the problem of missing values.

Some aspects of the lack of fit of Model 1 which were identified at the start of the project were later found to be attributed, at least in part, to outlying observations. In fact, a conclusion of this project is that in many respects Model 1 outperforms the more sophisticated models developed here. All the models considered here are strongly influenced by outlying observations. This fact makes it necessary to pay special attention to the quality of the historical record before attempting to fit a model.

In order to objectively determine which member of the family of models is most appropriate in a given situation a model selection criterion is used. The Akaike's Information Criterion (AIC) is proposed for this purpose.

Six sites were selected to evaluate the performance of the models considered in this report. The choice of the sites was, of course, constrained by the availability of suitable historical records. Within this constraint we attempted to represent, as well as possible, the various climate regions of South Africa. The sites chosen were:

Elsenburg — South West Cape

Kakamas — Northern Cape

Middelburg — Eastern Central Cape

Nelspruit — Eastern Transvaal

Cedara — Natal

Hoopstad — Orange Free State.

The aim of model validation is to establish that the models preserve the important properties of the historical records, at least to an appropriate degree, so that the generated sequences can be regarded as representative of the population of sequences which could occur.

An objective of this project was that the technicalities of the methodology should be

transparent to the user, that is, the results should be accessible to users with limited or no knowledge of statistics. Here one must distinguish between the person who fits the model to observations at a new site and the person who uses the fitted model to generate artificial climate sequences. A limited amount of training is required to apply the estimation techniques using the software that is provided, but once the parameters of the model have been estimated at a particular site, the software provided to generate artificial climate sequences is accessible to anyone who can run a computer program.

In order to make the methodology accessible to as wide a variety of users as possible, it was decided at the outset of the project that all the software developed would have to be such that it could be implemented on micro computers. Secondly, it was decided that no use should be made of licensed software which may not be available to some potential users. The programs listed at the back of this report are self-contained — no additional software is required either to estimate the parameters or to generate artificial climate sequences.

This report is structured as follows:

The preliminary statistical analysis of the data is described in Chapter 2. This includes a description of the data, the types of difficulties encountered in detecting and dealing with faults in the data and the statistics computed to identify the structures present in the climate sequence.

Chapter 3 gives a theoretical description of the five climate models which were investigated and of the methods used to estimate the model parameters. The contents of the chapter are technical and very detailed and thus rather demanding of the reader. Fortunately it is not necessary to absorb all this detail in order to understand the remainder of the report. We recommend that the reader who is not concerned with the mathematical development of the models simply skip over this chapter. Details on the implementation of the models to the historical records are given in Chapter 4. The algorithms for implementing the theory are listed in Chapter 5. These include algorithms for generating artificial climate sequences. These algorithms are intended to bridge the gap between the formulae given in Chapter 3 and the FORTRAN programs given in Appendix D. Extensive tests were performed on the fitted models in order to assess their performance in preserving the important properties of climate sequences. The results of this model validation investigation are summarized in Chapter 6. A summary of the findings of the study and the main conclusions are given in Chapter 7.

There are 5 Appendices: Appendix A explains the choice of the Fourier approximation,  $L$ , Appendix B describes the properties of the Fourier series approximation, Appendix C gives an algorithm, known as the Cholesky decomposition, which rewrites a matrix as a product of a triangular matrix with its transpose. This is needed to generate normal random numbers with a covariance matrix  $\Sigma$ . Appendix D gives information on where a list of the ANSI FORTRAN 77 programs used in this study can be obtained. Appendix E describes the EM algorithm, a very general iterative method for maximum likelihood estimation in incomplete data sets. The EM algorithm is used in this study to estimate and fill missing values in the climate data sets.

## CHAPTER 2

### THE DATA SET AND PRELIMINARY ANALYSIS

It is to be expected that data records collected over a long period of time will contain gaps, and usually the number of gaps increases in proportion with the size of the data set. The data sets considered in this study are no exception to this.

Gaps occur for two reasons. Firstly, a high proportion of the observations are missing. Although missing observations are relatively easy to detect, they lead to complications in the analysis. In particular, the multivariate time series models considered here require simultaneous observations of all the variables. Furthermore, the serial correlation structure in the series does not allow one to simply discard observations as one would do if the observations were serially independently distributed.

Secondly, some of the readings are incorrect (or incorrectly recorded). These are often quite difficult to detect, especially if the values fall within the feasible range of the variable under consideration. This problem is particularly difficult to deal with satisfactorily.

This chapter describes the general format of the data sets used, some of the problems encountered and the method used to overcome them. Finally, some preliminary analyses performed for initial model identification are discussed.

#### **The data set**

##### **(a) Format**

The climate variables of interest are:

- rainfall (mm)
- maximum temperature ( $^{\circ}C$ )
- minimum temperature ( $^{\circ}C$ )
- A pan evaporation (mm)
- sunshine duration (hours)
- windrun (km/day)
- maximum humidity (%)
- minimum humidity (%)

Not all stations have records for evaporation, however, as it is easily derived from other climate variables, it is not essential to include it in the models. Whenever readings are available, evaporation is kept purely to demonstrate the elasticity of the models in that variables can be omitted or incorporated without the model structure changing. The number of variables is simply increased.

The unit of measurement for each variable is shown above in brackets following the variable name.

Three properties of the time series (discussed later) determine how the final data set for parameter estimation must be constructed. Firstly, simultaneous observations for all variables are required as one is dealing with a multivariate time series. As data collection of some variables (humidity, for example) has only been started recently, only years for which measurements are available for all variables simultaneously can be used. The only exception is rainfall as it is modelled independently of the other variables.

Secondly, continuous data is required because of the seasonality and serial-correlation structure in the time series. Large gaps in the data caused by shutting down a station for a long period of time and then reopening it at a later stage, cannot be treated as missing values. Only the sequence previous to or following the closing (depending on which period is longer) may be used.

Finally, records should begin on 1 January and end on 31 December. This restriction simplifies the algorithms and the programming. However, it is not necessary to waste data in order to meet this requirement if only a few months are missing in a year. For example, if the original available record starts on 1.2.1978, then one should code the days 1.1.1978–31.1.1978 as missing and then regard the record as starting on 1.1.1978.

### **(b) Quality**

There is always a possibility of readings being recorded incorrectly. The type of recording error which can easily be identified is when the value recorded lies outside the permissible range, for example a recorded value for sunshine duration of 25 hours. In addition, a variable may have values which, although within a feasible region, are nevertheless incorrect. This cannot be established with any certainty. The fact that one is clearly dealing with data sets that are not “clean” and that preliminary work done showed the models to be extremely sensitive to “unclean” data, a thorough procedure to detect possible errors in the records is

necessary.

From model assumptions, the residual time series obtained after fitting Model 1 (any other model can be used) to the climate series, is normally distributed with mean zero and standard deviation of unity. Therefore, 99.7% of the residual values should lie within the interval  $(-3, 3)$ . In the present case, it is almost impossible to tell whether large residuals reflect model misfit or poor data, therefore large residuals were examined for possible occurrences of outliers. Preceding and succeeding values can give an indication whether or not these values should be considered as outliers. Observations across the variables at these times also show what patterns to expect. For example, Barry and Charley (1968) state that evaporation can be expressed by:

- duration of sunshine
- mean air temperature
- mean air humidity
- mean wind speed.

Thus, one would expect to see an increase of evaporation with an increase of sunshine duration.

Outliers are treated as missing values.

### **(c) Treatment of leap years**

Whenever a leap year occurs, the value observed on the 29 February is added to the value observed on 28 February for the variables

- rainfall and
- evaporation.

For the variables

- maximum temperature
- minimum temperature
- sunshine duration
- windrun
- maximum humidity, and
- minimum humidity,

the mean of the observed values of the 29 February and 28 February replaces the observed value of 28 February.

If the 28 February has a missing value then it is replaced by the observed value of 29 February.

### **Distinctive features of the time series useful for model identification**

The station chosen for preliminary analyses was Elsenburg in the Cape Province (Latitude  $33^{\circ}51'$ , Longitude  $18^{\circ}50'$ ). Any station could have been chosen for this purpose as they all display similar features.

#### **(a) Seasonality**

A simple moving average smooth was used to filter the series. This is given by

$$M_t = \frac{1}{2L+1} \sum_{\ell=-L}^L m_{t+\ell}$$

where  $m_t$  is the mean of the time series at time  $t$ , i.e.

$$m_t = \frac{1}{I} \sum_{i=1}^I x_{i,t}$$

where  $x_{i,t}$  is the observation made at time  $t$  of the  $i$ th year,

$i = 1, 2, \dots, I$ ,  $I$  being the number of years for which data is available,

$t = 1, 2, \dots, 365$

and  $L$  is the lag.

Lags of 10, 25, 50 and 100 days were applied.

Note that in the above equation, because  $m_t$  is cyclic one has that

$$m_{366} = m_1$$

$$m_{367} = m_2$$

$$\vdots$$

$$m_0 = m_{365}$$

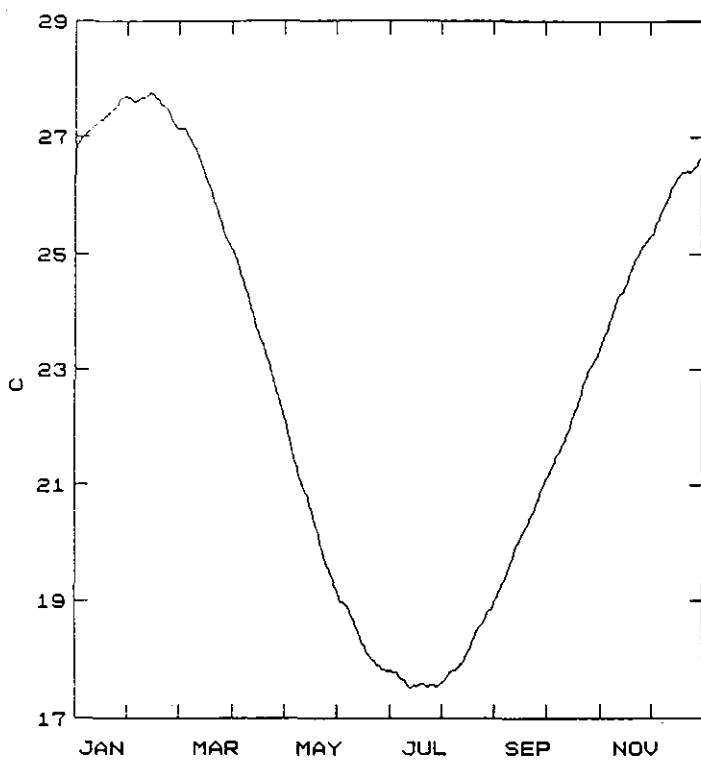
$$m_{-1} = m_{364}$$

and so on.

Figure 2.1 shows the smooth plots for the various variables.

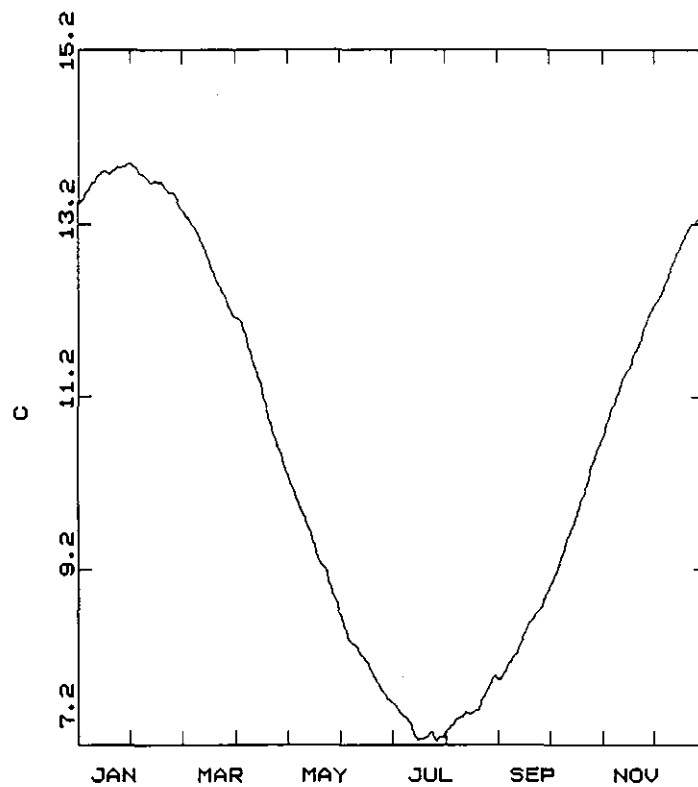
**FIGURE 2.1** Simple moving average smooth (lag=50) for all variables of Elsenburg

MAX TEMP



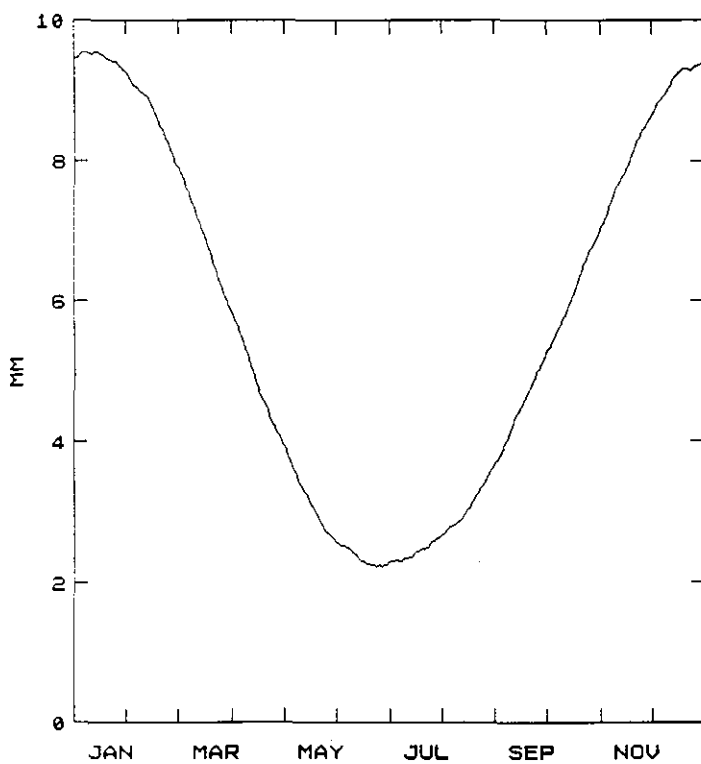
DAYS

MIN TEMP



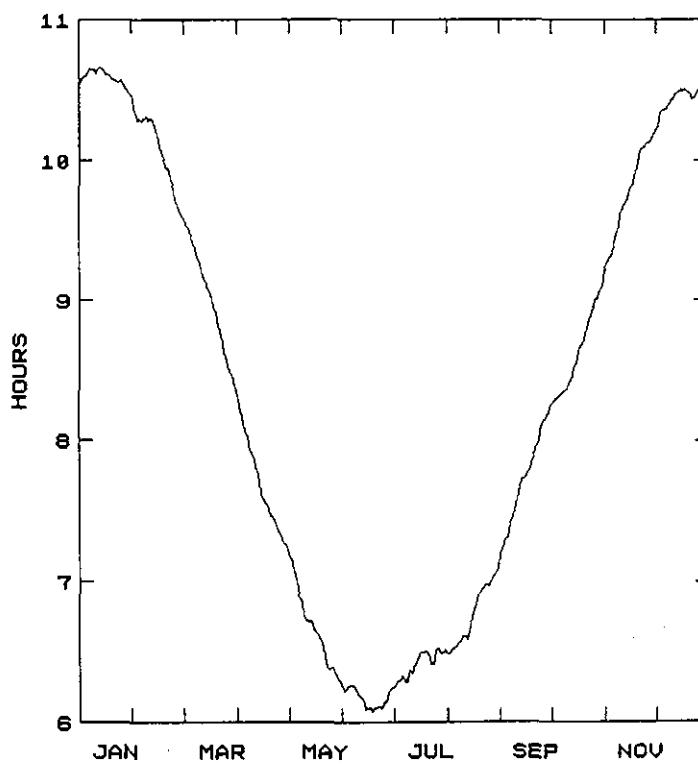
DAYS

EVAPO



DAYS

SUNSHINE

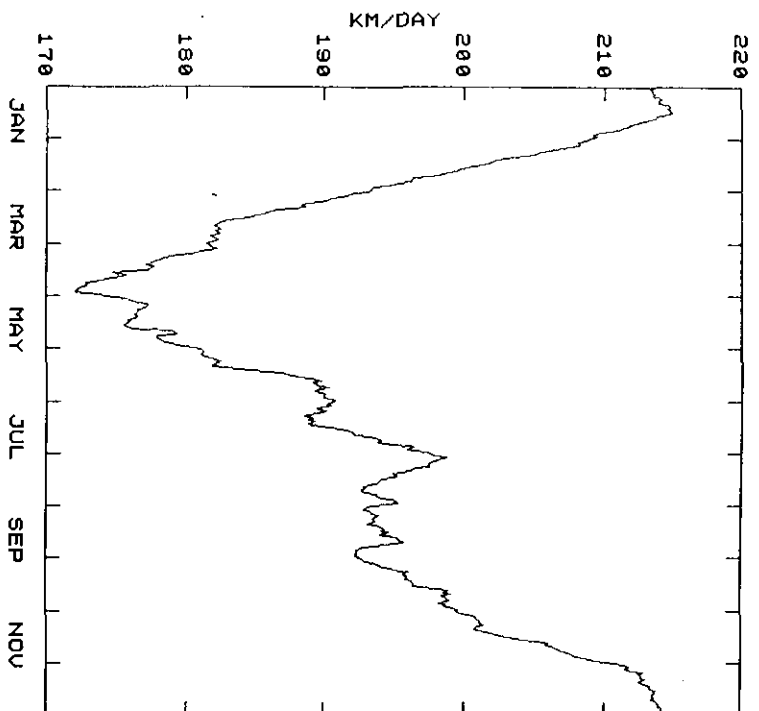


DAYS

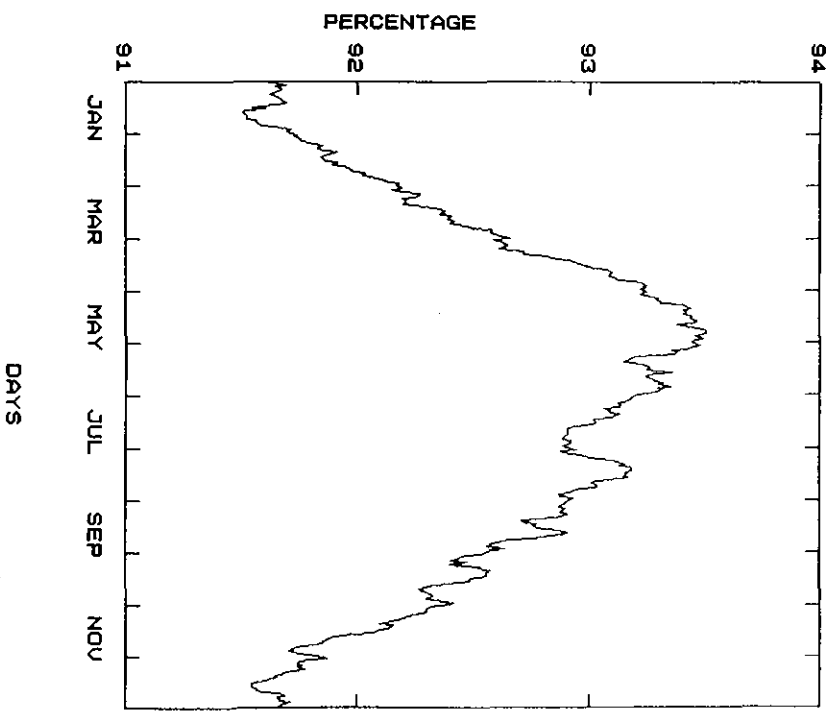
## CHAPTER 2

## *The data set and preliminary analysis*

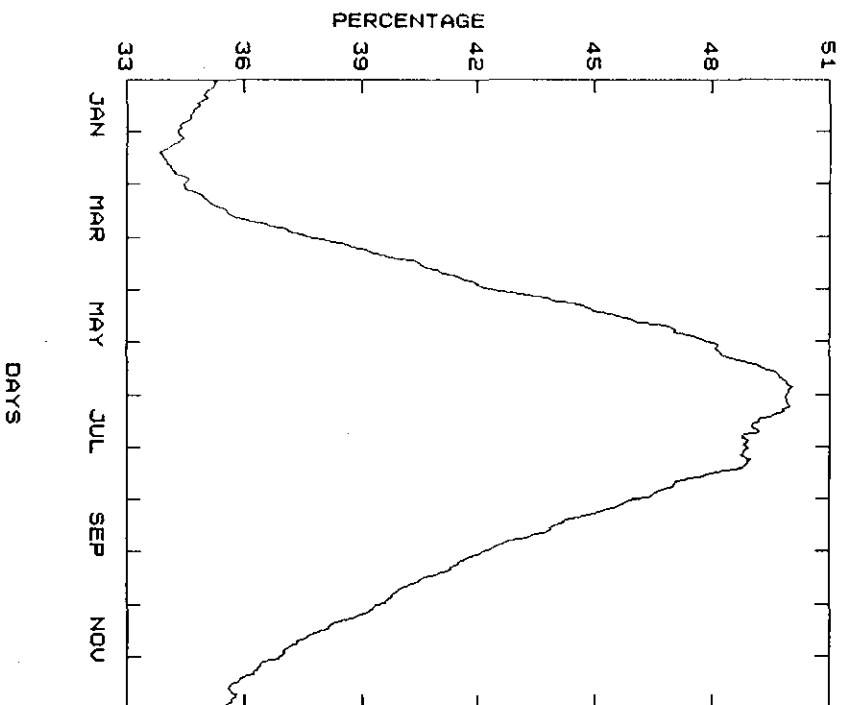
WINDRUN



MAX HUM



MIN HUM



From the smooth plots it can be concluded that each time series of the variables is seasonal, has a cyclic period of one year and has a sinusoidal shape.

### (b) Autocorrelation

Table 2.1 shows the autocorrelation for each variable up to lags of three.

The following abbreviations for each variable will be adopted in the annotation of tables and figures:

max temp — maximum temperature

min temp — minimum temperature

evapo — evaporation

sunshine — sunshine duration

max hum — maximum humidity

min hum — minimum humidity

From Table 2.1 it can be seen that the variables are autocorrelated, i.e. there is a short-term persistence within each variable.

---

**TABLE 2.1 Autocorrelation coefficients**

---

Variable	Lag 1	Lag2	Lag3
rainfall	0.23	0.08	0.05
max temp	0.77	0.59	0.51
min temp	0.69	0.54	0.48
evapo	0.76	0.70	0.66
sunshine	0.52	0.30	0.22
windrun	0.38	0.10	0.02
max hum	0.32	0.15	0.08
min hum	0.53	0.33	0.27

---

### (c) Cross-correlation

Intuitively, one would expect climate variables to be related to each other in some way,

for example one would expect the amount of evaporation to be related to the temperature. In fact, as already mentioned, evaporation can be expressed approximately in terms of the other variables. The interdependence among the variables was determined by computing the lag cross-correlation coefficients of the time series. These cross-correlation coefficients are shown in Table 2.2. These results confirm that the variables are indeed interdependent.

**TABLE 2.2** Cross-correlation coefficients between variables

Variables	Lag cross-correlation				
	$R_2(j,i)$	$R_1(j,i)$	$R_0(i,j)$	$R_1(i,j)$	$R_2(i,j)$
max temp – min temp	0.42	0.44	0.52	0.72	0.73
max temp – evapo	0.65	0.74	0.78	0.64	0.56
max temp – sunshine	0.52	0.65	0.63	0.34	0.24
max temp – wind	-0.07	-0.17	-0.14	0.07	0.11
max temp – max hum	-0.12	-0.21	-0.27	-0.19	-0.08
max temp – min hum	-0.42	-0.57	-0.71	-0.39	-0.26
min temp – evapo	0.65	0.61	0.51	0.50	0.48
min temp – sunshine	0.43	0.25	0.06	0.17	0.23
min temp – windrun	-0.01	0.22	0.26	0.14	0.07
min temp – max hum	-0.17	-0.18	-0.18	-0.05	-0.05
min temp – min hum	-0.44	-0.32	-0.07	-0.13	-0.20
evapo – sunshine	0.50	0.59	0.75	0.50	0.39
evapo – windrun	0.09	0.04	0.12	0.12	0.13
evapo – max hum	-0.12	-0.18	-0.29	-0.30	-0.17
evapo – min hum	-0.37	-0.46	-0.61	-0.47	-0.37
sunshine – windrun	0.03	-0.18	-0.20	-0.02	0.08
sunshine – max hum	0.00	-0.06	-0.19	-0.24	-0.15
sunshine – min hum	-0.16	-0.33	-0.70	-0.50	-0.32
windrun – max hum	-0.05	-0.08	-0.10	-0.12	-0.06
windrun – min hum	-0.12	-0.06	0.18	0.17	-0.01
max hum – min hum	0.16	0.32	0.32	0.17	0.07

(d) Time series observations differ depending on the wet or dry status of the day

It is known that on days that rain occurs, a marked change also occurs in other climatic variables, for example, temperature and sunshine duration are more likely to be below normal on rainy days than on dry days. Humidity, on the other hand will be above average on

a rainy day rather than on a dry day. This property of the climate variables was investigated to determine whether the difference was significantly distinct.

The observations of all variables were found to be significantly different depending on whether rain had or had not occurred in that time period. Figure 2.2 shows the mean time series of each variable conditioned on the wet or dry status of the day. Table 2.3 shows a comparison of the mean for each variable conditioned on the wet or dry status of the day.

---

**TABLE 2.3** Mean for conditioned time series

---

Variable	Dry state	Wet state
max temp	24.14	18.13
min temp	10.43	10.91
evapo	6.71	2.76
sunshine	9.73	4.15
windrun	177.7	245.8
max hum	91.9	94.3
min hum	36.8	55.1

---

Having concluded that climatic variables vary depending on whether rain or no rain has occurred, it remains to examine whether the amount of rainfall is related in any way to the observations of the climate variables. Figure 2.3 shows the graphs of rainfall versus each climate variable. From these plots it is concluded that there is no visible pattern to the values of the climate variables in relation to the amount of rainfall.

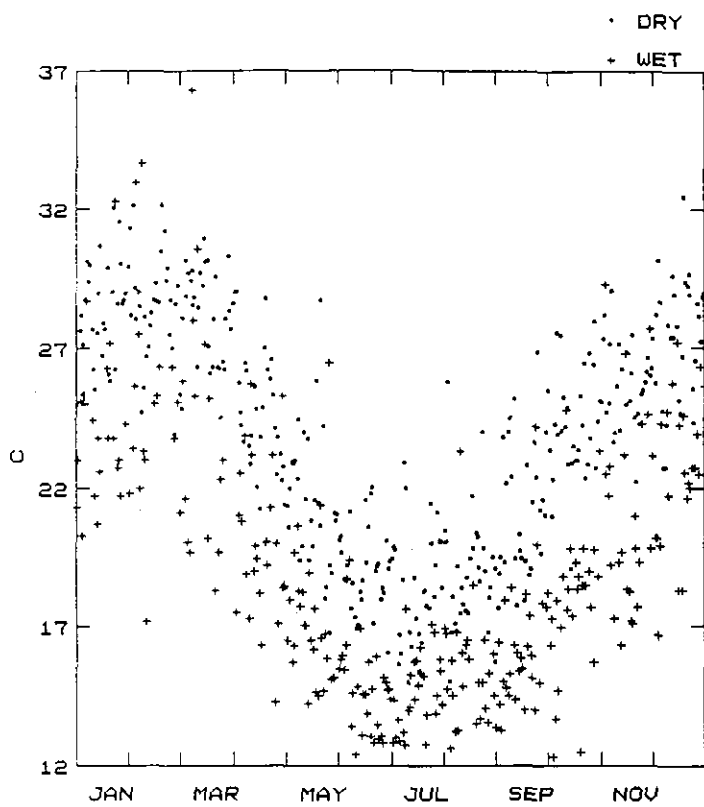
**(e) Rainfall is a “strange” variable**

The rainfall variable is somewhat unusual from a statistical point of view in the sense that it exhibits different properties from those of the other climatic variables. The distribution of rainfall is both discrete and continuous. The occurrence or non-occurrence of rainfall is considered as discrete, while on the times that it does rain, the depth of rainfall has a continuous distribution.

Another distinctive feature of rainfall is that especially in a country like South Africa, the proportion of rainy days is relatively small.

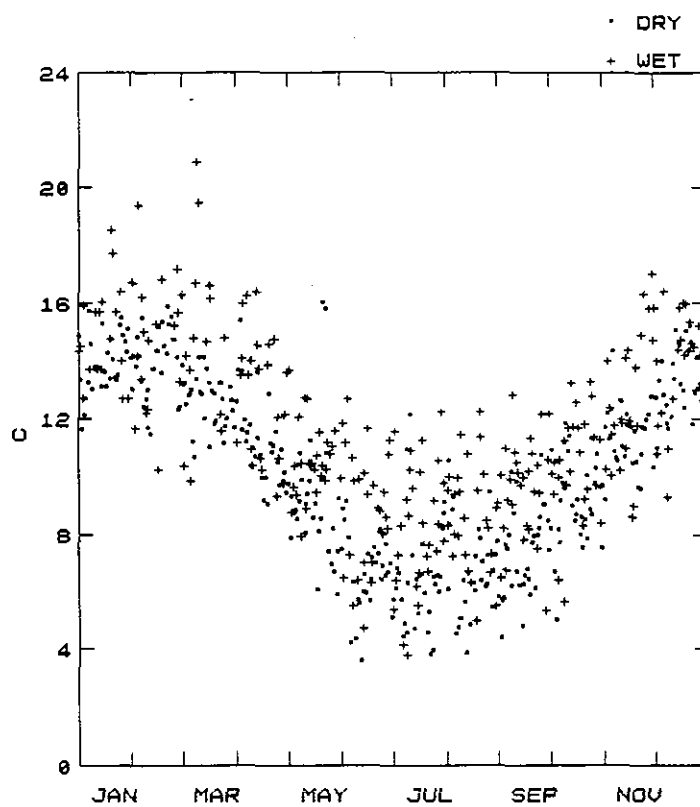
**FIGURE 2.2 Mean time series conditioned on status of day**

MAX TEMP



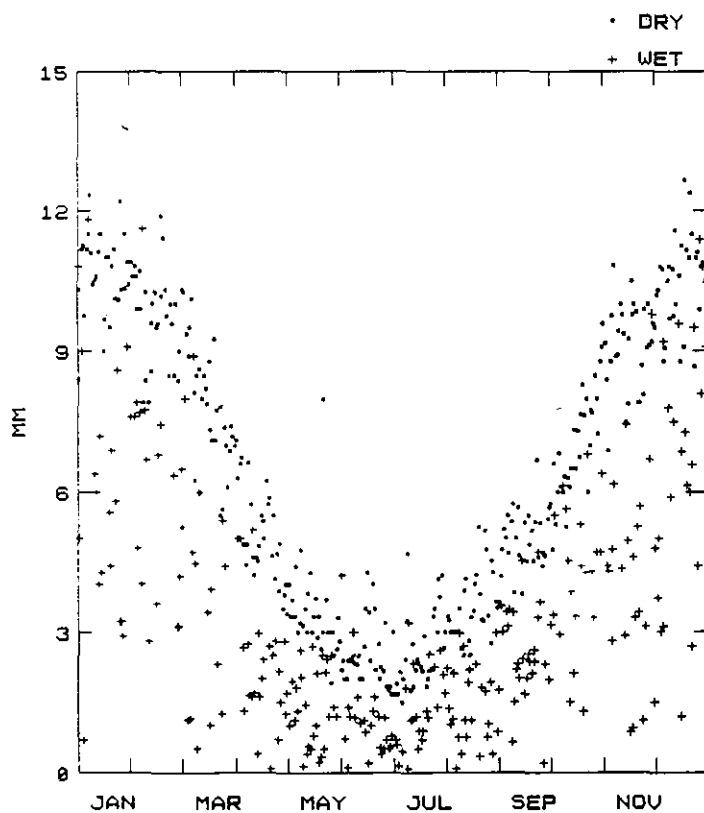
DAYS

MIN TEMP



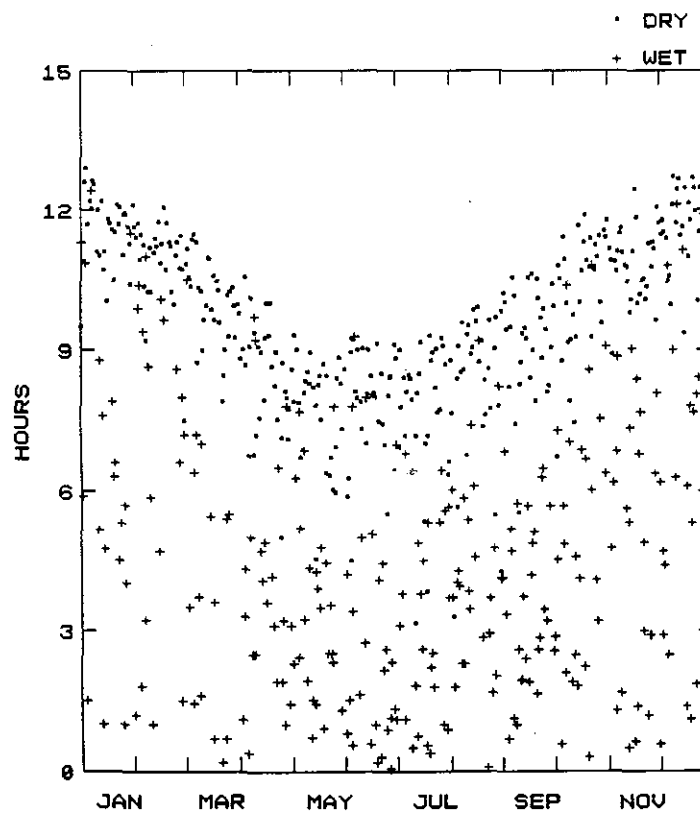
DAYS

EVAPO



DAYS

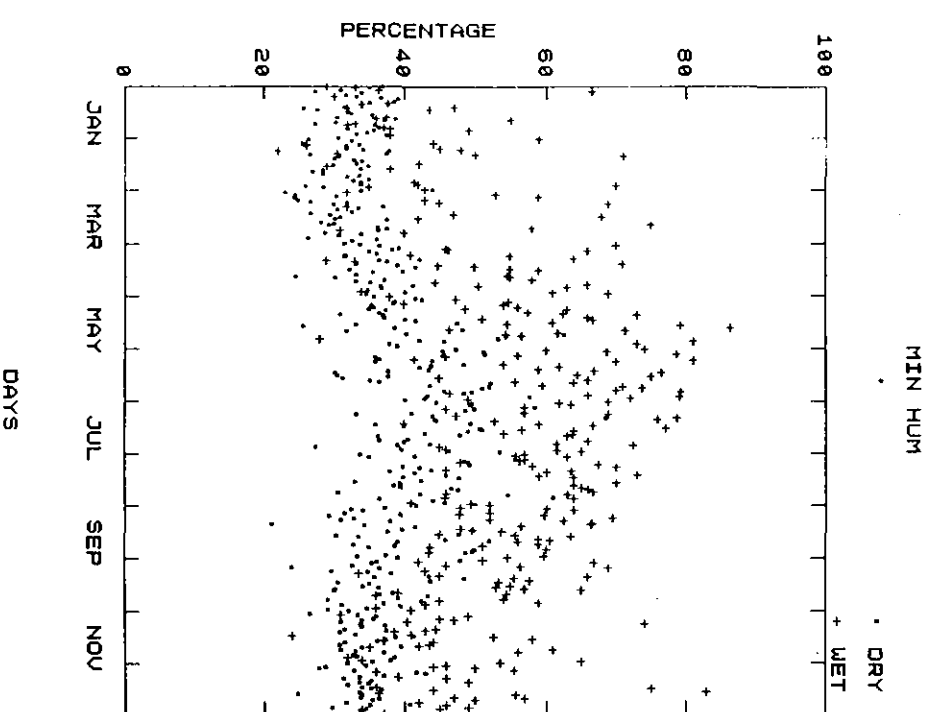
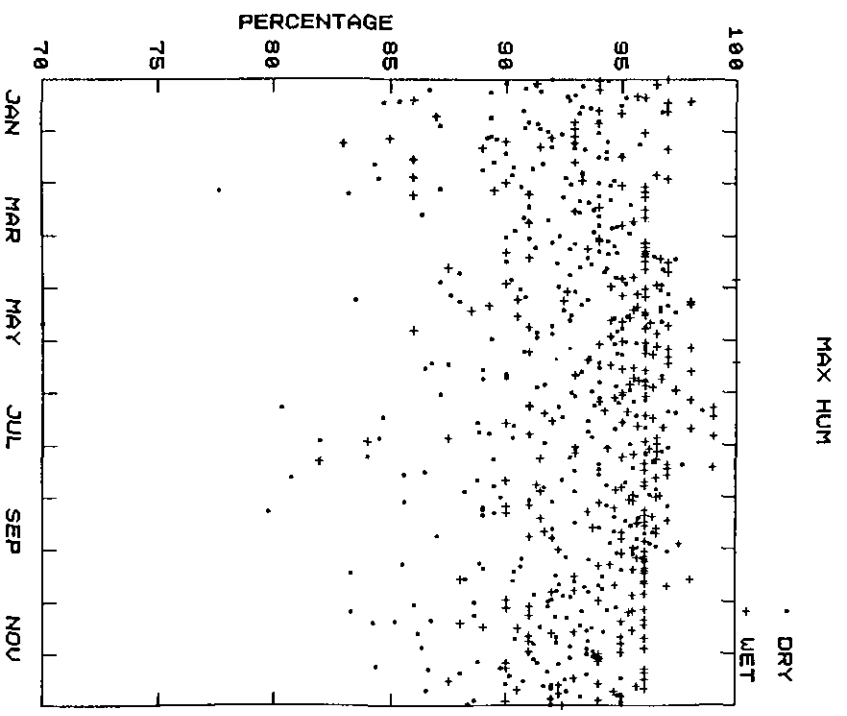
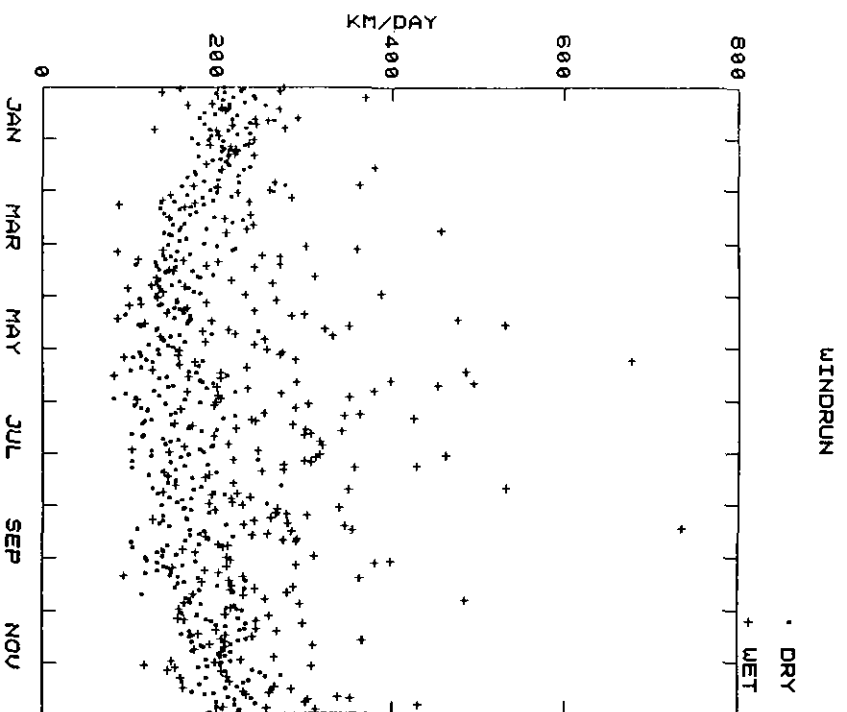
SUNSHINE



DAYS

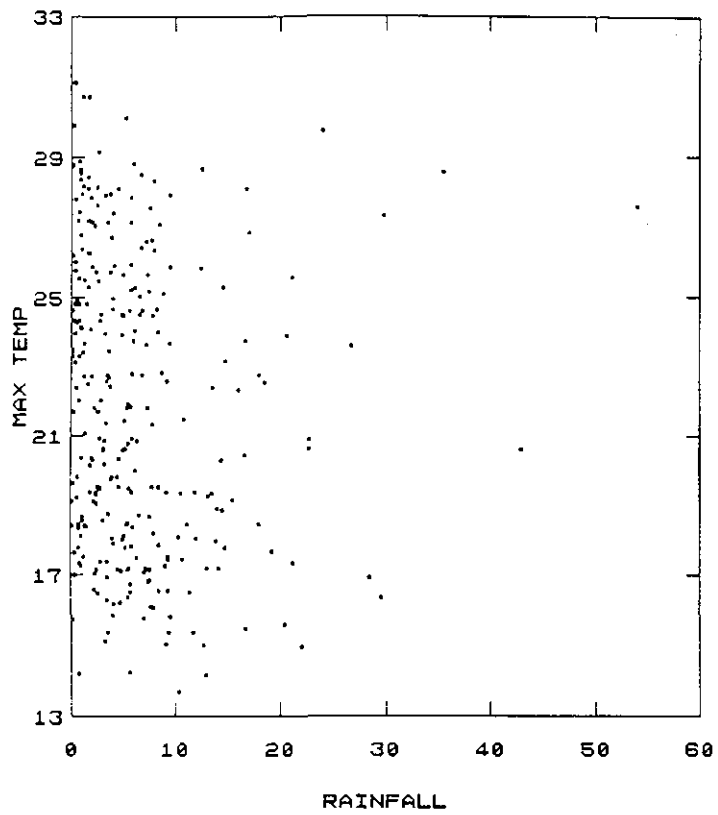
## CHAPTER 2

*The data set and preliminary analysis*

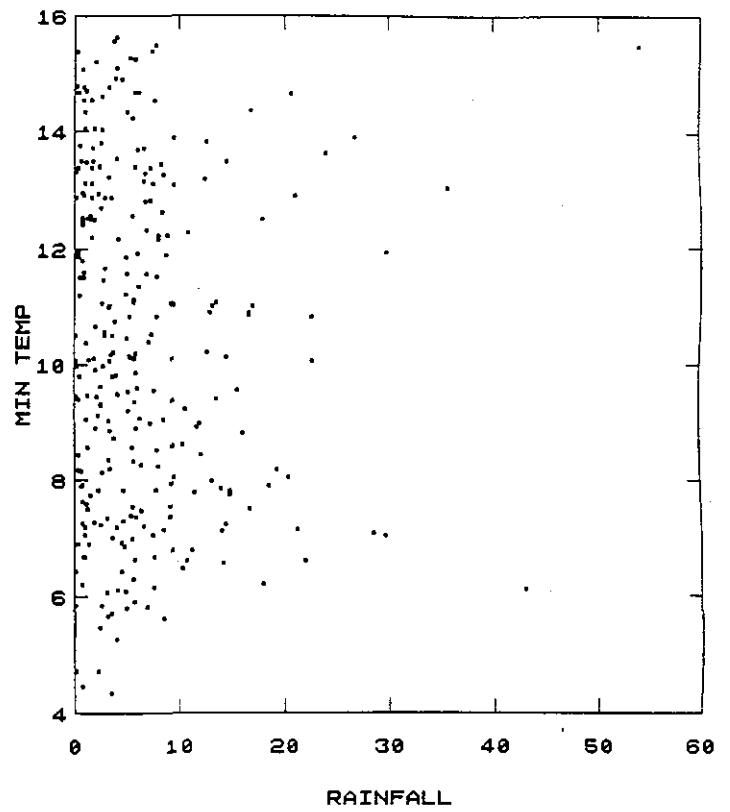


**FIGURE 2.3** Daily mean rainfall versus daily mean climate for all variables

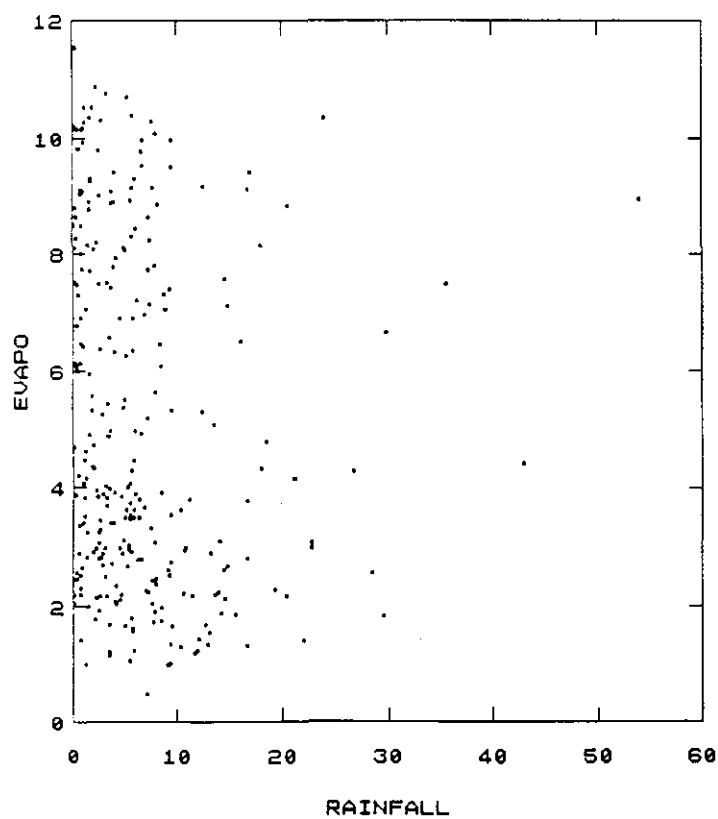
RAIN VS MAX TEMP



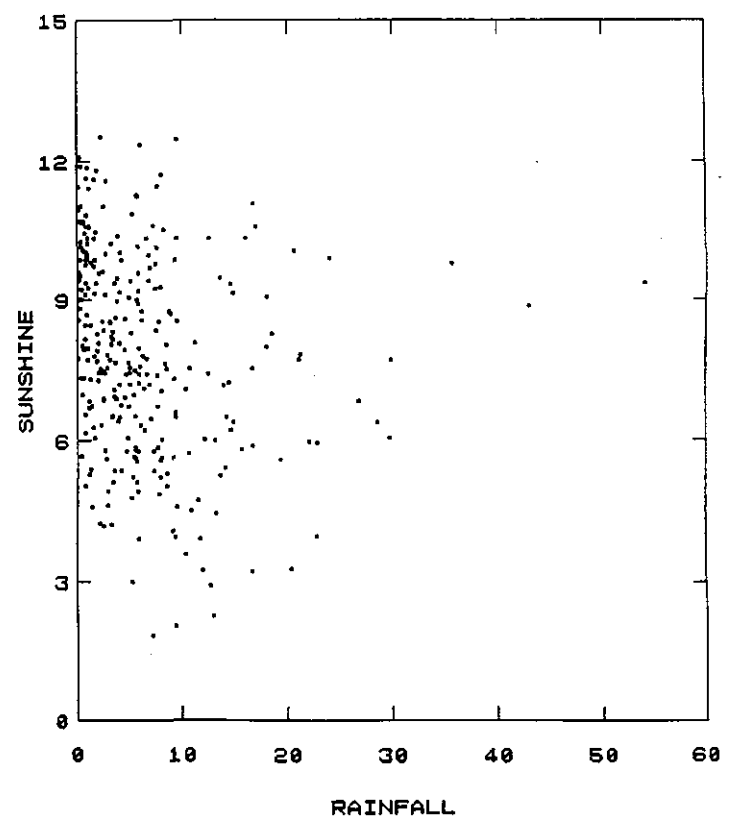
RAIN VS MIN TEMP



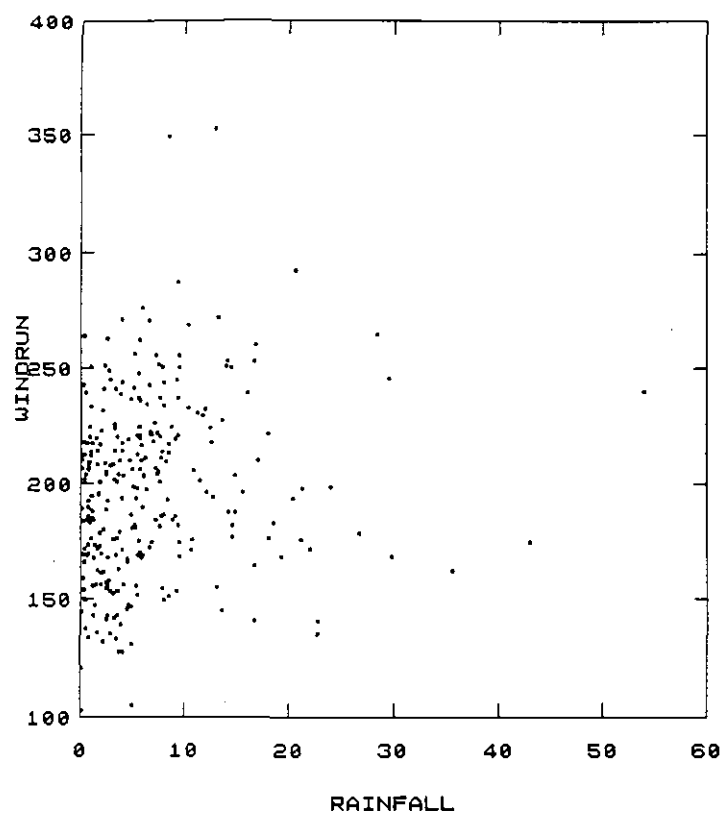
RAIN VS EVAPO



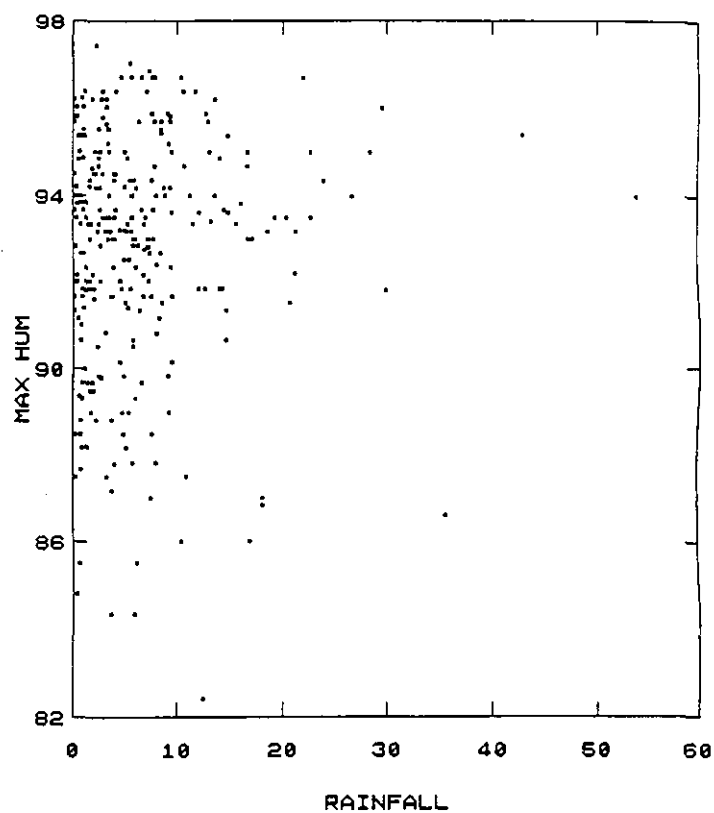
RAIN VS SUN



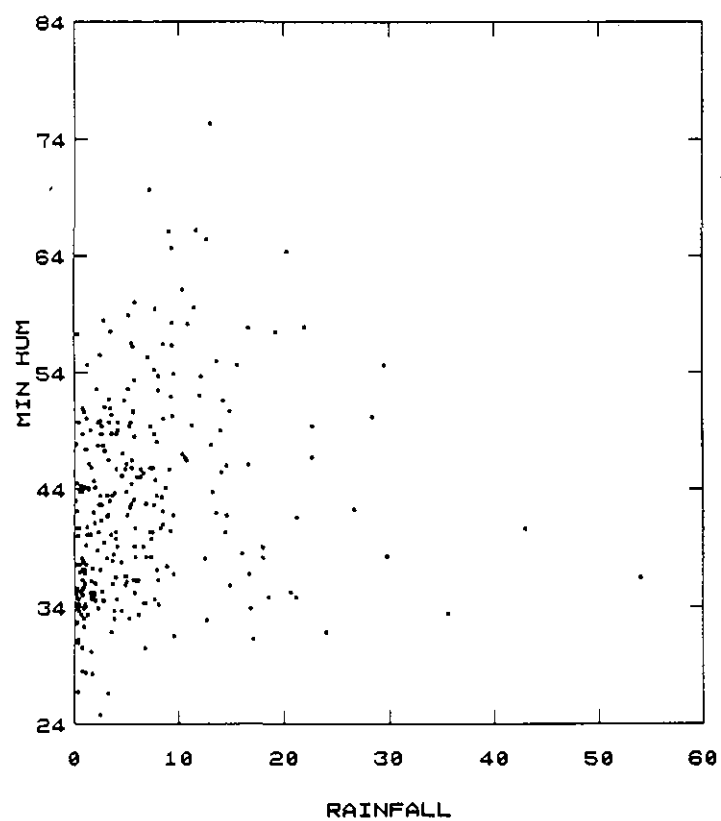
RAIN VS WIND



RAIN VS MAX HUM



RAIN VS MIN HUM



**Concluding remarks**

The above preliminary analysis establishes a number of facts. Firstly, the individual climate variables exhibit seasonal fluctuations and these fluctuations appear to follow a sinusoidal pattern. This would suggest that the mean function of each variable could be parsimoniously modelled by means of a truncated Fourier series. Secondly, the individual variables are serially correlated (even after this seasonal fluctuation has been taken into account). In other words, the individual climate variables constitute time series and have to be modelled as such. This preliminary analysis would suggest that an autoregressive model might be suitable to describe the autocorrelation structure of the variables. Here one has to keep in mind that the number of parameters in the final model must be kept to a minimum in order to avoid the usual statistical problems associated with estimating a large number of parameters. An autoregressive model is ideal in this respect.

Finally, the variables are cross-correlated, that is, they do not vary independently of each other. It follows that it is not possible to model climate by separately modelling its component variables; a multivariate time series model is required.

Seeing that the variable rainfall has some extra properties that have to be taken into consideration and that the remaining climatic variables differ depending on the state of the rainfall variable, it is proposed that the rainfall variable should be determined independently of the other variables and then to condition the other variables for a given day on whether the day was wet or dry.

As no pattern was found between different precipitation amounts and the climate observations, it was decided to consider a non-rainy day as one with a precipitation amount of zero and a rainy day as one with a rainfall depth greater than zero.

## CHAPTER 3

### THE MODELS

The preliminary analysis described in Chapter 2 established that sequences of climate variables exhibit a number of distinctive features. In particular the distribution of each climate variable varies seasonally, the variables are serially correlated, they are dependent, and finally the distributions of the variables depend on the wet or dry status of the day under consideration. Any useful model for the simultaneous description of climate sequences must of course preserve all these properties.

The models considered here are constructed in two stages. One begins by constructing a model for the rainfall process. This provides synthetic sequences of wet and dry days. The remaining variables are then modelled according to the wet or dry status of each day. Thus the joint distribution of all the variables other than rainfall changes not only with season but also with changes in wet or dry status.

The rainfall component of the five models to be discussed is common to all five models and is thus described first. The first of the five models is due to Richardson (1981), the remaining four are new.

#### The rainfall model

Several models have been proposed for simulating daily precipitation. (Gabriel and Neumann, 1962; Richardson, 1981; Roldan and Woolhiser, 1982; Stern and Coe, 1984; Zucchini and Adamson, 1984.) Most precipitation models are specified by a discrete occurrence process describing the sequence of wet and dry days, and a continuous distribution function for the amount of precipitation of days with rain. The parameters of the model are allowed to vary seasonally.

#### A model to describe the occurrence of wet and dry sequences of days

A first-order Markov chain is used to describe the occurrence of wet and dry days. By this one assumes that the state of day  $t$  depends on the state of the previous day,  $t-1$ . This does not imply that the state at time  $t$  is independent of the state on day  $t-2, t-3$ , etc ..., but rather that the information given by  $t-1$  is equivalent to all the information given by  $t-1, t-2$ , etc .... One also assumes that, except for the seasonality, the process is stationary.

A first-order Markov chain has been found to be an adequate model for precipitation occurrence in many different regions. (Gabriel and Neumann, 1962; Caskey, 1963; Weiss, 1964; Hopkins and Robillard, 1964; Haan *et al*, 1976; Smith and Schreiber, 1973; Woolhiser and Prengam, 1979; Richardson, 1981; Roldan and Woolhiser, 1982; Zucchini and Adamson, 1984.) The order of the Markov chain may of course be increased, but this has to be done at the cost of increasing complexity and the number of parameters in the model. A further problem arises if one attempts to increase the order of the Markov chain in arid areas, namely the estimation of the probability that a rain day follows two or more consecutive rain days. In arid areas there are relatively few runs of three or more consecutive rain days and thus there is hardly any data on which to base estimates of this conditional probability. (Note that this has to be estimated for each day of the year.) Finally, it was demonstrated in Zucchini and Adamson (1984) that a first order Markov chain provides an adequate description of the occurrence of wet and dry sequences of days in the complete range of South African conditions.

#### (a) Notation and preliminaries

The day will be used as the time unit. That is, the year is divided into  $NT (= 365)$  equal intervals, denoted by  $t = 1, 2, \dots, NT$ . A day with total rainfall greater than 0 mm is considered as a wet day.

The following notation will be used:

$R$  represents the occurrence of rain (i.e. wet day).

$\bar{R}$  represents the non-occurrence of rain (i.e. dry day).

For  $t = 1, 2, \dots, NT$

$NR(t)$  is the number of times it was wet in period  $t$ .

$N\bar{R}(t)$  is the number of times it was dry in period  $t$ .

$N\bar{R}R(t)$  is the number of times it was dry in period  $t - 1$  and wet in period  $t$ .

$N\bar{R}\bar{R}(t)$  is the number of times it was dry in period  $t - 1$  and dry in period  $t$ .

$NRR(t)$  is the number of times it was wet in period  $t - 1$  and wet in period  $t$ .

$ND(t) = N\bar{R}R(t) + N\bar{R}\bar{R}(t)$  is the number of times that it was dry in period  $t - 1$  and there was an observation (wet or dry) in period  $t$ .

$NW(t) = NRR(t) + N\bar{R}\bar{R}(t)$  is the number of times that it was wet in period  $t - 1$  and

there was an observation (wet or dry) in period  $t$ .

$\pi_{R/R}(t)$  the probability that period  $t$  is wet given that period  $t-1$  is wet.

$\pi_{\bar{R}/R}(t)$  the probability that period  $t$  is dry given that period  $t-1$  is wet.

$\pi_{R/\bar{R}}(t)$  the probability that period  $t$  is wet given that period  $t-1$  is dry.

$\pi_{\bar{R}/\bar{R}}(t)$  the probability that period  $t$  is dry given that period  $t-1$  is dry.

Then  $\pi_{R/R}(t) + \pi_{\bar{R}/R}(t) = 1$

$\pi_{\bar{R}/\bar{R}}(t) + \pi_{R/\bar{R}}(t) = 1$ .

Therefore the transition probabilities are fully defined given  $\pi_{R/R}(t), \pi_{R/\bar{R}}(t)$  and the wet or dry state on day  $t-1$ , and one only needs to estimate these two probabilities.

From elementary probability theory we have

$$NRR(t) \sim B(NW(t), \pi_{R/R}(t))$$

$$N\bar{R}\bar{R}(t) \sim B(ND(t), \pi_{R/\bar{R}}(t)), \quad t = 1, 2, \dots, NT$$

where  $B(N, \pi)$  denotes the binomial distribution with parameters  $N$  and  $\pi$ .

### (b) Estimation

The functions  $\pi_{R/R}(t)$  and  $\pi_{R/\bar{R}}(t)$  are estimated using the same method but different data. To simplify the notation in what follows, one makes use of the following generic names:

$$\text{Let } M(t) \sim B(MM(t), \pi(t)), \quad t = 1, 2, \dots, NT.$$

First we note that the binomial distribution belongs to the exponential family. Therefore we have a set of independent random variables  $M(t), t = 1, 2, \dots, NT$ , each with a distribution from the exponential family; each  $M(t)$  depends on a single parameter  $\pi(t)$  and the distributions of all  $M(t), t = 1, 2, \dots, NT$ , are of the same form (i.e. all binomial). Thus the properties of a generalized linear model are satisfied, and estimates of  $\pi(t)$  may be obtained by using the theory for estimation for generalized linear models. (Dobson, 1983.)

The probabilities  $\pi(t)$  are assumed to be functions of linear combinations of parameters  $\gamma_1, \gamma_2, \dots, \gamma_L, L < NT$ . That is

$$g(\pi(t)) = \lambda(t, L)$$

where  $g$  is the link function and  $\lambda(t, L)$  is a linear combination of the  $\gamma_i$ .

To ensure that the estimated values of  $\pi(t)$  are restricted to the interval  $[0, 1]$ , one uses the logit link function, given by

$$g(\pi(t)) = \log \left( \frac{\pi(t)}{1 - \pi(t)} \right) = \lambda(t, L).$$

To obtain the linear combination of the  $\gamma_i$ ,  $\lambda(t, L)$ , we look at some of the properties of  $\pi(t)$ , namely that it is a smooth, periodic and approximately sinusoidal shaped function. Transforming  $\pi(t)$ , using the logistic transformation, to a logit  $\lambda(t)$  given by

$$\lambda(t) = \log \left( \frac{\pi(t)}{1 - \pi(t)} \right),$$

one obtains a representation which still has the same properties as  $\pi(t)$ , and thus we can approximate  $\lambda(t)$  by the first few terms of its Fourier representation. This approximation has been used by Stern and Coe (1984) and Zucchini and Adamson (1984).

The exact Fourier representation of  $\lambda(t)$  is given by

$$\lambda(t) = \sum_{i=1}^{NT} \gamma_i \varphi_i(t), \quad t = 1, 2, \dots, NT$$

where

$$\varphi_i(t) = \begin{cases} \cos(\omega(t-1)i/2) & i = 2, 4, \dots \\ \sin(\omega(t-1)(i-1)/2) & i = 3, 5, \dots \end{cases}$$

$$\varphi_1(t) = 1; \quad t = 1, 2, \dots, NT,$$

and

$$\omega = \frac{2\pi}{NT}.$$

Define the function  $\lambda(t, L)$  by

$$\lambda(t, L) = \sum_{i=1}^L \gamma_i \varphi_i(t), \quad t = 1, 2, \dots, NT; \quad L \leq NT$$

where  $\varphi_i(t)$  is defined as before and  $L$  is the order of the Fourier series approximation. One is thus making the following approximation:

For some  $L < NT$

$$\lambda(t, L) \approx \lambda(t), \quad t = 1, 2, \dots, NT.$$

A procedure to choose the order of the Fourier series approximation (i.e. the value of  $L$ ) will be discussed later. Generally this approximation is accurate for small values of  $L$ . The number of parameters,  $L$ , is always chosen to be an odd number. The reason for this choice is given in Appendix A. An outline of the properties of the Fourier representation are given in Appendix B.

The log-likelihood function of the observed values as a function of the probabilities  $\pi(t)$ , is given by

$$\ell(\pi(t); M(t)) = \sum_{t=1}^{NT} \left[ M(t) \log \left( \frac{\pi(t)}{1 - \pi(t)} \right) + MM(t) \log(1 - \pi(t)) + \log \left( \frac{MM(t)}{M(t)} \right) \right].$$

Therefore, the log-likelihood function of the observed values as a function of the parameters  $\gamma_1, \gamma_2, \dots, \gamma_L$  is given by

$$\ell(\gamma; M(t)) = \sum_{t=1}^{NT} \left[ M(t) \lambda(t, L) - MM(t) \log(1 + e^{\lambda(t, L)}) + \log \left( \frac{MM(t)}{M(t)} \right) \right].$$

The score vector  $U$  with respect to  $\gamma_1, \gamma_2, \dots, \gamma_L$  has elements given by

$$\begin{aligned} U_j &= \frac{\partial \ell(\gamma; M(t))}{\partial \gamma_j} = \sum_{t=1}^{NT} \left[ M(t) - MM(t) \frac{e^{\lambda(t, L)}}{1 + e^{\lambda(t, L)}} \right] \varphi_j(t) \\ &= \sum_{t=1}^{NT} [M(t) - MM(t) \pi(t)] \varphi_j(t) \end{aligned}$$

since  $\text{Var}(M(t)) = MM(t) \frac{e^{\lambda(t, L)}}{(1 + e^{\lambda(t, L)})^2}$  and

$$\begin{aligned} E(M(t)) &= \frac{MM(t) e^{\lambda(t, L)}}{(1 + e^{\lambda(t, L)})} \quad \text{and so} \\ \frac{\partial E(M(t))}{\partial \lambda(t, L)} &= \frac{MM(t) e^{\lambda(t, L)}}{(1 + e^{\lambda(t, L)})^2} = \text{Var}(M(t)). \end{aligned}$$

Similarly, the information matrix  $\mathbf{I}_{L \times L}$  has elements given by

$$\mathbf{I}_{jk} = \sum_{t=1}^{NT} \varphi_j(t) \varphi_k(t) MM(t) \frac{e^{\lambda(t, L)}}{(1 + e^{\lambda(t, L)})^2}.$$

Since  $\frac{e^{\lambda(t, L)}}{(1 + e^{\lambda(t, L)})^2} = \pi(t)(1 - \pi(t))$  it follows that

$$\mathbf{I}_{jk} = \sum_{t=1}^{NT} \varphi_j(t) \varphi_k(t) MM(t) \pi(t)(1 - \pi(t)).$$

The maximum likelihood estimates for  $\gamma_1, \gamma_2, \dots, \gamma_L$  are then obtained by solving the iterative equation

$$\mathbf{I}^{(m-1)}\hat{\gamma}^{(m)} = \mathbf{I}^{(m-1)}\hat{\gamma}^{(m-1)} + U^{(m-1)}$$

where  $m$  indicates the  $m$ th approximation and  $\hat{\gamma}$  is the vector of estimates.

Some initial approximation  $\gamma^{(0)}$  is used to evaluate  $\mathbf{I}^{(0)}$  and  $U^{(0)}$ , then the iterative equation is solved to give  $\gamma^{(1)}$  which in turn is used to obtain better approximations for  $\mathbf{I}$  and  $U$ , and so on until adequate convergence is achieved. When the difference between successive approximations  $\gamma^{(m)}$  and  $\gamma^{(m-1)}$  is sufficiently small,  $\gamma^{(m)}$  is taken as the maximum likelihood estimate vector.

### (c) Model selection

Whenever a model is fitted to observed data, two types of discrepancy arise. The discrepancy due to approximation (the fewer the number of parameters fitted, the higher the value of this discrepancy) and the discrepancy due to estimation (the more parameters fitted, the higher the value of this discrepancy). When choosing the number of parameters to be fitted, one attempts to minimize the combined effect arising from the two discrepancies.

Selection of the number of parameters,  $L$ , may be done by using the criterion of the Kullbach–Leibler measure of discrepancy. (Linhart and Zucchini, 1982; Zucchini and Adamson, 1984.)

Under the assumption that for some  $L_0$ ,  $\lambda(t)$  is exactly fitted by  $L_0$  parameters, i.e.

$$\lambda(t) = \lambda(t, L_0), \quad L_0 < NT,$$

the above method leads to the Akaike Information Criterion where

$$AIC = -\ell(\gamma; M(t)) + L$$

where  $\ell(\gamma; M(t))$  is the log-likelihood function given before.

Each value of  $L$  leads to a different approximating model. The criterion is computed for  $L = 1, 3, 5, \dots$  and the model which leads to the smallest value of the criterion is selected.

The AIC criterion is much easier to compute than the full Kullbach–Leibler discrepancy and leads to almost identical results if the discrepancy due to approximation is small (which it is in this application).

### The distribution of rainfall on days when rain occurs

Several models have been proposed for the distribution of precipitation amounts given the occurrence of a wet day. These include the exponential (Todorovic and Woolhiser, 1975; Richardson, 1981); gamma (Ison *et al.*, 1981; Buishand, 1977; Stern and Coe, 1984); two-parameter gamma (Buishand, 1978); three-parameter mixed exponential (Woolhiser and Pegram, 1979); kappa (Mielke, 1973); lognormal and Weibull (Zucchini and Adamson, 1984).

Woolhiser and Roldan (1982b) found that out of the exponential, gamma and mixed exponential distributions, the latter fitted the model of precipitation amounts best. Zucchini and Adamson (1984) found that for stations in South Africa, the lognormal distribution did not fit some stations, while the Weibull seemed to provide better fits.

It is known that the distribution of precipitation depths when rain occurs is positively skewed (i.e. smaller amounts occurring more frequently than the larger amounts) and that it exhibits the same seasonal variability as found with the probabilities  $\pi(t)$ . To account for this seasonality, the simplest solution is to fit a family of distributions and then to allow the parameters to change over the year, where these parameters are expressed in terms of its Fourier series approximation.

The method of modelling precipitation amounts is adopted from Zucchini and Adamson (1984). Here one does not fit any model initially, the first two moment functions of the distribution are fitted instead. These are then used to estimate the parameters (by the method of moments) to any desired two-parameter model. Different families can be fitted to a single record, e.g. one for the rainy season and a second for the dry season.

#### (a) Notation

The year is divided into  $NT$  equal intervals denoted by  $t = 1, 2, \dots, NT$ .

$M(t)$  represents the number of times that it rained in period  $t$ .

$R(i, t)$  represents the rainfall depth on the  $i$ th year that it rained in period  $t$ , where  $i = 1, 2, \dots, M(t)$ .

$C$  represents the coefficient of variation which we assume to be constant for all  $t$  (Zucchini and Adamson, 1984).

$\mu(t)$  represents the mean rainfall per rainy day in period  $t = 1, 2, \dots, NT$ .

**(b) Estimating the mean and coefficient of variation**

As observed before  $\mu(t)$  can be approximated by its truncated Fourier series representation and thus reducing the number of parameters to be estimated. That is, we make the approximation:

$$\mu(t, L) \approx \mu(t), \quad t = 1, 2, \dots, NT; \quad L < NT$$

where  $\mu(t)$  is defined as

$$\mu(t) = \sum_{i=1}^{NT} \mu_i \varphi_i(t) \quad t = 1, 2, \dots, NT$$

and

$$\mu(t, L) = \sum_{i=1}^L \mu_i \varphi_i(t) \quad t = 1, 2, \dots, NT; \quad L \leq NT$$

and  $\varphi_i(t)$  is defined as before.

Define  $m(t)$  to be the observed means for each period, i.e.

$$m(t) = \frac{1}{M(t)} \sum_{i=1}^{M(t)} R(i, t), \quad t = 1, 2, \dots, NT; \quad i = 1, 2, \dots, M(t); \quad M(t) > 0$$

where  $m(t)$  is not defined when  $M(t) = 0$ , i.e. it never rained in period  $t$ .

We use the method of least squares on  $m(t)$  to estimate  $\mu_1, \mu_2, \dots, \mu_L$ , that is, minimize

$$\sum_{t=1}^{NT} (m(t) - \mu(t, L))^2 \quad (1)$$

with respect to the  $\mu_i$ ,  $i = 1, 2, \dots, L$ . Approximations to the least squares estimators when some of the  $M(t) = 0$ , something which occurs often in arid regions, are given by

$$\hat{\mu}_i = K(i) \sum_{\substack{t=1 \\ M(t) > 0}}^{NT} m(t) \varphi_i(t) \quad (2)$$

where

$$K(i) = \sum_{\substack{t=1 \\ M(t) > 0}}^{NT} \varphi_i(t)^2 \quad i = 1, 2, \dots, L.$$

The  $m(t)$  in (1) are given the same weight and so periods which had very little rainfall have a large influence in the estimates of  $\mu(t)$ . To overcome this difficulty, the following criterion is used instead:

Minimize

$$S(\mu) = \sum_{t=1}^{NT} \sum_{i=1}^{M(t)} (R(i, t) - \mu(t, L))^2 \quad (3)$$

with respect to  $\mu_i$ ,  $i = 1, 2, \dots, L$ .

By adding and subtracting  $m(t)$  inside the squared term of (3),  $S(\mu)$  can be rewritten as

$$S(\mu) = S + \sum_{t=1}^{NT} M(t)(m(t) - \mu(t, L))^2 \quad (4)$$

where

$$S = \sum_{t=1}^{NT} \sum_{i=1}^{M(t)} (R(i, t) - m(t))^2$$

and  $m(t)$  is defined as before if  $M(t) \neq 0$  and  $m(t) = 0$  if  $M(t) = 0$ .

To minimize (4) set its partial derivatives equal to zero:

$$\frac{\partial S(\mu)}{\partial \mu_i} = -2 \sum_{t=1}^{NT} M(t)(m(t) - \mu(t, L))\varphi_i(t), \quad i = 1, 2, \dots, L.$$

These  $L$  equations can be solved using the Newton-Raphson iteration method. For this, we need the second partial derivatives:

$$\frac{\partial^2 S(\mu)}{\partial \mu_i \partial \mu_j} = 2 \sum_{t=1}^{NT} M(t)\varphi_i(t)\varphi_j(t), \quad i, j = 1, 2, \dots, L.$$

Denote the  $i$ th element of the vector  $f^{(k)}$  by

$$f_i^{(k)} = \sum_{t=1}^{NT} M(t)(m(t) - \mu^{(k)}(t, L))\varphi_i(t), \quad i = 1, 2, \dots, L \quad (5)$$

and the  $(i, j)$ th element of the matrix  $F^{(k)}$  by

$$F_{ij}^{(k)} = \sum_{t=1}^{NT} M(t)\varphi_i(t)\varphi_j(t), \quad i, j = 1, 2, \dots, L \quad (6)$$

where  $k$  denotes the  $k$ th iteration.

Then an algorithm to estimate  $\mu_i$ ,  $i = 1, 2, \dots, L$  is given by:

**Step 1:** Obtain initial estimates  $\mu_1^{(0)}, \dots, \mu_L^{(0)}$  using (2) and compute  $\mu^{(0)}(t, L)$ .

**Step 2:** Compute  $f^{(k)}$  using (5) and  $F^{(k)}$  using (6).

**Step 3:** Compute the vector  $\delta^{(k)}$  which is the solution to the system of  $L$  linear equations given by

$$F^{(k)}\delta^{(k)} = f^{(k)}$$

**Step 4:** Set  $\mu^{(k+1)} = \mu^{(k)} - \delta^{(k)}$ .

**Step 5:** Test for convergence, e.g. if the elements of  $f^{(k)}$  are sufficiently close to zero. If the convergence criterion is met, stop, otherwise increase  $k$  by 1 and go to Step 2.

Note that  $F^{(k)}$  is symmetric. This fact can be used to reduce the number of computations performed.

An estimator of  $C$  is given by:

$$\hat{C} = \left[ \frac{\left[ \sum_{t=1}^{NT} \sum_{i=1}^{M(t)} (R(i,t) - \hat{\mu}(t))^2 \right]}{\left[ \sum_{t=1}^{NT} M(t) \hat{\mu}(t)^2 \right]} \right]^{\frac{1}{2}}.$$

(c) **Selecting the number of parameters**

$$\Delta(L) = \sum_{t=1}^{NT} (\mu(t) - E(\hat{\mu}(t, L)))^2, \quad L = 1, 3, 5, \dots$$

would be a suitable discrepancy on which to base the selection, except that some  $M(t)$  are zero and so only approximately unbiased estimators are available. The reliability of this criterion is therefore difficult to determine.

If one is prepared to make distributional assumptions, then selection criteria are relatively easy to derive, for example based on the Kullback–Leibler discrepancy.

A reasonable procedure is to select  $L$  for a parametric family of models and then use the same  $L$  in the estimation of  $\mu(t)$ .

(d) **Fitting the Weibull family**

Zucchini and Adamson (1984) found the Weibull family to fit the rainfall depth models for stations in South Africa and so this family was used to model the observed rainfall amounts on days that rain was recorded.

Having estimated the mean value function  $\mu(t)$  and the coefficient of variation,  $C$ , one can apply the method of moments to estimate the parameter functions of the Weibull distribution.

Denote the scale parameter by  $\alpha(t)$ ,  $t = 1, 2, \dots, NT$  and the shape parameter by  $\beta$ .

Now

$$C = \left\{ \frac{\Gamma(1 + 2/\beta)}{\Gamma(1 + 1/\beta)^2} - 1 \right\}^{\frac{1}{2}}$$

To obtain  $\beta$  as a function of  $C$  a rational function approximation has to be derived as no closed expression of this function is available.

The following approximation has been obtained from Zucchini and Adamson (1984):

$$\hat{\beta} = \frac{339.5410 + 148.445\hat{C} + 192.7492\hat{C}^2 + 22.4401\hat{C}^3}{1 + 257.1162\hat{C} + 287.8362\hat{C}^2 + 157.2230\hat{C}^3}.$$

Using the relationship

$$\mu(t) = \alpha(t)\Gamma(1 + 1/\beta) \quad t = 1, 2, \dots, NT$$

we obtain the estimator

$$\hat{\alpha}(t) = \frac{\hat{\mu}(t)}{\Gamma(1 + 1/\hat{\beta})} \quad t = 1, 2, \dots, NT.$$

## MODEL FOR CLIMATE SEQUENCES

Little attention has been given to stochastic modelling of climatic variables such as maximum and minimum temperature, evaporation, sunshine duration, windrun, and maximum and minimum humidity. Recently, though, there have been some models proposed to stochastically simulate possible sequences of maximum and minimum temperature and solar radiation. (Goh and Tan, 1977; Nicks and Harp, 1980; Richardson, 1981; Larsen and Pense, 1982.) Bruhn *et al* (1980) looked at minimum relative humidity as well.

Variables such as temperature, evaporation, sunshine duration, windrun and humidity are not as difficult to model statistically as precipitation because there is not a high proportion of zero observations and the distributions of these variables are not as skewed as the rainfall distribution.

In the models that follow, because the cross-correlation between the variables is non zero, the variables are considered to reflect a continuous multivariate stochastic process with the parameters conditioned on the wet or dry status of the day.

**Model 1: Multivariate model for climate data proposed by Richardson (1981)**

The approach taken here to model the climate variables is the method suggested by Richardson (1981). The weather variables evaporation, windrun, maximum and minimum humidity have been added to the multivariate process.

**(a) Notation**

Partition the year into  $NT(= 365)$  equal intervals, denoted by  $t = 1, 2, \dots, NT$ .

$NV$  is the number of variables.

$NY$  is the number of years observed.

$W$  represents the occurrence of rain.

$D$  represents the non-occurrence of rain.

$Y_{i,t}$  is the precipitation amount on period  $t$  of year  $i$ ,  $i = 1, 2, \dots, NY$ .

$S_{i,t}$  is the generic name for the observation at period  $t$  of the  $i$ th year.

$\mu_t^D$  is the generic name for the mean for a dry day on period  $t$  (i.e.  $Y_{i,t} = 0$ ).

$\mu_t^W$  is the generic name for the mean for a wet day on period  $t$  (i.e.  $Y_{i,t} > 0$ ).

$\sigma_t^D$  is the generic name for the standard deviation for a dry day on period  $t$ .

$\sigma_t^W$  is the generic name for the standard deviation for a wet day on period  $t$ .

$\chi_{i,t}$  is the generic name for the residual component at period  $t$  and year  $i$ .

$\rho_0(j, k)$  is the lag 0 cross-correlation coefficient between variables  $j$  and  $k$ .

$\rho_1(j, k)$  is the lag 1 cross-correlation coefficient between variables  $j$  and  $k$ .

$\rho_1(j)$  is the lag 1 serial correlation for variable  $j$ .

**(b) The model and assumptions**

Each variable is modelled in the same way. The procedure given below to model  $S_{i,t}$  is carried out once for each variable to be included in the multivariate model.

The distribution of  $S_{i,t}$  is seasonal and so its parameters, e.g. the mean and standard deviation, are allowed to vary seasonally. As was the case with the parameter functions of the rainfall model, it can be reasonably assumed that the parameter functions of the climate variables are smooth, periodic and sinusoidal in shape. This would again lead one to expect that they can be accurately approximated by the first few terms of their Fourier

representation.

The truncated Fourier representations for the daily means and standard deviations for wet days and for dry days are given by:

$$\left. \begin{aligned} \mu_t^W &= \sum_{i=1}^L \alpha_i^W \varphi_i(t) \\ \sigma_t^W &= \sum_{i=1}^L \xi_i^W \varphi_i(t) \end{aligned} \right\} \quad \text{if } Y_{i,t} > 0$$

$$\left. \begin{aligned} \mu_t^D &= \sum_{i=1}^L \alpha_i^D \varphi_i(t) \\ \sigma_t^D &= \sum_{i=1}^L \xi_i^D \varphi_i(t) \end{aligned} \right\} \quad \text{if } Y_{i,t} = 0$$

$$t = 1, 2, \dots, NT$$

where

$$\varphi_i(t) = \begin{cases} \cos(\omega(t-1)i/2), & i = 2, 4, \dots, L-1 \\ \sin(\omega(t-1)(i-1)/2), & i = 3, 5, \dots, L \end{cases}$$

$$\varphi_1(t) = 1,$$

$$\omega = 2\pi/NT \quad \text{and}$$

$\alpha_i^W, \alpha_i^D, \xi_i^W, \xi_i^D$  are the coefficients of the respective Fourier series and,  $L$  is the order of the Fourier series approximation, i.e. we assume that for some  $L < NT$

$$\mu_t = \sum_{i=1}^L \alpha_i \varphi_i(t) \approx \sum_{i=1}^{NT} \alpha_i \varphi_i(t)$$

and

$$\sigma_t = \sum_{i=1}^L \xi_i \varphi_i(t) \approx \sum_{i=1}^{NT} \xi_i \varphi_i(t)$$

where the above two equations hold for both wet and dry days. (Whenever  $W$  or  $D$  is omitted it means that the equation applies for both.)

The number of parameters,  $L$ , does not have to be the same in all instances, i.e. the number of parameters for the means of wet days can differ from that for dry days. The same applies for the standard deviations. To avoid complicating the notation, it will be assumed in what follows that  $L$  refers to the number of parameters of the particular parameter function under consideration.

The estimation of the Fourier coefficients will be discussed later.

The approach used by Richardson (1981) is to determine the daily means and standard deviations of each variable conditioned on the wet or dry status of each day where Fourier series is used to smooth their seasonality. The time series  $S_{i,t}$  is then reduced to a time series of residual elements by removing the periodic means and standard deviations. This residual time series is given by the equations:

$$\chi_{i,t} = \begin{cases} \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} & \text{if } Y_{i,t} = 0 \\ \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} & \text{if } Y_{i,t} > 0. \end{cases}$$

This standardization leads to a residual series for each variable that is stationary in the mean and standard deviation with mean zero and standard deviation of unity.

The serial correlation and cross-correlation coefficients are then calculated to describe the time dependence and the interdependence (respectively) of the residual series.

The model proposed for generating residual series for each variable is the weakly stationary process suggested by Matalas (1967) given by

$$\chi_{i,t} = A \chi_{i,t-1} + B \epsilon_{i,t} \quad (7)$$

where  $\epsilon_{i,t}$  is a  $(NV \times 1)$  matrix of independent random components that are normally distributed with mean zero and a variance of unity, i.e.

$$\epsilon_{i,t} \sim NID(0, 1).$$

$A$  and  $B$  are  $(NV \times NV)$  matrices whose elements are defined in such a way that the sequence generated will have the desired serial correlation and cross-correlation coefficients.

This model is based on the assumption that the residuals of the variables are normally distributed and that the serial correlation of each variable may be described by a first-order linear autoregressive model.

### (c) Estimation

Firstly, a method for estimating the matrices  $A$  and  $B$  will be considered.

From the properties of the distribution of  $\epsilon_{i,t}$  and  $\chi_{i,t}$  we have that

$$E(\epsilon_{i,t}) = 0$$

and

$$E(\chi_{i,t}) = E(\chi_{i,t-1}) = 0.$$

Postmultiplying (7) through by  $\chi_{i,t-1}^T$ , the transpose of  $\chi_{i,t-1}$ , and taking expectations we have

$$E[\chi_{i,t}\chi_{i,t-1}^T] = AE[\chi_{i,t-1}\chi_{i,t-1}^T] + BE[\epsilon_{i,t}\chi_{i,t-1}^T] \quad (8)$$

Define

$$M_0 = E[\chi_{i,t-1}\chi_{i,t-1}^T]$$

and

$$M_1 = E[\chi_{i,t}\chi_{i,t-1}^T].$$

$M_0$  is an  $(NV \times NV)$  matrix whose elements are the lag 0 cross-correlation coefficients and  $M_1$  is an  $(NV \times NV)$  matrix whose elements are the lag 1 cross-correlation coefficients.

The matrices may be written as

$$M_0 = \begin{pmatrix} 1 & \rho_0(1,2) & \dots & \rho_0(1,NV) \\ \rho_0(2,1) & 1 & \dots & \rho_0(2,NV) \\ \vdots & & \dots & \vdots \\ \rho_0(NV,1) & \dots & \dots & 1 \end{pmatrix}$$

and

$$M_1 = \begin{pmatrix} \rho_1(1) & \rho_1(1,2) & \dots & \rho_1(1,NV) \\ \rho_1(2,1) & \rho_1(2) & \dots & \rho_1(2,NV) \\ \vdots & & \dots & \vdots \\ \rho_1(NV,1) & \rho_1(NV,2) & \dots & \rho_1(NV) \end{pmatrix}$$

where  $\rho_0(j,k)$  is the lag 0 cross-correlation coefficient between variables  $j$  and  $k$ ,  $\rho_1(j,k)$  is the cross-correlation coefficient between variables  $j$  and  $k$  with variable  $k$  lagged 1 day in relation to variable  $j$ , and  $\rho_1(j)$  is the lag 1 serial correlation for variable  $j$ . We can thus rewrite (8) as

$$M_1 = AM_0 \quad \text{since} \quad E[\epsilon_{i,t}\epsilon_{i,t-1}^T] = 0.$$

Since  $M_0$  is a variance covariance matrix, it is non-singular, and therefore its inverse exists.

The matrix  $A$  is given by

$$A = M_1M_0^{-1}.$$

Postmultiplying (7) through by  $\chi_{i,t}^T$  and taking expectations one gets

$$M_0 = AM_1^T + BB^T$$

since  $E[\epsilon_{i,t}\epsilon_{i,t}^T] = I$ , the identity matrix.

Therefore, the matrix  $B$  is given by the solution to

$$BB^T = M_0 - M_1M_0^{-1}M_1^T.$$

The Cholesky decomposition (Appendix C) can be used to solve for  $B$ .

Now, we will discuss the method to obtain parameter estimates for the coefficients of the truncated Fourier series.

The functions  $\mu_t$  and  $\sigma_t$  are estimated using the same method but different data sets. The theory will thus be discussed for the mean function  $\mu_t$  only.

Let  $\bar{S}_t$  be the daily mean vector for  $S_{i,t}$  and assume that it is given by the linear model

$$\bar{S}_t = x_t^T \beta + e_t, \quad t = 1, 2, \dots, NT$$

with

$$e_t \sim NID(0, \sigma_t^2).$$

This is a special case of a generalized linear model because the elements  $\bar{S}_t$  are independent with distributions  $N(\mu_t, \sigma_t^2)$  where

$$\mu_t = x_t^T \beta.$$

Also the normal distribution is a member of the exponential family (provided the  $\sigma_t^2$  are regarded as known). In this case the link function,  $g$ , is the identity function, i.e.

$$g(\mu_t) = \mu_t = \sum_{i=1}^L \alpha_i \varphi_i(t) = \eta_t$$

where  $\sum_{i=1}^L \alpha_i \varphi_i(t)$  represents the truncated Fourier series of the mean function  $\mu_t$ , and  $\varphi_i(t)$  is defined as before.

The log-likelihood function of the observed mean values as a function of the mean function  $\mu_t$  is given by:

$$\ell(\mu_t; \bar{S}_t) = \left[ \sum_{t=1}^{NT} -\frac{\bar{S}_t^2}{2\sigma_t^2} + \frac{\bar{S}_t \mu_t}{\sigma_t^2} - \frac{\mu_t^2}{2\sigma_t^2} - \frac{1}{2} \log(2\pi\sigma_t^2) \right].$$

Therefore, the log-likelihood function of the observed values as a function of the parameters  $\alpha_1, \alpha_2, \dots, \alpha_L$  is given by

$$\ell(\alpha; \bar{S}_t) = \sum_{t=1}^{NT} \left[ -\frac{\bar{S}_t^2}{2\sigma_t^2} + \frac{\bar{S}_t \sum_{i=1}^L \alpha_i \varphi_i(t)}{\sigma_t^2} - \frac{\sum_{i=1}^L \alpha_i \varphi_i(t)}{2\sigma_t^2} - \frac{1}{2} \log(2\pi\sigma_t^2) \right]$$

The score vector  $U$  with respect to  $\alpha_1, \alpha_2, \dots, \alpha_L$  has elements given by

$$U_j = \frac{\partial \ell(\alpha; \bar{S}_t)}{\partial \alpha_j} = \sum_{t=1}^{NT} \left[ \frac{(\bar{S}_t - \mu_t)}{\sigma_t^2} \varphi_j(t) 1 \right]$$

since

$$\begin{aligned} E(\bar{S}_t) &= \mu_t, \\ \text{Var}(\bar{S}_t) &= \sigma_t^2, \quad \text{and} \\ \partial \mu_t / \partial \eta_t &= 1. \end{aligned}$$

Similarly, the information matrix  $\mathbf{I}_{L \times L}$  has elements given by

$$\mathbf{I}_{jk} = \sum_{t=1}^{NT} \frac{\varphi_j(t) \varphi_k(t)}{\sigma_t^2} 1^2.$$

The maximum likelihood estimates for  $\alpha_1, \alpha_2, \dots, \alpha_L$  are then obtained by solving the iterative equation

$$\mathbf{I}^{(m-1)} \hat{\alpha}^{(m)} = \mathbf{I}^{(m-1)} \hat{\alpha}^{(m-1)} + U^{(m-1)}$$

where  $m$  indicates the  $m$ th approximation and  $\hat{\alpha}$  is the vector of estimates.

When the difference between successive approximations  $\hat{\alpha}^{(m)}$  and  $\hat{\alpha}^{(m-1)}$  is sufficiently small,  $\hat{\alpha}^{(m)}$  is taken as the maximum likelihood estimate vector.

#### (d) Model selection

The order of the Fourier series approximation,  $L$ , for the conditioned mean function and for the conditioned standard deviation function is selected by Akaike Information Criterion, AIC, where

$$\text{AIC} = -\ell(\alpha; \theta) + L$$

where  $\ell(\alpha; \theta)$  is the log-likelihood function of the model. Each value of  $L$  leads to a different approximating model. The criterion is computed for  $\alpha = 1, 3, 5, \dots$  and the model which leads to the smallest value of the criterion is selected.

**Model 2: Multivariate model for climate data**

Although Model 1 appeared to be satisfactory in many respects, it performed poorly in some respects. In particular the annual standard deviation for windrun, maximum and minimum humidity were systematically underestimated. The (lagged) cross-correlations between some of the variables (e.g. maximum temperature and minimum temperature) were not preserved by the model. However the most noticeable deficiency was found to be that the model did not preserve the serial correlation structure of many of the variables. This was attributed to the lack of flexibility of Model 1 in this respect. In particular the model is based on the assumption that the serial correlation function does not depend on the wet/dry status of the days in question. In fact the correlation between variables on two successive days depends on whether the two days are both wet, both dry, wet followed by dry or dry followed by wet. It was therefore decided to develop a model which incorporates additional flexibility in its autocorrelation function, that is, a model which allows for the serial correlations between variables on successive days to depend on their wet/dry status.

Model 2 was developed as a prototype to Models 3, 4 and 5. It attempts to deal with the mentioned deficiency in Model 1.

**(a) Notation**

Partition the year into  $NT (= 365)$  equal intervals, denoted by  $t = 1, 2, \dots, NT$ .

$NV$  is the number of variables.

$NY$  is the number of years observed.

$W$  represents the occurrence of rain.

$D$  represents the non-occurrence of rain.

$N(D)$  is the set of time periods  $t$  such that period  $t$  was dry.

$N(W)$  is the set of time periods  $t$  such that period  $t$  was wet.

$Y_{i,t}$  is the precipitation amount on period  $t$  of year  $i$ ,  $i = 1, 2, \dots, NY$ .

$S_{i,t}$  is the generic name for the observation at time  $t$  of the  $i$ th year.

$\mu_t^D$  is the generic name for the mean for a dry day on period  $t$  (i.e.  $Y_{i,t} = 0$ ).

$\mu_t^W$  is the generic name for the mean for a wet day on period  $t$  (i.e.  $Y_{i,t} > 0$ ).

$\sigma^D$  is the generic name for the standard deviation for a dry day.

$\sigma^W$  is the generic name for the standard deviation for a wet day.

$\theta^D$  is the coefficient of the AR(1) process, given a dry day.

$\theta^W$  is the coefficient of the AR(1) process given a wet day.

$C(D)$  denotes the number of elements in the set  $N(D)$ .

$C(W)$  denotes the number of elements in the set  $N(W)$ .

Then  $T = C(D) + C(W)$ .

Since all variables are modelled in the same way, the representation will be given for modelling one variable. The same procedure is then repeated for each of the remaining climate variables.

### (b) Model and assumptions

The Model under consideration is given by:

$$\chi_{i,t} = \begin{cases} \frac{S_{i,t} - \mu_t^D}{\sigma^D} & \text{if } Y_{i,t} = 0 \\ \frac{S_{i,t} - \mu_t^W}{\sigma^W} & \text{if } Y_{i,t} > 0 \end{cases}$$

where  $i = 1, 2, \dots, NY$  and  $t = 1, 2, \dots, NT$ .

That is, the residual time series  $\chi_{i,t}$  is obtained by removing the periodic mean and the standard deviation from the observed time series  $S_{i,t}$ . The resulting time series thus has a mean of zero and unit variance.

Assume  $\chi_{i,t}$  is generated by an autoregressive process of order  $p$  (AR(p)) defined as

$$\chi_{i,t} = \theta_1 \chi_{i,t-1} + \theta_2 \chi_{i,t-2} + \dots + \theta_p \chi_{i,t-p} + e_{i,t}$$

where  $\{e_{i,t}\}$  is a set of independent, normally distributed variables with mean zero and variance of unity, i.e.

$$e_{i,t} \sim NID(0, 1).$$

That is,  $\chi_{i,t}$  is regressed on past values of  $\chi_{i,t}$  instead of on independent variables as on the classical multiple regression.

The assumption that  $\chi_{i,t}$  is described by an autoregressive process can be substantiated by arguments put forward by Cochrane and Orcutt (1949). The sources of autocorrelation in the error term can be:

- (i) When modelling climatic variables, errors in modelling arise from faulty descriptions of these variables. Since these variables are themselves autocorrelated, this type of error will be autocorrelated.
- (ii) Error terms may arise from omitting variables from the analysis because these variables are either not available or their importance is not realised or because the influence they have is so small that it is not convenient to insert them. As already indicated these variables are autocorrelated and, therefore, one may expect the resulting error terms to be also autocorrelated.

An autoregressive process of order 1, AR(1), sometimes called the Markov process, was chosen to describe  $\chi_{i,t}$ . The reason for this choice will be discussed later. To simplify the formula, the theory will only be shown for an AR(1) process from now on. The order of the process can be increased to any order desired, but this has to be done at the cost of increasing both the complexity and the number of parameters to be estimated.

The form of  $\chi_{i,t}$  is thus given by:

$$\chi_{i,t} = \theta \chi_{i,t-1} + e_{i,t}$$

where  $e_{i,t} \sim NID(0,1)$   $i = 1, 2, \dots, NY$ ;  $t = 1, 2, \dots, NT$ .

The model that incorporates the different wet and dry sequences is given by:

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D} \quad \text{for a dry sequence}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^W}{\sigma^W} - \theta^W \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W} \quad \text{for a wet sequence.}$$

The seasonal mean function,  $\mu_t$ , is approximated by its truncated Fourier series representation, i.e.

$$\mu_t^D = \sum_{i=1}^L \alpha_i^D \varphi_i(t) \quad \text{if } t \text{ dry}$$

and

$$\mu_t^W = \sum_{i=1}^L \alpha_i^W \varphi_i(t) \quad \text{if } t \text{ wet}$$

where  $\varphi_i(t)$  is defined as before and  $L$  is the order of the Fourier series representation.

## (c) Estimation

The procedure to estimate the parameters  $\alpha_j^D, \alpha_j^W, \theta^D, \theta^W, \sigma^D$  and  $\sigma^W$ ;  $j = 1, 2, \dots, L$  is now discussed.

Since  $e_{i,t} \sim NID(0, 1)$ , the density function of  $e_{i,t}$  is given by

$$f(e_{i,t}) = \frac{1}{\sqrt{2\pi}} \exp\left(-\frac{1}{2}e_{i,t}^2\right)$$

Therefore the joint likelihood function, conditioned on the wet and dry status of the day is given by

$$\begin{aligned} L(\psi) &= L(\alpha_j^D, \alpha_j^W, \theta^D, \theta^W, \sigma^D, \sigma^W; e_{i,t}) \\ &= \prod_{t \in N(D)} f(e_{i,t}|D) \prod_{t \in N(W)} f(e_{i,t}|W) \end{aligned}$$

where  $f(e_{i,t}|D)$  represents the density function of  $e_{i,t}$  given that a dry day has been observed, and  $f(e_{i,t}|W)$  represents the density function of  $e_{i,t}$  given that a wet day has been observed.

Substituting the density function, one obtains

$$L(\psi) = \left(\frac{1}{\sqrt{2\pi}}\right)^T \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(DD)} (e_{i,t}|D)^2 + \sum_{t \in N(WW)} (e_{i,t}|W)^2 \right] \right\}.$$

Making the following transformation

$$e_{i,t} = \frac{S_{i,t} - \mu_t}{\sigma} - \theta \frac{S_{i,t-1} - \mu_{t-1}}{\sigma}$$

where the Jacobian of the transformation is given by

$$\begin{aligned} &\left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right|, \quad i = 1, 2, \dots, NY; \quad t = 1, 2, \dots, NT; \quad p = 1, 2, \dots, NT \\ &= \left| \frac{\partial e_k}{\partial S_n} \right|, \quad k = 1, 2, \dots, T; \quad n = 1, 2, \dots, T \\ &= \begin{vmatrix} 1/\sigma & 0 & \dots & \dots & 0 \\ -\theta/\sigma & 1/\sigma & 0 & \dots & 0 \\ \vdots & & & & \vdots \\ 0 & \dots & \dots & -\theta/\sigma & 1/\sigma \end{vmatrix} = \prod_{k=1}^T 1/\sigma = \prod_{i=1}^{NY} \prod_{t=1}^{NT} 1/\sigma \end{aligned}$$

since we are dealing with a triangular matrix. Taking into account the dry and wet status of the day, the Jacobian can be rewritten as

$$\left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right| = \prod_{t \in N(D)} \frac{1}{\sigma^D} \prod_{t \in N(W)} \frac{1}{\sigma^W},$$

then the joint likelihood function is given by

$$\begin{aligned} L(\psi) = & \left( \frac{1}{\sqrt{2\pi}} \right)^T \left( \frac{1}{\sigma^D} \right)^{C(D)} \left( \frac{1}{\sigma^W} \right)^{C(W)} \\ & \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(D)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D} \right)^2 \right. \right. \\ & \left. \left. + \sum_{t \in N(W)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^W} - \theta^W \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W} \right)^2 \right] \right\} \end{aligned}$$

and the log likelihood function is given by

$$\begin{aligned} \ell(\psi) = & -\frac{T}{2} \log(2\pi) - C(D) \log(\sigma^D) - C(W) \log(\sigma^W) \\ & - \frac{1}{2} \left[ \sum_{t \in N(D)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D} \right)^2 \right. \\ & \left. + \sum_{t \in N(W)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^W} - \theta^W \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W} \right)^2 \right]. \end{aligned}$$

Maximum likelihood estimates can be obtained by minimising  $\ell(\psi)$ . That is, the first partial derivatives with respect to the parameters are set to zero.

The first partial derivatives with respect to the parameters are given by

$$\begin{aligned} \frac{\partial \ell(\psi)}{\partial \theta^D} &= \sum_{t \in N(D)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D} \right) \\ & \quad \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D} \right) \\ \frac{\partial \ell(\psi)}{\partial \theta^W} &= \sum_{t \in N(W)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^W} - \theta^W \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W} \right) \\ & \quad \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W} \right) \\ \frac{\partial \ell(\psi)}{\partial \alpha_j^D} &= - \left[ \sum_{t \in N(D)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D} \right) \right. \\ & \quad \left. \left( \frac{-\varphi_j(t)}{\sigma^D} + \frac{\theta^D \varphi_j(t-1)}{\sigma^D} \right) \right] \end{aligned}$$

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \alpha_j^W} &= - \left[ \sum_{t \in N(W)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^W} - \theta^W \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W} \right) \right. \\
&\quad \left. \left( \frac{-\varphi_j(t)}{\sigma^W} + \frac{\theta^W \varphi_j(t-1)}{\sigma^W} \right) \right] \\
\frac{\partial \ell(\psi)}{\partial \sigma^D} &= -\frac{C(D)}{\sigma^D} - \sum_{t \in N(D)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D} \right) \\
&\quad \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma^D)^2} + \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^D)^2} \right) \\
\frac{\partial \ell(\psi)}{\partial \sigma^W} &= -\frac{C(W)}{\sigma^W} - \sum_{t \in N(W)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^W} - \theta^W \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W} \right) \\
&\quad \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma^W)^2} + \theta^W \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^W)^2} \right)
\end{aligned}$$

The parameter estimates are given by

$$\hat{\theta}^D = \frac{\sum_{t \in N(D)} (S_{i,t} - \hat{\mu}_t^D)(S_{i,t-1} - \hat{\mu}_{t-1}^D)}{\sum_{t \in N(D)} (S_{i,t-1} - \hat{\mu}_{t-1}^D)^2} \quad (1)$$

$$\hat{\theta}^W = \frac{\sum_{t \in N(W)} (S_{i,t} - \hat{\mu}_t^W)(S_{i,t-1} - \hat{\mu}_{t-1}^W)}{\sum_{t \in N(W)} (S_{i,t-1} - \hat{\mu}_{t-1}^W)^2} \quad (2)$$

$$\hat{\sigma}^D = \left( \frac{1}{C(D)} \sum_{t \in N(D)} (S_{i,t} - \hat{\mu}_t^D - \hat{\theta}^D(S_{i,t-1} - \hat{\mu}_{t-1}^D))^2 \right)^{\frac{1}{2}} \quad (3)$$

$$\hat{\sigma}^W = \left( \frac{1}{C(W)} \sum_{t \in N(W)} (S_{i,t} - \hat{\mu}_t^W - \hat{\theta}^W(S_{i,t-1} - \hat{\mu}_{t-1}^W))^2 \right)^{\frac{1}{2}} \quad (4)$$

$$\hat{\alpha}_j^D = \left\{ \sum_{t \in N(D)} (\varphi_j(t) - \hat{\theta}^D \varphi_j(t-1))^2 \right\}^{-1} [A - M] \quad (5)$$

where

$$A = \sum_{t \in N(D)} (S_{i,t} - \hat{\theta}^D S_{i,t-1})(\varphi_j(t) - \hat{\theta}^D \varphi_j(t-1))$$

and

$$M = \sum_{t \in N(D)} \left[ \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t) \right) - \hat{\theta}^D \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t-1) \right) \right] \\ [\varphi_j(t) - \hat{\theta}^D \varphi_j(t-1)]$$

$$\hat{\alpha}_j^W = \left\{ \sum_{t \in N(W)} (\varphi_j(t) - \hat{\theta}^W \varphi_j(t-1))^2 \right\}^{-1} [A_2 - M_2] \quad (6)$$

where

$$A_2 = \sum_{t \in N(W)} (S_{i,t} - \hat{\theta}^W S_{i,t-1})(\varphi_j(t) - \hat{\theta}^W \varphi_j(t-1))$$

and

$$M_2 = \sum_{t \in N(W)} \left[ \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t) \right) - \hat{\theta}^W \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t-1) \right) \right] \\ [\varphi_j(t) - \hat{\theta}^W \varphi_j(t-1)]$$

These equations cannot be solved explicitly and therefore have to be solved iteratively.

Note that

$\mu_t$  is a function of the  $\alpha_j$  where the  $\alpha_j$  are functions of  $\theta$

$\theta$  is a function of  $\mu_t$  and

$\sigma$  is a function of  $\mu_t$  and  $\theta$ .

The following algorithm can be performed to estimate the parameters.

#### Algorithm

**Step 1:** Estimate initial  $\mu_t$  by approximating by it Fourier series transformation and estimating the parameters  $\alpha_i$  by the method mentioned in the previous models.

**Step 2:** Estimate  $\theta$  using (1) and (2)

**Step 3:** Estimate  $\sigma$  using (3) and (4)

**Step 4:** Estimate  $\mu_t$  using (5) and (6)

**Step 5:** Test for convergence of all parameters, i.e. when the percentage change in parameter estimates is sufficiently small. If convergence is not met, return to Step 2.

To model the multivariate time series the covariance matrix of the residuals of the different climate variables is needed.

The cross-correlation matrix,  $\hat{\Sigma}$ , has elements  $R_{jk}$ , where

$$R_{jk} = \frac{\frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} e_{i,t}^{(k)} - \frac{1}{T^2} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)}}{\left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(j)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \right)^2 \right]^{\frac{1}{2}} \left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(k)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)} \right)^2 \right]^{\frac{1}{2}}}$$

where  $e_{i,t}^{(j)}$  denotes the residual time series of variable  $j$ ,  $j = 1, 2, \dots, NV$  and  $e_{i,t}^{(k)}$  denotes the residual time series of variable  $k$ ,  $k = 1, 2, \dots, NV$ .

#### (d) Model Selection

The order of the autoregressive process is selected in the same way as in the previous models, as is the order of the Fourier series approximation.

A major problem was encountered in Model 2, in that a high proportion of information is discarded. The reason why this problem occurs is explained by means of an example.

Suppose the following sequence has been observed

day $t$	1	2	3	4	5	6	7	8	9	10	11	12
status of day	D	D	W	D	W	W	D	W	D	D	D	W

By definition,  $N(D)$  and  $N(W)$  represent the sets of time periods  $t$  such that  $t$  was dry or wet respectively. Thus,  $N(D)$  consists of the elements

$$\{1, 2, 4, 7, 9, 10, 11\},$$

and  $N(W)$  of

$$\{3, 5, 6, 8, 12\}.$$

In Model 2, one is only interested in conditioning the parameters into dry and wet sequences, therefore one is restricting the status of time period  $t - 1$  to be the same as that of  $t$ . Thus, given that day  $t$  was dry, the model is given by

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta^D \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^D}$$

and one does not consider the case of

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^W}.$$

Similarly for when  $t$  is wet.

Therefore, when constructing  $N(D)$  and  $N(W)$ , only sequences of at least two dry (or two wet) consecutive days can be used. In this case,  $N(D)$  has elements

$$\{2, 10, 11\}$$

and  $N(W)$

$$\{6\}.$$

Thus, a high proportion of the observations are discarded. This led to the development of Model 3 and therefore Model 2 is of no further interest.

**Model 3: Multivariate model for climate data**

The two previous models condition the parameters of the model on the wet or dry status of the day. When generating climate sequences these models represent conditions in which a wet day follows a wet day and a dry day follows a dry day but fail to explain the relationship between conditions such as a wet day following a dry day or a dry day following a wet day.

To generate representative climate sequences these sequences must be related to the sequences of rain and no-rain days. To achieve this relationship, the parameters of the model must be conditioned on the four possible sequences in the rainfall variable:

1. a dry day follows a dry day,
2. a wet day follows a wet day,
3. a wet day follows a dry day,
4. a dry day follows a wet day.

**(a) Notation**

Partition the year into  $NT(= 365)$  equal intervals, denoted by  $t = 1, 2, \dots, NT$

$NV$  is the number of variables.

$NY$  is the number of years observed.

$W$  represents the occurrence of rain.

$D$  represents the non-occurrence of rain.

$DD$  represents the sequence when day  $t - 1$  was dry and day  $t$  was dry.

$WW$  represents the sequence when day  $t - 1$  was wet and day  $t$  was wet.

$DW$  represents the sequence when day  $t - 1$  was dry and day  $t$  was wet.

$WD$  represents the sequence when day  $t - 1$  was wet and day  $t$  was dry.

$T$  represents the total number of observations, i.e.  $NYNT$ .

$N(DD)$  is the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was dry,  $t = 1, 2, \dots, T$ .

$N(WW)$  is the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was wet.

$N(DW)$  is the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was dry.

$N(WD)$  is the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was wet.

$Y_{i,t}$  is the precipitation amount of period  $t$  of year  $i$ ,  $i = 1, 2, \dots, NY$ .

$S_{i,t}$  is the generic name for the observation at time  $t$  of the  $i$ th year.

$\mu_t^D$  is the generic name for the mean for a dry day on period  $t$ .

$\mu_t^W$  is the generic name for the mean for a wet day on period  $t$ .

$\sigma^{DD}$  is the generic name for the standard deviation given sequence  $DD$ .

$\sigma^{WW}$  is the generic name for the standard deviation given sequence  $WW$ .

$\sigma^{DW}$  is the generic name for the standard deviation given sequence  $DW$ .

$\sigma^{WD}$  is the generic name for the standard deviation given sequence  $WD$ .

$\theta^{DD}$  is the coefficient of the AR(1) process given sequence  $DD$ .

$\theta^{WW}$  is the coefficient of the AR(1) process given sequence  $WW$ .

$\theta^{DW}$  is the coefficient of the AR(1) process given sequence  $DW$ .

$\theta^{WD}$  is the coefficient of the AR(1) process given sequence  $WD$ .

$C(DD)$  is the number of elements in the set  $N(DD)$ .

$C(WW)$  is the number of elements in the set  $N(WW)$ .

$C(DW)$  is the number of elements in the set  $N(DW)$ .

$C(WD)$  is the number of elements in the set  $N(WD)$ .

Then  $T = C(DD) + C(WW) + C(DW) + C(WD)$ .

### (b) Model and assumptions

The time series  $S_{i,t}$  is reduced to a time series of residual elements,  $\chi_{i,t}$ , by removing the periodic means and the standard deviation for the appropriate sequence, i.e.

$$\chi_{i,t} = \frac{S_{i,t} - \mu_t}{\sigma}$$

This results in a time series with zero mean and standard deviation of unity, which is assumed to follow an AR(1) process, i.e.

$$\chi_{i,t} = \theta \chi_{i,t-1} + e_{i,t}$$

where  $e_{i,t} \sim NID(0,1)$   $i = 1, 2, \dots, NY$ ;  $t = 1, 2, \dots, NT$ .

The model that incorporates the different wet-dry sequences is then given by:

$$e_{i,t} = \frac{S_{i,t} - \mu_t^{DD}}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^{DD}}{\sigma^{DD}} \quad \text{if day } t-1 \text{ was dry and day } t \text{ was dry.}$$

$$e_{i,t} = \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \quad \text{if day } t-1 \text{ was wet and day } t \text{ was wet.}$$

$$e_{i,t} = \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \quad \text{if day } t-1 \text{ was dry and day } t \text{ was wet.}$$

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \quad \text{if day } t-1 \text{ was wet and day } t \text{ was dry.}$$

The parameter  $\mu_t$  for this model, and for the models following, was only conditioned on the rain and no-rain status of the day. This was done to simplify the model, otherwise for each sequence one would have two equations. For example, if sequence  $DD$  has occurred, then

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DD}} \quad \text{if the sequence } DDD \text{ was observed}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^{DD}}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^{WD}}{\sigma^{DD}} \quad \text{if the sequence } WDD \text{ was observed}$$

This not only increases the number of parameters to be estimated but in addition we are no longer assuming that all the information we need of previous values of the model is given by the value of the previous day. The state of the second previous day is also required.

As already discussed, it is reasonable to approximate the mean function  $\mu_t$  by its truncated Fourier representation, i.e.

$$\mu_t^D = \sum_{i=1}^L \alpha_i^D \varphi_i(t) \quad \text{if } t \text{ dry}$$

and

$$\mu_t^W = \sum_{i=1}^L \alpha_i^W \varphi_i(t) \quad \text{if } t \text{ wet}$$

where  $\varphi_i(t)$  is defined as before and  $L$  is the order of the Fourier series approximation.

**(c) Estimation**

The procedure to estimate the parameters  $\alpha_j^D, \alpha_j^W, \theta^{DD}, \theta^{WW}, \theta^{WD}, \theta^{DW}, \sigma^{DD}, \sigma^{WW}, \sigma^{DW}, \sigma^{WD}$ ;  $j = 1, 2, \dots, L$  is now discussed.

Maximum likelihood estimates can be obtained by observing that since  $e_{i,t} \sim NID(0, 1)$ , then the density function of  $e_{i,t}$  is given by

$$f(e_{i,t}) = \frac{1}{\sqrt{2\pi}} \exp\left(-\frac{1}{2} e_{i,t}^2\right)$$

Therefore the joint likelihood function conditioned on the four different sequences is given by

$$\begin{aligned} L(\psi) &= L(\alpha_j^D, \alpha_j^W, \theta^{DD}, \theta^{WW}, \theta^{WD}, \theta^{DW}, \sigma^{DD}, \sigma^{WW}, \sigma^{DW}, \sigma^{WD}; e_{i,t}) \\ &= \prod_{t \in N(DD)} f(e_{i,t}|DD) \prod_{t \in N(WW)} f(e_{i,t}|WW) \prod_{t \in N(DW)} f(e_{i,t}|DW) \prod_{t \in N(WD)} f(e_{i,t}|WD) \end{aligned}$$

where  $f(e_{i,t}|DD)$  represents the density function of  $e_{i,t}$  given that the sequence  $DD$  has been observed, and similarly for the other density functions.

$$\begin{aligned} &= \left( \frac{1}{\sqrt{2\pi}} \right)^T \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(DD)} (e_{i,t}|DD)^2 + \sum_{t \in N(WW)} (e_{i,t}|WW)^2 \right. \right. \\ &\quad \left. \left. + \sum_{t \in N(DW)} (e_{i,t}|DW)^2 + \sum_{t \in N(WD)} (e_{i,t}|WD)^2 \right] \right\} \end{aligned}$$

One now makes the following transformation

$$e_{i,t} = \frac{S_{i,t} - \mu_t}{\sigma} - \theta \frac{S_{i,t-1} - \mu_{t-1}}{\sigma}.$$

The Jacobian of the transformation is given by

$$\begin{aligned} &\left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right|, i = 1, 2, \dots, NY; \quad t = 1, 2, \dots, NT; \quad p = 1, 2, \dots, NT. \\ &= \left| \frac{\partial e_k}{\partial S_n} \right|, k = 1, 2, \dots, T; \quad n = 1, 2, \dots, T. \\ &= \begin{vmatrix} 1/\sigma & 0 & \dots & \dots & 0 \\ -\theta/\sigma & 1/\sigma & 0 & \dots & 0 \\ \vdots & & & & \\ 0 & \dots & \dots & -\theta/\sigma & 1/\sigma \end{vmatrix} = \prod_{k=1}^T 1/\sigma = \prod_{i=1}^{NY} \prod_{t=1}^{NT} 1/\sigma \end{aligned}$$

since we are dealing with a triangular matrix.

Taking into account the conditional sequences imposed on  $e_{i,t}$ , the Jacobian is given by

$$\left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right| = \prod_{t \in N(DD)} \frac{1}{\sigma^{DD}} \prod_{t \in N(WW)} \frac{1}{\sigma^{WW}} \prod_{t \in N(DW)} \frac{1}{\sigma^{DW}} \prod_{t \in N(WD)} \frac{1}{\sigma^{WD}}.$$

Then the joint probability density function of  $S_{i,t}$  is given by

$$\begin{aligned} L(\psi) = & \left( \frac{1}{\sqrt{2\pi}} \right)^T \left( \frac{1}{\sigma^{DD}} \right)^{C(DD)} \left( \frac{1}{\sigma^{WW}} \right)^{C(WW)} \left( \frac{1}{\sigma^{DW}} \right)^{C(DW)} \left( \frac{1}{\sigma^{WD}} \right)^{C(WD)} \\ & \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^{DD}}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^{DD}}{\sigma^{DD}} \right)^2 \right. \right. \\ & + \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right)^2 \\ & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right)^2 \\ & \left. \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right)^2 \right] \right\} \end{aligned}$$

and the log likelihood is given by:

$$\begin{aligned} \ell(\psi) = & \frac{-T}{2} \log(2\pi) - C(DD) \log(\sigma^{DD}) - C(WW) \log(\sigma^{WW}) \\ & - C(DW) \log(\sigma^{DW}) - C(WD) \log(\sigma^{WD}) \\ & - \frac{1}{2} \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^{DD}}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^{DD}}{\sigma^{DD}} \right)^2 \right. \\ & + \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right)^2 \\ & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right)^2 \\ & \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right)^2 \right]. \end{aligned}$$

Maximum likelihood estimates can be obtained by minimising  $\ell(\psi)$  and this is achieved by setting its partial derivatives with respect to the parameters equal to zero.

The first partial derivatives with respect to the parameteers are given by:

$$\frac{\partial \ell(\psi)}{\partial \theta^{DD}} = \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^{DD}}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^{DD}}{\sigma^{DD}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DD}} \right)$$

$$\frac{\partial \ell(\psi)}{\partial \theta^{WW}} = \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right)$$

$$\frac{\partial \ell(\psi)}{\partial \theta^{DW}} = \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right)$$

$$\frac{\partial \ell(\psi)}{\partial \theta^{WD}} = \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right)$$

$$\begin{aligned} \frac{\partial \ell(\psi)}{\partial \alpha_j^D} = & - \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^{DD}}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^{DD}}{\sigma^{DD}} \right) \right. \\ & \left( \frac{-\varphi_j(t)}{\sigma^{DD}} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma^{DD}} \right) \\ & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma^{DW}} \right) \\ & \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \left( \frac{-\varphi_j(t)}{\sigma^{WD}} \right) \right] \end{aligned}$$

$$\begin{aligned} \frac{\partial \ell(\psi)}{\partial \alpha_j^W} = & - \left[ \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right) \right. \\ & \left( \frac{-\varphi_j(t)}{\sigma^{WW}} + \frac{\theta^{WW} \varphi_j(t-1)}{\sigma^{WW}} \right) \\ & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \left( \frac{-\varphi_j(t)}{\sigma^{DW}} \right) \\ & \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \left( \frac{\theta^{WD} \varphi_j(t-1)}{\sigma^{WD}} \right) \right] \end{aligned}$$

$$\begin{aligned} \frac{\partial \ell(\psi)}{\partial \sigma^{DD}} = & - \frac{C(DD)}{\sigma^{DD}} - \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^{DD}}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^{DD}}{\sigma^{DD}} \right) \\ & \left( - \frac{(S_{i,t} - \mu_t^D)}{(\sigma^{DD})^2} + \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DD})^2} \right) \end{aligned}$$

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \sigma^{WW}} &= -\frac{C(WW)}{\sigma^{WW}} - \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right) \\
&\quad \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma^{WW})^2} + \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WW})^2} \right) \\
\frac{\partial \ell(\psi)}{\partial \sigma^{DW}} &= -\frac{C(DW)}{\sigma^{DW}} - \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \\
&\quad \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma^{DW})^2} + \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DW})^2} \right) \\
\frac{\partial \ell(\psi)}{\partial \sigma^{WD}} &= -\frac{C(WD)}{\sigma^{WD}} - \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \\
&\quad \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma^{WD})^2} + \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WD})^2} \right)
\end{aligned}$$

The parameter estimates are given by:

$$\begin{aligned}\hat{\theta}^{DD} &= \frac{\sum_{t \in N(DD)} (S_{i,t} - \hat{\mu}_t^D)(S_{i,t-1} - \hat{\mu}_{t-1}^D)}{\sum_{t \in N(DD)} (S_{i,t-1} - \hat{\mu}_{t-1}^D)^2} \\ \hat{\theta}^{WW} &= \frac{\sum_{t \in N(WW)} (S_{i,t} - \hat{\mu}_t^W)(S_{i,t-1} - \hat{\mu}_{t-1}^W)}{\sum_{t \in N(WW)} (S_{i,t-1} - \hat{\mu}_{t-1}^W)^2} \\ \hat{\theta}^{DW} &= \frac{\sum_{t \in N(DW)} (S_{i,t} - \hat{\mu}_t^W)(S_{i,t-1} - \hat{\mu}_{t-1}^D)}{\sum_{t \in N(DW)} (S_{i,t-1} - \hat{\mu}_{t-1}^D)^2} \\ \hat{\theta}^{WD} &= \frac{\sum_{t \in N(WD)} (S_{i,t} - \hat{\mu}_t^D)(S_{i,t-1} - \hat{\mu}_{t-1}^W)}{\sum_{t \in N(WD)} (S_{i,t-1} - \hat{\mu}_{t-1}^W)^2}\end{aligned}$$

$$\begin{aligned}\hat{\sigma}^{DD} &= \left( \frac{1}{C(DD)} \sum_{t \in N(DD)} (S_{i,t} - \hat{\mu}_t^D - \hat{\theta}^{DD}(S_{i,t-1} - \hat{\mu}_{t-1}^D))^2 \right)^{\frac{1}{2}} \\ \hat{\sigma}^{WW} &= \left( \frac{1}{C(WW)} \sum_{t \in N(WW)} (S_{i,t} - \hat{\mu}_t^W - \hat{\theta}^{WW}(S_{i,t-1} - \hat{\mu}_{t-1}^W))^2 \right)^{\frac{1}{2}} \\ \hat{\sigma}^{DW} &= \left( \frac{1}{C(DW)} \sum_{t \in N(DW)} (S_{i,t} - \hat{\mu}_t^W - \hat{\theta}^{DW}(S_{i,t-1} - \hat{\mu}_{t-1}^D))^2 \right)^{\frac{1}{2}} \\ \hat{\sigma}^{WD} &= \left( \frac{1}{C(WD)} \sum_{t \in N(WD)} (S_{i,t} - \hat{\mu}_t^D - \hat{\theta}^{WD}(S_{i,t-1} - \hat{\mu}_{t-1}^W))^2 \right)^{\frac{1}{2}}\end{aligned}$$

$$\begin{aligned}\hat{\alpha}_J^D &= \left\{ \frac{1}{(\hat{\sigma}^{DD})^2} \sum_{t \in N(DD)} \varphi_j(t)^2 - \frac{2}{(\hat{\sigma}^{DD})^2} \hat{\theta}^{DD} \sum_{t \in N(DD)} \varphi_j(t) \varphi_j(t-1) \right. \\ &\quad + \frac{1}{(\hat{\sigma}^{DD})^2} (\hat{\theta}^{DD})^2 \sum_{t \in N(DD)} \varphi_j(t-1)^2 \\ &\quad + \frac{1}{(\hat{\sigma}^{DW})^2} (\hat{\theta}^{DW})^2 \sum_{t \in N(DW)} \varphi_j(t-1)^2 \\ &\quad \left. + \frac{1}{(\hat{\sigma}^{WD})^2} \sum_{t \in N(WD)} \varphi_j(t)^2 \right\}^{-1} [-A - M]\end{aligned}$$

where

$$\begin{aligned}
A = & -\frac{1}{(\hat{\sigma}^{DD})^2} \sum_{t \in N(DD)} (S_{i,t} - \hat{\theta}^{DD} S_{i,t-1}) \varphi_j(t) \\
& + \frac{1}{(\hat{\sigma}^{DD})^2} \hat{\theta}^{DD} \sum_{t \in N(DD)} (S_{i,t} - \hat{\theta}^{DD} S_{i,t-1}) \varphi_j(t-1) \\
& + \frac{1}{(\hat{\sigma}^{DW})^2} \hat{\theta}^{DW} \sum_{t \in N(DW)} (S_{i,t} - \hat{\theta}^{DW} S_{i,t-1}) \varphi_j(t-1) \\
& - \frac{1}{(\hat{\sigma}^{WD})^2} \sum_{t \in N(WD)} (S_{i,t} - \hat{\theta}^{WD} S_{i,t-1}) \varphi_j(t) \\
& - \frac{1}{(\hat{\sigma}^{DW})^2} \hat{\theta}^{DW} \sum_{t \in N(DW)} \hat{\mu}_t^W \varphi_j(t-1) \\
& - \frac{1}{(\hat{\sigma}^{WD})^2} \hat{\theta}^{WD} \sum_{t \in N(WD)} \hat{\mu}_{t-1}^W \varphi_j(t)
\end{aligned}$$

and

$$\begin{aligned}
M = & -\frac{1}{(\hat{\sigma}^{DD})^2} \hat{\theta}^{DD} \sum_{t \in N(DD)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t-1) \right) \varphi_j(t) \\
& - \frac{1}{(\hat{\sigma}^{DD})^2} \hat{\theta}^{DD} \sum_{t \in N(DD)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t) \right) \varphi_j(t-1) \\
& + \frac{1}{(\hat{\sigma}^{DD})^2} \sum_{t \in N(DD)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t) \right) \varphi_j(t) \\
& + \frac{1}{(\hat{\sigma}^{DD})^2} (\hat{\theta}^{DD})^2 \sum_{t \in N(DD)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t-1) \right) \varphi_j(t-1) \\
& + \frac{1}{(\hat{\sigma}^{DW})^2} (\hat{\theta}^{DW})^2 \sum_{t \in N(DW)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t-1) \right) \varphi_j(t-1) \\
& + \frac{1}{(\hat{\sigma}^{WD})^2} \sum_{t \in N(WD)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^D \varphi_k(t) \right) \varphi_j(t).
\end{aligned}$$

The estimate for  $\alpha_j^W$  is given by

$$\begin{aligned}\hat{\alpha}_j^W = & \left\{ \frac{1}{(\hat{\sigma}^{WW})^2} \sum_{t \in N(WW)} (\varphi_j(t))^2 - \frac{2}{(\hat{\sigma}^{WW})^2} \hat{\theta}^{WW} \sum_{t \in N(WW)} \varphi_j(t-1) \varphi_j(t) \right. \\ & + \frac{1}{(\hat{\sigma}^{WW})^2} (\hat{\theta}^{WW})^2 \sum_{t \in N(WW)} (\varphi_j(t-1))^2 \\ & + \frac{1}{(\hat{\sigma}^{DW})^2} \sum_{t \in N(DW)} (\varphi_j(t))^2 \\ & \left. + \frac{1}{(\hat{\sigma}^{WD})^2} (\hat{\theta}^{WD})^2 \sum_{t \in N(WD)} (\varphi_j(t-1))^2 \right\}^{-1} [-A_2 - M_2]\end{aligned}$$

where

$$\begin{aligned}A_2 = & -\frac{1}{(\hat{\sigma}^{WW})^2} \sum_{t \in N(WW)} (S_{i,t} - \hat{\theta}^{WW} S_{i,t-1}) \varphi_j(t) \\ & + \frac{1}{(\hat{\sigma}^{WW})^2} \hat{\theta}^{WW} \sum_{t \in N(WW)} (S_{i,t} - \hat{\theta}^{WW} S_{i,t-1}) \varphi_j(t-1) \\ & - \frac{1}{(\hat{\sigma}^{DW})^2} \sum_{t \in N(DW)} (S_{i,t} - \hat{\theta}^{DW} S_{i,t-1}) \varphi_j(t) \\ & - \frac{1}{(\hat{\sigma}^{DW})^2} \hat{\theta}^{DW} \sum_{t \in N(DW)} \hat{\mu}_{i-1}^D \varphi_j(t) \\ & + \frac{1}{(\hat{\sigma}^{WD})^2} \hat{\theta}^{WD} \sum_{t \in N(WD)} (S_{i,t} - \hat{\theta}^{WD} S_{i,t-1}) \varphi_j(t-1) \\ & - \frac{1}{(\hat{\sigma}^{WD})^2} \hat{\theta}^{WD} \sum_{t \in N(WD)} (\hat{\mu}_t^D) \varphi_j(t-1)\end{aligned}$$

and

$$\begin{aligned}
M_2 = & \frac{1}{(\hat{\sigma}^{WW})^2} \sum_{t \in N(WW)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t) \right) \varphi_j(t) \\
& - \frac{1}{(\hat{\sigma}^{WW})^2} \hat{\theta}^{WW} \left[ \sum_{t \in N(WW)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t-1) \right) \varphi_j(t) \right. \\
& \left. + \sum_{t \in N(WW)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t) \right) \varphi_j(t-1) \right] \\
& + \frac{1}{(\hat{\sigma}^{WW})^2} (\hat{\theta}^{WW})^2 \sum_{t \in N(WW)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t-1) \right) \varphi_j(t-1) \\
& + \frac{1}{(\hat{\sigma}^{DW})^2} \sum_{t \in N(DW)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t) \right) \varphi_j(t) \\
& + \frac{1}{(\hat{\sigma}^{WD})^2} (\hat{\theta}^{WD})^2 \sum_{t \in N(WD)} \left( \sum_{\substack{k=1 \\ k \neq j}}^L \hat{\alpha}_k^W \varphi_k(t-1) \right) \varphi_j(t-1).
\end{aligned}$$

These equations cannot be solved explicitly and therefore the Newton-Raphson iteration method is used to solve them. The second partial derivatives are required to use this method and these are given by

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{DD}} &= - \sum_{t \in N(DD)} \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DD}} \right)^2 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{WW}} &= \frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{DW}} = \frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{WD}} = 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \alpha_j^D} &= \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DD}} \right) \left( \frac{-\varphi_j(t-1)}{\sigma^{DD}} \right) \right. \\
&\quad \left. + \left( \frac{-\varphi_j(t)}{\sigma^{DD}} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma^{DD}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DD}} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \alpha_j^W} &= 0
\end{aligned}$$

$$\begin{aligned}
& \left\{ \left( \frac{MM^0}{1-i^2\eta - 1-i^2S} \right) \left( \frac{z(MM^0)}{1-i^2\eta - 1-i^2S} MM\theta + \frac{z(MM^0)}{(1-i^2\eta - i^2S)} - \right) + \right. \\
& \left. \left( \frac{z(MM^0)}{(1-i^2\eta - 1-i^2S)} - \right) \left( \frac{MM^0}{1-i^2\eta - 1-i^2S} MM\theta - \frac{MM^0}{1-i^2\eta - i^2S} \right) \right\} \sum^{(MM)N\exists i} = \frac{MM^0\partial MM\theta\partial}{(\phi)\partial_z\partial} \\
& 0 = \frac{\partial MM^0\partial MM\theta\partial}{(\phi)\partial_z\partial} = \frac{MM^0\partial MM\theta\partial}{(\phi)\partial_z\partial} = \frac{\partial\partial MM^0\partial MM\theta\partial}{(\phi)\partial_z\partial} \\
& \left\{ \left( \frac{MM^0}{1-i^2\eta - 1-i^2S} \right) \left( \frac{MM^0}{(1-i^2)\partial} MM\theta + \frac{MM^0}{(i^2)\partial} - \right) + \right. \\
& \left. \left( \frac{MM^0}{(1-i^2)\partial} - \right) \left( \frac{MM^0}{1-i^2\eta - 1-i^2S} MM\theta - \frac{MM^0}{1-i^2\eta - i^2S} \right) \right\} \sum^{(MM)N\exists i} = \frac{MM^0\partial MM\theta\partial}{(\phi)\partial_z\partial}
\end{aligned}$$

$$\begin{aligned}
0 &= \frac{\partial^2\partial MM\theta\partial}{(\phi)\partial_z\partial} = \frac{\partial MM\theta\partial MM\theta\partial}{(\phi)\partial_z\partial} = \frac{MM\theta\partial MM\theta\partial}{(\phi)\partial_z\partial} \\
& \left( \frac{MM^0}{1-i^2\eta - 1-i^2S} \right) \sum^{(MM)N\exists i} - = \frac{MM\theta\partial MM\theta\partial}{(\phi)\partial_z\partial} \\
0 &= \frac{\partial MM^0\partial\partial\partial\partial}{(\phi)\partial_z\partial} = \frac{MM^0\partial\partial\partial\partial}{(\phi)\partial_z\partial} = \frac{MM^0\partial\partial\partial\partial}{(\phi)\partial_z\partial} \\
& \left\{ \left( \frac{\partial\partial^0}{1-i^2\eta - 1-i^2S} \right) \left( \frac{z(\partial\partial^0)}{1-i^2\eta - 1-i^2S} \partial\partial\theta + \frac{z(\partial\partial^0)}{(1-i^2\eta - i^2S)} - \right) + \right. \\
& \left. \left( \frac{z(\partial\partial^0)}{(1-i^2\eta - 1-i^2S)} - \right) \left( \frac{\partial\partial^0}{1-i^2\eta - 1-i^2S} \partial\partial\theta - \frac{\partial\partial^0}{1-i^2\eta - i^2S} \right) \right\} \sum^{(\partial\partial)N\exists i} = \frac{\partial\partial^0\partial\partial\partial\partial}{(\phi)\partial_z\partial}
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \theta^{DW}} &= - \sum_{t \in N(DW)} \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right)^2 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \theta^{WD}} &= 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \alpha_j^D} &= \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \left( \frac{-\varphi_j(t-1)}{\sigma^{DW}} \right) \right. \\
&\quad \left. + \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma^{DW}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \alpha_j^W} &= \sum_{t \in N(DW)} \left( \frac{-\varphi_j(t)}{\sigma^{DW}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \sigma^{DD}} &= \frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \sigma^{WW}} = \frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \sigma^{WD}} = 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \sigma^{DW}} &= \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \left( -\frac{(S_{i,t-1} - \mu_{t-1}^D)}{(\sigma^{DW})^2} \right) \right. \\
&\quad \left. + \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma^{DW})^2} + \theta^{DW} \frac{(S_{i,t-1} - \mu_{t-1}^D)}{(\sigma^{DW})^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \theta^{WD}} &= - \sum_{t \in N(WD)} \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right)^2 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \alpha_j^D} &= \sum_{t \in N(WD)} \left( \frac{-\varphi_j(t)}{\sigma^{WD}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right)
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \alpha_j^W} &= \sum_{t \in N(WD)} \left\{ \left( \frac{\varphi_j(t-1) \theta^{WD}}{\sigma^{WD}} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \right. \\
&\quad \left. + \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \left( \frac{-\varphi_j(t-1)}{\sigma^{WD}} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \sigma^{DD}} &= \frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \sigma^{WW}} = \frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \sigma^{DW}} = 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \sigma^{WD}} &= \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \left( -\frac{(S_{i,t-1} - \mu_{t-1}^W)}{(\sigma^{WD})^2} \right) \right. \\
&\quad \left. + \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma^{WD})^2} + \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WD})^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \right\}
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \alpha_k^D} &= - \left[ \sum_{t \in N(DD)} \left( \frac{-\varphi_j(t)}{\sigma^{DD}} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma^{DD}} \right) \left( \frac{-\varphi_k(t)}{\sigma^{DD}} + \frac{\theta^{DD} \varphi_k(t-1)}{\sigma^{DD}} \right) \right. \\
&\quad + \sum_{t \in N(DW)} \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma^{DW}} \right) \left( \frac{\theta^{DW} \varphi_k(t-1)}{\sigma^{DW}} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{-\varphi_j(t)}{\sigma^{WD}} \right) \left( \frac{-\varphi_k(t)}{\sigma^{WD}} \right) \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \alpha_k^W} &= - \left[ \sum_{t \in N(DW)} \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma^{DW}} \right) \left( \frac{-\varphi_k(t)}{\sigma^{DW}} \right) \right. \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{-\varphi_j(t)}{\sigma^{WD}} \right) \left( \frac{\theta^{WD} \varphi_k(t-1)}{\sigma^{WD}} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \sigma^{DD}} &= - \left[ \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DD}} \right) \right. \right. \\
&\quad \left. \left( \frac{\varphi_j(t)}{(\sigma^{DD})^2} - \frac{\theta^{DD} \varphi_j(t-1)}{(\sigma^{DD})^2} \right) \right. \\
&\quad \left. + \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma^{DD})^2} + \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DD})^2} \right) \left( \frac{-\varphi_j(t)}{\sigma^{DD}} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma^{DD}} \right) \right\} \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \sigma^{WW}} &= 0
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \sigma^{DW}} &= - \left[ \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \left( \frac{-\theta^{DW} \varphi_j(t-1)}{(\sigma^{DW})^2} \right) \right. \right. \\
&\quad \left. \left. + \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma^{DW}} \right) \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma^{DW})^2} + \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DW})^2} \right) \right\} \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \sigma^{WD}} &= - \left[ \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \left( \frac{\varphi_j(t)}{(\sigma^{WD})^2} \right) \right. \right. \\
&\quad \left. \left. + \left( \frac{-\varphi_j(t)}{\sigma^{WD}} \right) \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma^{WD})^2} + \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WD})^2} \right) \right\} \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \alpha_k^W} &= - \left[ \sum_{t \in N(WW)} \left( \frac{-\varphi_k(t)}{\sigma^{WW}} + \frac{\varphi_k(t-1)\theta^{WW}}{\sigma^{WW}} \right) \right. \\
&\quad \left( \frac{-\varphi_j(t)}{\sigma^{WW}} + \frac{\theta^{WW}\varphi_j(t-1)}{\sigma^{WW}} \right) \\
&\quad + \sum_{t \in N(DW)} \left( \frac{-\varphi_k(t)}{\sigma^{WW}} \right) \left( \frac{-\varphi_j(t)}{\sigma^{WW}} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{\theta^{WD}\varphi_k(t-1)}{\sigma^{WD}} \right) \left( \frac{\theta^{WD}\varphi_j(t-1)}{\sigma^{WD}} \right) \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \sigma^{DD}} &= 0
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \sigma^{WW}} &= - \left[ \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right) \right. \right. \\
&\quad \left( \frac{\varphi_j(t)}{(\sigma^{WW})^2} - \frac{\theta^{WW}\varphi_j(t-1)}{(\sigma^{WW})^2} \right) + \left( \frac{-(S_{i,t} - \mu_t^W)}{(\sigma^{WW})^2} + \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WW})^2} \right) \\
&\quad \left. \left( \frac{-\varphi_j(t)}{\sigma^{WW}} + \frac{\theta^{WW}\varphi_j(t-1)}{\sigma^{WW}} \right) \right\} \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \sigma^{DW}} &= - \left[ \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \left( \frac{\varphi_j(t)}{(\sigma^{DW})^2} \right) \right. \\
&\quad \left. + \left( \frac{-(S_{i,t} - \mu_t^W)}{(\sigma^{DW})^2} + \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DW})^2} \right) \left( \frac{-\varphi_j(t)}{\sigma^{DW}} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \sigma^{WD}} &= - \left[ \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \left( \frac{-\theta^{WD}\varphi_j(t-1)}{(\sigma^{WD})^2} \right) \right. \\
&\quad \left. + \left( \frac{-(S_{i,t} - \mu_t^D)}{(\sigma^{WD})^2} + \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WD})^2} \right) \left( \frac{\theta^{WD}\varphi_j(t-1)}{\sigma^{WD}} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \sigma^{DD} \partial \sigma^{DD}} &= \frac{C(DD)}{(\sigma^{DD})^2} - \left[ \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{DD}} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DD}} \right) \right. \right. \\
&\quad \left( 2 \frac{S_{i,t} - \mu_t^D}{(\sigma^{DD})^3} - 2\theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DD})^3} \right) \\
&\quad \left. \left. + \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma^{DD})^2} + \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DD})^2} \right)^2 \right\} \right]
\end{aligned}$$

$$\frac{\partial^2 \ell(\psi)}{\partial \sigma^{DD} \partial \sigma^{WW}} = \frac{\partial^2 \ell(\psi)}{\partial \sigma^{DD} \partial \sigma^{DW}} = \frac{\partial^2 \ell(\psi)}{\partial \sigma^{DD} \partial \sigma^{WD}} = 0$$

$$\begin{aligned} \frac{\partial^2 \ell(\psi)}{\partial \sigma^{WW} \partial \sigma^{WW}} &= \frac{C(WW)}{(\sigma^{WW})^2} - \left[ \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{WW}} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WW}} \right) \right. \right. \\ &\quad \left. \left( 2 \frac{S_{i,t} - \mu_t^W}{(\sigma^{WW})^3} - 2\theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WW})^3} \right) \right. \\ &\quad \left. \left. + \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma^{WW})^2} + \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WW})^2} \right)^2 \right\} \right] \end{aligned}$$

$$\frac{\partial^2 \ell(\psi)}{\partial \sigma^{WW} \partial \sigma^{DW}} = \frac{\partial^2 \ell(\psi)}{\partial \sigma^{WW} \partial \sigma^{WD}} = 0$$

$$\begin{aligned} \frac{\partial^2 \ell(\psi)}{\partial \sigma^{DW} \partial \sigma^{DW}} &= \frac{C(DW)}{(\sigma^{DW})^2} - \left[ \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma^{DW}} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma^{DW}} \right) \right. \right. \\ &\quad \left. \left( 2 \frac{S_{i,t} - \mu_t^W}{(\sigma^{DW})^3} - 2\theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DW})^3} \right) \right. \\ &\quad \left. \left. + \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma^{DW})^2} + \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma^{DW})^2} \right)^2 \right\} \right] \end{aligned}$$

$$\frac{\partial^2 \ell(\psi)}{\partial \sigma^{DW} \partial \sigma^{WD}} = 0$$

$$\begin{aligned} \frac{\partial^2 \ell(\psi)}{\partial \sigma^{WD} \partial \sigma^{WD}} &= \frac{C(WD)}{(\sigma^{WD})^2} - \left[ \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma^{WD}} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma^{WD}} \right) \right. \right. \\ &\quad \left. \left( 2 \frac{S_{i,t} - \mu_t^D}{(\sigma^{WD})^3} - 2\theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WD})^3} \right) \right. \\ &\quad \left. \left. + \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma^{WD})^2} + \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma^{WD})^2} \right)^2 \right\} \right] \end{aligned}$$

The following algorithm is used to estimate the parameters.

#### Algorithm

**Step 1:** Estimate initial  $\hat{\mu}_t$  by approximating by its Fourier series representation and estimating the parameters  $\alpha_i$  by the method mentioned in the previous models.

**Step 2:** Estimate initial  $\hat{\theta}^{DD}, \hat{\theta}^{WW}, \hat{\theta}^{DW}$  and  $\hat{\theta}^{WD}$  using the following formula:

$$\hat{\theta}^{DD} = \frac{\sum_{t \in N(DD)} (S_{i,t} - \hat{\mu}_t^D)(S_{i,t-1} - \hat{\mu}_{t-1}^D)}{\sum_{t \in N(DD)} (S_{i,t-1} - \hat{\mu}_{t-1}^D)^2}$$

Similarly for  $\hat{\theta}^{WW}, \hat{\theta}^{DW}$  and  $\hat{\theta}^{WD}$ .

**Step 3:** Estimate initial  $\hat{\sigma}^{DD}, \hat{\sigma}^{WW}, \hat{\sigma}^{DW}$  and  $\hat{\sigma}^{WD}$  using the following formula:

$$\hat{\sigma}^{DD} = \left[ \frac{1}{C(DD) - 1} \sum_{t \in N(DD)} (S_{i,t} - \hat{\mu}_t^D - \hat{\theta}^{DD}(S_{i,t-1} - \hat{\mu}_{t-1}^D))^2 \right]^{\frac{1}{2}}$$

Similarly for  $\hat{\sigma}^{WW}, \hat{\sigma}^{DW}$  and  $\hat{\sigma}^{WD}$ .

**Step 4:** Compute  $f^{(k)}$  and  $F^{(k)}$ , where  $f^{(k)}$  is the vector of first partial derivatives and  $F^{(k)}$  is the matrix of second partial derivatives, computed at the  $k$ th iteration.

**Step 5:** Compute the vector  $\delta^{(k)}$  which is the solution to the system of NP linear equations

$$F^{(k)} \delta^{(k)} = f^{(k)},$$

where NP represents the number of parameters.

**Step 6:** Set  $\beta^{(k+1)} = \beta^{(k)} - \delta^{(k)}$ , where  $\beta^{(k)}$  contains the parameter estimates at the  $k$ th iteration.

**Step 7:** Test for convergence, for example, if the elements of  $f^{(k)}$  are sufficiently close to zero. If the convergence criterion is met then stop, otherwise increase  $k$  by 1 and return to step 4.

The cross-correlation matrix,  $\hat{\Sigma}$ , has elements given by:

$$R_{jk} = \frac{\frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} e_{i,t}^{(k)} - \frac{1}{T^2} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)}}{\left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(j)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \right)^2 \right]^{\frac{1}{2}} \left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(k)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)} \right)^2 \right]^{\frac{1}{2}}}$$

where  $e_{i,t}^{(j)}$  denotes the residual time series of variable  $j$ ,  $j = 1, 2, \dots, NV$

and  $e_{i,t}^{(k)}$  denotes the residual time series of variable  $k$ ,  $k = 1, 2, \dots, NV$ .

#### (d) Model Selection

The order of the autoregressive process is selected in the same way as in the previous models, as is the order of the Fourier series approximation.

**Model 4: Multivariate model for climate data****(a) Notation**

Partition the year into  $NT(= 365)$  equal intervals, denoted by  $t = 1, 2, \dots, NT$ .

$NV$  is the number of variables.

$NY$  is the number of years observed.

$W$  represents the occurrence of rain.

$D$  represents the non-occurrence of rain.

$DD$  represents the sequence when day  $t - 1$  was dry and day  $t$  was dry.

$WW$  represents the sequence when day  $t - 1$  was wet and day  $t$  was wet.

$DW$  represents the sequence when day  $t - 1$  was dry and day  $t$  was wet.

$WD$  represents the sequence when day  $t - 1$  was wet and day  $t$  was dry.

$T$  represents the total number of observations, i.e.  $NT NY$ .

$N(DD)$  is the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was dry,  
 $t = 1, 2, \dots, T$ .

$N(WW)$  is the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was wet.

$N(DW)$  is the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was dry.

$N(WD)$  is the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was wet.

$Y_{i,t}$  is the precipitation amount on period  $t$  of year  $i$ ,  $i = 1, 2, \dots, NY$ .

$S_{i,t}$  is the generic name for the observation at time  $t$  of the  $i$ th year.

$\mu_t^D$  is the generic name for the mean for a dry day on period  $t$ .

$\mu_t^W$  is the generic name for the mean for a wet day on period  $t$ .

$\sigma_t^D$  is the generic name for the standard deviation for a dry day on period  $t$ .

$\sigma_t^W$  is the generic name for the standard deviation for a wet day on period  $t$ .

$\theta$  is the coefficient of the AR(1) process.

**(b) Model and assumptions**

Following the procedure suggested by Richardson (1981), the time series  $S_{i,t}$  is reduced to a time series of residual elements,  $\chi_{i,t}$ , by removing the periodic means and standard

deviations, i.e.

$$\chi_{i,t} = \frac{S_{i,t} - \mu_t}{\sigma_t}.$$

This standardization leads to a time series for each variable that is stationary in the mean and standard deviation with mean zero and standard deviation of unity.

The model proposed is assumed to follow an AR(1) process, i.e.

$$\chi_{i,t} = \theta \chi_{i,t-1} + e_{i,t}$$

where  $e_{i,t} \sim NID(0,1)$   $i = 1, 2, \dots, NY$ ;  $t = 1, 2, \dots, NT$ .

The model that incorporates the different rain sequences is then given by:

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \quad \text{if day } t-1 \text{ was dry and day } t \text{ was dry.}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \quad \text{if day } t-1 \text{ was wet and day } t \text{ was wet.}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \quad \text{if day } t-1 \text{ was dry and day } t \text{ was wet.}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \quad \text{if day } t-1 \text{ was wet and day } t \text{ was dry.}$$

Here again the mean and standard deviation functions,  $\mu_t$  and  $\sigma_t$ , are approximated by their respective truncated Fourier representation, i.e.

$$\left. \begin{aligned} \mu_t^D &= \sum_{i=1}^L \alpha_i^D \varphi_i(t) \\ \sigma_t^D &= \sum_{i=1}^L \xi_i^D \varphi_i(t) \end{aligned} \right\} \quad \text{if } t \text{ dry}$$

$$\left. \begin{aligned} \mu_t^W &= \sum_{i=1}^L \alpha_i^W \varphi_i(t) \\ \sigma_t^W &= \sum_{i=1}^L \xi_i^W \varphi_i(t) \end{aligned} \right\} \quad \text{if } t \text{ wet}$$

where  $\varphi_i(t)$  is defined as before and  $L$  is the order of the Fourier series approximation.

### (c) Estimation

Since  $e_{i,t} \sim NID(0,1)$ , the density function of  $e_{i,t}$  is given by

$$f(e_{i,t}) = \frac{1}{\sqrt{2\pi}} \exp\left(-\frac{1}{2}e_{i,t}^2\right).$$

The joint likelihood function, conditioned on the four different sequences is given by:

$$\begin{aligned} L(\psi) &= L(\alpha_j^D, \alpha_j^W, \xi_j^D, \xi_j^W, \theta; e_{i,t}) \\ &= \prod_{t \in N(DD)} f(e_{i,t}|DD) \prod_{t \in N(WW)} f(e_{i,t}|WW) \\ &\quad \prod_{t \in N(DW)} f(e_{i,t}|DW) \prod_{t \in N(WD)} f(e_{i,t}|WD) \end{aligned}$$

where  $f(e_{i,t}|DD)$  represents the density function of  $e_{i,t}$  given that the sequence  $DD$  has been observed, and similarly for the others.

$$\begin{aligned} &= \left(\frac{1}{\sqrt{2\pi}}\right)^T \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(DD)} (e_{i,t}|DD)^2 + \sum_{t \in N(WW)} (e_{i,t}|WW)^2 \right. \right. \\ &\quad \left. \left. + \sum_{t \in N(DW)} (e_{i,t}|DW)^2 + \sum_{t \in N(WD)} (e_{i,t}|WD)^2 \right] \right\}. \end{aligned}$$

One now makes the following transformation:

$$e_{i,t} = \frac{S_{i,t} - \mu_t}{\sigma_t} - \theta \frac{S_{i,t-1} - \mu_{t-1}}{\sigma_{t-1}}.$$

The Jacobian of the transformation is given by

$$\begin{aligned} \left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right| &= \begin{vmatrix} 1/\sigma_1 & 0 & & & 0 \\ -\theta/\sigma_1 & 1/\sigma_2 & 0 & \dots & 0 \\ \vdots & & & & \vdots \\ 0 & \dots & \dots & -\theta/\sigma_{364} & 1/\sigma_{365} \\ \vdots & & & & \vdots \\ 1/\sigma_1 & 0 & \dots & \dots & 0 \\ -\theta/\sigma_1 & 1/\sigma_2 & 0 & \dots & 0 \\ \vdots & & & & \vdots \\ 0 & \dots & \dots & -\theta/\sigma_{364} & 1/\sigma_{365} \end{vmatrix} \\ &= \prod_{i=1}^{NY} \prod_{t=1}^{NT} \frac{1}{\sigma_t} \end{aligned}$$

Taking into account the conditional sequences imposed on  $e_{i,t}$ , the Jacobian is then given by

$$\left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right| = \prod_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{1}{\sigma_t^D} \prod_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{1}{\sigma_t^W}.$$

The joint probability density function is thus given by:

$$\begin{aligned}
 L(\psi) = & \left( \frac{1}{\sqrt{2\pi}} \right)^T \prod_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{1}{\sigma_t^D} \prod_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{1}{\sigma_t^W} \\
 & \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \right. \right. \\
 & + \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \\
 & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \\
 & \left. \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \right] \right\}
 \end{aligned}$$

and the log-likelihood is given by:

$$\begin{aligned}
 \ell(\psi) = & -\frac{T}{2} \log(2\pi) - \sum_{\substack{t \in N(DD) \\ t \in N(WD)}} \log(\sigma_t^D) - \sum_{\substack{t \in N(WW) \\ t \in N(DW)}} \log(\sigma_t^W) \\
 & - \frac{1}{2} \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \right. \\
 & + \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \\
 & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \\
 & \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \right].
 \end{aligned}$$

Maximum likelihood estimates can be obtained by minimising  $\ell(\psi)$  and this is achieved by setting its partial derivatives with respect to the parameters equal to zero.

The first partial derivatives with respect to the parameters are given by:

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \theta} = & \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
& + \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \\
& + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
& + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right).
\end{aligned}$$

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \alpha_j^D} = & - \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right. \\
& + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \\
& \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \alpha_j^W} = & - \left[ \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right. \\
& + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \\
& \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \xi_j^D} = & - \sum_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{\varphi_j(t)}{\sigma_t^D} - \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \\
& \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma_t^D)^2} \varphi_j(t) + \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \varphi_j(t-1) \right) \\
& + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \varphi_j(t-1) \right) \\
& \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( -\frac{(S_{i,t} - \mu_t^D)}{(\sigma_t^D)^2} \varphi_j(t) \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \xi_j^W} = & - \sum_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{\varphi_j(t)}{\sigma_t^W} - \left[ \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \\
& \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma_t^W)^2} \varphi_j(t) + \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \varphi_j(t-1) \right) \\
& + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( -\frac{(S_{i,t} - \mu_t^W)}{(\sigma_t^W)^2} \varphi_j(t) \right) \\
& \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \varphi_j(t-1) \right) \right]
\end{aligned}$$

The equations obtained when the partial derivatives are set to zero can be solved using the Newton-Raphson iteration method. For this, the second partial derivatives are required. These are given by:

$$\frac{\partial^2 \ell(\psi)}{\partial \theta \partial \theta} = - \left[ \sum_{\substack{t \in N(DD) \\ t \in N(DW)}} \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 + \sum_{\substack{t \in N(WW) \\ t \in N(WD)}} \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \right]$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta \partial \alpha_j^D} = & \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right. \\
& + \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left. \right\} \\
& + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right. \\
& + \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left. \right\} \\
& + \sum_{t \in N(WD)} \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta \partial \alpha_j^W} &= \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right. \\
&\quad \left. + \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\} \\
&\quad + \sum_{t \in N(DW)} \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
&\quad + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right. \\
&\quad \left. + \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\}
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta \partial \xi_j^D} &= \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right. \\
&\quad \left. + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right\} \\
&\quad + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right. \\
&\quad \left. + \left( \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right\} \\
&\quad + \sum_{t \in N(WD)} \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta \partial \xi_j^W} &= \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right. \\
&\quad \left. + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\} \\
&\quad + \sum_{t \in N(DW)} \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
&\quad + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right. \\
&\quad \left. + \left( \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\}
\end{aligned}$$

$$\begin{aligned} \frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \alpha_k^D} = & - \left[ \sum_{t \in N(DD)} \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_k(t)}{\sigma_t^D} + \frac{\theta \varphi_k(t-1)}{\sigma_{t-1}^D} \right) \right. \\ & + \sum_{t \in N(DW)} \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{\theta \varphi_k(t-1)}{\sigma_{t-1}^D} \right) \\ & \left. + \sum_{t \in N(WD)} \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \left( \frac{-\varphi_k(t)}{\sigma_t^D} \right) \right] \end{aligned}$$

$$\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \alpha_k^W} = - \left[ \sum_{t \in N(DW)} \left( \frac{-\varphi_k(t)}{\sigma_t^W} \right) \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) + \sum_{t \in N(WD)} \left( \frac{\theta \varphi_k(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right]$$

$$\begin{aligned} \frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \xi_k^D} = & - \left[ \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \right. \\ & \left. \left( \frac{\varphi_k(t) \varphi_j(t)}{(\sigma_t^D)^2} - \frac{\theta \varphi_k(t-1) \varphi_j(t-1)}{(\sigma_{t-1}^D)^2} \right) \right. \\ & + \left. \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right\} \\ & + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\theta \varphi_k(t-1) \varphi_j(t-1)}{(\sigma_{t-1}^D)^2} \right) \right. \\ & + \left. \left( \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right\} \\ & + \sum_{t \in N(WD)} \left\{ \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right. \\ & + \left. \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^D)^2} \right) \right\} \Bigg] \\ \frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \xi_k^W} = & - \left[ \sum_{t \in N(DW)} \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right. \\ & \left. + \sum_{t \in N(WD)} \left( \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right] \end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \alpha_k^W} &= - \left[ \sum_{t \in N(WW)} \left( \frac{-\varphi_k(t)}{\sigma_t^W} + \frac{\theta \varphi_k(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right. \\
&\quad + \sum_{t \in N(DW)} \left( \frac{-\varphi_k(t)}{\sigma_t^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{\theta \varphi_k(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \xi_k^D} &= - \left[ \sum_{t \in N(DW)} \left( \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \right. \\
&\quad \left. + \sum_{t \in N(WD)} \left( -\varphi_k(t) \frac{S_{i,t-1} - \mu_t^D}{(\sigma_t^D)^2} \right) \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \xi_k^W} &= - \left[ \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \right. \\
&\quad \left( \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^W)^2} - \frac{\theta \varphi_j(t-1) \varphi_k(t-1)}{(\sigma_{t-1}^W)^2} \right) \\
&\quad + \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \Big\} \\
&\quad + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^W)^2} \right) \right. \\
&\quad + \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \Big\} \\
&\quad + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\theta \varphi_j(t-1) \varphi_k(t-1)}{(\sigma_{t-1}^W)^2} \right) \right. \\
&\quad \left. + \left( \frac{\theta \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right\} \Big]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \xi_j^D \partial \xi_k^D} &= \sum_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^D)^2} - \left[ \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \right. \\
&\quad \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^3} - 2\theta \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^3} \right) \\
&\quad + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \\
&\quad \left. \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right\} \\
&\quad + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \\
&\quad \left( -2\theta \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^3} \right) \\
&\quad + \left( \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right)^2 \varphi_j(t-1) \varphi_k(t-1) \left. \right\} \\
&\quad + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^3} \right) \right. \\
&\quad + \left. \left( \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right)^2 \varphi_j(t) \varphi_k(t) \right\} \left. \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \xi_j^D \partial \xi_k^W} &= - \left[ \sum_{t \in N(DW)} \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left( \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right. \\
&\quad + \sum_{t \in N(WD)} \left( \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \left. \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \xi_j^W \partial \xi_k^W} = & \sum_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^W)^2} - \left[ \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \right. \\
& \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^3} - 2\theta \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^3} \right) \\
& + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \\
& \left. \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right\} \\
& + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^3} \right) \right. \\
& + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left. \right\} \\
& + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \\
& \left( -2\theta \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^3} \right) \\
& + \left( \theta \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( \theta \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left. \right\} \left. \right]
\end{aligned}$$

The following algorithm is used to estimate the parameters.

#### Algorithm

**Step 1:** Estimate initial  $\hat{\mu}_t$  and  $\hat{\sigma}_t$  by aproximating by its Fourier series representation and estimating the parameters  $\alpha_i$  and  $\xi_i$  by the method mentioned in the previous models.

**Step 2:** Estimate initial  $\hat{\theta}$  using the following formula:

$$\hat{\theta} = \frac{T \sum_{t=2}^T (S_{i,t} - \hat{\mu}_t)(S_{i,t-1} - \hat{\mu}_{t-1})}{(T-1) \sum_{t=2}^{T+1} (S_{i,t-1} - \hat{\mu}_{t-1})^2}$$

where  $\hat{\mu}_t$  depends on the status of day  $t$  and  $\hat{\mu}_{t-1}$  depends on the status of day  $t-1$ .

**Step 3:** Compute  $f^{(k)}$  and  $F^{(k)}$  where  $f^{(k)}$  is the vector of first partial derivatives and  $F^{(k)}$  is the matrix of second partial derivatives, computed at the  $k$ th iteration.

**Step 4:** Compute the vector  $\delta^{(k)}$  which is the solution to the system of NP linear equations

$$F^{(k)} \delta^{(k)} = f^{(k)}$$

where NP is the number of parameters in the model.

**Step 5:** Set  $\beta^{(k+1)} = \beta^{(k)} - \delta^{(k)}$ , where  $\beta^{(k)}$  contains the parameter estimates computed at the  $k$ th iteration.

**Step 6** Test for convergence, for example, if the elements of  $f^{(k)}$  are sufficiently close to zero. If the convergence criterion is met then stop, otherwise increase  $k$  by 1 and return to step 3.

The cross-correlation matrix,  $\hat{\Sigma}$ , has elements given by

$$R_{jk} = \frac{\frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} e_{i,t}^{(k)} - \frac{1}{T^2} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)}}{\left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(j)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \right)^2 \right]^{\frac{1}{2}} \left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(k)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)} \right)^2 \right]^{\frac{1}{2}}}$$

where

$e_{i,t}^{(j)}$  denotes the residual time series of variable  $j$ ;  $j = 1, 2, \dots, NV$

and

$e_{i,t}^{(k)}$  denotes the residual time series of variable  $k$ ;  $k = 1, 2, \dots, NV$ .

#### (d) Model Selection

The order of the autoregressive process is chosen in the same way as in the previous model as is the order of the Fourier series approximation.

**Model 5: Multivariate model for climate data****(a) Notation**

Partition the year into  $NT(= 365)$  equal intervals, denoted by  $t = 1, 2, \dots, NT$ .

$NV$  is the number of variables.

$NY$  is the number of years observed.

$W$  represents the occurrence of rain.

$D$  represents the non-occurrence of rain.

$DD$  represents the sequence when day  $t - 1$  was dry and day  $t$  was dry.

$WW$  represents the sequence when day  $t - 1$  was wet and day  $t$  was wet.

$DW$  represents the sequence when day  $t - 1$  was dry and day  $t$  was wet.

$WD$  represents the sequence when day  $t - 1$  was wet and day  $t$  was dry.

$T$  represents the total number of observations, i.e.  $NT NY$ .

$N(DD)$  is the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was dry,  $t = 1, 2, \dots, T$ .

$N(WW)$  is the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was wet.

$N(DW)$  is the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was dry.

$N(WD)$  is the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was wet.

$Y_{i,t}$  is the precipitation amount on period  $t$  of year  $i$ ,  $i = 1, 2, \dots, NY$ .

$S_{i,t}$  is the generic name for the observation at time  $t$  of the  $i$ th year.

$\mu_t^D$  is the generic name for the mean for a dry day on period  $t$ .

$\mu_t^W$  is the generic name for the mean for a wet day on period  $t$ .

$\sigma_t^D$  is the generic name for the standard deviation for a dry day on period  $t$ .

$\sigma_t^W$  is the generic name for the standard deviation for a wet day on period  $t$ .

$\theta^{DD}$  is the coefficient of the AR(1) process, given sequence  $DD$ .

$\theta^{WW}$  is the coefficient of the AR(1) process given sequence  $WW$ .

$\theta^{DW}$  is the coefficient of the AR(1) process given sequence  $DW$ .

$\theta^{WD}$  is the coefficient of the AR(1) process given sequence  $WD$ .

**(b) Model and assumptions**

Model 5 is formulated in the same manner as that of Model 4, except that here it is assumed that the coefficient of the AR(1) process,  $\theta$ , varies according to the wet/dry status of the present and previous day.

Therefore, the time series  $S_{i,t}$  is once again reduced to a residual time series  $\chi_{i,t}$  by removing the periodic mean and standard deviation, i.e.

$$\chi_{i,t} = \frac{S_{i,t} - \mu_t}{\sigma_t}.$$

Assume that this residual time series follows an AR(1) process, i.e.

$$\chi_{i,t} = \theta \chi_{i,t-1} + e_{i,t}$$

where  $e_{i,t} \sim NID(0,1)$   $i = 1, 2, \dots, NY$ ;  $t = 1, 2, \dots, NT$ .

The model incorporating the different wet/dry sequences is given by:

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \quad \text{if day } t-1 \text{ was dry and day } t \text{ was dry.}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \quad \text{if day } t-1 \text{ was wet and day } t \text{ was wet.}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \quad \text{if day } t-1 \text{ was dry and day } t \text{ was wet.}$$

or

$$e_{i,t} = \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \quad \text{if day } t-1 \text{ was wet and day } t \text{ was dry.}$$

The mean and standard deviation functions are approximated by their respective truncated Fourier representation, i.e.

$$\left. \begin{aligned} \mu_t^D &= \sum_{i=1}^L \alpha_i^D \varphi_i(t) \\ \sigma_t^D &= \sum_{i=1}^L \xi_i^D \varphi_i(t) \end{aligned} \right\} \quad \text{if } t \text{ dry}$$

$$\left. \begin{aligned} \mu_t^W &= \sum_{i=1}^L \alpha_i^W \varphi_i(t) \\ \sigma_t^W &= \sum_{i=1}^L \xi_i^W \varphi_i(t) \end{aligned} \right\} \quad \text{if } t \text{ wet}$$

where  $\varphi_i(t)$  is defined as before and  $L$  is the order of the Fourier series approximation.

(c) **Estimation**

The density function of  $e_{i,t}$  is given by

$$f(e_{i,t}) = \frac{1}{\sqrt{2\pi}} \exp\left(-\frac{1}{2}e_{i,t}^2\right)$$

since  $e_{i,t} \sim NID(0, 1)$ .

The joint likelihood function, conditioned on the four different sequences is given by:

$$\begin{aligned} L(\psi) &= L(\alpha_j^D, \alpha_j^W, \xi_j^D, \xi_j^W, \theta^{DD}, \theta^{WW}, \theta^{DW}, \theta^{WD}; e_{i,t}) \\ &= \prod_{t \in N(DD)} f(e_{i,t}|DD) \prod_{t \in N(WW)} f(e_{i,t}|WW) \\ &\quad \prod_{t \in N(DW)} f(e_{i,t}|DW) \prod_{t \in N(WD)} f(e_{i,t}|WD) \\ &= \left(\frac{1}{\sqrt{2\pi}}\right)^T \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(DD)} (e_{i,t}|DD)^2 \right. \right. \\ &\quad + \sum_{t \in N(WW)} (e_{i,t}|WW)^2 + \sum_{t \in N(DW)} (e_{i,t}|DW)^2 \\ &\quad \left. \left. + \sum_{t \in N(WD)} (e_{i,t}|WD)^2 \right] \right\}. \end{aligned}$$

Make the following transformation

$$e_{i,t} = \frac{S_{i,t} - \mu_t}{\sigma_t} - \theta \frac{S_{i,t-1} - \mu_{t-1}}{\sigma_{t-1}}.$$

The Jacobian of the transformation is

$$\begin{aligned} \left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right| &= \begin{vmatrix} 1/\sigma_1 & 0 & & & 0 \\ -\theta/\sigma_1 & 1/\sigma_2 & 0 & \dots & 0 \\ \vdots & & & & \vdots \\ 0 & \dots & 0 & -\theta/\sigma_{364} & 1/\sigma_{365} \\ \vdots & & & & \vdots \\ 1/\sigma_1 & 0 & \dots & \dots & 0 \\ -\theta/\sigma_1 & 1/\sigma_2 & 0 & \dots & 0 \\ \vdots & & & & \vdots \\ 0 & \dots & 0 & -\theta/\sigma_{364} & 1/\sigma_{365} \end{vmatrix} \\ &= \prod_{i=1}^{NY} \prod_{t=1}^{NT} \frac{1}{\sigma_t} \end{aligned}$$

Taking into account the conditional sequences imposed on  $e_{i,t}$ , the Jacobian is then given by

$$\left| \frac{\partial e_{i,t}}{\partial S_{i,p}} \right| = \prod_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{1}{\sigma_t^D} \prod_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{1}{\sigma_t^W}.$$

The joint probability density function is thus given by:

$$\begin{aligned} L(\psi) = & \left( \frac{1}{\sqrt{2\pi}} \right)^T \prod_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{1}{\sigma_t^D} \prod_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{1}{\sigma_t^W} \\ & \exp \left\{ -\frac{1}{2} \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \right. \right. \\ & + \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \\ & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \\ & \left. \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \right] \right\} \end{aligned}$$

and the log-likelihood is given by:

$$\begin{aligned} \ell(\psi) = & -\frac{T}{2} \log(2\pi) - \sum_{\substack{t \in N(DD) \\ t \in N(WD)}} \log(\sigma_t^D) - \sum_{\substack{t \in N(WW) \\ t \in N(DW)}} \log(\sigma_t^W) \\ & - \frac{1}{2} \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \right. \\ & + \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \\ & + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \\ & \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \right]. \end{aligned}$$

To obtain maximum likelihood estimates for the parameters,  $\ell(\psi)$  is minimized. To minimize  $\ell(\psi)$ , its first partial derivatives with respect to the parameters are set to zero.

The first partial derivatives with respect to the parameters are given by

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \theta^{DD}} &= \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
\frac{\partial \ell(\psi)}{\partial \theta^{WW}} &= \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \\
\frac{\partial \ell(\psi)}{\partial \theta^{DW}} &= \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
\frac{\partial \ell(\psi)}{\partial \theta^{WD}} &= \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \\
\\
\frac{\partial \ell(\psi)}{\partial \alpha_j^D} &= - \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right. \\
&\quad + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right] \\
\\
\frac{\partial \ell(\psi)}{\partial \alpha_j^W} &= - \left[ \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \\
&\quad \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta^{WW} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \\
&\quad + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{\theta^{WD} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right] \\
\\
\frac{\partial \ell(\psi)}{\partial \xi_j^D} &= - \sum_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{\varphi_j(t)}{\sigma_t^D} - \left[ \sum_{t \in N(DD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \\
&\quad \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta^{DD} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \\
&\quad + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \theta^{DW} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial \ell(\psi)}{\partial \xi_j^W} = & - \sum_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{\varphi_j(t)}{\sigma_t^W} - \left[ \sum_{t \in N(WW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \\
& \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta^{WW} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \\
& + \sum_{t \in N(DW)} \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \\
& \left. + \sum_{t \in N(WD)} \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \theta^{WD} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right]
\end{aligned}$$

The Newton-Raphson method is used to solve the system of equation. For this the second partial derivatives are required and these are given by

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{DD}} &= - \sum_{t \in N(DD)} \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{WW}} &= \frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{DW}} = \frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \theta^{WD}} = 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \alpha_j^D} &= \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right. \\
& \quad \left. + \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \alpha_j^W} &= \frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \xi_j^W} = 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DD} \partial \xi_j^D} &= \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right. \\
& \quad \left. + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta^{DD} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WW} \partial \theta^{WW}} &= - \sum_{t \in N(WW)} \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WW} \partial \theta^{DW}} &= \frac{\partial^2 \ell(\psi)}{\partial \theta^{WW} \partial \theta^{WD}} = \frac{\partial^2 \ell(\psi)}{\partial \theta^{WW} \partial \alpha_j^D} = 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WW} \partial \alpha_j^W} &= \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right. \\
& \quad \left. + \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta^{WW} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WW} \partial \xi_j^D} &= 0
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WW} \partial \xi_j^W} &= \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \\
&\quad \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \\
&\quad \left. + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta^{WW} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \theta^{DW}} &= - \sum_{t \in N(DW)} \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right)^2 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \theta^{WD}} &= 0 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \alpha_j^D} &= \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right. \\
&\quad \left. + \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \alpha_j^W} &= \sum_{t \in N(DW)} \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \xi_j^D} &= \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right. \\
&\quad \left. + \left( \theta^{DW} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{DW} \partial \xi_j^W} &= \sum_{t \in N(DW)} \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \theta^{WD}} &= - \sum_{t \in N(WD)} \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right)^2 \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \alpha_j^D} &= \sum_{t \in N(WD)} \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \alpha_j^W} &= \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right. \\
&\quad \left. + \left( \frac{\theta^{WD} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\} \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \xi_j^D} &= \sum_{t \in N(WD)} \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \\
\frac{\partial^2 \ell(\psi)}{\partial \theta^{WD} \partial \xi_j^W} &= \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( -\varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right. \\
&\quad \left. + \left( \theta^{WD} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right\}
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \alpha_k^D} &= - \left[ \sum_{t \in N(DD)} \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{-\varphi_k(t)}{\sigma_t^D} + \frac{\theta^{DD} \varphi_k(t-1)}{\sigma_{t-1}^D} \right) \right. \\
&\quad + \sum_{t \in N(DW)} \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \frac{\theta^{DW} \varphi_k(t-1)}{\sigma_{t-1}^D} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{-\varphi_k(t)}{\sigma_t^D} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \alpha_k^W} &= - \left[ \sum_{t \in N(DW)} \left( \frac{-\varphi_k(t)}{\sigma_t^W} \right) \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \right. \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{\theta^{WD} \varphi_k(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \xi_k^D} &= - \left[ \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \right. \\
&\quad \left( \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^D)^2} - \frac{\theta^{DD} \varphi_j(t-1) \varphi_k(t-1)}{(\sigma_{t-1}^D)^2} \right) \\
&\quad + \left( \frac{-\varphi_j(t)}{\sigma_t^D} + \frac{\theta^{DD} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right. \\
&\quad \left. \left. + \theta^{DD} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right\} \\
&\quad + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{-\theta^{DW} \varphi_k(t-1) \varphi_j(t-1)}{(\sigma_{t-1}^D)^2} \right) \right. \\
&\quad + \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( \theta^{DW} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left. \right\} \\
&\quad + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^D)^2} \right) \right. \\
&\quad \left. + \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \right\} \left. \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^D \partial \xi_k^W} &= - \left[ \sum_{t \in N(DW)} \left( \frac{\theta^{DW} \varphi_j(t-1)}{\sigma_{t-1}^D} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \right. \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{-\varphi_j(t)}{\sigma_t^D} \right) \left( \theta^{WD} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \alpha_k^W} &= - \left[ \sum_{t \in N(WW)} \left( \frac{-\varphi_k(t)}{\sigma_t^W} + \frac{\theta^{WW} \varphi_k(t-1)}{\sigma_{t-1}^W} \right) \right. \\
&\quad \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta^{WW} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \\
&\quad + \sum_{t \in N(DW)} \left( \frac{-\varphi_k(t)}{\sigma_t^W} \right) \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{\theta^{WD} \varphi_k(t-1)}{\sigma_{t-1}^W} \right) \left( \frac{\theta^{WD} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \xi_k^D} &= - \left[ \sum_{t \in N(DW)} \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \left( \theta^{DW} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right. \\
&\quad \left. + \sum_{t \in N(WD)} \left( \frac{\theta^{WD} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \alpha_j^W \partial \xi_k^W} &= - \left[ \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \right. \\
&\quad \left( \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^W)^2} - \frac{\theta^{WW} \varphi_j(t-1) \varphi_k(t-1)}{(\sigma_{t-1}^W)^2} \right) \\
&\quad + \left( \frac{-\varphi_j(t)}{\sigma_t^W} + \frac{\theta^{WW} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \\
&\quad \left. \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta^{WW} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right\} \\
&\quad + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^W)^2} \right) \right. \\
&\quad + \left( \frac{-\varphi_j(t)}{\sigma_t^W} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left. \right\} \\
&\quad + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( \frac{-\theta^{WD} \varphi_j(t-1) \varphi_k(t-1)}{(\sigma_{t-1}^W)^2} \right) \right. \\
&\quad \left. + \left( \frac{\theta^{WD} \varphi_j(t-1)}{\sigma_{t-1}^W} \right) \left( \theta^{WD} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right\} \left. \right]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \xi_j^D \partial \xi_k^D} = & \sum_{\substack{t \in N(DD) \\ t \in N(WD)}} \frac{\varphi_k(t) \varphi_j(t)}{(\sigma_t^D)^2} - \left[ \sum_{t \in N(DD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{DD} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \right. \\
& \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^3} - 2\theta^{DD} \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^3} \right) \\
& + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta^{DD} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \\
& \left. \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} + \theta^{DD} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \right\} \\
& + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{DW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \right. \\
& \left( -2\theta^{DW} \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^3} \right) \\
& + \left( \theta^{DW} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( \theta^{DW} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \Big\} \\
& + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^3} \right) \right. \\
& + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \Big\} \Big]
\end{aligned}$$

$$\begin{aligned}
\frac{\partial^2 \ell(\psi)}{\partial \xi_j^D \partial \xi_k^W} &= - \left[ \sum_{t \in N(DW)} \left( \theta^{DW} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^D}{(\sigma_{t-1}^D)^2} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \right. \\
&\quad \left. + \sum_{t \in N(WD)} \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^D}{(\sigma_t^D)^2} \right) \left( \theta^{WD} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right] \\
\frac{\partial^2 \ell(\psi)}{\partial \xi_j^W \partial \xi_k^W} &= \sum_{\substack{t \in N(WW) \\ t \in N(DW)}} \frac{\varphi_j(t) \varphi_k(t)}{(\sigma_t^W)^2} \\
&\quad - \left[ \sum_{t \in N(WW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \right. \\
&\quad \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^3} - 2\theta^{WW} \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^3} \right) \\
&\quad \left. + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta^{WW} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right. \\
&\quad \left. \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} + \theta^{WW} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right\} \\
&\quad + \sum_{t \in N(DW)} \left\{ \left( \frac{S_{i,t} - \mu_t^W}{\sigma_t^W} - \theta^{WW} \frac{S_{i,t-1} - \mu_{t-1}^D}{\sigma_{t-1}^D} \right) \left( 2\varphi_j(t) \varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^3} \right) \right. \\
&\quad \left. + \left( -\varphi_j(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \left( -\varphi_k(t) \frac{S_{i,t} - \mu_t^W}{(\sigma_t^W)^2} \right) \right\} \\
&\quad + \sum_{t \in N(WD)} \left\{ \left( \frac{S_{i,t} - \mu_t^D}{\sigma_t^D} - \theta^{WD} \frac{S_{i,t-1} - \mu_{t-1}^W}{\sigma_{t-1}^W} \right) \right. \\
&\quad \left( -2\theta^{WD} \varphi_j(t-1) \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^3} \right) \\
&\quad \left. + \left( \theta^{WD} \varphi_j(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \left( \theta^{WD} \varphi_k(t-1) \frac{S_{i,t-1} - \mu_{t-1}^W}{(\sigma_{t-1}^W)^2} \right) \right\} \Bigg]
\end{aligned}$$

The following algorithm is used to estimate the parameters

#### Algorithm

Step 1: Estimate initial  $\hat{\mu}_t$  and  $\hat{\sigma}_t$  by approximating by its Fourier series representation and estimating the parameters  $\alpha_i$  and  $\xi_i$  by the method mentioned in the previous models.

Step 2: Estimate initial  $\hat{\theta}^{DD}, \hat{\theta}^{WW}, \hat{\theta}^{DW}$  and  $\hat{\theta}^{WD}$  using the following formula:

$$\hat{\theta}^{DD} = \frac{\sum_{t \in N(DD)} (S_{i,t} - \hat{\mu}_t^D)(S_{i,t-1} - \hat{\mu}_{t-1}^D)}{\sum_{t \in N(DD)} (S_{i,t-1} - \hat{\mu}_{t-1}^D)^2}$$

Similarly for  $\hat{\theta}^{WW}, \hat{\theta}^{DW}$  and  $\hat{\theta}^{WD}$ .

Step 3: Compute  $f^{(k)}$  and  $F^{(k)}$ , where  $f^{(k)}$  is the vector of first partial derivatives and  $F^{(k)}$  is the matrix of second partial derivatives, computed at the  $k$ th iteration.

Step 4: Compute the vector  $\delta^{(k)}$  which is the solution to the system of NP linear equations

$$F^{(k)} \delta^{(k)} = f^{(k)}$$

where NP represents the number of parameters.

Step 5: Set  $\beta^{(k+1)} = \beta^{(k)} - \delta^{(k)}$ , where  $\beta^{(k)}$  contains the parameter estimates at the  $k$ th iteration.

Step 6: Test for convergence, for example, if the elements of  $f^{(k)}$  are sufficiently close to zero. If the convergence criterion is met then stop, otherwise increase  $k$  by 1 and return to step 3.

The cross-correlation matrix,  $\hat{\Sigma}$ , has elements given by

$$R_{jk} = \frac{\frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} e_{i,t}^{(k)} - \frac{1}{T^2} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)}}{\left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(j)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \right)^2 \right]^{\frac{1}{2}} \left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(k)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)} \right)^2 \right]^{\frac{1}{2}}}$$

where  $e_{i,t}^{(j)}$  denotes the residual time series of variable  $j$ ,  $j = 1, 2, \dots, NV$ .

and  $e_{i,t}^{(k)}$  denotes the residual time series of variable  $k$ ,  $k = 1, 2, \dots, NV$ .

#### (d) Model Selection

The order of the autoregressive process is selected in the same way as in the previous models as is the order of the Fourier series approximation.

## CHAPTER 4

### MODEL IMPLEMENTATION

This chapter gives details of the implementation of the proposed time series models to describe historical climate series. In particular, the model selection process is described step by step, the parameter estimates are given and the results of tests to check the model assumptions are discussed.

Six stations were chosen for study which broadly represent the various climate regions of South Africa. Table 4.1 lists them together with the years for which simultaneous observations of all climate variables were recorded. The stations marked with an asterisk indicate those stations for which the variable evaporation was not available.

Considerable difficulties were experienced in obtaining suitable data sets for model implementation. This refers to problems in obtaining stations for which all the climate variables of interest are recorded as well as to the quantity and the quality of the available data. Thus, one is restricted by the stations one can fit the climate models to, and the quantity and quality of the historical records determines the performance of the models. As already mentioned, the models are sensitive to “unclean” data records and relatively short historical records lead to three problems. Firstly, the precision of the estimates decreases as a large number of parameters are estimated using very few data values. Secondly, the effective record length for the conditioned estimates is further reduced as the models separate the sequences into wet and dry sequences. Thirdly, the fact that the record length of the stations are quite small, combined with the fact that there are missing observations in the records means that the historical data might not wholly be representative of the long term climate for that particular location.

Since rainfall was considered to be the primary variable and all other variables are conditioned on whether a given day was wet or dry, it was modelled independently of all other variables.

**Simple Markov chain to describe the occurrence of wet and dry sequences of days.**

The logit transformation of the probabilities  $\pi(t)$ ,  $t = 1, 2, \dots, NT$  is given by

$$\lambda(t) = \log \left( \frac{\pi(t)}{1 - \pi(t)} \right)$$

**Table 4.1 Climate Stations**

Station	Province	Years available
Elsenburg	Cape	1979–1984
Kakamas	Cape	1975–1986
Middelburg	Cape	1977–1986
Nelspruit*	Transvaal	1981–1987
Cedara*	Natal	1980–1989
Hoopstad*	Orange Free State	1981–1989

where  $\lambda(t)$  is represented by a Fourier series approximation, i.e.

$$\lambda(t) = \sum_{i=1}^L \gamma_i \phi_i(t), \quad t = 1, 2, \dots, NT$$

and  $\phi_i(t)$  is defined as in Chapter 3.

The parameters  $\gamma_i$  have to be estimated for the probability that a wet day is preceded by a wet day ( $P(R|R)$ ) and for the probability that a wet day is preceded by a dry day ( $P(R|\bar{R})$ ). For each of the probabilities the order of the Fourier series approximation,  $L$ , has to be selected.

The selection of the appropriate  $L$  was based on Akaike's Information Criterion, where

$$AIC = -\ell(\gamma; M(t)) + L$$

where  $\ell(\gamma; M(t))$  is the log likelihood function of a particular model. The criterion is computed for  $L = 1, 3, 5, \dots$  and the model which leads to the smallest value of the criterion is selected.

Table 4.2 gives the optimal number of parameters for  $P(R|R)$  and  $P(R|\bar{R})$ . The values of  $L$  for  $P(R|R)$  and  $P(R|\bar{R})$  ranged between 1 and 3 and between 1 and 5 respectively, with modes 3 and 5. A choice of 3 parameters for both models was decided upon for the following reasons. Firstly, the method of model selection employed here is less stringent than conventional tests of hypotheses, and therefore generally leads to a selection of more parameters. Thus a choice of 3 parameters would be preferable to 5. Secondly, the length of the historical record plays a role in determining  $\alpha$  and it must be kept in mind

that the results here have been obtained with a relatively small data set, thus making the selection of fewer parameters inevitable. Zucchini and Adamson (1984a) chose 5 parameters for both models, but their data sets (typically 40 years) were large enough to warrant that number of parameters.

**TABLE 4.2 Optimal number of parameters to estimate  $P(R|R)$  and  $P(R|\bar{R})$**

Station	Model	
	$P(R R)$	$P(R \bar{R})$
Elsenburg	3	5
Kakamas	—	1
Middelburg	1	3
Nelspruit	3	5
Cedara	3	5
Hoopstad	1	3

The station Kakamas presented a problem in obtaining convergence when estimating the parameters for  $P(R|R)$ . This can be explained by the rare occurrence of rainfall, and in particular that of consecutive days of rainfall in Kakamas. Moreover, the few years of records available for estimation intensify this problem. That is, when preparing the array  $NRR(t)$  required for parameter estimation, where  $NRR(t)$  represents the number of times it was wet in period  $t - 1$  and wet in period  $t$ , most of the entries are zero and therefore there are very few values on which to compute parameter estimates leading to difficulties in achieving convergence.

Zucchini and Adamson (1984b) have computed parameter estimates for Kakamas, and as rainfall is modelled independently of the other climate variables, these estimates were used.

It is important to note that the readings were recorded by multiplying each value by ten, i.e. a record of 10.2 is given as 102. This convention was used throughout the study and applied to all results given in this report with the exceptions indicated below. This does not affect the generation of climate sequences which can be easily converted to the original

units by dividing by ten. The only exceptions to this were the variables wind run, maximum humidity and minimum humidity for the stations Nelspruit, Cedara and Hoopstad.

The parameter estimates for the probability that a wet day follows a wet day and that a wet day follows a dry day are given in Table 4.3.

#### **The distribution for rainfall on days when rain occurs.**

The mean rainfall per rainy day in period  $t, \mu(t)$ , can be approximated by its truncated Fourier series representation

$$\mu(t) = \sum_{i=1}^L \mu_i \phi_i(t), \quad t = 1, 2, \dots, NT$$

where  $\phi_i(t)$  is defined as in Chapter 3.

The parameters  $\mu_i$  need to be estimated and the order of the Fourier series approximation selected. A 3-term Fourier series approximation was chosen following arguments similar to those in the previous section.

Table 4.4 shows the parameter estimates for mean rainfall and the estimate for the coefficient of variation. It is sometimes easier to work with the Fourier series coefficients in their polar form, therefore the amplitude and phase representation of the mean rainfall is also given. From these parameter estimates, parameters of the corresponding Weibull distribution can then be estimated by the method of moments (see Zucchini and Adamson 1984a)).

### **MODEL FOR CLIMATE SEQUENCES**

#### **Transforming the data set**

Preliminary work carried out to assess the feasibility of modelling climate on a daily basis highlighted some weaknesses in the models. Firstly, although the models satisfactorily preserved the mean and standard deviation, they failed to preserve the extreme values. This problem arises because some climate variables lie within permissible boundaries with some variables having a high frequency of values near or on an upper or lower limit so that it is expected that simulated sequences will occasionally have values that exceed these boundaries. Secondly, some minimum temperature values were slightly higher than the corresponding maximum temperature value. The same occurred with the humidity variable.

The problem that generated values fall outside their respective admissible range can of course be easily overcome by simply setting the generated values to the appropriate boundary

**TABLE 4.3** Parameter estimates for  $P(R|R)$  and  $P(R|\bar{R})$ 

Station	Variable	Parameters			Polar Form				
		$\hat{\gamma}_1$	$\hat{\gamma}_2$	$\hat{\gamma}_3$	amplitudes			phases	
					(0)	(1)	(2)	(1)	(2)
Elsenburg	$P(R R)$	-0.143	-0.398	-0.148	-0.143	0.425		203.19	
	$P(R \bar{R})$	-1.593	-0.487	-0.249	-1.593	0.547		210.00	
Kakamas	$P(R R)$				-1.194	0.241	0.221	106.59	84.75
	$P(R \bar{R})$				-3.367	0.810	0.321	51.69	92.34
Middelburg	$P(R R)$	-0.281	0.175	0.032	-0.281	0.178		10.57	
	$P(R \bar{R})$	-2.054	0.558	0.195	-2.054	0.591		19.54	
Nelspruit	$P(R R)$	-0.204	0.391	-0.133	-0.204	0.413		345.94	
	$P(R \bar{R})$	-1.567	1.294	-0.037	-1.567	1.295		363.34	
Cedara	$P(R R)$	0.293	0.918	-0.180	0.293	0.935		353.76	
	$P(R \bar{R})$	-0.888	1.488	-0.139	-0.888	1.494		359.61	
Hoopstad	$P(R R)$	-0.192	0.251	-0.017	-0.192	0.252		361.03	
	$P(R \bar{R})$	-1.927	1.201	0.190	-1.927	1.216		9.11	

value whenever they fall outside the range. Such a procedure is easy to implement but it does change the parameter functions of the generated process (for example the mean), unless the percentage of such points is quite small in which case the resultant bias will be small. Alternatively, one can transform the data. The transformation used ensures that the generated climate sequences lie within the admissible regions and that maximum temperature/humidity values will be greater than minimum temperature/humidity, while the characteristics displayed by the climate series remain unchanged. No transformation was performed on variables which the models described adequately.

**TABLE 4.4** Parameter estimates for the distribution of rainfall on days when rain occurs

Station	Variable	Parameters			Polar Form				
		$\hat{\mu}_1$	$\hat{\mu}_2$	$\hat{\mu}_3$	amplitudes			phases	
					(0)	(1)	(2)	(1)	(2)
Elsenburg	mean	64.73	-12.65	14.64	64.73	19.35		132.66	
	coeff.var	1.2216							
Kakamas	mean				62.56	20.86	2.60	30.99	128.25
	coeff.var	1.0637							
Middelburg	mean	52.94	17.85	-1.01	52.94	17.88		361.72	
	coeff.var	1.3688							
Nelspruit	mean	64.63	20.35	6.34	64.63	21.31		17.54	
	coeff.var	1.5305							
Cedara	mean	53.98	5.01	-0.21	53.98	5.01		362.60	
	coeff.var	2.1594							
Hoopstad	mean	63.91	0.45	-4.94	63.91	4.96		279.01	
	coeff.var	1.4534							

The general transformation used is of the form

$$V_{TF} = \log \left( \frac{a - V_{NTF}}{V_{NTF} - b} \right)$$

where  $a$  is the upper bound of the variable and  $b$  is the lower bound.  $V_{TF}$  represents the variable in its transformed state and  $V_{NTF}$  represents the variable in its original form.

The models are then implemented on the transformed time series. The simulated sequences are easily changed back to the original units by reversing the transformation,

that is

$$S_{NTF} = \frac{a + b e^{S_{TF}}}{e^{S_{TF}} + 1}$$

where  $S_{NTF}$  represents the simulated series in the original form and  $S_{TF}$  represents the simulated series in a transformed state.

The above transformation has the property that

$$a > S_{NTF} > b.$$

By a suitable choice of  $a$  and  $b$ , one can prevent maximum temperature being less than minimum temperature. Similarly for humidity. For example, specifying  $a = \max \text{ temp}_{NTF}$  when transforming minimum temperature (i.e. condition minimum temperature on maximum temperature), one obtains that

$$\max \text{ temp} > \min \text{ temp} > b.$$

Alternatively, one can condition maximum temperature on minimum temperature by specifying  $b = \min \text{ temp}_{NTF}$  when transforming maximum temperature, obtaining

$$a > \max \text{ temp} > \min \text{ temp}.$$

Unfortunately the choice of which variable should be conditioned is not obvious. An option can only be verified by implementing the model and then examining the simulated sequences to check whether the properties of the climate sequences are being preserved. Usually one can get an indication of which variable to condition when one fits the model to the untransformed time series. If for one variable it is noted that the properties are not being retained as well as for its corresponding variable, then it would be advisable to first try the transformation where the “worse behaving” variable is conditioned.

In the case of sunshine duration, the upper bound was allowed to vary seasonally with time instead of being a constant. Define the upper limit by  $B(t)$ , where

$$B(t) = \text{ave} + 5 + \left(\frac{\text{amp}}{2}\right) \cos\left(\left(\frac{2\pi}{NT}\right)(t + 11)\right), \quad t = 1, 2, \dots, NT$$

where

$$\text{ave} = \frac{\text{smax} + \text{smin}}{2},$$

$$\text{amp} = \text{smax} - \text{smin} \quad \text{and}$$

$B(t)$  and  $B_{\min}$  are chosen so that  $B(t) \geq B_{\min}$  sunshine duration observed at time  $t$ .

Care must be taken that one does not divide by zero, which can happen at times when the lower limit is zero and a zero observation occurs. This problem can be overcome by adding a small value (e.g. 0.01) to all observations.

**Model 1: Multivariate model for climate data proposed by Richardson (1981).**

Table 4.5 shows the transformations used for each station. Only the lower and upper bounds are given in the table as the form of the transformation is given above.

The historical data for each of the climate variables was conditioned on the wet or dry status of the day, thus obtaining a mean function and a standard deviation function for each of the conditioned data sets. The mean and standard deviation were both approximated by a truncated Fourier series representation. That is

$$\mu_t = \sum_{i=1}^L \alpha_i \phi_i(t) \quad \text{and}$$

$$\sigma_t = \sum_{i=1}^L \xi_i \phi_i(t), \quad t = 1, 2, \dots, NT$$

where  $\phi_i(t)$  is defined as in Chapter 3 and where  $L$  does not have to be of the same order for both of the mean and the standard deviation function.

For the purposes of model selection the truncation level  $L$ , which determines the family of approximating models being fitted, was varied and the fit in each case was examined. The decision on which order of approximation to use was based on Akaike's Information Criterion (AIC). Tables 4.6 – 4.9 show the value of AIC and the choice of the order of approximation is given in square brackets. The percentage decrease of the criterion is given in parentheses whenever the value of AIC continued to decrease after five parameters had already been fitted. Here the number of parameters selected is based on the model which leads to a decrease in the criterion of more than 5 percent. This decision was taken for reasons mentioned in the previous section on the undesirability of fitting a large number of parameters to the models.

The values of  $L$  ranged between 1 and 5, with a mode of 3 for both the mean function given a dry day and for the mean function given a wet day. Therefore, a 3-term Fourier series approximation is estimated to be appropriate. Different  $L$  values for each variable for a particular station were not chosen in order to simplify the implementation and interpretation of the complete (multivariate time series) model.

**TABLE 4.5 Transformations for Model 1**

Variable	S t a t i o n					
	Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp unchanged	a=490	a=400	a=420	a=400	a=410	
	b=min temp	b=min temp	b=min temp	b=min temp	b=min temp	b=min temp
Min Temp	a=max temp	a=320	a=250	a=250	unchanged	a=230
	b=0	b=-50	b=-90	b=0		b=-100
Evapo	square	a=300	square	N/A	N/A	N/A
	root	b=0	root	N/A	N/A	N/A
Sun	a=B(t)	a=B(t)	a=B(t)	a=B(t)	a=B(t)	a=B(t)
	smax=134	smax=136	smax=139	smax=130	smax=132	smax=135
	smin=94	smin=100	smin=100	smin=110	smin=102	smin=110
	b=0	b=0	b=0	b=0	b=0	b=0
Wind	a=10000	a=10000	a=10000	a=1000	a=1000	a=1000
	b=0	b=0	b=0	b=0	b=0	b=0
Max Hum	a=1001	a=1001	a=1001	a=101	a=101	a=101
	b=min hum	b=min hum	b=min hum	b=0	b=0	b=min hum
Min Hum	a=1000	a=1000	a=1000	a=max hum	a=max hum	a=100
	b=0	b=0	b=0	b=0	b=0	b=0

**TABLE 4.6 Model selection criteria (AIC) for the mean function for non-rainy days**

Variable	L	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	1	327261	345	357	341	339	351
	3	78867	331	335	343	341	338
	5	72626	333	337			339
	7	72133(0.4%)					
Selected	[5]		[3]	[3]	[1]	[1]	[3]
Min Temp	1	355	407	397	465	350496	464
	3	346	333	336	344	49924	339
	5	348	333	337	340(1%)	34256	338(0.3%)
	7					33719(2%)	
Selected	[3]		[3]	[3]	[3]	[5]	[3]
Evapo	1	1098	414	737			
	3	391	332	370			
	5	390(0.3%)	334	379			
Selected	[3]		[3]	[3]			
Sun	1	403	343	370	528	494	359
	3	399	345	372	479	475	359
	5	400			473	476	
Selected	[3]		[1]	[1]	[3]	[3]	[1]
Wind	1	340	339	334	331	330	328
	3	337	332	334	332	329	330
	5	338	334			330	
Selected	[3]		[3]	[1]	[1]	[3]	[1]
Max Hum	1	376	383	385	347	389	389
	3	375	356	364	343	371	345
	5	377	358	365	344	373	345
Selected	[3]		[3]	[3]	[3]	[3]	[3]
Min Hum	1	345	338	336	361	370	343
	3	341	332	333	341	346	340
	5	342	333	335	342	348	342
Selected	[3]		[3]	[3]	[3]	[3]	[3]

**TABLE 4.7** Model selection criteria (AIC) for the mean function for rain days

Variable	L	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	1	285827	341	372	357	347	354
	3	107685	341	347	351	346	340
	5	104150(3%)		349	354	346	343
Selected		[3]	[1]	[3]	[3]	[3]	[3]
Min Temp	1	340	340	348	349	215468	348
	3	342	333	332	330	42050	330
	5		335	333	332	36230	332
	7					35482(2%)	
Selected		[1]	[3]	[3]	[3]	[5]	[3]
Evapo	1	1196	355	1154			
	3	756	353	841			
	5	748(1%)	355	832(1%)			
Selected		[3]	[3]	[3]			
Sun	1	951	457	970	1541	1251	661
	3	865	458	947	1354	1177	657
	5	855(1%)		940(0.7%)	1331(2%)	1130(4%)	654(0.5%)
Selected		[3]	[1]	[3]	[3]	[3]	[3]
Wind	1	338	327	333	328	328	337
	3	340	330	334	330	330	337
Selected		[1]	[1]	[1]	[1]	[1]	[1]
Max Hum	1	371	504	446	346	352	366
	3	375	504	438	347	352	359
	5			439			356(0.8%)
Selected		[1]	[1]	[3]	[1]	[1]	[3]
Min Hum	1	339	337	349	372	392	360
	3	337	339	350	375	388	360
	5	338				389	
Selected		[3]	[1]	[1]	[1]	[3]	[1]

**TABLE 4.8 Model selection criteria (AIC) for the standard deviation function for non-rainy days**

Variable	L	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	1	30600	327	327	327	321	327
	3	29683	329	329	329	323	329
	5	29003(2%)					
Selected		[3]	[1]	[1]	[1]	[1]	[1]
Min Temp	1	327	327	327	332	20944	327
	3	329	329	329	333	19404	329
	5					18588(4%)	
Selected		[1]	[1]	[1]	[1]	[3]	[1]
Evapo	1	339	327	328			
	3	340	329	329			
Selected		[1]	[1]	[1]			
Sun	1	367	327	354	504	442	333
	3	365	329	355	489	439	331
	5	367			484(1%)	441	333
Selected		[3]	[1]	[1]	[3]	[3]	[3]
Wind	1	327	327	327	327	321	327
	3	329	329	329	329	323	329
Selected		[1]	[1]	[1]	[1]	[1]	[1]
Max Hum	1	348	333	332	327	321	327
	3	342	334	331	329	324	329
	5	344		332			
Selected		[3]	[1]	[3]	[1]	[1]	[1]
Min Hum	1	327	327	327	327	321	327
	3	329	329	329	329	323	329
Selected		[1]	[1]	[1]	[1]	[1]	[1]

**TABLE 4.9** Model selection criteria (AIC) for the standard deviation function for rain days

Variable	no of parameters	S t a t i o n					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	1	51315	332	329	329	327	329
	3	48654	334	331	331	329	331
	5	47971(1%)					
Selected		[3]	[1]	[1]	[1]	[1]	[1]
Min Temp	1	327	327	328	327	15941	327
	3	329	329	329	329	15330	329
	5					15221(1%)	
Selected		[1]	[1]	[1]	[1]	[3]	[1]
Evapo	1	457	332	497			
	3	457	333	497			
Selected		[1]	[1]	[1]			
Sun	1	619	382	702	744	665	572
	3	547	381	691	686	625	572
	5	545(0.4%)	383	689(0.3%)	683(0.4%)	602(4%)	
Selected		[3]	[3]	[3]	[3]	[3]	[1]
Wind	1	329	327	327	327	327	330
	3	330	329	329	329	329	332
Selected		[1]	[1]	[1]	[1]	[1]	[1]
Max Hum	1	360	392	357	328	328	333
	3	355	392	360	330	332	335
	5	355					
Selected		[3]	[1]	[1]	[1]	[1]	[1]
Min Hum	1	327	329	329	342	338	333
	3	330	331	332	342	336	335
	5					338	
Selected		[1]	[1]	[1]	[1]	[3]	[1]

The values of  $L$  ranged between 1 and 3, with a mode of 1, for both the standard deviation function given a dry day and for the standard deviation function given a wet day.

Again a 3-term Fourier series approximation was chosen to simplify programming by having a common approximation order.

Tables 4.10–4.15 show the parameter estimates for the mean function and for the standard deviation function, both conditioned on the wet or dry status of period  $t$ .

The resulting time series obtained by subtracting the fitted mean function and by dividing through by the fitted standard deviation function should be a time series with a mean of zero and a standard deviation of unity. Since the mean value functions and the standard deviation functions which were fitted are based on truncated Fourier series, that is, on approximating models, the means of the residual series would not be exactly zero and the standard deviations would not be exactly one. However, deviations in this respect were found to be quite small. (Table 4.16.)

**TABLE 4.10** Parameter estimates for the mean and standard deviation function for Nelspruit

Variable	Day Status	Mean Function			Standard deviation function		
		$\hat{\alpha}_1$	$\hat{\alpha}_2$	$\hat{\alpha}_3$	$\hat{\xi}_1$	$\hat{\xi}_2$	$\hat{\xi}_3$
Max Temp	Dry	1.1406	0.0491	0.0021	0.5492	0.0722	-0.1045
	Wet	0.7439	-0.3820	-0.2312	0.7346	-0.0129	-0.1027
Min Temp	Dry	-0.0768	-1.1420	-0.2058	0.4849	-0.1264	0.0172
	Wet	-0.3966	-0.8404	-0.2213	0.3666	-0.0469	0.0026
Sun	Dry	-0.8305	0.6966	-0.2708	1.5700	0.4211	-0.2742
	Wet	2.2995	-1.2331	-1.1324	3.1566	-0.8673	-0.5615
Wind	Dry	1.9637	0.0005	0.1367	0.2404	0.0175	-0.0155
	Wet	2.0005	-0.0579	0.1554	0.2789	-0.0069	-0.0338
Max Hum	Dry	-1.4305	-0.0110	-0.2614	0.5754	-0.0945	-0.0809
	Wet	-2.1826	0.1480	-0.0467	0.7114	0.0075	-0.0685
Min Hum	Dry	-0.2314	-0.4867	-0.0223	0.5471	-0.0377	-0.0490
	Wet	-0.9915	0.0100	0.1507	0.7902	-0.1771	-0.0605

**TABLE 4.11** Parameter estimates for the mean and standard deviation function for Kakamas

Variable	Day Status	Mean Function			Standard deviation function		
		$\hat{\alpha}_1$	$\hat{\alpha}_2$	$\hat{\alpha}_3$	$\hat{\xi}_1$	$\hat{\xi}_2$	$\hat{\xi}_3$
Max Temp	Dry	0.1820	-0.4165	-0.0206	0.3912	0.0299	-0.0215
	Wet	0.7169	-0.1589	-0.1176	0.5849	0.0534	-0.1207
Min Temp	Dry	0.0561	-0.8786	-0.2398	0.4488	0.0249	-0.0311
	Wet	-0.2287	-0.6914	-0.2603	0.3976	0.0583	-0.1232
Evapo	Dry	0.9503	-0.9578	0.0770	0.4293	0.0004	-0.0106
	Wet	1.4806	-0.3786	-0.1060	0.7570	-0.0814	-0.1750
Sun	Dry	-1.6980	-0.0062	0.0217	0.8455	0.0029	0.1077
	Wet	0.5187	-0.0863	-0.1922	1.3270	-0.2649	0.1156
Wind	Dry	1.4858	-0.3014	0.0995	0.4098	-0.0406	-0.0231
	Wet	1.2941	0.0066	0.0779	0.3349	-0.0073	-0.0103
Max Hum	Dry	-0.0944	0.5496	-0.1002	1.3666	-0.2549	0.0597
	Wet	-0.7814	0.0296	-0.2719	1.4821	-0.2785	-0.1268
Min Hum	Dry	1.1816	0.2893	-0.0566	0.4990	0.0249	-0.0066
	Wet	0.4772	-0.0064	-0.0173	0.5523	-0.0303	-0.0673

**TABLE 4.12** Parameter estimates for the mean and standard deviation function for Middelburg

Variable	Day Status	Mean Function			Standard deviation function		
		$\hat{\alpha}_1$	$\hat{\alpha}_2$	$\hat{\alpha}_3$	$\hat{\xi}_1$	$\hat{\xi}_2$	$\hat{\xi}_3$
Max Temp	Dry	-0.1424	-0.5116	-0.0217	0.5336	0.0718	-0.0156
	Wet	0.4120	-0.7261	-0.1572	0.6608	0.0509	-0.0500
Min Temp	Dry	0.2575	-0.8132	-0.1849	0.5279	-0.1576	-0.0616
	Wet	-0.1811	-0.7344	-0.2009	0.4172	-0.0803	-0.0499
Evapo	Dry	7.9960	1.9682	-0.4032	1.3287	-0.1210	-0.1113
	Wet	6.1770	2.1295	0.0773	2.0379	-0.2645	-0.2099
Sun	Dry	-2.0452	0.0167	0.0293	1.2232	0.0837	0.0880
	Wet	0.6383	-0.5181	-0.3849	1.9803	-0.4339	-0.2604
Wind	Dry	1.4977	-0.0610	0.1297	0.4592	-0.0855	-0.0228
	Wet	1.4042	0.0481	0.1328	0.4821	-0.1459	-0.0122
Max Hum	Dry	-1.3432	-0.4132	-0.2987	1.2128	-0.2591	-0.1682
	Wet	-1.2860	-0.3606	-0.1579	1.1107	-0.1790	-0.0987
Min Hum	Dry	1.1904	0.1649	-0.1501	0.5000	-0.0057	-0.0314
	Wet	0.5364	0.2806	0.0152	0.6164	-0.0767	-0.0799

**TABLE 4.13** Parameter estimates for the mean and standard deviation function for Elsenburg

Variable	Day Status	Mean Function			Standard deviation function		
		$\hat{\alpha}_1$	$\hat{\alpha}_2$	$\hat{\alpha}_3$	$\hat{\xi}_1$	$\hat{\xi}_2$	$\hat{\xi}_3$
Max Temp	Dry	238.22	48.57	19.04	38.38	3.179	-0.0134
	Wet	193.53	46.05	21.63	25.71	5.296	3.317
Min Temp	Dry	0.3177	-0.2871	-0.1851	0.5278	-0.1475	-0.0490
	Wet	-0.4954	-0.0572	-0.1013	0.5552	-0.1558	-0.0066
Evapo	Dry	7.680	2.787	0.0901	1.0026	-0.0536	0.0270
	Wet	4.961	2.506	0.1371	1.8950	0.0165	-0.0899
Sun	Dry	-1.6113	-0.2528	-0.0514	0.9575	-0.2123	0.0104
	Wet	1.1820	-1.0978	0.2018	1.9234	-0.9789	0.3166
Wind	Dry	1.6157	-0.2502	0.0099	0.3923	-0.0539	-0.0276
	Wet	1.2347	0.0298	0.1007	0.4737	-0.1934	-0.0040
Max Hum	Dry	-2.2258	0.1167	-0.1301	0.8527	-0.1969	0.1334
	Wet	-2.0879	0.0697	-0.0156	0.4737	-0.1934	-0.0040
Min Hum	Dry	0.5553	0.2438	0.0887	0.4858	-0.0842	-0.0051
	Wet	-0.1166	0.4444	0.0339	0.6925	-0.3514	0.0798

**TABLE 4.14** Parameter estimates for the mean and standard deviation function for Cedara

Variable	Day Status	Mean Function			Standard deviation function		
		$\hat{\alpha}_1$	$\hat{\alpha}_2$	$\hat{\alpha}_3$	$\hat{\xi}_1$	$\hat{\xi}_2$	$\hat{\xi}_3$
Max Temp	Dry	0.0677	-0.1084	0.0099	0.5117	0.0419	-0.0775
	Wet	0.8100	-0.1902	-0.1313	0.8119	0.0835	-0.1253
Min Temp	Dry	96.367	56.733	11.619	26.661	-3.558	-2.103
	Wet	111.379	46.644	13.587	22.168	-0.926	-2.553
Sun	Dry	-1.383	0.4607	-0.1772	1.3118	0.2792	-0.1978
	Wet	2.0787	0.2088	-0.9158	3.3718	0.3074	-0.6914
Wind	Dry	1.7602	-0.0816	0.1776	0.2919	-0.0758	-0.0153
	Wet	1.6711	-0.0878	0.1973	0.3219	-0.0127	-0.0353
Max Hum	Dry	-2.0807	-0.3614	-0.2952	0.9552	-0.1621	-0.0217
	Wet	-3.0247	-0.2189	-0.1216	0.8009	0.1149	0.0001
Min Hum	Dry	-0.0543	-0.5256	-0.0813	0.6235	-0.0317	-0.0688
	Wet	-1.1142	-0.2502	0.0721	1.0571	-0.0468	-0.1675

**TABLE 4.15** Parameter estimates for the mean and standard deviation function for Hoopstad

Variable	Day Status	Mean Function			Standard deviation function		
		$\hat{\alpha}_1$	$\hat{\alpha}_2$	$\hat{\alpha}_3$	$\hat{\xi}_1$	$\hat{\xi}_2$	$\hat{\xi}_3$
Max Temp	Dry	-0.3482	-0.4120	0.0216	0.4415	0.1293	-0.0081
	Wet	0.2494	-0.5653	-0.0547	0.6324	0.0222	0.0372
Min Temp	Dry	-1.4750	-1.1625	-0.2069	0.4701	-0.0122	-0.0174
	Wet	-0.6601	-0.8644	-0.1761	0.3764	-0.0442	-0.0151
Sun	Dry	-1.7544	0.1099	0.1206	0.7901	0.2790	0.0360
	Wet	0.6600	-0.4369	-0.0254	1.8220	-0.1887	0.1484
Wind	Dry	2.1065	-0.2612	0.2238	0.5323	-0.0128	-0.0153
	Wet	1.8836	0.0616	0.3142	0.4281	0.1357	0.0514
Max Hum	Dry	-0.4700	0.5763	-0.4229	0.7188	-0.1011	0.0364
	Wet	-0.7645	0.2807	-0.3057	0.7501	-0.0202	0.0191
Min Hum	Dry	0.9651	0.0142	-0.2402	0.5565	0.0808	-0.0154
	Wet	0.0503	0.1797	-0.0793	0.7935	-0.0143	0.0135

**TABLE 4.16** Mean and standard deviation of residual time series obtained by standardizing the data

Variable		S t a t i o n					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	M	-0.010	0.002	-0.007	0.0007	-0.02	0.006
	SD	1.09	1.04	1.05	1.08	1.08	1.06
Min Temp	M	0.004	-0.004	0.002	-0.0005	-0.01	-0.002
	SD	1.08	1.03	1.07	1.10	1.08	1.06
Evapo	M	-0.003	-0.0004	0.006			
	SD	1.12	1.05	1.05			
Sun	M	-0.007	0.006	-0.002	0.011	0.02	0.0009
	SD	1.16	1.06	1.12	1.22	1.17	1.13
Wind	M	-0.015	-0.003	0.0002	0.192	-0.007	0.013
	SD	1.09	1.03	1.05	1.12	1.11	1.15
Max Hum	M	0.012	-0.002	0.002	0.004	-0.009	-0.004
	SD	1.19	1.05	1.06	1.07	1.06	1.09
Min Hum	M	-0.010	-0.0003	0.004	0.003	-0.01	-0.006
	SD	1.10	1.05	1.06	1.10	1.08	1.07

Another assumption made by the model is that the residual time series follows an autoregressive process of order 1. If this is true then  $\rho_k = \rho_1^k$  where  $\rho_k$  is the autocorrelation with lag  $k$ . This assumption (or more precisely, this approximation) was checked by comparing  $\hat{\rho}_k$ ,  $k = 1, 2, 3, 4$  with  $\hat{\rho}_1^k$ , and was found to be reasonable except for a few cases (Table 4.17). It is possible to increase the order of the autoregressive process to these cases, but this has to be done at cost of increasing the complexity and number of parameters in the model, and therefore not advisable.

The results of the above checks would suggest that the residual series do seem to satisfy the required assumptions of the model. It is therefore reasonable to approximate each of the seven series by the sum of a seasonal component and a residual component, to approximate the seasonal component by a 3-term Fourier series and finally to approximate the standard deviation of the residual series by a 3-term Fourier approximation.

TABLE 4.17 Comparison of  $\hat{\rho}_k$  and  $\hat{\rho}_1^*$  of the residual series

Station	Max Temp $\hat{\rho}_k$	Max Temp $\hat{\rho}_1^*$	Min Temp $\hat{\rho}_k$	Min Temp $\hat{\rho}_1^*$	Evapo $\hat{\rho}_k$	Evapo $\hat{\rho}_1^*$	Sun $\hat{\rho}_k$	Sun $\hat{\rho}_1^*$	Wind $\hat{\rho}_k$	Wind $\hat{\rho}_1^*$	Max Hum $\hat{\rho}_k$	Max Hum $\hat{\rho}_1^*$	Min Hum $\hat{\rho}_k$	Min Hum $\hat{\rho}_1^*$
Eisenburg	0.42	0.42	0.14	0.14	0.19	0.19	0.21	0.21	0.26	0.26	0.28	0.28	0.38	0.38
	0.15	0.18	0.01	0.02	0.11	0.04	0.06	0.04	0.04	0.07	0.19	0.08	0.23	0.14
	0.04	0.07	-0.01	0.00	0.05	0.01	0.01	0.01	0.02	0.13	0.02	0.20	0.05	
	0.01	0.03	-0.03	0.00	0.02	0.00	0.00	0.00	0.01	0.01	0.14	0.01	0.18	0.02
Kakamas	0.43	0.43	0.66	0.66	0.29	0.29	0.30	0.30	0.26	0.26	0.46	0.46	0.62	0.62
	0.20	0.18	0.42	0.44	0.23	0.08	0.12	0.09	0.03	0.07	0.29	0.21	0.49	0.38
	0.12	0.08	0.29	0.29	0.19	0.02	0.08	0.03	0.03	0.02	0.23	0.10	0.45	0.24
	0.08	0.03	0.21	0.19	0.16	0.01	0.05	0.01	0.02	0.00	0.21	0.04	0.42	0.15
Middelburg	0.34	0.34	0.34	0.34	0.36	0.36	0.22	0.22	0.31	0.31	0.24	0.24	0.46	0.46
	0.12	0.12	0.14	0.12	0.11	0.13	0.10	0.05	0.03	0.10	0.15	0.06	0.33	0.21
	0.09	0.04	0.07	0.04	0.10	0.05	0.07	0.01	0.03	0.03	0.14	0.01	0.30	0.10
	0.03	0.01	0.04	0.01	0.07	0.02	0.02	0.00	0.03	0.01	0.09	0.00	0.26	0.04
Nelspruit	0.24	0.24	0.45	0.45			0.15	0.15	0.78	0.78	0.24	0.24	0.17	0.17
	0.04	0.06	0.31	0.20			0.02	0.02	0.66	0.61	0.09	0.06	0.06	0.03
	0.01	0.01	0.25	0.09			-0.01	0.00	0.56	0.47	0.06	0.01	0.05	0.01
	0.02	0.00	0.18	0.04			0.02	0.00	0.48	0.37	0.02	0.00	0.02	0.00
Cedara	0.13	0.13	0.48	0.48			0.10	0.10	0.76	0.76	0.12	0.12	0.14	0.14
	0.00	0.02	0.29	0.23			-0.01	0.01	0.62	0.58	0.02	0.01	0.04	0.02
	-0.01	0.00	0.20	0.11			-0.02	0.00	0.52	0.44	0.03	0.00	0.01	0.00
	-0.01	0.00	0.15	0.05			0.00	0.00	0.43	0.33	0.02	0.00	0.03	0.00
Hoopstad	0.50	0.50	0.50	0.50			0.23	0.23	0.75	0.75	0.40	0.40	0.51	0.51
	0.31	0.25	0.35	0.25			0.14	0.05	0.63	0.56	0.31	0.16	0.35	0.26
	0.22	0.13	0.28	0.13			0.09	0.01	0.54	0.42	0.23	0.06	0.26	0.13
	0.18	0.06	0.22	0.06			0.07	0.00	0.46	0.32	0.21	0.03	0.18	0.07

Model 1, proposed by Richardson (1981) is given by:

$$\chi_{i,t} = A \chi_{i,t-1} + B \epsilon_{i,t}$$

where  $\chi_{i,t}$  is the residual series at time period  $t$  of year  $i$ . Display 4.1 gives the estimated  $A$  matrix for the various stations and Display 4.2 gives the estimated  $B$  matrix. The order of the climate variables in these displays is as follows: maximum temperature, minimum temperature, evaporation, sunshine duration, windrun, maximum humidity and finally minimum humidity.

#### **Model T: Multivariate model for climate data**

Models 3, 4 and 5 were developed as an alternative to Model 1 in an attempt to deal with a deficiency in Model 1, namely the assumption that the autocorrelation function of each variable is assumed invariant with respect to wet/dry and dry/wet day boundaries. Each model varies in complexity and emphasizes a slightly different aspect of the joint distribution of the variables. Table 4.18 shows the fundamental assumptions of each model.

The models depicted here are complex, describing several distinguishing properties of the climate series. No one model will be “best” in all respects or for all sites. In general, simpler models can be expected to outperform the more complex ones when the historical record at the site is small, whereas the opposite will be true when the record is long. Statistically one can select the appropriate model for each variable using Akaike’s Information Criterion (AIC) where

$$AIC = -\ell(\psi; e_{i,t}) + L$$

where  $\ell(\psi; e_{i,t})$  is the loglikelihood function of the model and  $L$  is the number of parameters. The model producing the lowest AIC value is chosen as the model that best described that particular climate variable. It must be noted that Models 3, 4 and 5 are not hierarchical and therefore a model with a larger number of parameters does not imply that the value of its loglikelihood function will be smaller.

Elsenburg	[	0.69	-0.49	-0.02	0.05	-0.03	-0.08	-0.04]
		0.17	0.09	-0.03	0.01	-0.03	-0.05	0.02]
		0.26	-0.25	-0.02	-0.03	-0.08	0.14	0.19]
		-0.15	-0.07	0.02	0.13	0.03	-0.01	-0.05]
		0.22	0.19	-0.01	0.07	0.16	-0.19	-0.22]
		0.04	-0.02	0.01	0.07	0.00	0.26	0.02]
		0.12	-0.21	-0.07	0.08	-0.08	0.08	0.51]
Kakamas	[	0.40	0.25	0.05	-0.07	0.06	-0.10	0.04]
		-0.09	0.74	0.02	0.04	0.00	0.17	-0.07]
		-0.07	0.13	0.26	-0.01	0.02	-0.14	0.07]
		0.16	-0.20	0.06	0.18	-0.01	-0.15	0.03]
		-0.45	-0.06	0.09	0.06	0.17	-0.02	-0.02]
		0.06	-0.26	0.01	0.00	-0.03	0.32	0.13]
		0.11	-0.05	0.10	0.09	-0.06	0.24	0.64]
Middelburg	[	0.40	0.35	-0.01	-0.06	0.03	-0.18	-0.03]
		-0.04	0.39	-0.07	0.01	-0.04	0.19	0.07]
		0.23	-0.18	0.25	-0.05	0.03	0.27	0.11]
		0.11	-0.08	-0.02	0.15	0.03	-0.16	0.01]
		-0.58	0.05	-0.09	0.14	0.12	-0.23	-0.12]
		0.11	-0.01	0.04	0.00	0.05	0.24	0.06]
		0.10	-0.02	-0.08	0.10	-0.04	0.28	0.59]
Nelspruit	[	0.43	0.23	-0.15	0.38	-0.08	0.07]	
		0.13	0.46	0.00	-0.03	0.06	0.04]	
		0.22	0.07	0.03	0.38	0.03	0.01]	
		-0.04	-0.07	-0.02	0.15	-0.06	-0.02]	
		-0.03	0.02	0.06	-0.30	0.16	0.02]	
		-0.10	-0.04	0.04	-0.27	0.17	0.08]	
Cedara	[	0.16	-0.28	-0.10	0.16	-0.05	-0.07]	
		-0.10	0.45	0.05	-0.14	-0.07	-0.04]	
		0.07	-0.03	0.03	0.14	-0.02	-0.07]	
		-0.03	0.13	0.02	0.22	0.05	0.01]	
		0.01	-0.05	0.08	-0.18	0.09	0.09]	
		-0.03	0.08	0.12	-0.15	0.06	0.19]	
Hoopstad	[	0.40	0.24	-0.07	0.12	-0.19	-0.21]	
		0.07	0.52	0.06	0.11	0.04	0.04]	
		0.09	-0.12	0.09	-0.03	-0.10	-0.10]	
		-0.01	-0.11	0.02	0.43	0.05	0.07]	
		-0.05	-0.07	0.06	-0.05	0.36	0.13]	
		-0.09	-0.01	0.11	-0.13	0.23	0.47]	

DISPLAY 4.1 The estimated matrix  $A$

$$\begin{array}{l}
 \text{Elsenburg} \begin{bmatrix} 0.76 \\ 0.46 & 0.86 \\ 0.36 & 0.00 & 0.84 \\ -0.34 & -0.33 & -0.34 & 0.75 \\ 0.20 & 0.18 & -0.35 & -0.03 & 0.80 \\ 0.05 & -0.19 & 0.07 & 0.05 & -0.05 & 0.93 \\ 0.57 & 0.27 & 0.14 & -0.05 & 0.04 & -0.08 & 0.60 \end{bmatrix} \\
 \\
 \text{Kakamas} \begin{bmatrix} 0.80 \\ -0.15 & 0.73 \\ 0.40 & 0.31 & 0.80 \\ 0.31 & -0.19 & 0.15 & 0.83 \\ 0.01 & 0.30 & 0.41 & -0.09 & 0.74 \\ -0.25 & -0.27 & -0.05 & 0.05 & -0.06 & 0.75 \\ -0.43 & -0.05 & -0.08 & -0.08 & 0.02 & 0.10 & 0.59 \end{bmatrix} \\
 \\
 \text{Middelburg} \begin{bmatrix} 0.77 \\ -0.34 & 0.87 \\ -0.36 & -0.34 & 0.68 \\ 0.44 & -0.07 & -0.32 & 0.77 \\ 0.00 & 0.30 & -0.49 & -0.11 & 0.58 \\ 0.00 & -0.41 & 0.20 & 0.01 & -0.11 & 0.84 \\ -0.56 & -0.10 & 0.20 & -0.09 & 0.08 & -0.02 & 0.55 \end{bmatrix} \\
 \\
 \text{Nelspruit} \begin{bmatrix} 0.84 \\ -0.40 & 0.80 \\ 0.61 & 0.00 & 0.66 \\ -0.03 & -0.02 & -0.08 & 0.98 \\ -0.31 & 0.00 & -0.08 & -0.27 & 0.83 \\ -0.66 & -0.02 & -0.07 & 0.02 & -0.08 & 0.63 \end{bmatrix} \\
 \\
 \text{Cedara} \begin{bmatrix} 0.92 \\ 0.20 & 0.83 \\ 0.66 & 0.08 & 0.72 \\ -0.22 & -0.17 & -0.02 & 0.95 \\ -0.45 & 0.06 & 0.04 & -0.10 & 0.87 \\ -0.71 & 0.01 & -0.16 & -0.03 & 0.01 & 0.63 \end{bmatrix} \\
 \\
 \text{Hoopstad} \begin{bmatrix} 0.77 \\ -0.34 & 0.80 \\ 0.40 & -0.20 & 0.89 \\ -0.21 & 0.07 & 0.01 & 0.84 \\ -0.05 & -0.05 & -0.04 & -0.06 & 0.93 \\ -0.56 & 0.04 & -0.12 & -0.01 & 0.04 & 0.57 \end{bmatrix}
 \end{array}$$

DISPLAY 4.2 The estimated matrix  $B$ 

Thus, one does not restrict the generation of the climate variables to any particular

**TABLE 4.18 Assumption of the different models**

Model	Assumptions
Model 1	<ul style="list-style-type: none"> <li>– seasonal mean function</li> <li>– seasonal standard deviation function</li> <li>– constant autocorrelation coefficient across different rain – no rain sequences</li> <li>– conditioned on wet and dry sequences only</li> </ul>
Model 3	<ul style="list-style-type: none"> <li>– seasonal mean function</li> <li>– constant standard deviation function</li> <li>– different autocorrelation coefficients across different rain – no rain sequences</li> <li>– conditioned on wet/wet, dry/dry, wet/dry and dry/wet sequences</li> </ul>
Model 4	<ul style="list-style-type: none"> <li>– seasonal mean function</li> <li>– seasonal standard deviation function</li> <li>– constant autocorrelation coefficients across different rain – no rain sequences</li> <li>– conditioned on wet/wet, dry/dry, wet/dry and dry/wet sequences</li> </ul>
Model 5	<ul style="list-style-type: none"> <li>– seasonal mean function</li> <li>– seasonal standard deviation function</li> <li>– different autocorrelation coefficients across different rain – no rain sequences</li> <li>– conditioned on wet/wet, dry/dry, wet/dry and dry/wet sequences.</li> </ul>

model, but each variable is generated according to the model that “best” describes it. The multivariate model to generate simultaneous daily climate sequences will be referred to as Model T, where, for each climate variable, Model T is constructed by selecting between Models 3, 4 and 5 for the model that produces the lowest AIC.

The following algorithm is used to implement Model T.

#### Algorithm

- Step 1: Implement Model 3 to obtain parameter estimates and AIC for each variable.
- Step 2: Implement Model 4 to obtain parameter estimates and AIC for each variable.

Step 3: Implement Model 5 to obtain parameter estimates and AIC for each variable.

Step 4: Construct Model T by choosing for each variable the model producing the lowest AIC.

Step 5: Obtain the estimated cross-correlation matrix,  $\hat{\Sigma}$ , whose elements are given by:

$$R_{jk} = \frac{\frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} e_{i,t}^{(k)} - \frac{1}{T^2} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)}}{\left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(j)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(j)} \right)^2 \right]^{\frac{1}{2}} \left[ \frac{1}{T} \sum_{i=1}^{NY} \sum_{t=1}^{NT} (e_{i,t}^{(k)})^2 - \frac{1}{T^2} \left( \sum_{i=1}^{NY} \sum_{t=1}^{NT} e_{i,t}^{(k)} \right)^2 \right]^{\frac{1}{2}}}$$

where  $e_{i,t}^{(j)}$  denotes the residual time series of variable  $j$ ,  $j = 1, 2, \dots, NV$  and where for each  $j$  the residual series is the series obtained after the model producing the lowest AIC for variable  $j$  has been fitted and  $e_{i,t}^{(k)}$  denotes the residual time series of variable  $k$ ,  $k = 1, 2, \dots, NV$ , the residual series obtained in the same way as above.

#### Implementing Models 3, 4 and 5

The transformations applied to these models are the same as those in the previous model except for the variables maximum temperature and minimum temperature of the station Nelspruit. Here the bounds are given by

$$a = 420, \quad \text{and}$$

$$b = 0$$

for maximum temperature, and

$$a = \text{max temp}$$

$$b = 0$$

for minimum temperature.

Models 3, 4 and 5 are implemented by following the respective algorithms given in Chapter 3. The initial estimates for the mean function and for the seasonal standard deviation function are the same as the estimates of Model 1 and therefore need not be recomputed. Only when a different transformation to that applied in Model 1 is used, is it necessary to compute initial estimates for the mean function and the seasonal standard deviation function.

The selection of the order of the Fourier series approximation for the mean function and the standard deviation function was based on the initial estimates of these functions and so a choice of a 3-term Fourier series was made for both functions.

The estimation of the parameters is accomplished by iteration. The procedure described in the algorithms is that of Newton–Raphson. This method, although converging within a few iterations, was found to be sensitive to the initial values given. On occasions where convergence was not reached, a conjugate gradient method was used for parameter estimation. The computer programs for this procedure are given in Appendix D. The advantages of this method is that one gets convergence and only the vector of first derivatives needs to be computed. The disadvantage is that it is time consuming and takes a large number of iterations to converge.

In our implementation a particular parameter estimate was deemed to have converged when its value changed by less than 0.01% in successive iterations. The estimation procedure was deemed to have converged when all estimates had converged.

Akaike's Information Criterion for the selection of a model for each variable is given in Table 4.19. The lowest AIC value is shown in bold and the corresponding model is selected to generate climate sequences for that variable. This model selection is performed for each of the stations.

Values with an asterisk indicate the model that was finally selected, although it did not produce the lowest AIC. This choice was necessary in some instances because of the relatively few occurrences of rainfall in some stations. More importantly, at some sites consecutive rainy days seldom occur. Thus, in Models 3 and 5 the estimation of the autocorrelation coefficient given that the sequence WW was observed, is based on very few observations, and consequently, it is possible to obtain inadmissible estimates. Similarly, when the sequence WD or DW was observed. In such cases, generally Model 4 was chosen as it gave acceptable estimates since here the autocorrelation function is not conditioned and all observations are used in the estimation.

Parameter estimates for the mean function, the standard deviation function and the coefficient of the autoregressive process of order 1 are given in Tables 4.20–4.25 for each station.

**Table 4.19 Akaike's Information Criterion for Models 3, 4 and 5**

Variable	Model	S t a t i o n					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	Mod 3	21190	<b>3393</b>	5751	2151	5781	<b>2610</b>
	Mod 4	21174	3537	5796	2030	5745	2988
	Mod 5	<b>21160</b>	3502	<b>5749</b>	<b>2006</b>	<b>5726</b>	2773
Min Temp	Mod 3	3579	2994	5474	4684	26594	3191
	Mod 4	3484	2992	5178	4502	26594	3193
	Mod 5	<b>3467</b>	<b>2962</b>	<b>5162</b>	<b>4472</b>	<b>26562</b>	<b>3182</b>
Evapo	Mod 3	7102	<b>5018</b>	12721			
	Mod 4	7124	5124	12710			
	Mod 5	<b>7091</b>	5080	<b>12697</b>			
Sun	Mod 3	7458	<b>10528</b>	<b>12716</b>	<b>10742</b>	<b>12645</b>	<b>8458</b>
	Mod 4	7263	10669	12841	10951	12786	8511
	Mod 5	<b>7241</b>	10665	12822	10953	12758	8505
Wind	Mod 3	2557	4528	4664	<b>493</b>	2105	3973
	Mod 4	<b>2477</b>	<b>4499</b>	<b>4536</b>	607	2162	4463
	Mod 5	2483	4501	4539	565	<b>2057</b>	<b>3760</b>
Max Hum	Mod 3	5799	13644	11890	4900	8101	6579
	Mod 4	5544	13799*	<b>11712</b>	4879	8090	6614
	Mod 5	<b>5542</b>	<b>13587</b>	11717	<b>4859</b>	<b>8078</b>	<b>6573</b>
Min Hum	Mod 3	3012	3951*	5653	4961	<b>7047</b>	<b>4266</b>
	Mod 4	2952	3972	<b>4876</b>	<b>4903</b>	7067	4541
	Mod 5	<b>2926</b>	<b>3929</b>	5668	4904	7054	4312

The estimate of the cross-correlation matrix,  $\hat{\Sigma}$ , is obtained by following Step 5 of the algorithm given in this chapter. Problems arise when computing this formula when missing observations occur in the residual series. A simple approach to estimate the cross-correlation matrix in the presence of missing values, is to restrict the analysis to time periods  $t$  with all variables observed. However, this method discards a considerable amount of data and the estimate obtained is biased. A more efficient approach is to estimate the missing values and to replace them by their estimate.

Makhuvha (1988) investigated several methods of estimating the missing values in rain-fall records. She concluded that of the methods compared, the EM algorithm is the most

TABLE 4.20 Parameter estimates of Model T for Elsenburg

Variable	Sta- tus	$\hat{\alpha}_1$	Mean $\hat{\alpha}_2$	$\hat{\alpha}_3$	Standard Deviation			(WD)	Autocorrelation coefficient			
					$\hat{\xi}_1$ (DD)	$\hat{\xi}_2$ (WW)	$\hat{\xi}_3$ (DW)		(DD)	(WW)	(DW)	(WD)
Max Temp	D	231.18	48.72	19.01	33.55	1.8024	0.2573		0.5794	0.3459	0.3872	0.5942
	W	204.50	53.68	22.74	29.03	6.6864	3.1137					
Min Temp	D	0.3022	-0.2920	-0.1657	0.5212	-0.1354	-0.0546		0.2342	-0.0720	0.2671	0.0520
	W	-0.4868	-0.0160	-0.0984	0.6016	-0.1326	-0.0079					
Evapo	D	7.6192	2.7414	0.0900	1.0359	-0.0949	0.0635		0.3227	-0.0491	0.1168	0.0920
	W	5.0406	2.5638	-0.1071	2.1324	0.2682	-0.0206					
Sun	D	-1.6281	-0.2424	-0.0765	1.0764	-0.2902	0.0525		0.0009	0.1173	0.2898	0.1036
	W	0.9495	-1.1692	0.0193	2.2477	-1.0441	0.2785					
Wind	D	1.6106	-0.2549	0.0180	0.3955	-0.0433	-0.0349		0.2780			
	W	1.2669	-0.0216	0.0709	0.5073	-0.1575	-0.0129					
Max Hum	D	-2.2759	0.0996	-0.1112	0.8724	-0.2045	0.1318		0.3436	0.2144	0.2482	0.2042
	W	-2.0439	0.0681	0.0147	0.8779	-0.3449	0.1664					
Min Hum	D	0.5042	0.2483	0.0683	0.4446	-0.0893	0.0071		0.5027	0.1687	0.3148	0.3223
	W	-0.0733	0.4691	0.0451	0.5791	-0.0833	0.0936					

TABLE 4.21 Parameter estimates of Model T for Kakamas

Variable	Sta- tus	$\hat{\alpha}_1$	Mean	Standard Deviation				Autocorrelation coefficient				
			$\hat{\alpha}_2$	$\hat{\alpha}_3$	$\hat{\xi}_1$ (DD)	$\hat{\xi}_2$ (WW)	$\hat{\xi}_3$ (DW)	(WD)	(DD)	(WW)	(DW)	(WD)
Max Temp	D	0.1627	-0.4202	-0.0276	0.3367	1.0537	0.6734	0.6167	0.4781	0.4354	0.4171	0.4551
	W	0.3463	-0.1978	-0.0721								
Min Temp	D	0.0488	-0.8637	-0.2536	0.3342	0.0154	-0.0237		0.6947	0.3330	0.7680	0.4741
	W	-0.0969	-0.8725	-0.2746	0.3321	0.0307	-0.0707					
Evapo	D	0.9368	-0.9685	0.0712	0.4108	1.2381	0.7372	0.6606	0.2930	0.6665	0.5014	0.3466
	W	1.1645	-0.7191	0.0049								
Sun	D	-1.7212	0.0064	-0.0003	0.7934	1.4352	1.8488	1.6322	0.3372	0.0280	0.0585	-0.0044
	W	0.5558	-0.0440	-0.1465								
Wind	D	1.4873	-0.3020	0.1002	0.3989	-0.0255	-0.0251		0.2629			
	W	1.3139	-0.0322	0.0648	0.3927	-0.0213	-0.0013					
Max Hum	D	-0.1501	0.4878	-0.1316	1.1771	-0.2292	0.0032		0.5277			
	W	0.6801	0.1684	-0.1280	1.6941	-0.4360	0.0031					
Min Hum	D	1.1912	0.2865	-0.0486	0.3697	1.0398	0.5222	0.5177	0.6781	0.7747	0.7354	0.6003
	W	0.9306	0.2392	-0.0246								

TABLE 4.22 Parameter estimates of Model T for Middelburg

Variable	Sta- tus	$\hat{\alpha}_1$	Mean $\hat{\alpha}_2$	$\hat{\alpha}_3$	Standard Deviation			Autocorrelation coefficient				
					$\hat{\xi}_1$ (DD)	$\hat{\xi}_2$ (WW)	$\hat{\xi}_3$ (DW)	(WD)	(DD)	(WW)	(DW)	(WD)
Max Temp	D	-0.1750	-0.5214	-0.0207	0.4972	0.0700	-0.0271		0.3351	0.6246	0.2688	0.7525
	W	-0.0466	-0.5426	-0.0733	0.7664	-0.0019	-0.0946					
Min Temp	D	0.2309	-0.7890	-0.2229	0.5073	-0.1471	-0.0584		0.3855	0.1991	0.4919	0.2023
	W	-0.1527	-0.7246	-0.1868	0.4616	-0.1031	-0.0569					
Evapo	D	7.9724	1.9607	-0.3904	1.2470	-0.0886	-0.1185		0.3733	0.4681	0.2609	0.5932
	W	6.9672	1.8639	-0.2865	2.2131	-0.1846	-0.2240					
Sun	D	-2.0811	0.0179	0.0070	1.1915	3.0527	1.8217	1.7494	0.2798	0.2948	0.1615	0.1473
	W	0.0118	-0.3549	-0.2383								
Wind	D	1.5048	-0.0641	0.1327	0.4486	-0.0718	-0.0270		0.3153			
	W	1.3822	0.0314	0.1695	0.4822	-0.1150	-0.0313					
Max Hum	D	-1.3835	-0.4246	-0.3093	1.2037	-0.2473	-0.1653		0.2501			
	W	-1.0464	-0.4250	-0.2633	1.2606	-0.1815	-0.1403					
Min Hum	D	1.1584	0.1583	-0.1487	0.4467	0.0138	-0.0290		0.4710			
	W	0.7990	0.1551	-0.0416	0.6503	-0.1175	-0.0614					

TABLE 4.23 Parameter estimates of Model T for Nelspruit

Variable	Sta- tus	$\hat{\alpha}_1$	Mean	$\hat{\alpha}_3$	Standard Deviation			(WD)	Autocorrelation coefficient			
			$\hat{\alpha}_2$		$\hat{\xi}_1$ (DD)	$\hat{\xi}_2$ (WW)	$\hat{\xi}_3$ (DW)		(DD)	(WW)	(DW)	(WD)
Max Temp	D	-0.5530	-0.2991	-0.0618	0.3496	0.0415	-0.0825		0.5130	0.1863	0.3670	0.2989
	W	-0.3810	-0.4587	-0.1973	0.4045	0.0255	-0.0425					
Min Temp	D	0.0386	-0.7716	-0.1403	0.5612	-0.1514	-0.0713		0.3813	0.0797	0.1556	0.0717
	W	-0.5269	-0.2497	-0.0071	0.7003	-0.2251	-0.0918					
Sun	D	-1.0839	0.5190	-0.1531	1.4358	3.9054	2.7535	3.2377	0.1700	0.2423	0.1597	0.1180
	W	1.4841	-0.9117	-1.0326								
Wind	D	2.0032	-0.0014	0.1453	0.2386	0.3743	0.2590	0.3336	0.0001	0.3037	0.1856	0.1613
	W	2.0175	-0.0235	0.1640								
Max Hum	D	-1.4927	-0.0301	-0.2512	0.5749	-0.0730	-0.0896		0.3661	0.1056	0.1539	0.2178
	W	-2.0923	0.1254	-0.0859	0.7741	-0.0250	-0.0894					
Min Hum	D	-0.2500	-0.4921	-0.0472	0.5716	-0.0267	-0.0580		0.1953			
	W	-0.8259	-0.1266	0.1330	0.9221	-0.2392	-0.0932					

TABLE 4.24 Parameter estimates of Model T for Cedara

Variable	Sta- tus	$\hat{\alpha}_1$	Mean $\hat{\alpha}_2$	$\hat{\alpha}_3$	Standard Deviation			(WD)	Autocorrelation coefficient			
					$\hat{\xi}_1$ (DD)	$\hat{\xi}_2$ (WW)	$\hat{\xi}_3$ (DW)		(DD)	(WW)	(DW)	(WD)
Max Temp	D	0.0817	-0.1296	0.0067	0.5511	0.1035	-0.0839		0.2398	0.1117	0.0036	0.0444
	W	0.7722	-0.1986	-0.0917	0.8740	0.0132	-0.0792					
Min Temp	D	97.70	51.43	14.11	23.82	-2.624	-0.781		0.6092	0.3651	0.5537	0.3714
	W	110.28	46.56	14.04	21.60	-2.705	-1.667					
Sun	D	-1.5666	0.2819	-0.2128	1.2486	4.2278	2.8030	2.1725	0.1643	0.1914	-0.1618	-0.0123
	W	1.4590	0.1536	-0.6015								
Wind	D	1.7531	-0.0856	0.1766	0.3146	-0.0541	-0.0150		0.2607	0.0794	0.0000	0.1909
	W	1.6761	-0.0900	0.1703	0.3782	-0.0339	-0.0351					
Max Hum	D	-2.1008	-0.3668	-0.3021	1.0044	-0.1101	-0.0460		0.1672	0.1115	-0.0599	0.1598
	W	-3.0014	-0.2290	-0.0672	0.8444	0.0607	-0.0111					
Min Hum	D	-0.0160	-0.4920	-0.0928	0.6204	1.2461	0.9200	0.7907	0.2142	0.2060	-0.0295	0.0743
	W	-0.8994	-0.2833	0.0296								

TABLE 4.25 Parameter estimates of Model T for Hoopstad

Variable	Sta- tus	$\hat{\alpha}_1$	Mean $\hat{\alpha}_2$	$\hat{\alpha}_3$	Standard Deviation			Autocorrelation coefficient				
					$\hat{\xi}_1$ (DD)	$\hat{\xi}_2$ (WW)	$\hat{\xi}_3$ (DW)	(WD)	(DD)	(WW)	(DW)	(WD)
Max Temp	D	-0.5078	-0.5004	0.0254	0.2893	0.8638	0.4752	0.5262	0.5808	0.8066	0.7517	0.6951
	W	-0.3431	-0.5375	-0.0210								
Min Temp	D	-0.2257	-1.1626	-0.2329	0.3919	-0.0033	-0.0244		0.5982	0.4282	0.5456	0.3609
	W	-0.5434	-1.0062	-0.2076	0.3728	-0.0538	0.0099					
Sun	D	-1.8242	0.0955	0.0975	0.6753	2.9114	1.6379	1.2875	0.2845	0.2908	0.0003	0.1339
	W	0.0924	-0.3002	-0.0487								
Wind	D	2.0979	-0.2117	0.2139	0.4114	0.0001	-0.0514		0.4733	0.5789	0.0003	0.5387
	W	2.0284	-0.0710	0.2737	0.4468	0.0452	-0.0037					
Max Hum	D	-0.5581	0.4988	-0.4000	0.6267	-0.0181	-0.0186		0.5338	0.2678	0.2790	0.2456
	W	-0.7064	0.2786	-0.2623	0.8009	-0.0600	-0.0189					
Min Hum	D	1.1107	0.0944	-0.2305	0.3903	1.0292	0.6553	0.6209	0.5566	0.8081	0.7663	0.6498
	W	0.7927	0.1593	-0.1760								

efficient method that can be applied for the estimation of missing records, and in terms of accuracy, it performs at least as well as the other methods. The EM algorithm is a very general iterative method for maximum likelihood estimation in incomplete data sets. It comprises of the following steps:

1. Missing values are replaced by estimated values.
2. Parameters are estimated.
3. Missing values are re-estimated assuming that the new parameter estimates are correct.
4. Parameters are re-estimated and so forth, iterating until convergence.

A detailed explanation and the theory of the EM algorithm is given in Appendix E.

The estimates of the cross-correlation matrix for each station are given in Display 4.3. The matrices are symmetrical, therefore only the upper triangle is given. The order of the climate variables in the display is as follows:

maximum temperature, minimum temperature, evaporation, sunshine duration, windrun, maximum humidity and finally minimum humidity.

The results described in this chapter would suggest that the models are not inconsistent with the historical record.

The selected models have the following number of parameters:

The model for rainfall occurrences: has 6 parameters.

The model for rainfall depth: has 4 parameters.

Model 1: has 161 parameters.

Model 3: has 126 parameters.

Model 4: has 119 parameters.

Model 5: has 140 parameters.

Of course the tests described in this chapter cover only some limited aspects of the fit. The issue of model validation is considered more exhaustively in Chapter 6.

Elsenburg	[	1.000	0.606	0.440	-0.430	0.235	0.073	0.711
			1.000	0.219	-0.507	0.281	-0.154	0.614
				1.000	-0.550	-0.291	0.128	0.451
					1.000	-0.066	0.025	-0.466
						1.000	-0.127	0.162
							1.000	-0.072
								1.000
Kakamas	[	1.000	-0.038	0.385	0.193	0.000	0.006	-0.027
			1.000	0.294	-0.234	0.012	-0.010	0.028
				1.000	0.102	-0.012	-0.021	-0.002
					1.000	-0.007	0.112	0.062
						1.000	0.040	0.188
							1.000	0.307
								1.000
Middelburg	[	1.000	-0.354	-0.488	0.413	0.036	0.014	0.002
			1.000	-0.170	-0.270	0.141	-0.015	0.013
				1.000	-0.386	-0.349	-0.027	0.002
					1.000	0.056	0.007	-0.002
						1.000	-0.187	-0.251
							1.000	-0.231
								1.000
Nelspruit	[	1.000	-0.650	0.620	-0.155	-0.328	-0.718	
			1.000	-0.668	0.082	0.307	0.637	
				1.000	-0.123	-0.349	-0.594	
					1.000	-0.225	0.059	
						1.000	0.231	
							1.000	
Cedara	[	1.000	0.060	0.689	-0.236	-0.461	-0.480	
			1.000	0.130	-0.112	0.029	-0.028	
				1.000	-0.164	-0.305	-0.403	
					1.000	-0.039	0.095	
						1.000	0.212	
							1.000	
Hoopstad	[	-1.000	-0.343	0.351	-0.398	0.086	0.026	
			1.000	-0.375	0.165	-0.082	-0.013	
				1.000	-0.139	-0.043	0.023	
					1.000	-0.063	-0.007	
						1.000	-0.002	
							1.000	

DISPLAY 4.3 Estimated cross-correlation matrices for each station

## CHAPTER 5

### ALGORITHMS

This chapter describes the various procedures to be followed during model implementation and later during generation of climate sequences.

“Custom built” computer programs to carry out the preliminary analysis, to fit the models, to validate the models and finally to generate climate sequences have been written in ANSI 77 FORTRAN. The programs written conform to the full ANSI standard except for programs 6 and 8 where the array CLIMA is dimensioned using the HUGE attribute, which is an extension to the full ANSI standard. This was necessary when the climate data sets consisted of more than 9 years of daily data. Standard Fortran programs without this attribute should be no problem on a mainframe. Appendix D gives information where a listing of these programs (referred to in the algorithms below) can be obtained. The algorithms described here were all implemented on an IBM compatible PC micro-computer.

The following algorithms are discussed in this chapter:

- Algorithm for fitting the rainfall model.
- Algorithm for generating artificial rainfall sequences.
- Algorithm for fitting Model 1 to climate sequences.
- Algorithm for generating climate sequences using Model 1.
- Algorithm for fitting Model 3 to climate sequences.
- Algorithm for fitting Model 4 to climate sequences.
- Algorithm for fitting Model 5 to climate sequences.
- Algorithm for implementing Model T.
- Algorithm for generating climate sequences using Model T.

#### **Algorithm for implementing the rainfall model**

The following information is required for the parameter estimation programs and must be computed from the historical record:

*NT* the number of periods in the year (e.g. 365 for daily data).

$NY$  the number of years of data (including the missing values).

For each  $t = 1, 2, \dots, NT$ .

$NW(t)$  the number of times it was wet in period  $t - 1$  and there was an observation in period  $t$ .

$NRR(t)$  the number of times it was wet in period  $t - 1$  and wet in period  $t$ .

$ND(t)$  the number of times it was dry in period  $t - 1$  and there was an observation in period  $t$ .

$N\overline{RR}(t)$  the number of times it was dry in period  $t - 1$  and wet in period  $t$ .

$R(i, t)$  the  $i$ th non-zero rainfall depth in period  $t$ ,  $i = 1, 2, \dots, NR(t)$ .

$NR(t)$  the number of times it was wet in period  $t$ .

#### **Algorithm for estimating the probabilities of wet and dry sequences**

Step 1: Prepare data sets  $NW(t)$  and  $NRR(t)$ .

- Program 1
- if any of the  $NW(t)$  are equal to zero, then delete time period  $t$  from data set.

Step 2: Estimate the parameters for the probability that a wet period follows a wet period.

- Program 2.

Step 3: Prepare data sets  $ND(t)$  and  $N\overline{RR}(t)$ .

- Program 1
- if any of the  $ND(t)$  are equal to zero, delete time period  $t$  from the data set.

Step 4: Estimate the parameters for the probability that a wet day follows a dry period.

- Program 2.

#### **Algorithm to estimate the mean rainfall in wet periods**

Step 1: Prepare the data sets  $NR(t)$  and  $R(i, t)$ .

- Program 1.

Step 2: Estimate the parameters of the mean.

Step3: Estimate the coefficient of variation.

– Program 3 (does Steps 2 and 3).

#### **Algorithm for generating artificial rainfall sequences**

Step 1: Set initial state of day to be dry.

Step 2: Generate uniform random number between 0 and 1, inclusive ( $U(0,1)$ ).

Step 3: If  $U(0,1)$  random number is less than the probability of a wet day following a day with the status of the previous time period then

– the status of the present time period is wet.

Otherwise

– the status of the present time period is dry.

Step 4: If present state is wet than determine the rainfall depth.

Step 5: Repeat steps from Step 2 until enough rainfall sequences have been generated.

– Program 4.

#### **Algorithm for implementing Model 1 to climate sequences**

The following information is required for the parameter estimation programs and must be computed from the historical records:

$NT$  the number of periods in the year.

$NY$  the number of years of data.

$NV$  the number of variables in the model.

For each  $t = 1, 2, \dots, NT$  and for each variable

$m(t)$  the mean of the climate variable at time  $t$ .

$s(t)$  the standard deviation of the climate variable at time  $t$ .

For each variable do:

Step 1: Condition data set according to the wet or dry status of the day. That is, a record is kept of the time periods that had rain and the time periods that had no rain.

– Program 5.

For each conditioned data set do Step 2 – Step 6:

Step 2: Compute the daily mean vector,  $m(t)$  .

– Program 6.

Step 3: Estimate the parameters of the mean.

– Program 7.

Step 4: Compute the daily standard deviation vector,  $s(t)$  .

– Program 8.

Step 5: Estimate the parameters of the standard deviation.

– Program 7.

Step 6: Obtain the standardized residual series by subtracting the estimated daily mean function and dividing by the estimated daily standard deviation function.

– Program 9.

Step 7: Once the residual time series has been calculated for each variable, estimate the lag 0 and lag 1 cross-correlation coefficients.

– Program 10.

Step 8: Using estimates obtained in Step 7 compute the matrices  $A$  and  $B$  .

– Program 11.

#### Algorithm for generating artificial climate sequences using Model 1

Step 1: Generate rainfall sequence (algorithm given above)

For each variable do:

Step 2: Generate a normal random number from a distribution with a mean of zero and a standard deviation of unity ( $N(0,1)$ ) .

Step 3: Generate residual time series by:

$$\chi_{i,t} = \hat{A} \chi_{i,t-1} + \hat{B} \epsilon_{i,t}.$$

– the initial condition of the residual time series is taken to be equal to zero, i.e.

$$\chi_{1,0} = 0 .$$

Step 4: Generate climate sequences by:

$$S_{i,t} = \begin{cases} \chi_{i,t} \hat{\sigma}_t^W + \hat{\mu}_t^W & \text{if wet} \\ \chi_{i,t} \hat{\sigma}_t^D + \hat{\mu}_t^D & \text{if dry.} \end{cases}$$

Step 5: Repeat all of the above steps until the desired amount of climate sequences have been generated. So that the generating process has a chance of stabilizing itself, the first year of data generated is ignored.

– Program 12.

### Algorithm for implementing Model 3 to climate sequences

The following information is required for the parameter estimation programs and must be computed from the historical records:

$NT$  the number of periods in the year.

$NY$  the number of years of data.

$NV$  the number of variables in the model.

$T$  the total number of observations.

$N(DD)$  the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was dry,  $t = 1, 2, \dots, T$ .

$N(WW)$  the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was wet.

$N(DW)$  the set of time periods  $t$  such that period  $t$  was wet and period  $t - 1$  was dry.

$N(WD)$  the set of time periods  $t$  such that period  $t$  was dry and period  $t - 1$  was wet.

$C(DD)$  number of elements in the set  $N(DD)$ .

$C(WW)$  number of elements in the set  $N(WW)$ .

$C(DW)$  number of elements in the set  $N(DW)$ .

$C(WD)$  number of elements in the set  $N(WD)$ .

For each variable do:

Step 1: Estimate initial parameters of the mean function by performing Step 1 through to Step 3 of the algorithm for parameter estimation of Model 1.

Step 2: Prepare the data sets of possible sequences, i.e.  $N(DD)$ ,  $N(WW)$ ,  $N(DW)$  and  $N(WD)$ . Compute  $C(DD)$ ,  $C(WW)$ ,  $C(DW)$  and  $C(WD)$ .

– Program 13.

Step: 3: Estimate initial autocorrelation coefficients for each of the possible sequences.

– Program 14.

Step 4: Estimate initial standard deviation function for each of the possible sequences.

– Program 15.

Step 5: Estimate parameters of the mean function, standard deviation function and the autocorrelation coefficients, iterating until convergence is met by all parameters.

– Program 16 (or Program 17).

Step 6: Obtain residual time series by

$$e_{i,t} = \frac{S_{i,t} - \hat{\mu}_t}{\hat{\sigma}} - \hat{\theta} \frac{S_{i,t-1} - \hat{\mu}_{t-1}}{\hat{\sigma}}$$

where  $\hat{\mu}_t, \hat{\mu}_{t-1}, \hat{\theta}$  and  $\hat{\sigma}$  are chosen depending on which sequence the time periods  $t$  and  $t-1$  satisfy.

#### Algorithm for implementing Model 4 to climate sequences

The information necessary for parameter estimation programs is the same as for Model 3.

For each variable do:

Step1: Estimate initial parameters of the mean function by performing Step 1 through to Step 3 of the algorithm for parameter estimation of Model 1.

Step 2: Estimate initial parameters of the standard deviation function by performing Step 4 and Step 5 of the algorithm for implementing Model 1.

Step 3: Estimate initial autocorrelation coefficient.

– Program 18.

Step 4: Prepare the data sets of possible sequences,  $N(DD), N(WW), N(DW)$  and  $N(WD)$ . Compute  $C(DD), C(WW), C(DW)$  and  $C(WD)$ .

– Program 13.

Step 5: Estimate parameters of the mean function, standard deviation function and the autocorrelation coefficient, iterating until convergence is met by all parameters.

– Program 19 (or Program 20).

Step 6: Obtain residual time series by:

$$e_{i,t} = \frac{S_{i,t} - \hat{\mu}_t}{\hat{\sigma}_t} - \hat{\theta} \frac{S_{i,t-1} - \hat{\mu}_{t-1}}{\hat{\sigma}_{t-1}}$$

where  $\hat{\mu}_t, \hat{\mu}_{t-1}, \hat{\sigma}_t$  and  $\hat{\sigma}_{t-1}$  are chosen depending on which sequence the time periods  $t$  and  $t-1$  satisfy.

### Algorithm for implementing Model 5 to climate sequences

The information necessary for parameter estimation programs is the same as for Model 3.

For each variable do:

Step 1: Estimate initial parameters of the mean function and of the standard deviation function by performing Step 1 through to Step 5 of the algorithm for implementing Model 1.

Step 2: Estimate initial autocorrelation coefficients for each of the possible sequences by performing Step 3 of the algorithm for implementing Model 3.

Step 3: Prepare the data sets of possible sequences,  $N(DD), N(WW), N(DW)$  and  $N(WD)$ . Compute  $C(DD), C(WW), C(DW)$  and  $C(WD)$ .

– Program 13.

Step 4: Estimate parameters of the mean function, standard deviation function and autocorrelation coefficients, iterating until convergence is met by all parameters.

– Program 21 (or Program 22).

Step 5: Obtain residual time series by:

$$e_{i,t} = \frac{S_{i,t} - \hat{\mu}_t}{\hat{\sigma}_t} - \hat{\theta} \frac{S_{i,t-1} - \hat{\mu}_{t-1}}{\hat{\sigma}_{t-1}}$$

where  $\hat{\mu}_t, \hat{\mu}_{t-1}, \hat{\sigma}_t, \hat{\sigma}_{t-1}$  and  $\hat{\theta}$  are chosen depending on which sequence the time periods  $t$  and  $t-1$  satisfy.

### Algorithm for implementing Model T

Step 1: From each residual time series obtained after fitting Models 3, 4 and 5, select for each variable the residual series from the model which produced the lowest Akaike's Information Criterion.

Step 2: Record, for each variable the time periods  $t$  for which a missing observation occurs.

– Program 23.

Step 3: Use the EM algorithm to estimate and replace missing values by this estimate.

– Program 24.

Step 4: Estimate the cross-correlation matrix,  $\Sigma$ .

– Program 25.

### Algorithm for generating artificial climate sequences using Model T

Step 1: Generate rainfall sequence.

For each variable do:

Step 2: Generate  $N(0, \hat{\Sigma})$  random number.

Step 3: Generate climate values according to the model chosen for that variable. For example, if Model 3 is chosen, then

$$S_{i,t} = \hat{\mu}_t + \hat{\sigma} \left[ e_{i,t} + \hat{\theta} \frac{S_{i,t-1} - \hat{\mu}_{t-1}}{\hat{\sigma}} \right]$$

where  $\hat{\mu}_t, \hat{\mu}_{t-1}, \hat{\sigma}$  and  $\hat{\theta}$  are chosen depending on the sequence  $t$  and  $t-1$  satisfy.

If Model 4 is chosen, then

$$S_{i,t} = \hat{\mu}_t + \hat{\sigma}_t \left[ e_{i,t} + \hat{\theta} \frac{S_{i,t-1} - \hat{\mu}_{t-1}}{\hat{\sigma}_{t-1}} \right]$$

where  $\hat{\mu}_t, \hat{\mu}_{t-1}, \hat{\sigma}_t$  and  $\hat{\sigma}_{t-1}$  are chosen depending on the sequence  $t$  and  $t-1$  satisfy.

If Model 5 is chosen then

$$S_{i,t} = \hat{\mu}_t + \hat{\sigma}_t \left[ e_{i,t} + \hat{\theta} \frac{S_{i,t-1} - \hat{\mu}_{t-1}}{\hat{\sigma}_{t-1}} \right]$$

where  $\hat{\mu}_t, \hat{\mu}_{t-1}, \hat{\sigma}_t, \hat{\sigma}_{t-1}$  and  $\hat{\theta}$  are chosen depending on the sequence  $t$  and  $t-1$  satisfy.

Step 4: Repeat above steps until the desired amount of climate sequences have been generated.

– Program 26.

## CHAPTER 6

### GOODNESS OF FIT

Once a model has been identified and the parameters estimated, it remains to decide whether the model is adequate. Model validation is applied with the object of assessing the performance of the model and to uncover any possible lack of fit. In particular one wants to assess whether the model proposed and parameters estimated preserve the properties of the process being examined. This chapter summarizes the results of the checks carried out on Model 1 and Model T described in Chapter 3.

#### Validation of rainfall model

The rainfall model has been shown to be satisfactory in the various regions of South Africa (Zucchini and Adamson, 1984). They performed extensive checks on the properties of the model such as:

- (a) the annual mean and standard deviation and the distribution of annual totals and sum of  $k$  running totals,  $k = 1, 2, \dots, 5$ ,
- (b) the monthly means and standard deviations,
- (c) the expected number of wet days at different times of the year,
- (d) the distribution of runs of wet and dry days,
- (e) the distribution of  $n$ -day extreme rainfall.

The Markov chain/Weibull model adopted was found to preserve these properties. A number of these checks were repeated in this study. For a more complete model validation procedure see Zucchini and Adamson (1984).

Historical data (daily observations) were obtained for the weather stations Elsenburg, Kakamas, Middelburg, Nelspruit, Cedara and Hoopstad. More information on these records was given in Chapter 4.

Fifty years of simulated daily data were compared with the historical data on an annual, monthly and daily basis.

Table 6.1 gives both the historical and simulated annual mean number of wet days. This property has been adequately preserved by the model.

TABLE 6.1 Mean number of wet days per year

Station	Historical	Simulated
Elsenburg	91	92
Kakamas	16	19
Middelburg	63	63
Nelspruit	96	95
Cedara	150	149
Hoopstad	72	80

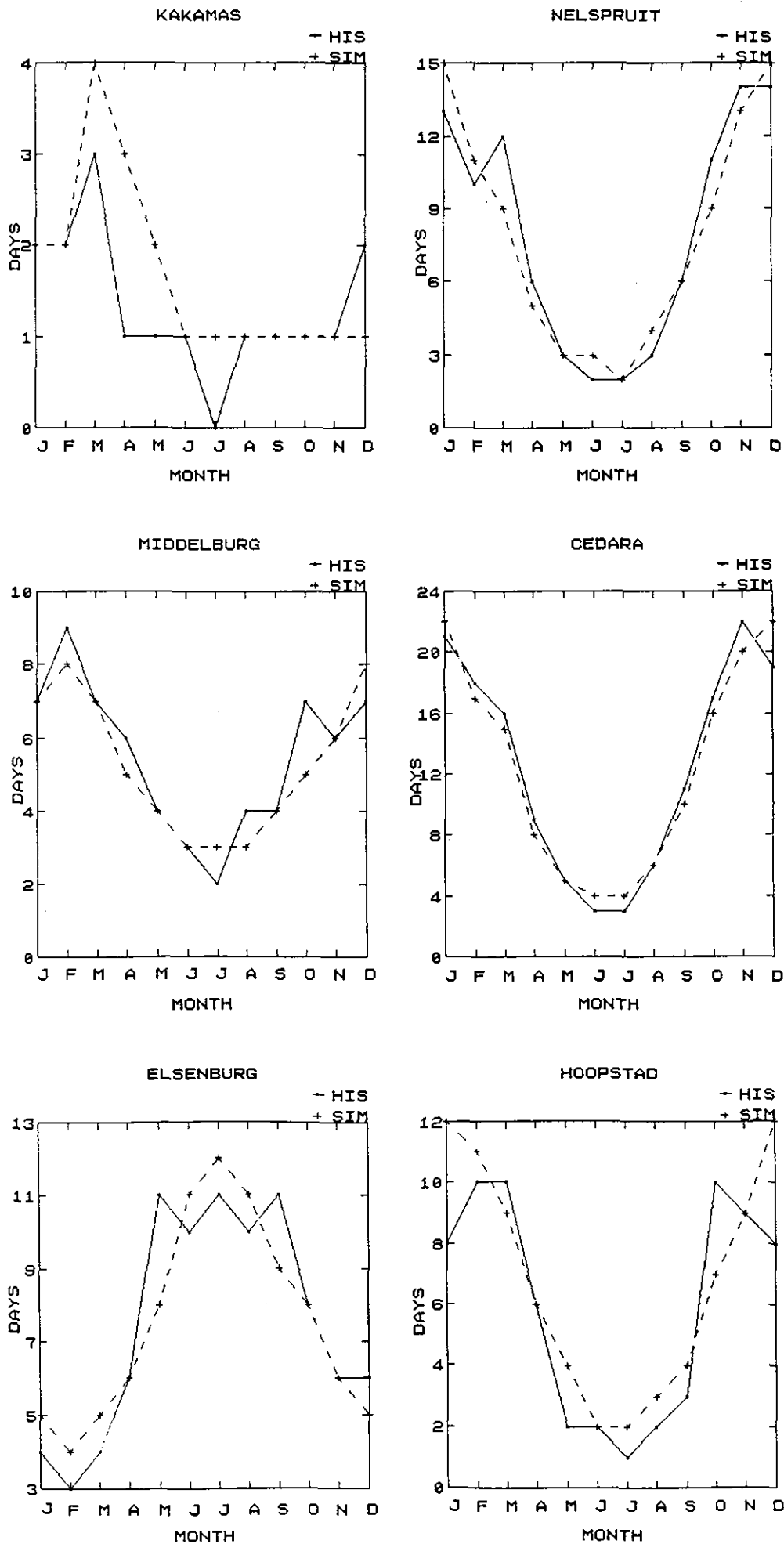
The mean number of wet days for each month has also been adequately preserved by the model (Figure 6.1)

It is especially important that the occurrence of wet days by season be adequately modelled as the generation of the other climate variables is conditioned on the occurrence of wet or dry days. The above results indicate that the Markov chain/Weibull model preserves the properties of the rainfall sequence at those locations.

The fits of the truncated Fourier series for the probability of having a wet day given a preceding wet day, and for the probability of having a wet day given the preceding one was dry, for each station, are shown in Figures 6.2 – 6.7. The fits are generally good.

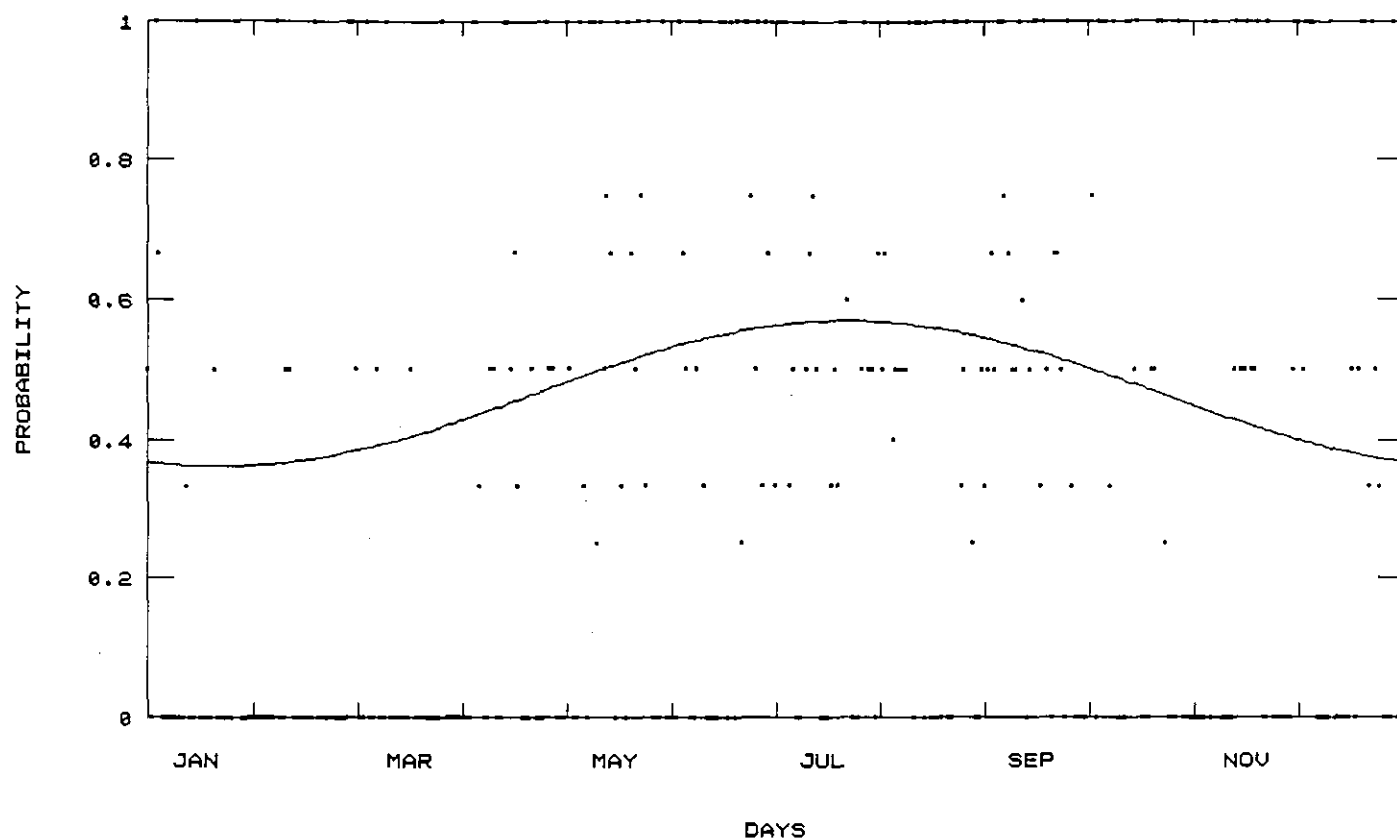
The interpretation of these figures requires some explanation. These are not ordinary regression equation fits with normally distributed residuals. The smooth line indicates the fitted probability for a binomial random variable where the number of trials is also random. The outcomes are *discrete* values representing the number of successes in a series of Bernoulli trials. This is analogous to a situation in which a coin, which has a probability  $p$  of landing heads, is tossed  $n$  times and the fraction of times the coin landed heads is recorded. The smooth line would then represent the (smoothly varying probability) and the points on the graph the proportion of heads. The visual impression that one gets from such a diagram might suggest that the fit is poor (because one is used to interpreting regressions with normally distributed residuals, that is continuous random residuals) when in fact the fit is very good. The latter is the case in Figures 6.2 – 6.7.

FIGURE 6.1 Historical and simulated mean number of wet days



**FIGURE 6.2** Empirical probabilities and estimates based on a 3 parameter model for  $P(W|W)$  and  $P(W|D)$  for Elsenburg

ELSENBURG -  $P(W|W)$



ELSENBURG -  $P(W|D)$

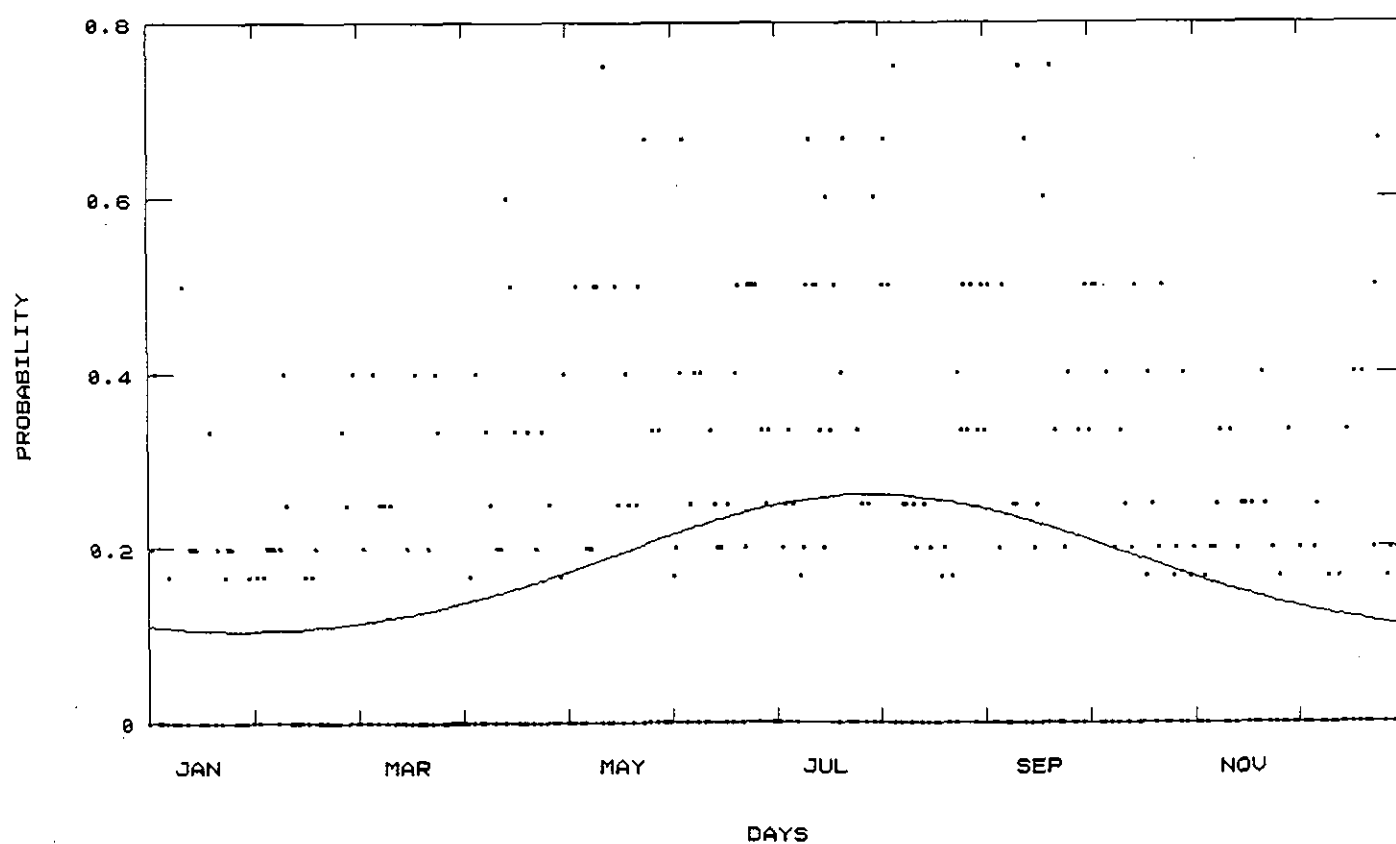
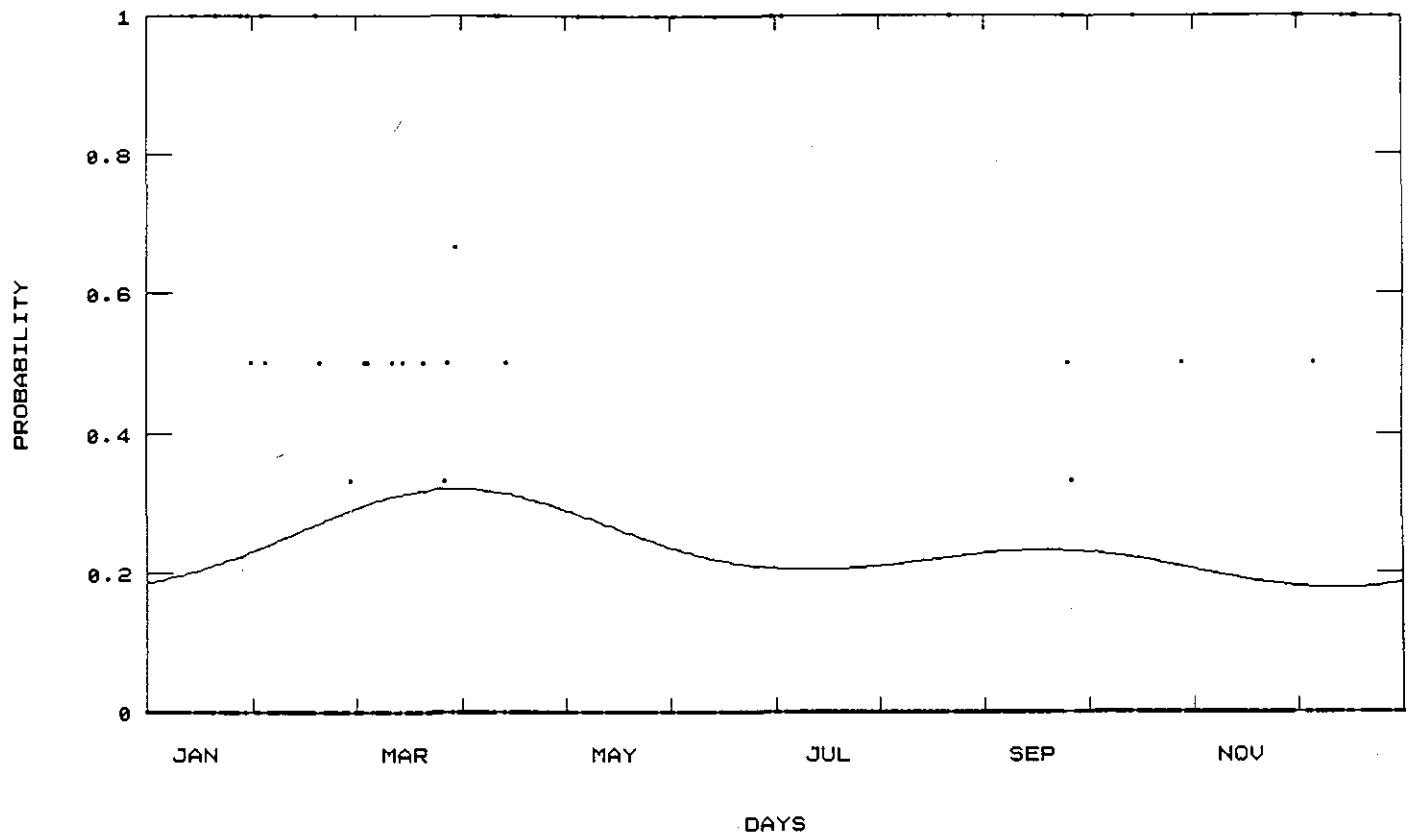
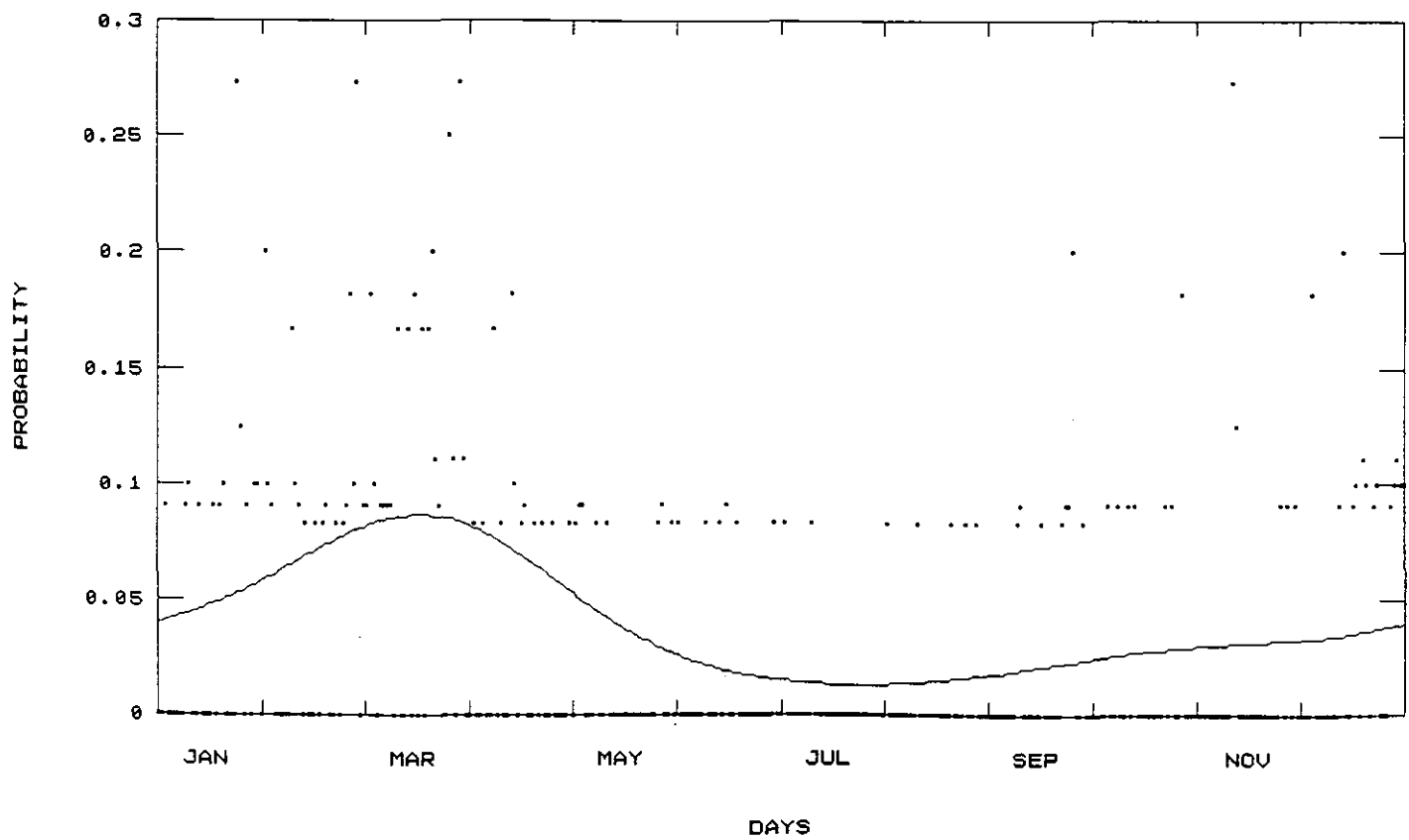


FIGURE 6.3 Empirical probabilities and estimates based on a 3 parameter model for  $P(W|W)$  and  $P(W|D)$  for Kakamas

KAKAMAS -  $P(W|W)$

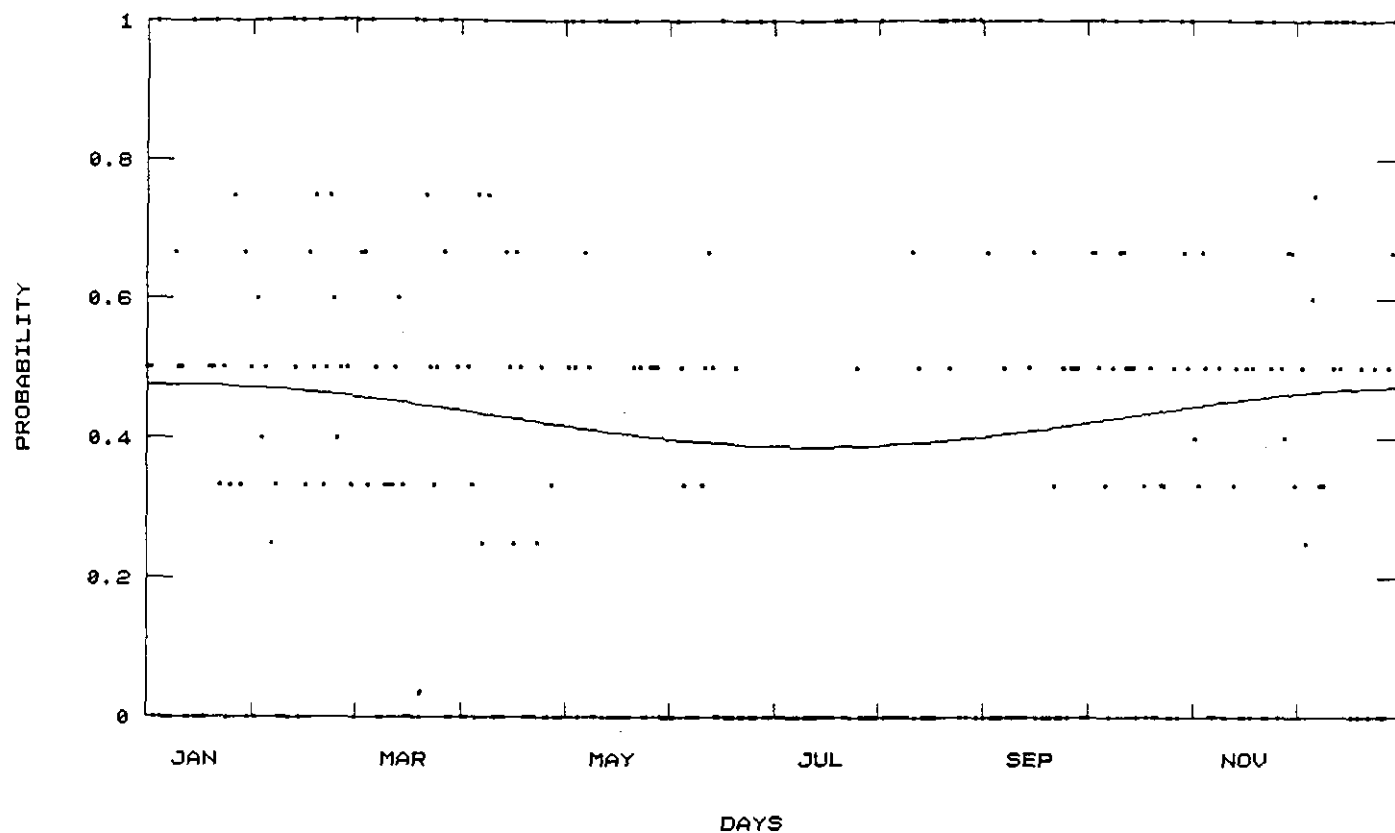


KAKAMAS -  $P(W|D)$

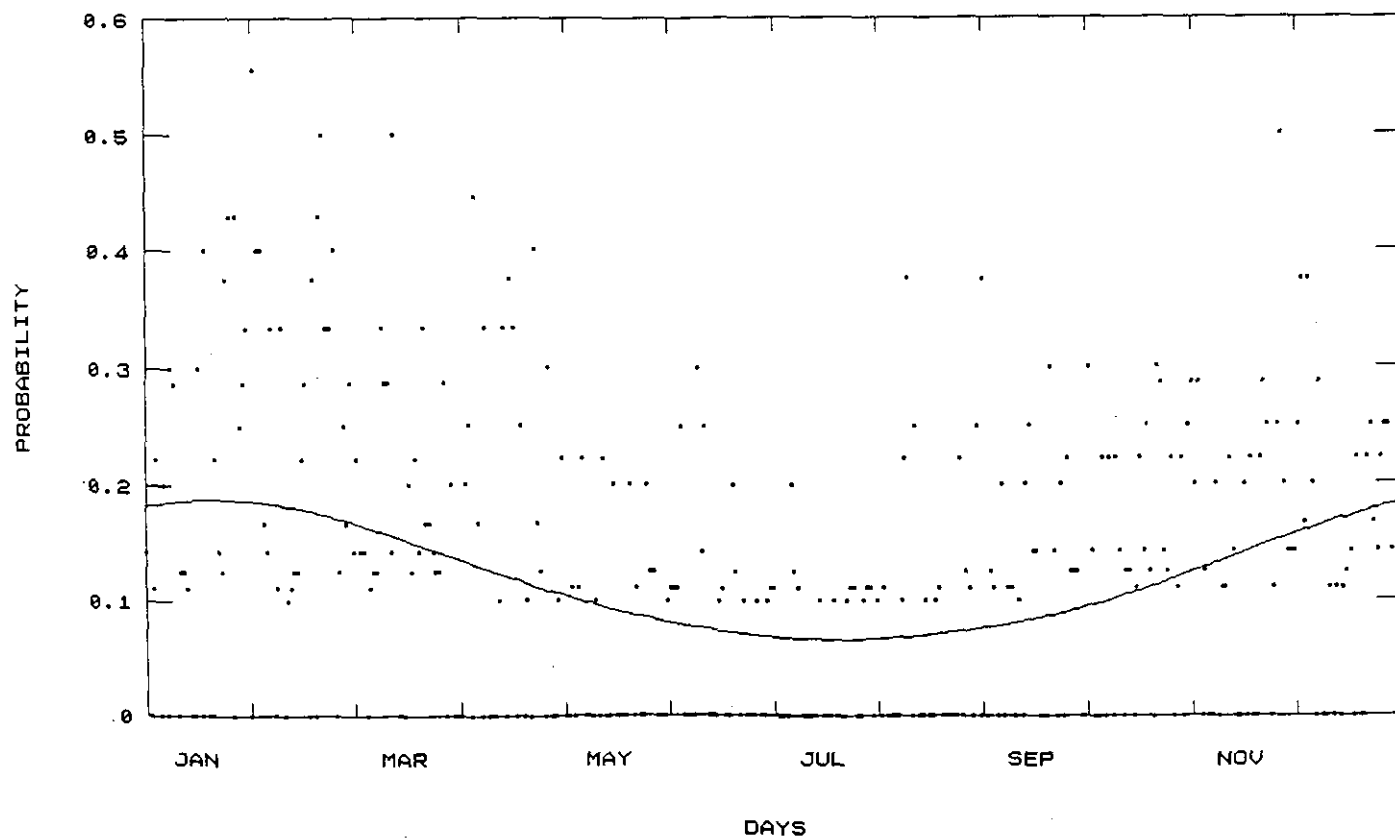


**FIGURE 6.4** Empirical probabilities and estimates based on a 3 parameter model for  $P(W|W)$  and  $P(W|D)$  for Middelburg

MIDDELBURG -  $P(W|W)$

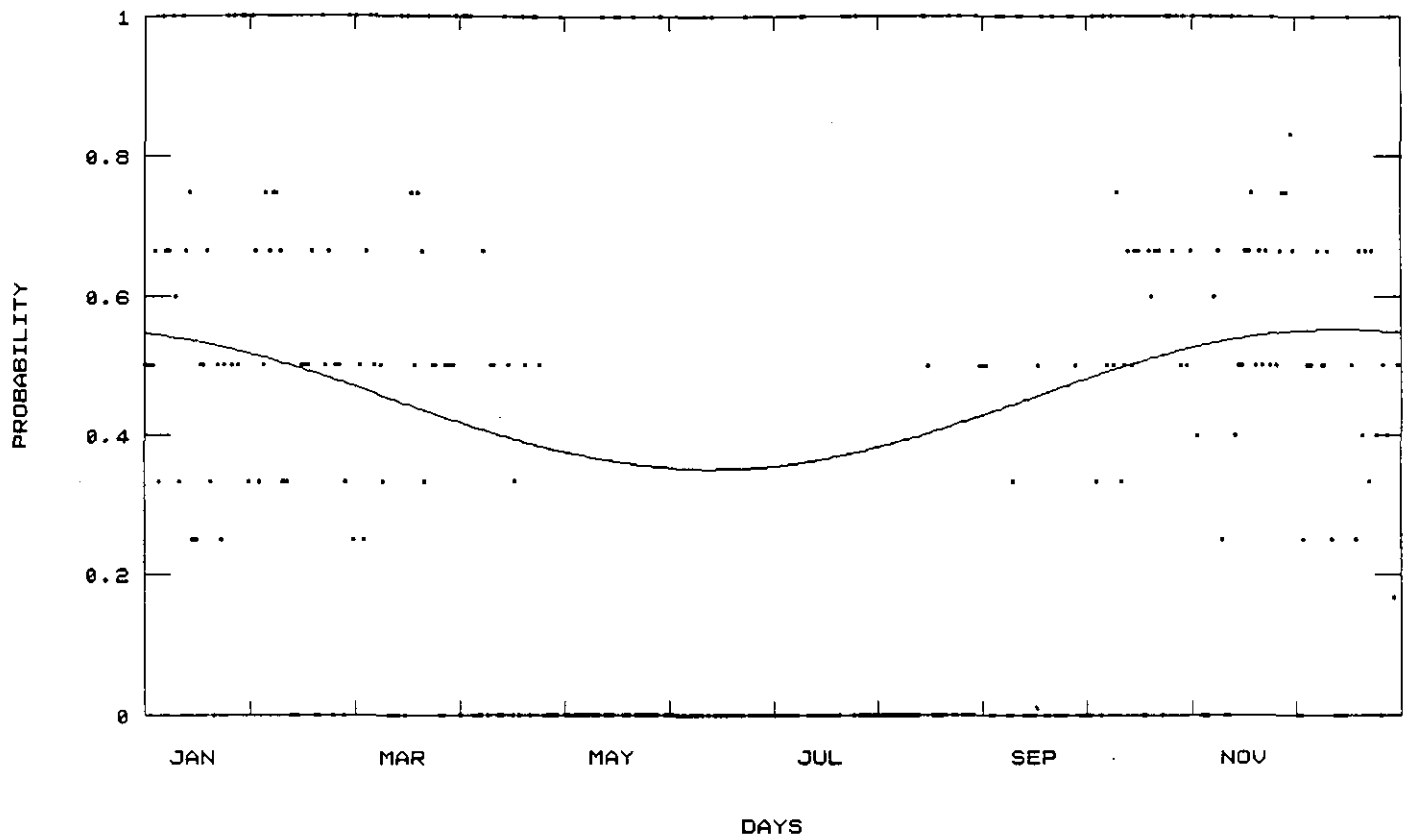


MIDDELBURG -  $P(W|D)$



**FIGURE 6.5** Empirical probabilities and estimates based on a 3 parameter model for  $P(W|W)$  and  $P(W|D)$  for Nelspruit

NELSPRUIT -  $P(W|W)$



NELSPRUIT -  $P(W|D)$

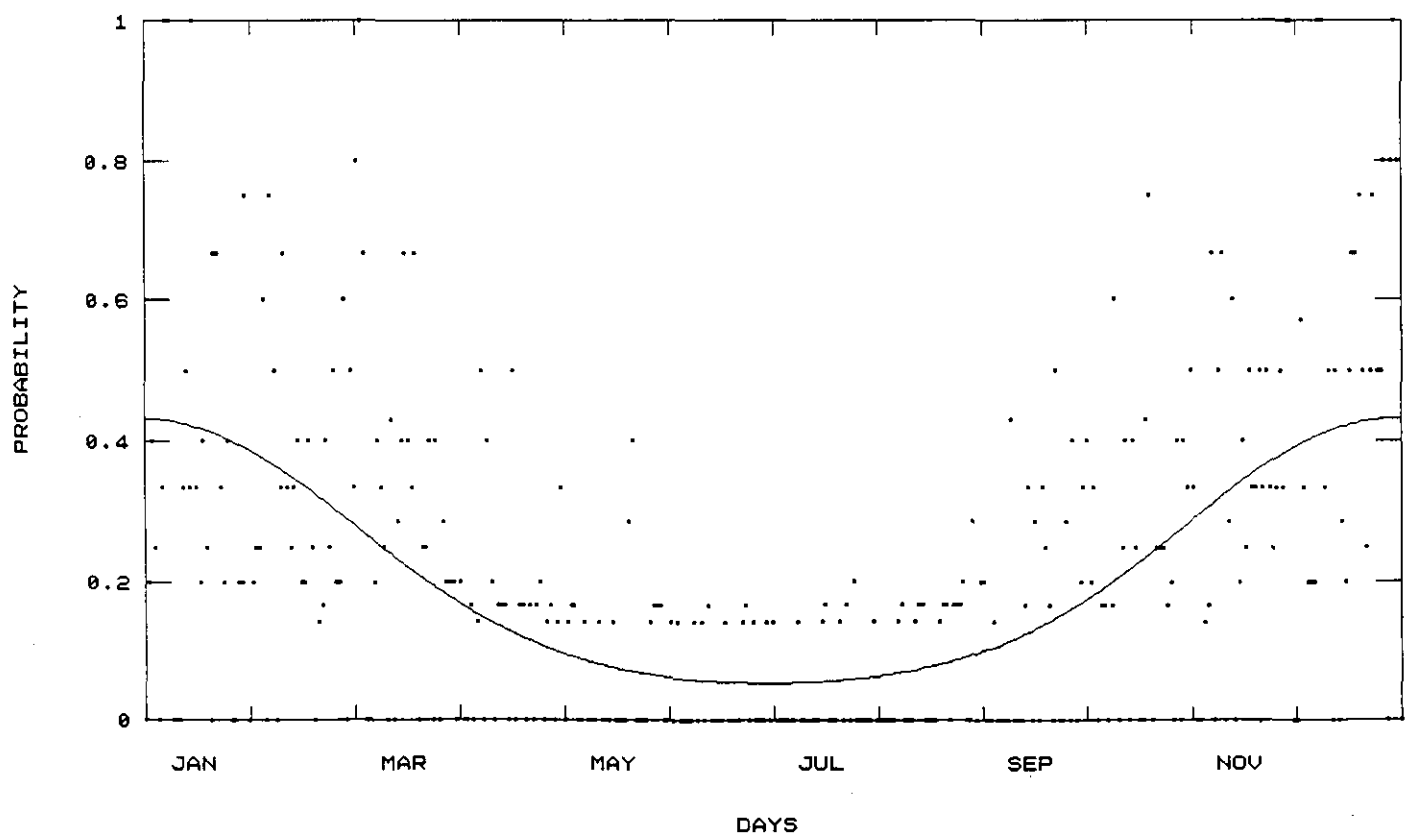
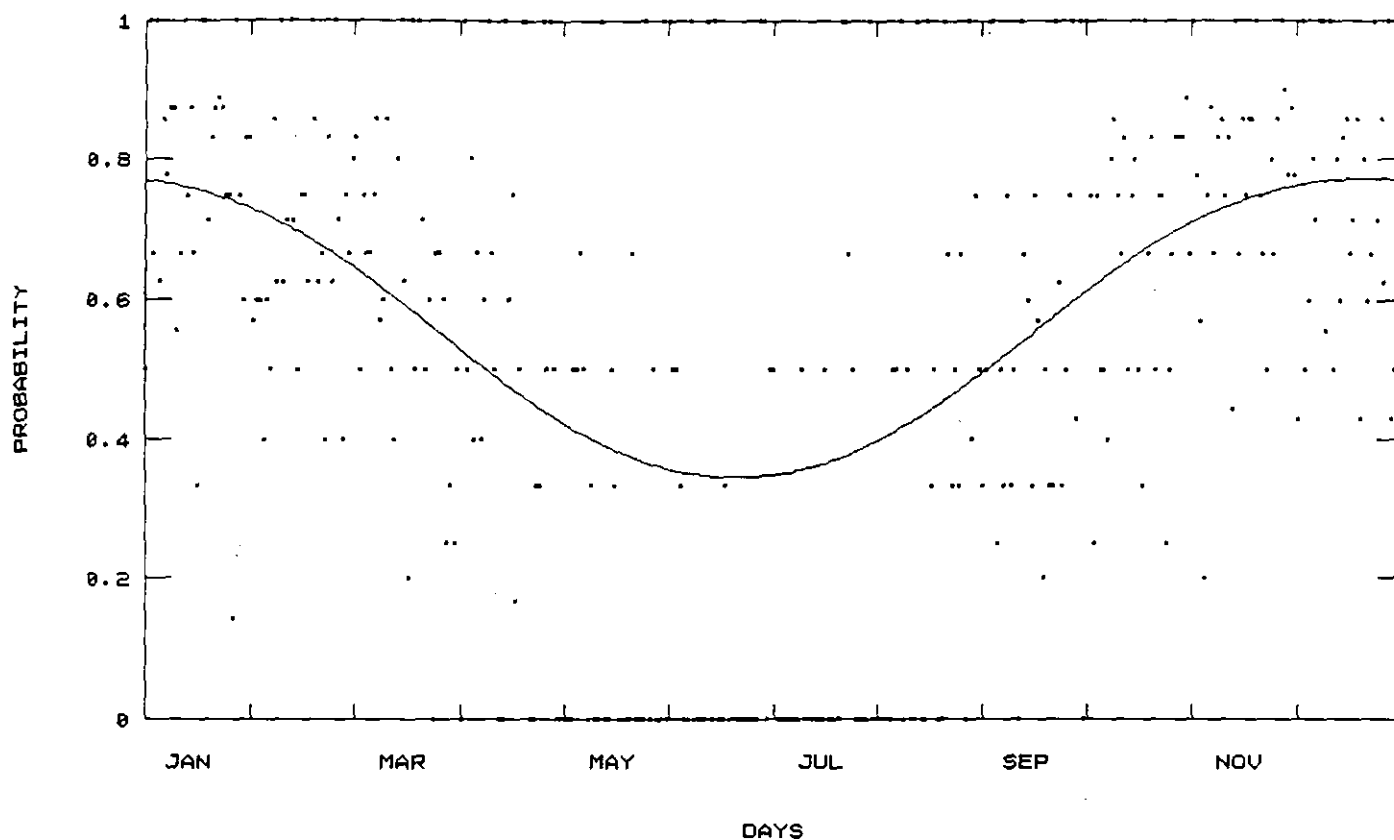
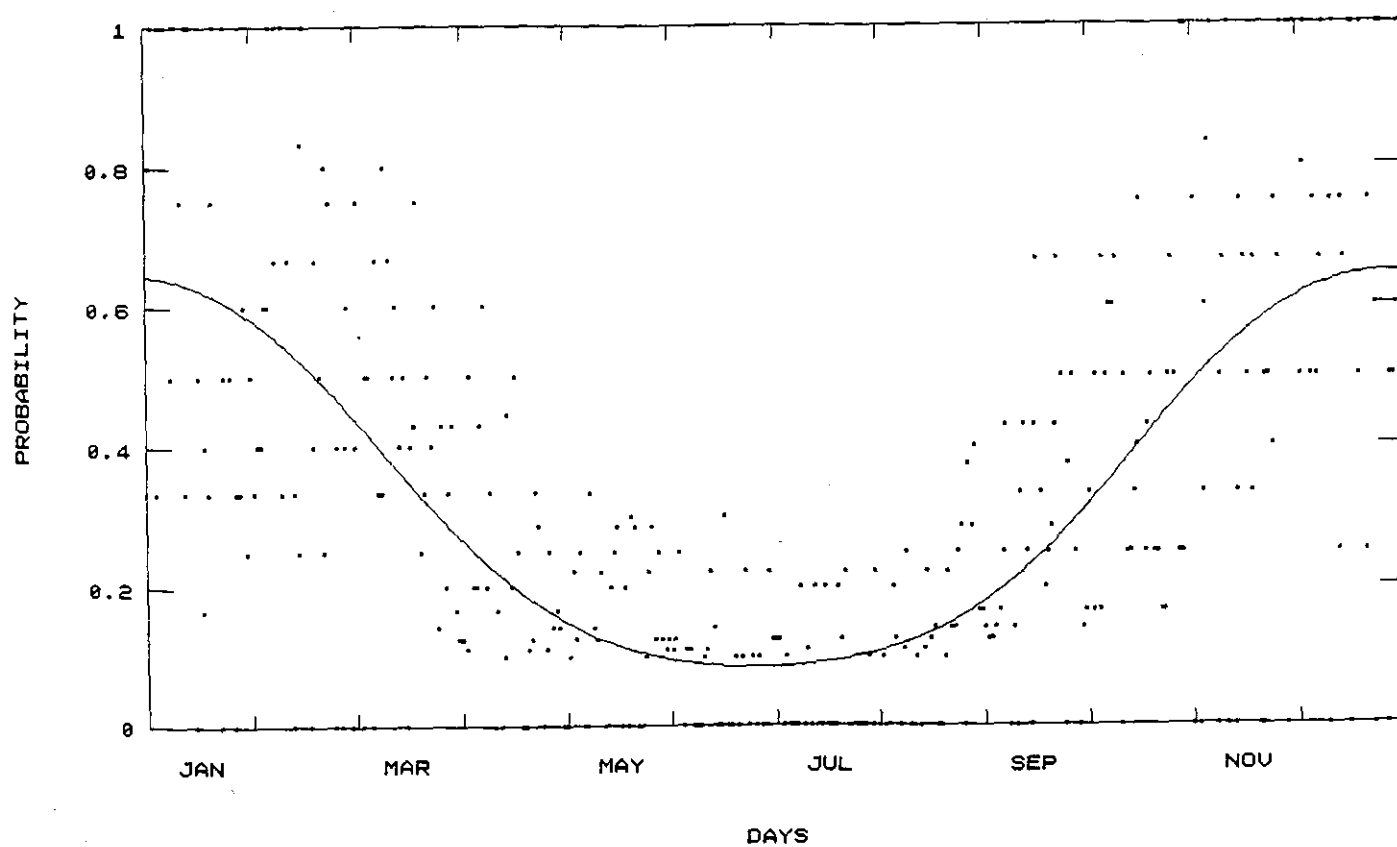


FIGURE 6.6 Empirical probabilities and estimates based on a 3 parameter model for  $P(W|W)$  and  $P(W|D)$  for Cedara

CEDARA -  $P(W|W)$

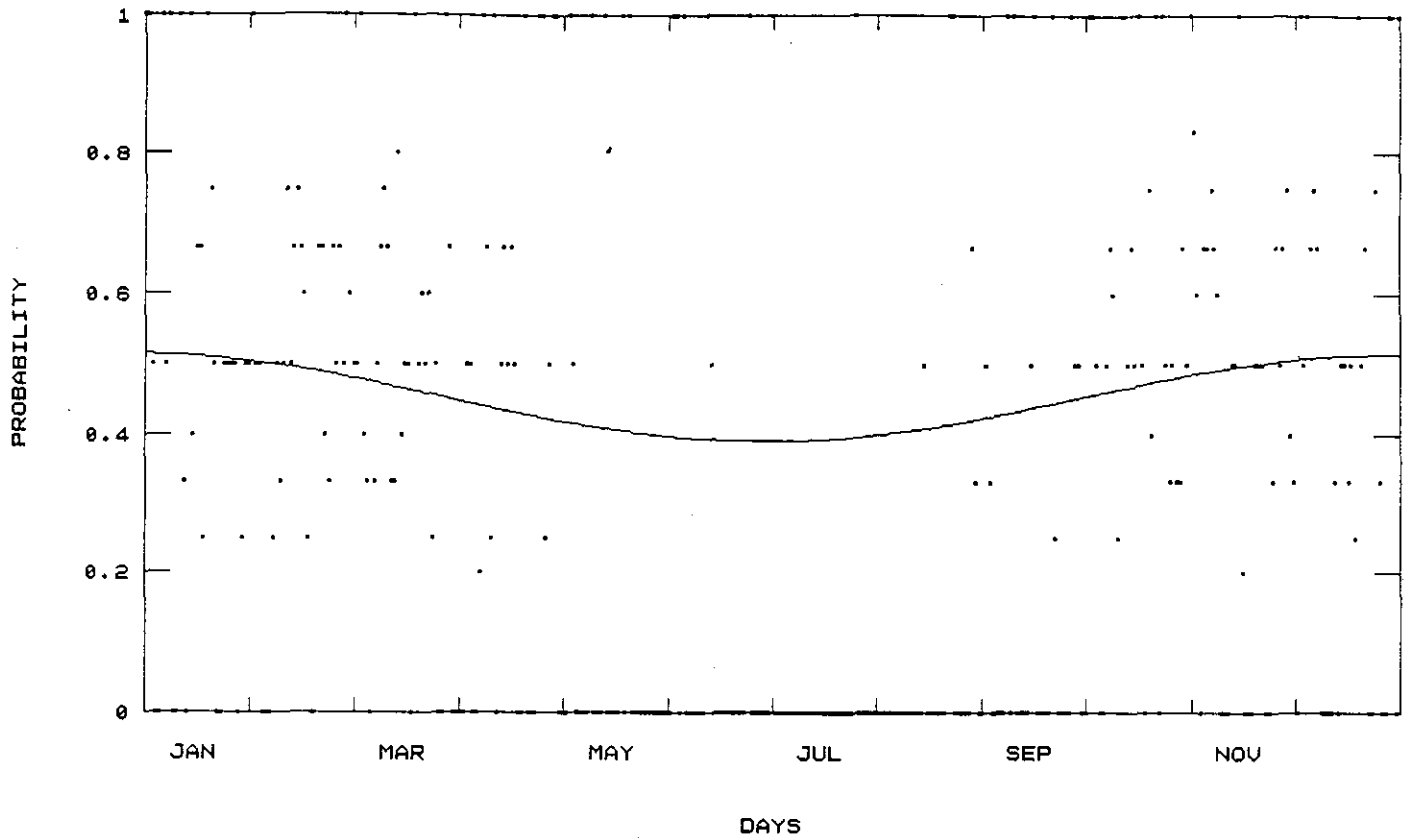


CEDARA -  $P(W|D)$

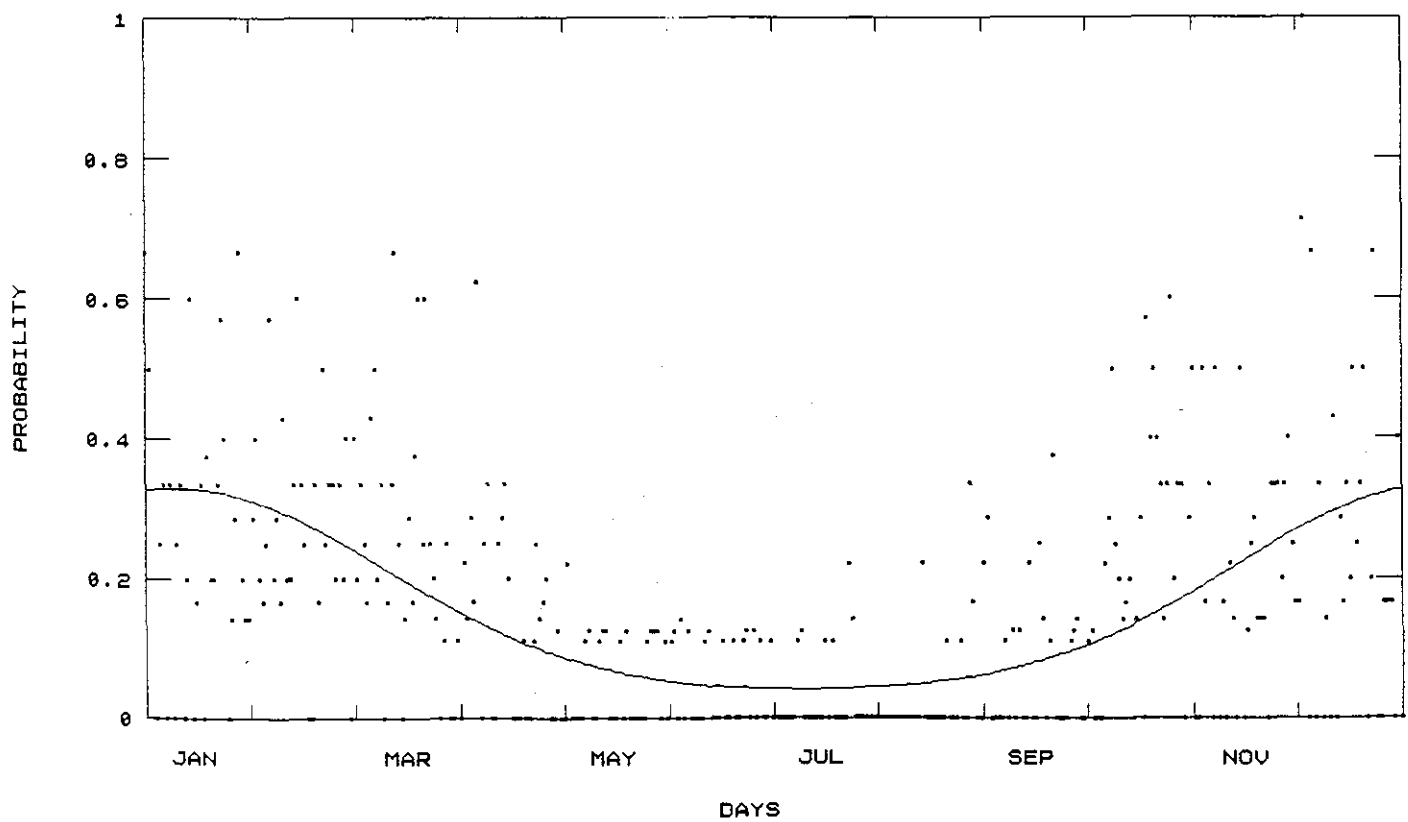


**FIGURE 6.7** Empirical probabilities and estimates based on a 3 parameter model for  $P(W|W)$  and  $P(W|D)$  for Hoopstad

HOOPSTAD -  $P(W|W)$



HOOPSTAD -  $P(W|D)$



### Validation of Climate Model

To consider the climate model as adequate in preserving the characteristics of the climate, the multivariate properties of the weather variables must be investigated as well as the univariate characteristics of each individual variable.

The following parameters and parameter functions must be preserved if one is to consider the climate model as satisfactory:

- (a) the annual mean and standard deviation for each climate variable for the unconditioned data and the data conditioned on the wet or dry status of the day,
- (b) the monthly means and standard deviations for each variable for the unconditioned data and the data conditioned on the wet or dry status of the day,
- (c) the extreme values of each climate variable, i.e. maximum and minimum daily values,
- (d) the autocorrelation within each variable for the unconditioned data and the data conditioned on the wet or dry status of the day,
- (e) the cross-correlation over all climate variables.

The checks above test either the multivariate part of the climate model, e.g. the cross-correlation over all variables, or the individual characteristics of each variable, e.g. the monthly means and standard deviations for each variable.

Again fifty years of simulated daily climate sequences were compared with the historical data on an annual, monthly and daily basis.

The following abbreviations are used in tables and figures:

Max Temp — Maximum Temperature

Min Temp — Minimum Temperature

Evapo — Evaporation

Sun — Sunshine Duration

Wind — Wind run

Max Hum — Maximum Humidity

Min Hum — Minimum Humidity

His — Historical Data

Mod 1 — Simulated data using Model 1

Mod T — Simulated data using Model T.

### **Validation of annual properties**

Table 6.2 shows the comparison of historical and simulated annual means for each variable and each station. This statistic has been adequately preserved by both models when the variables are conditioned on a wet day and when they are conditioned on the dry status of the day (Tables 6.3 – 6.4). There is however a slight underestimation of the annual mean for wet sequences by Model T for the variables wind run, maximum and minimum humidity at some of the stations. For Middelburg, the annual mean of wind run is slightly overestimated by Model T.

**TABLE 6.2 Comparison of historical and simulated annual mean**

Variable	Data	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	22.7	28.9	23.7	26.4	22.8	25.9
	Mod 1	22.7	28.9	23.5	26.3	22.8	26.1
	Mod T	22.4	29.2	24.4	26.1	22.9	27.9
Min Temp	His	10.5	13.1	6.8	13.3	10.1	7.9
	Mod 1	10.7	13.1	6.7	13.3	10.1	8.1
	Mod T	10.5	13.2	6.8	13.4	10.2	8.4
Evapo	His	5.7	8.9	6.4			
	Mod 1	5.7	8.9	6.3			
	Mod T	5.6	9.1	6.5			
Sun	His	8.3	9.8	9.2	7.3	6.9	9.2
	Mod 1	8.2	9.8	9.1	6.7	6.5	9.0
	Mod T	8.3	9.8	9.3	7.2	6.9	9.4
Wind	His	194.7	194.9	195.8	121.1	158.4	123.8
	Mod 1	193.0	196.0	195.4	125.6	157.5	123.1
	Mod T	192.5	195.0	195.4	122.6	158.3	119.8
Max Hum	His	92.5	62.2	80.3	81.8	88.8	74.0
	Mod 1	92.8	62.9	80.4	82.0	89.0	74.0
	Mod T	92.8	63.2	80.9	82.2	89.0	73.8
Min Hum	His	41.4	25.6	26.7	47.9	52.6	32.9
	Mod 1	41.2	24.9	26.6	48.4	53.2	32.9
	Mod T	41.9	24.9	26.4	48.0	51.1	28.1

**TABLE 6.3 Comparison of historical and simulated annual mean given a wet day**

Variable	Data	S t a t i o n					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	18.1	28.9	24.0	26.2	22.3	26.8
	Mod 1	18.1	27.8	23.5	25.9	22.3	27.1
	Mod T	19.0	30.5	26.5	26.4	22.5	30.2
Min Temp	His	10.9	17.8	10.9	16.6	13.0	13.6
	Mod 1	11.0	16.6	10.8	16.6	12.9	13.7
	Mod T	11.4	16.7	10.6	16.6	12.8	13.2
Evapo	His	2.8	7.2	5.2			
	Mod 1	2.6	6.4	5.0			
	Mod T	2.8	8.8	6.1			
Sun	His	4.2	5.7	6.0	4.6	4.4	6.0
	Mod 1	3.8	5.5	5.3	3.8	3.7	5.4
	Mod T	4.3	5.3	6.3	4.7	4.3	6.5
Wind	His	245.8	219.2	205.1	121.0	171.1	142.8
	Mod 1	242.9	215.0	203.6	126.2	169.9	138.7
	Mod T	235.4	213.5	211.3	123.8	170.3	129.7
Max Hum	His	94.3	75.8	83.3	87.8	95.0	81.7
	Mod 1	94.2	77.8	84.4	88.3	95.2	80.8
	Mod T	93.4	56.2	80.3	87.0	94.9	76.1
Min Hum	His	55.1	39.2	36.4	61.6	68.6	47.3
	Mod 1	55.0	38.4	36.6	62.9	69.5	47.3
	Mod T	54.0	29.7	31.5	59.5	65.8	34.2

**TABLE 6.4** Comparison of historical and simulated annual mean given a dry day

Variable	Data	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	24.1	28.9	23.6	26.5	23.3	25.6
	Mod 1	24.2	28.9	23.5	26.5	23.2	25.8
	Mod T	23.5	29.2	23.9	26.0	23.3	27.2
Min Temp	His	10.4	12.9	5.0	12.2	8.1	6.3
	Mod 1	10.5	12.9	5.8	12.1	8.1	6.6
	Mod T	10.2	13.0	6.0	12.2	8.4	7.1
Evapo	His	6.7	9.0	6.6			
	Mod 1	6.7	9.1	6.6			
	Mod T	6.6	9.1	6.6			
Sun	His	9.7	10.0	9.9	8.2	8.7	10.0
	Mod 1	9.7	10.0	9.9	7.8	8.5	10.0
	Mod T	9.6	10.0	9.9	8.1	8.6	10.1
Wind	His	117.7	193.7	193.9	121.2	149.4	118.6
	Mod 1	176.3	195.0	193.7	125.4	148.7	118.8
	Mod T	178.1	194.1	192.2	122.2	149.9	117.1
Max Hum	His	91.9	61.6	79.6	79.6	84.5	72.1
	Mod 1	92.3	62.1	79.5	79.8	84.7	72.2
	Mod T	92.6	63.5	81.1	80.5	84.9	73.2
Min Hum	His	36.8	25.0	24.7	43.1	41.3	29.1
	Mod 1	36.6	24.2	24.5	43.2	41.6	29.0
	Mod T	37.9	24.6	25.3	43.9	40.9	26.5

We note that there is a much smaller number of wet days in the year than there are dry days at the stations in this study. In particular Kakamas where observations of rainfall constitute only 4% of the data record. It would be therefore (statistically) surprising if all the parameter functions associated with wet days fitted the historical record very closely.

The annual standard deviation has been well described by both models for the cases when the variables are conditioned on the wet and dry status of the day as well as for the case when they are not (Tables 6.5 – 6.7). Again it is seen that for some stations, the standard deviation statistic for the simulated sequences of wind run and minimum humidity differ slightly from that of the historical record. In these instances, Model T generally performs better than Model 1.

**TABLE 6.5 Comparison of historical and simulated annual standard deviation**

Variable	Data	S t a t i o n					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	5.9	6.8	6.5	4.4	5.2	5.8
	Mod 1	5.7	6.8	6.5	4.5	5.1	5.8
	Mod T	5.7	6.7	6.3	4.4	5.2	6.0
Min Temp	His	3.8	6.6	6.1	5.2	5.1	7.1
	Mod 1	3.7	6.6	6.1	5.2	4.9	7.1
	Mod T	3.6	6.5	6.2	5.2	4.8	7.0
Evapo	His	3.7	4.6	3.2			
	Mod 1	3.7	4.7	3.2			
	Mod T	3.7	4.8	3.3			
Sun	His	3.6	2.4	3.0	3.6	3.7	2.7
	Mod 1	3.6	2.2	3.0	3.8	4.1	2.9
	Mod T	3.6	2.2	3.0	3.9	4.1	2.8
Wind	His	86.7	75.2	82.6	32.2	54.5	63.3
	Mod 1	76.7	71.1	74.6	30.1	48.1	61.3
	Mod T	78.6	71.7	76.6	32.4	51.1	57.0
Max Hum	His	6.9	22.0	17.5	10.7	12.5	15.9
	Mod 1	5.7	21.8	16.3	10.4	12.2	15.4
	Mod T	6.1	22.1	15.3	10.8	12.4	14.3
Min Hum	His	15.2	10.4	11.9	16.9	22.6	15.6
	Mod 1	14.4	7.6	11.2	16.7	22.5	15.0
	Mod T	14.7	10.4	10.7	16.6	22.3	14.0

**TABLE 6.6 Comparison of historical and simulated annual standard deviation given a wet day**

Variable	Data	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	4.5	6.2	7.1	5.1	5.6	5.6
	Mod 1	4.2	5.8	7.1	5.1	5.6	5.8
	Mod T	5.0	6.9	6.6	4.9	5.7	6.0
Min Temp	His	3.3	5.4	5.1	3.3	3.5	4.2
	Mod 1	3.4	5.1	5.1	3.5	3.6	4.3
	Mod T	3.5	5.5	5.4	4.0	3.8	5.0
Evapo	His	2.6	3.9	3.2			
	Mod 1	2.5	3.4	3.2			
	Mod T	2.8	5.4	4.0			
Sun	His	3.4	3.3	3.6	3.6	3.6	3.5
	Mod 1	3.4	3.2	3.7	4.0	4.3	3.8
	Mod T	3.8	3.6	4.3	4.6	4.6	4.2
Wind	His	113.7	68.2	85.3	37.7	56.5	65.3
	Mod 1	96.6	54.4	75.0	33.7	50.4	60.6
	Mod T	100.8	64.3	82.6	37.4	54.8	62.1
Max Hum	His	4.1	20.0	16.1	8.0	5.5	13.1
	Mod 1	4.5	16.6	12.3	8.0	5.0	12.9
	Mod T	6.1	25.7	14.7	9.5	5.5	14.5
Min Hum	His	15.4	14.6	15.4	15.1	18.8	19.1
	Mod 1	13.4	7.6	13.6	14.6	18.5	17.6
	Mod T	15.4	15.6	14.2	16.9	20.2	20.4

One of the difficulties that arose when modelling climate variables was that the variables are bounded with values lying outside these boundaries being inadmissible, for example, having negative sunshine. Also, some variables have a high frequency of values near or on an upper or lower limit so that it is expected that simulated sequences will occasionally have values that exceed these boundaries. Transformations were applied to the climate variables to overcome this problem. To verify that the simulated sequences of climate variables were adequately restrained within their boundaries and at the same time that extreme values simulated closely resemble those of the historical record, the maximum and minimum values simulated for each variable were compared with those observed in the

**TABLE 6.7 Comparison of historical and simulated annual standard deviation given a dry day**

Variable	Data	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	5.5	6.8	6.3	4.1	4.8	5.8
	Mod 1	5.2	6.8	6.4	4.2	4.7	5.8
	Mod T	5.4	6.6	6.2	4.2	4.8	5.9
Min Temp	His	3.9	6.6	5.9	5.2	5.0	6.9
	Mod 1	3.9	6.7	6.0	5.1	4.7	6.9
	Mod T	3.6	6.5	6.0	5.1	4.6	6.9
Evapo	His	3.5	4.6	3.2			
	Mod 1	3.4	4.7	3.1			
	Mod T	3.4	4.8	3.1			
Sun	His	2.5	2.1	2.3	3.0	2.4	1.8
	Mod 1	2.2	1.8	2.1	3.1	2.3	1.5
	Mod T	2.4	1.9	2.3	3.2	2.5	1.5
Wind	His	67.5	75.3	81.9	29.9	51.2	61.6
	Mod 1	60.1	71.7	74.3	28.6	44.3	60.9
	Mod T	63.4	72.0	74.9	30.3	46.6	55.2
Max Hum	His	7.5	21.9	17.7	10.7	14.0	15.9
	Mod 1	6.0	21.8	16.9	10.2	13.9	15.5
	Mod T	6.1	21.8	15.4	10.7	14.1	14.2
Min Hum	His	12.0	9.7	9.9	14.6	17.8	11.9
	Mod 1	11.5	7.0	9.4	14.2	17.3	11.4
	Mod T	11.9	9.9	9.5	14.5	17.4	11.2

historical record. Tables 6.8 and 6.9 show these comparisons.

As can be seen from the tables, the extreme values simulated compare favourably with those observed in the historical record. Even for those variables that show a slight difference in the extreme values, when a count was taken of those values of the simulated sequence that lay either above the maximum or below the minimum values observed in the historical data, the percentage of such values was found to be negligible, that is, the highest percentage observed was 0.4%.

**TABLE 6.8** Comparison of historical and simulated minimum values

Variable	Data	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	10.0	9.5	5.0	13.1	7.8	5.6
	Mod 1	6.9	7.4	2.6	8.5	4.5	5.4
	Mod T	4.8	7.2	1.1	8.6	3.7	3.9
Min Temp	His	1.7	-2.5	-8.0	0.0	-4.0	-8.1
	Mod 1	1.1	-2.3	-8.2	0.9	-10.4	-7.8
	Mod T	1.3	-3.0	-8.3	0.7	-5.5	-8.3
Evapo	His	0.0	0.5	0.0			
	Mod 1	0.0	0.5	0.0			
	Mod T	0.0	0.1	0.0			
Sun	His	0.0	0.0	0.0	0.0	0.0	0.0
	Mod 1	0.0	0.1	0.0	0.0	0.0	0.0
	Mod T	0.0	0.0	0.0	0.0	0.0	0.0
Wind	His	40.0	35.1	62.2	26.0	22.0	14.0
	Mod 1	17.0	34.0	17.3	45.3	41.9	14.0
	Mod T	19.3	32.8	23.0	34.4	37.2	15.7
Max Hum	His	50.0	12.0	18.0	30.0	22.0	20.0
	Mod 1	38.4	11.3	8.9	19.8	10.4	20.3
	Mod T	34.5	1.9	17.5	18.7	6.7	18.9
Min Hum	His	12.0	1.0	5.0	9.0	2.0	7.0
	Mod 1	8.7	7.1	3.7	3.1	1.2	3.6
	Mod T	8.6	0.6	3.0	3.9	0.6	0.3

TABLE 6.9 Comparison of historical and simulated maximum values

Variable	Data	Station					
		Elsenburg	Kakamas	Middelburg	Nelspruit	Cedara	Hoopstad
Max Temp	His	40.8	43.8	38.0	39.8	37.3	39.0
	Mod 1	42.1	44.6	37.8	39.5	38.0	39.5
	Mod T	42.3	47.4	39.3	40.1	38.2	40.6
Min Temp	His	20.9	29.8	22.5	23.3	21.1	21.4
	Mod 1	25.3	28.6	21.1	22.9	23.1	21.1
	Mod T	24.1	28.8	21.6	26.9	22.9	21.1
Evapo	His	18.5	24.0	18.0			
	Mod 1	18.8	23.7	20.2			
	Mod T	22.8	28.2	28.3			
Sun	His	13.3	13.5	13.6	12.9	13.0	13.4
	Mod 1	13.7	14.0	13.9	13.0	13.2	13.4
	Mod T	13.7	14.0	13.9	13.0	13.2	13.5
Wind	His	733.3	531.0	583.1	420.0	453.0	396.0
	Mod 1	705.1	583.7	614.1	301.7	468.3	584.3
	Mod T	681.3	564.6	640.7	341.5	449.7	592.3
Max Hum	His	100.0	100.0	100.0	100.0	100.0	100.0
	Mod 1	100.0	100.0	100.0	100.0	100.0	100.0
	Mod T	100.0	100.0	100.0	100.0	100.0	100.0
Min Hum	His	95.0	80.0	85.0	97.0	100.0	97.0
	Mod 1	93.9	62.5	86.2	94.5	99.7	93.5
	Mod T	95.9	90.2	86.7	98.2	99.3	99.3

**Validation of monthly properties**

It is important that the monthly characteristics of each climate variable, mainly the mean and standard deviation, be adequately described by the models. The monthly means and standard deviations of the simulated sequences for each station were compared to those of the respective historical record. Figures 6.8 – 6.14 show the monthly means of each station for the various climate variables.

From the figures it can be seen that the monthly means have been successfully preserved by both models. Model T slightly overestimates the monthly means of maximum temperature for the station Hoopstad, but the highest difference between the means of the simulated sequence and that of the historical data still lies within  $3^{\circ}\text{C}$  of the observed monthly mean. Model T fails to preserve the monthly means for the variable minimum humidity of the station Hoopstad. Here the monthly means are underestimated. Model 1 fits the data reasonably well for this variable.

Figures 6.15 – 6.21 show the monthly standard deviations of each station for the various climate variables. The monthly standard deviations have been preserved by both models. The variables wind run, maximum humidity and minimum humidity show the greatest differences between the standard deviations of the observed sequence and those of the generated sequence. For these variables the models tend to slightly underestimate the monthly standard deviations. Looking at the original sequence of these three variables we see that they do not follow an approximate sinusoidal shape, one of the assumptions made when fitting the mean by a truncated Fourier series, so it seems that these observable differences may be accountable for this.

The mean and standard deviation functions of each variable differ significantly depending on the wet or dry status of the day. It is therefore necessary that monthly means and standard deviations should also be preserved by the models when the climate variables are conditioned on wet and dry days. Figures 6.22 – 6.28 show the monthly means for each station when the climate variables are conditioned on wet days. Figures 6.29 – 6.35 gives the monthly means when the climate variables are conditioned on the dry days.

FIGURE 6.8 Monthly means for maximum temperature

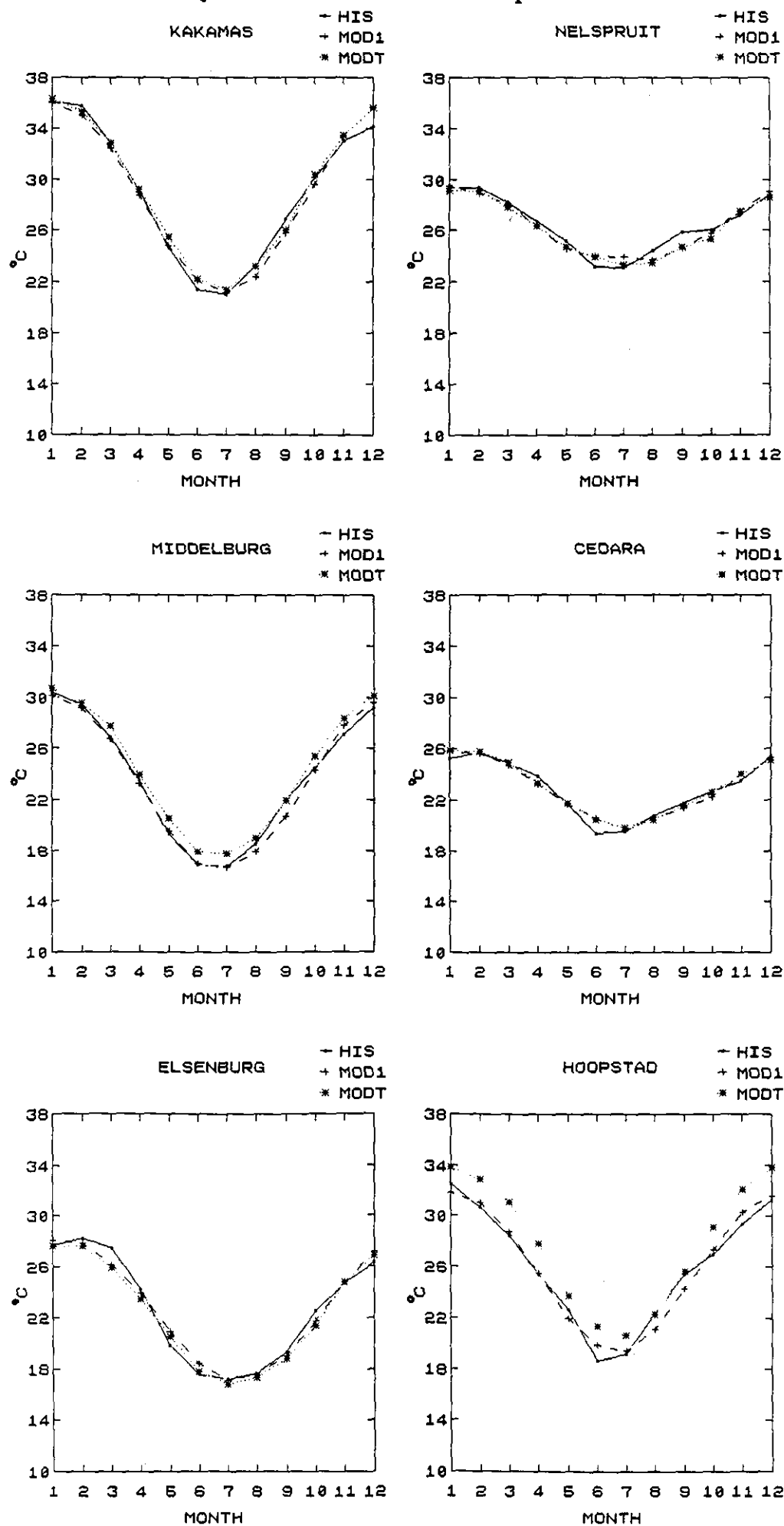


FIGURE 6.9 Monthly means for minimum temperature

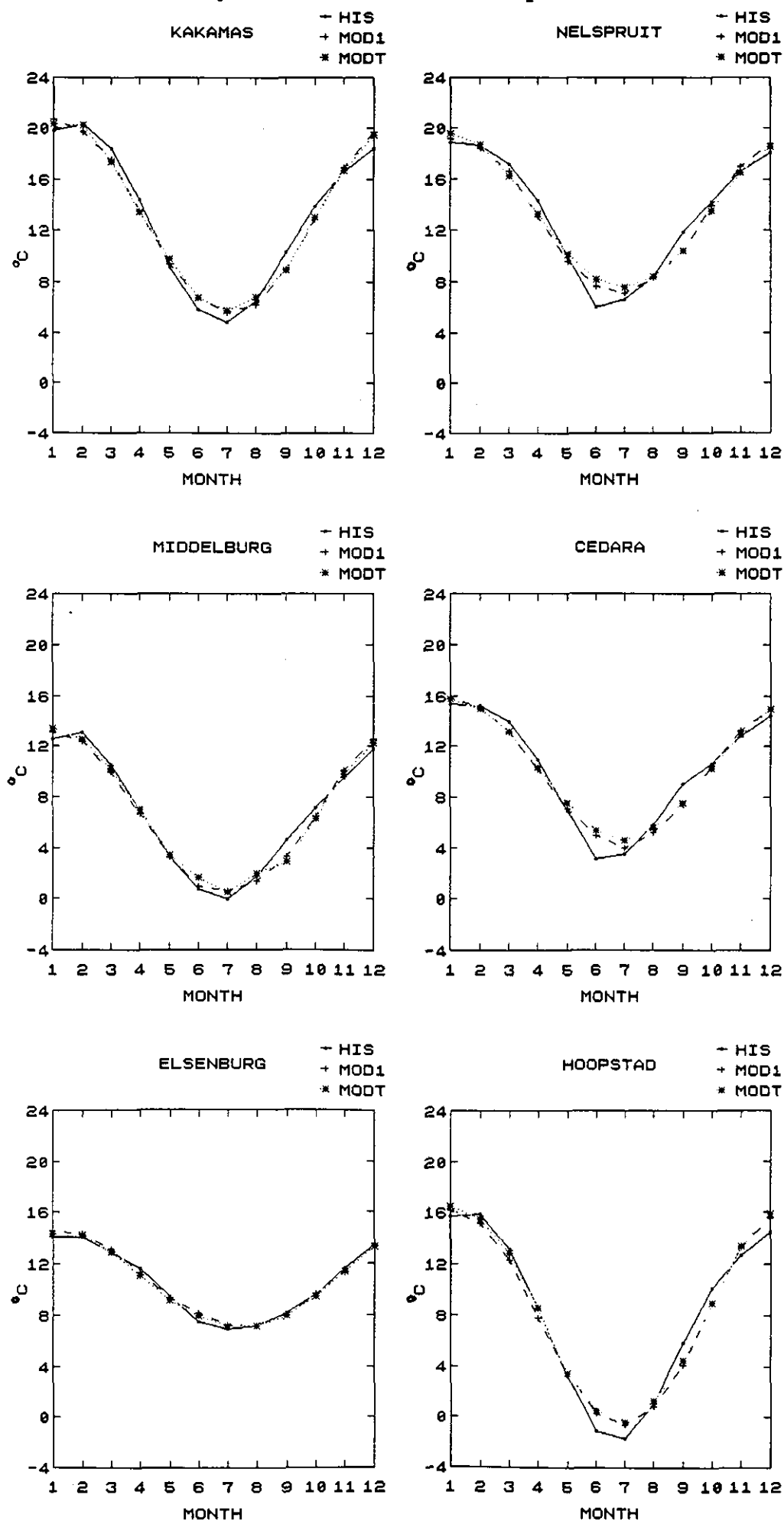


FIGURE 6.10 Monthly means for evaporation

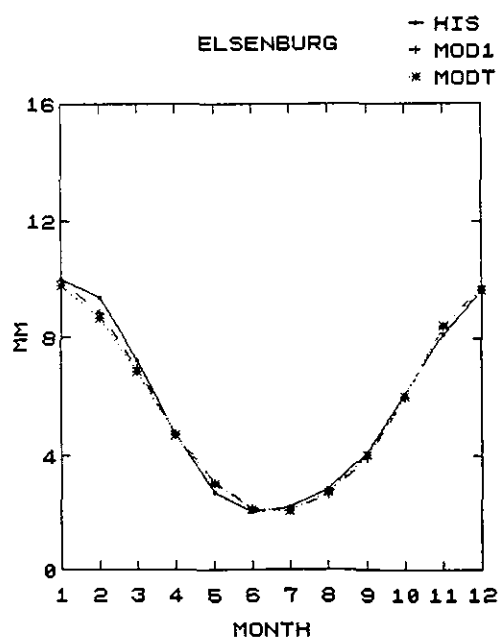
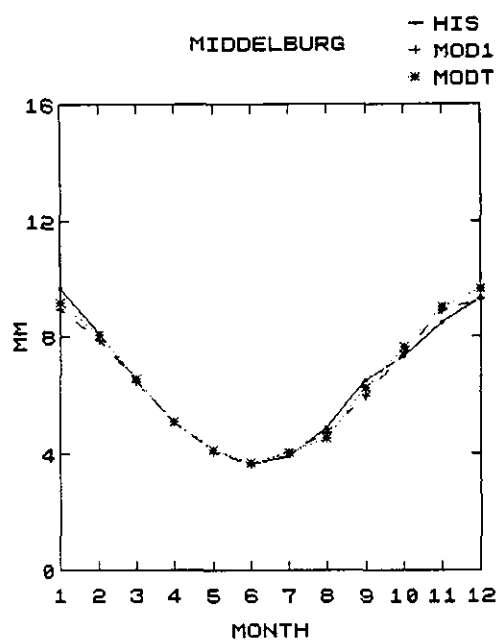
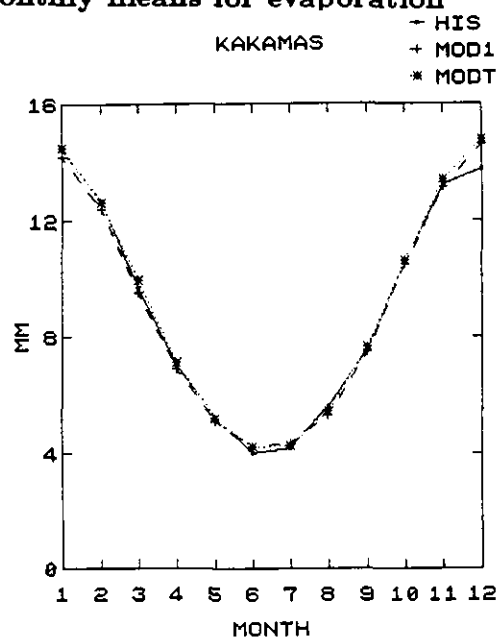
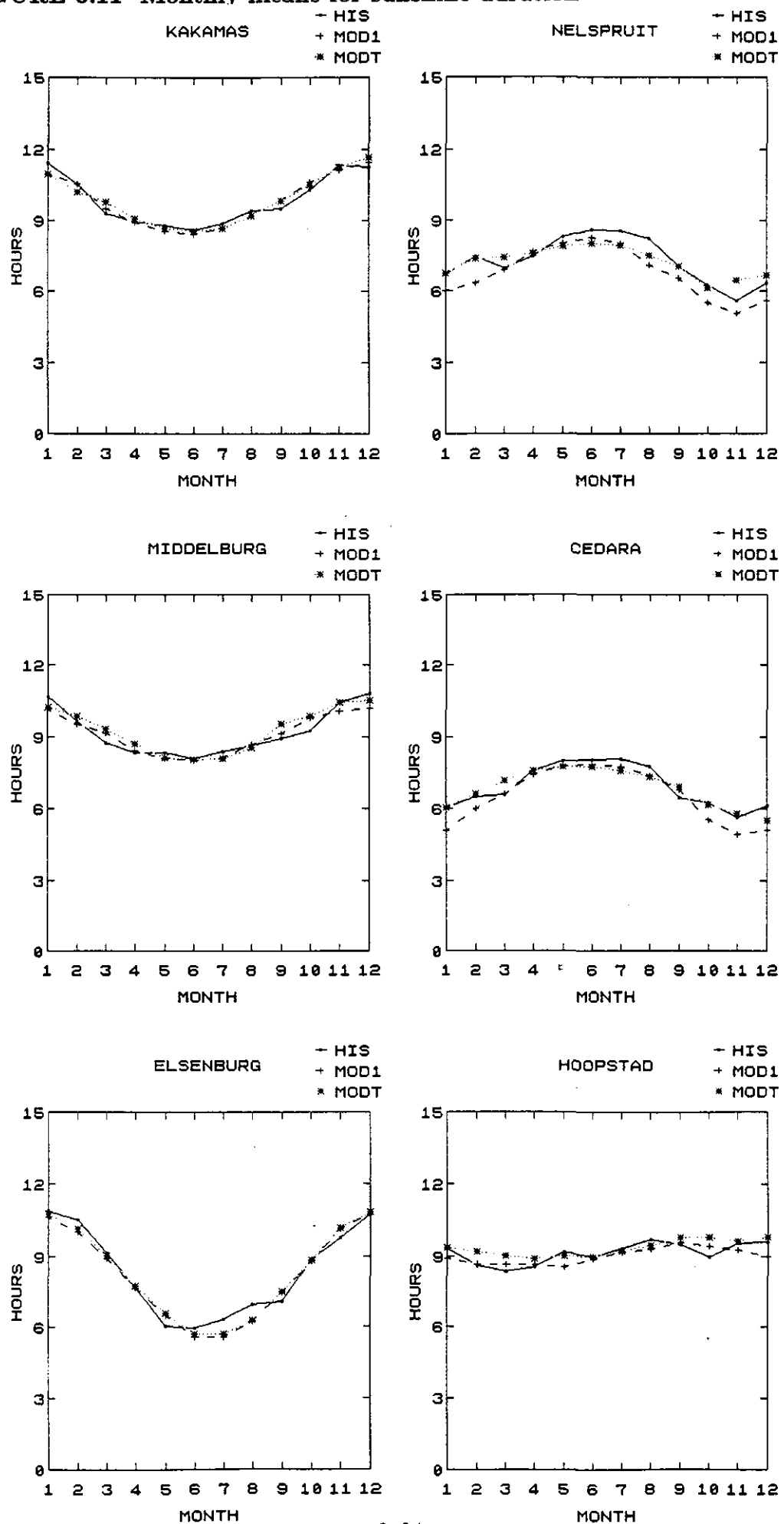


FIGURE 6.11 Monthly means for sunshine duration



**FIGURE 6.12 Monthly means for wind run**

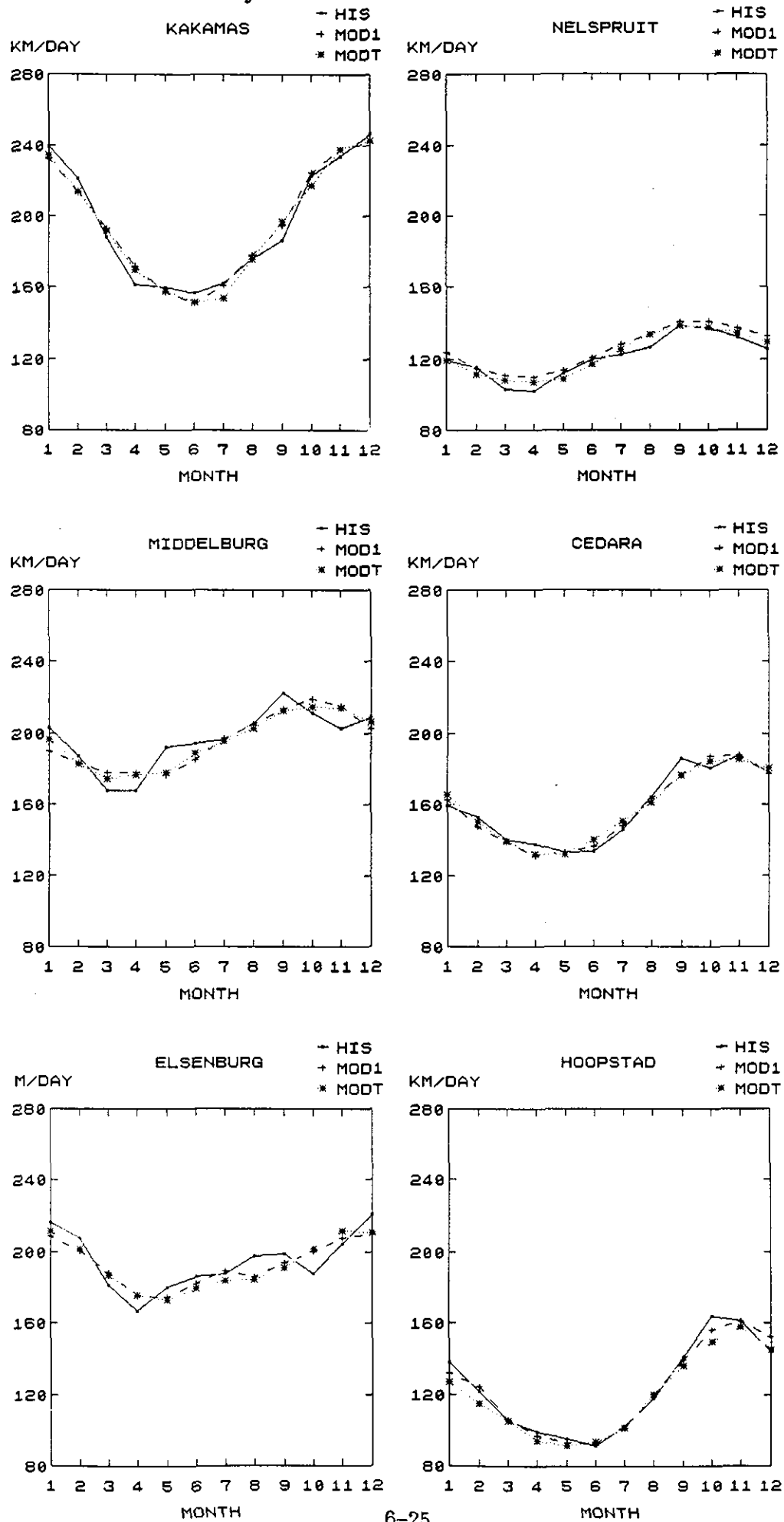


FIGURE 6.13 Monthly means for maximum humidity

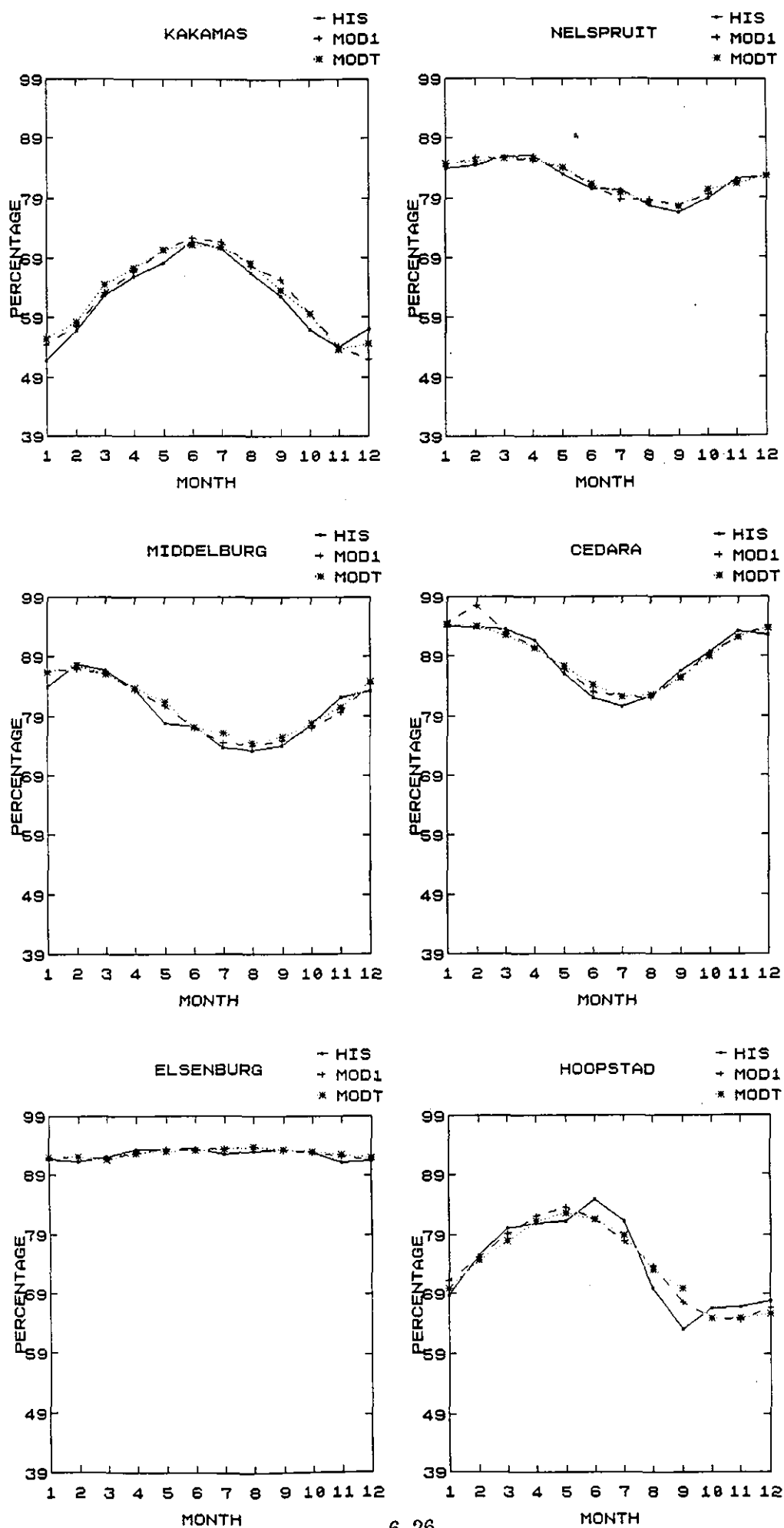


FIGURE 6.14 Monthly means for minimum humidity

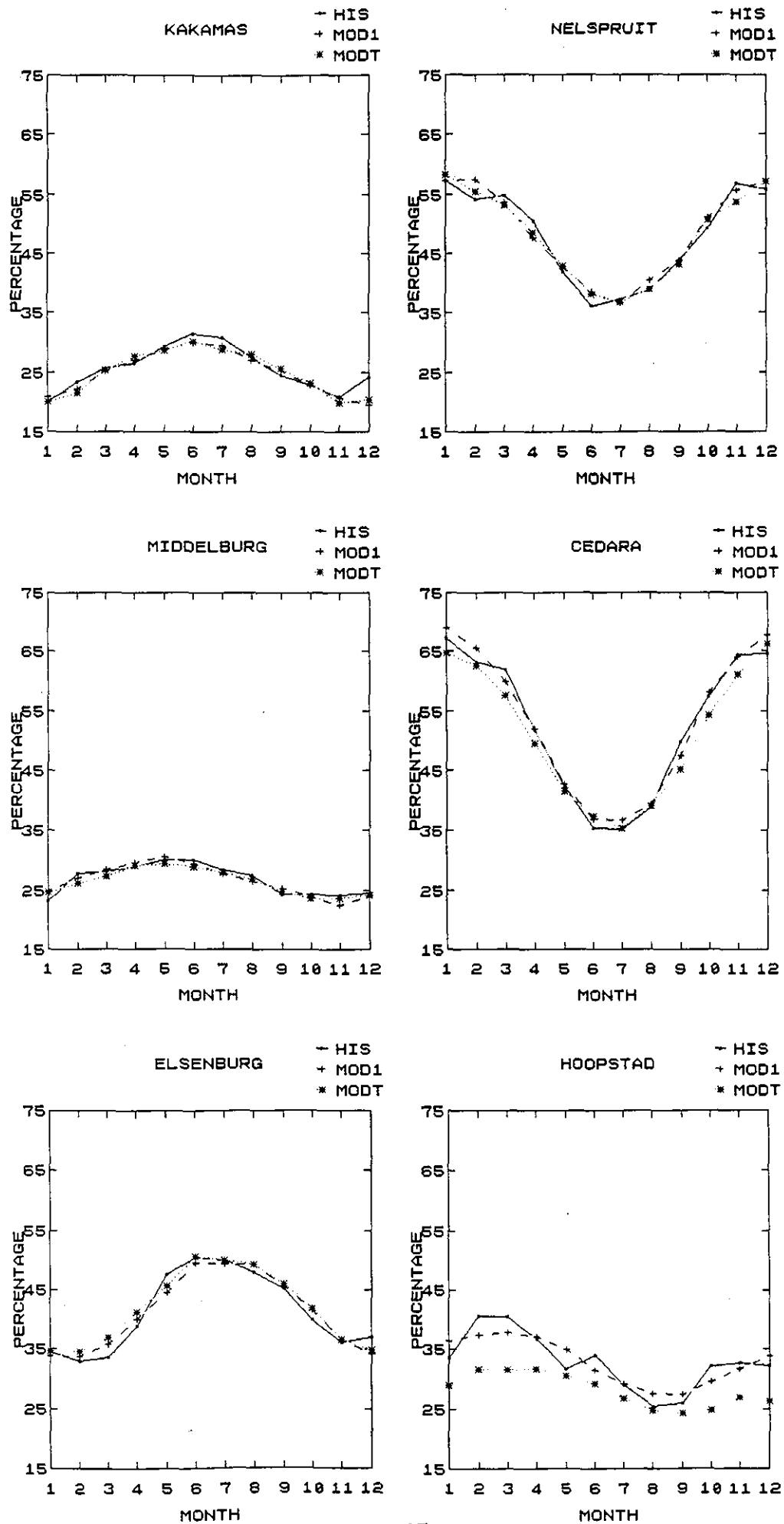


FIGURE 6.15 Monthly standard deviations for maximum temperature

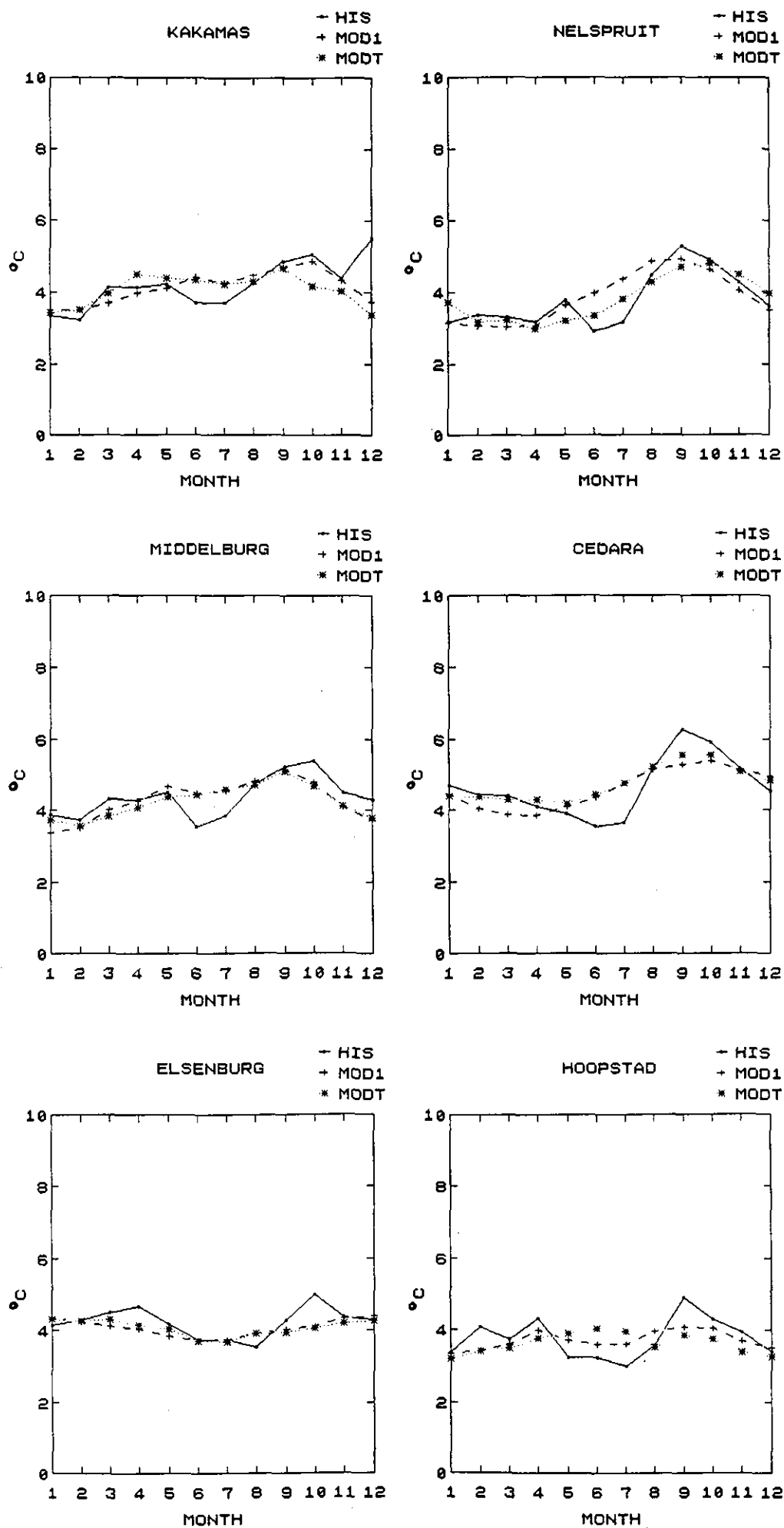
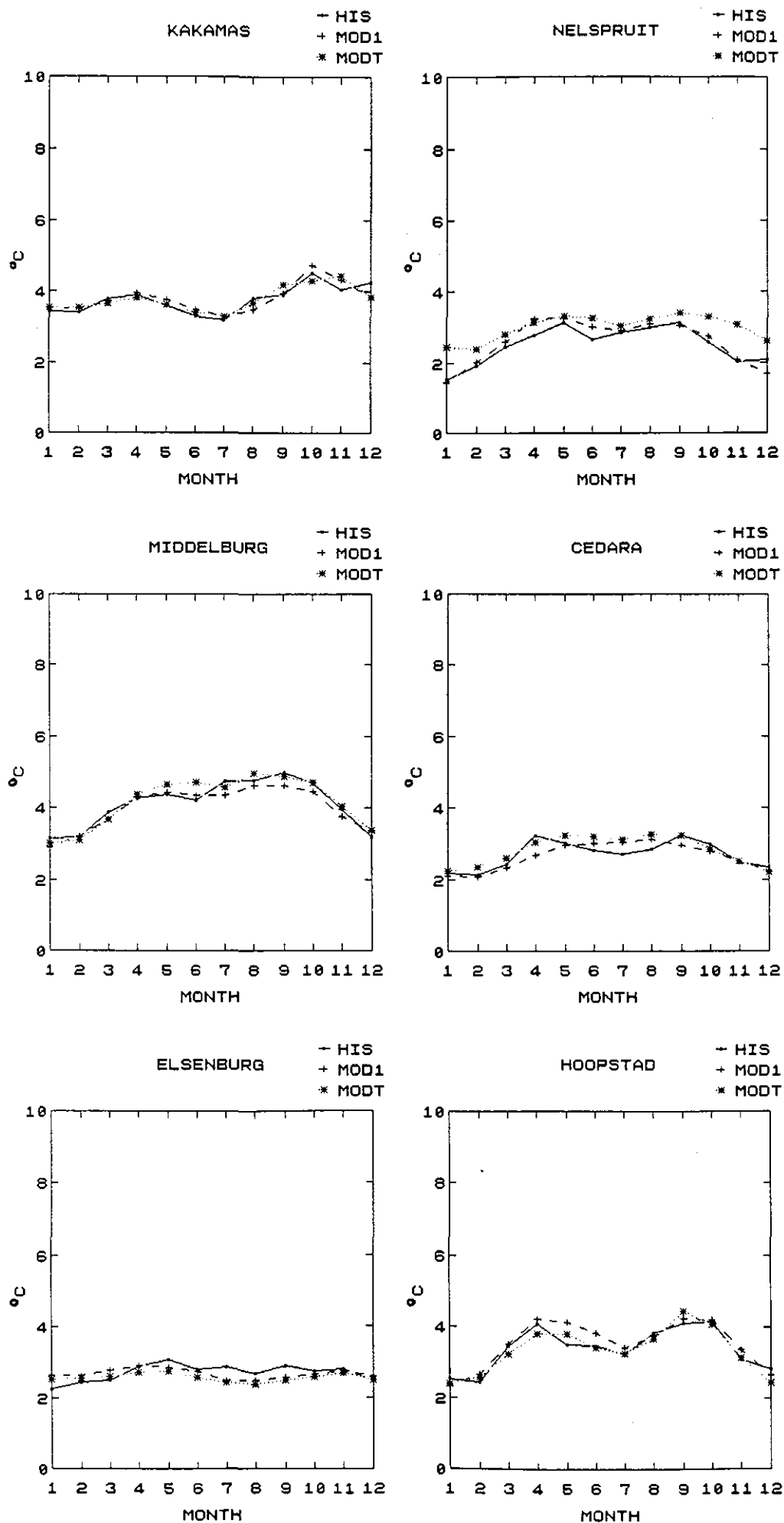


FIGURE 6.16 Monthly standard deviations for minimum temperature



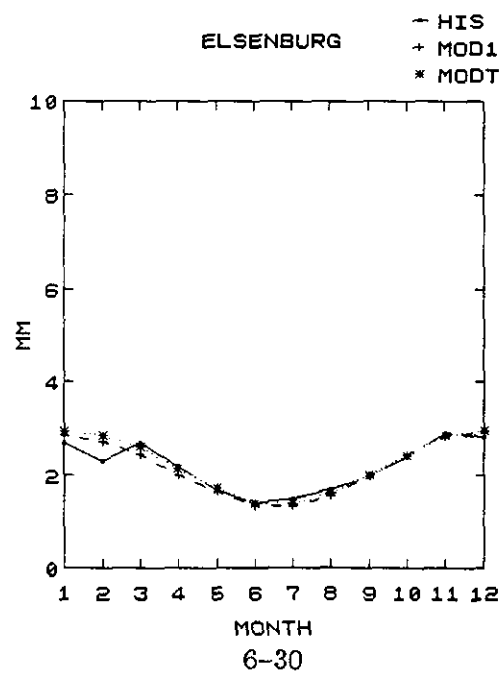
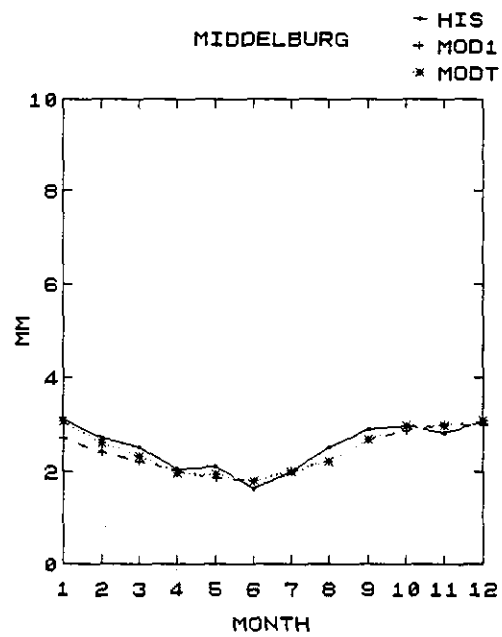
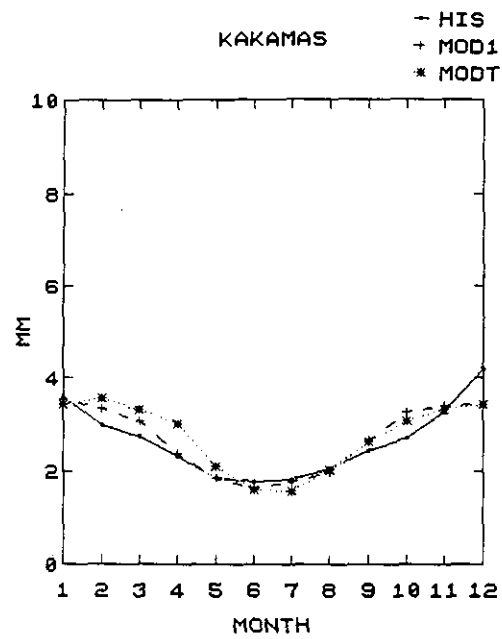
**FIGURE 6.17** Monthly standard deviations for evaporation

FIGURE 6.18 Monthly standard deviations for sunshine duration

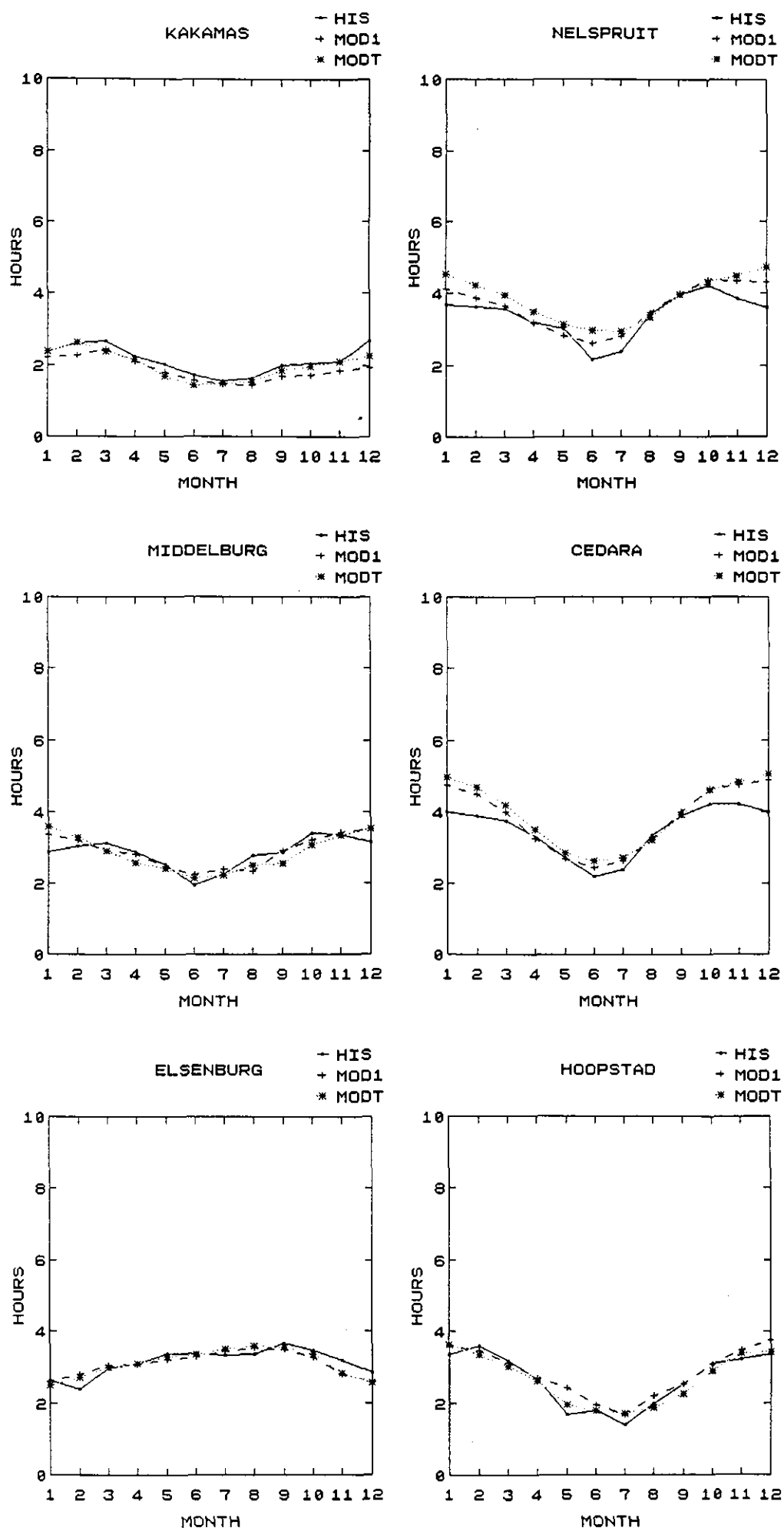


FIGURE 6.19 Monthly standard deviations for wind run

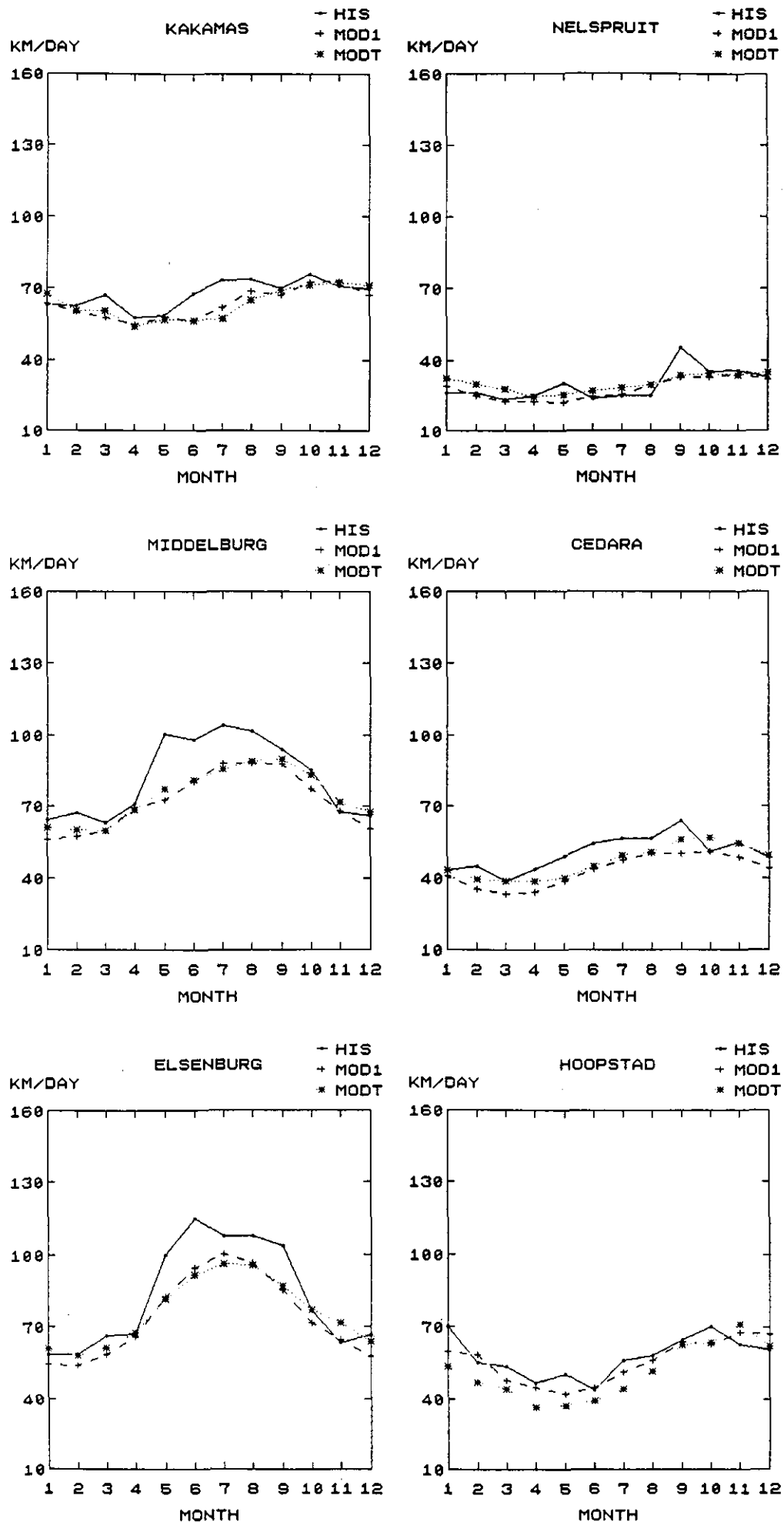


FIGURE 6.20 Monthly standard deviations for maximum humidity

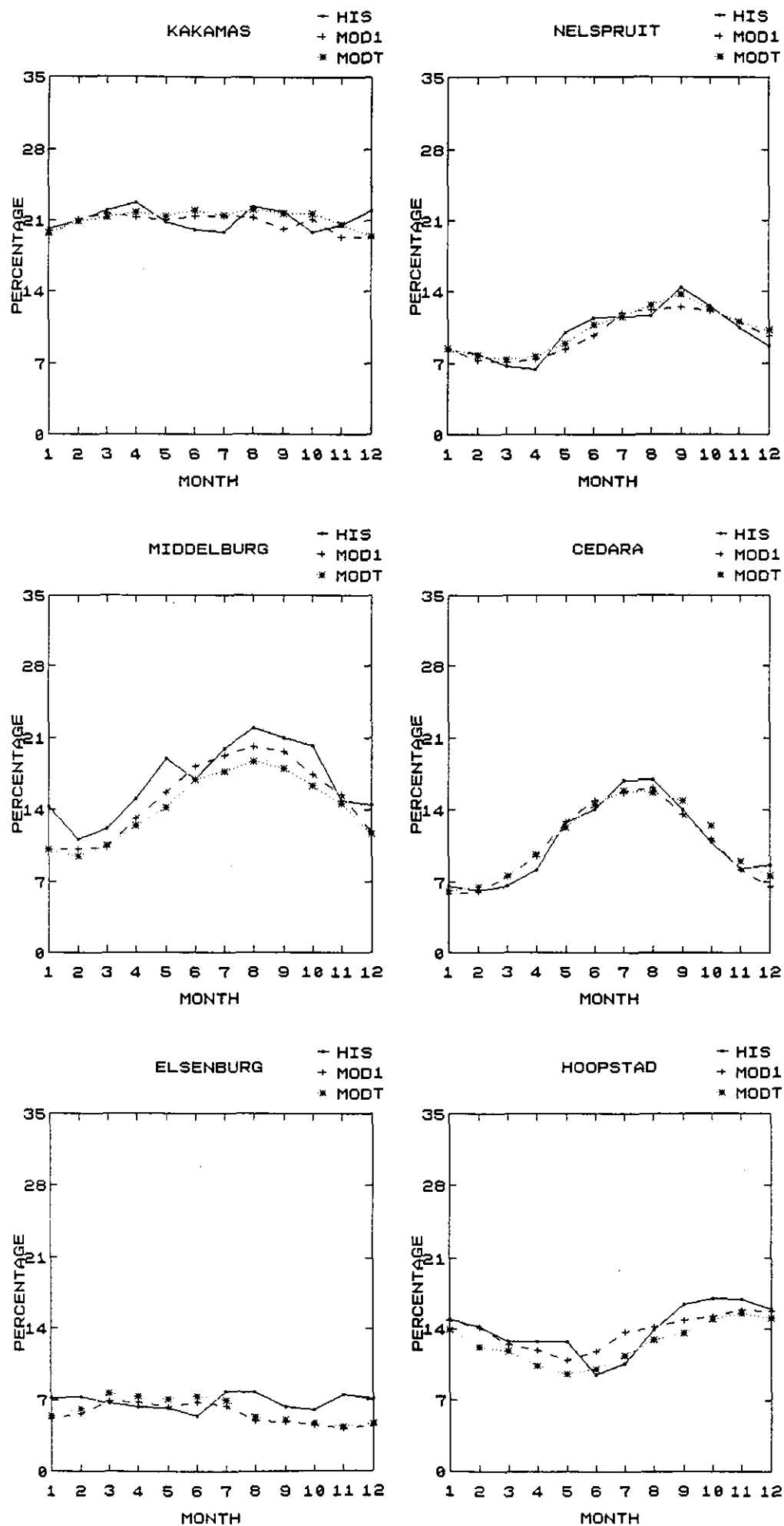


FIGURE 6.21 Monthly standard deviations for minimum humidity

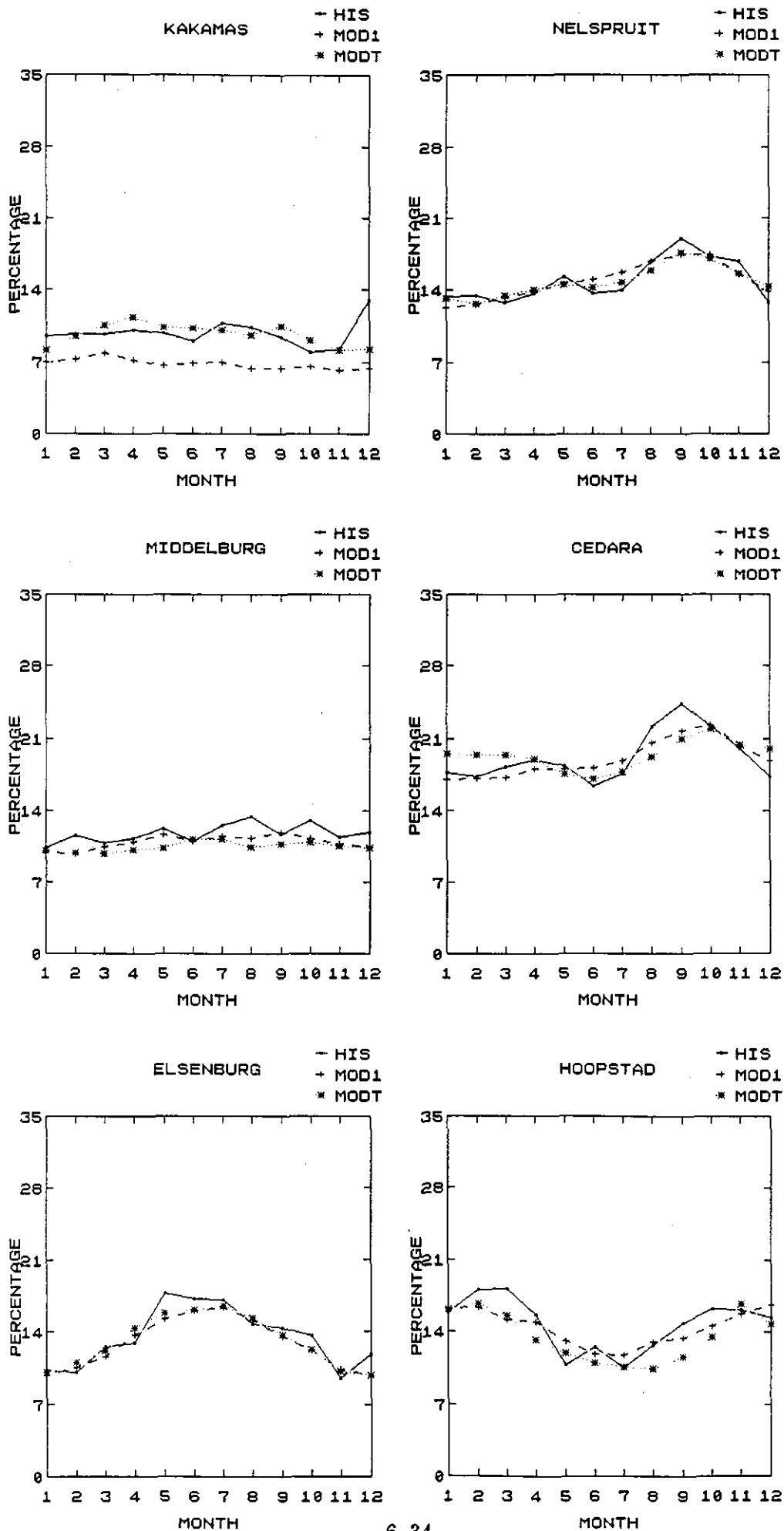
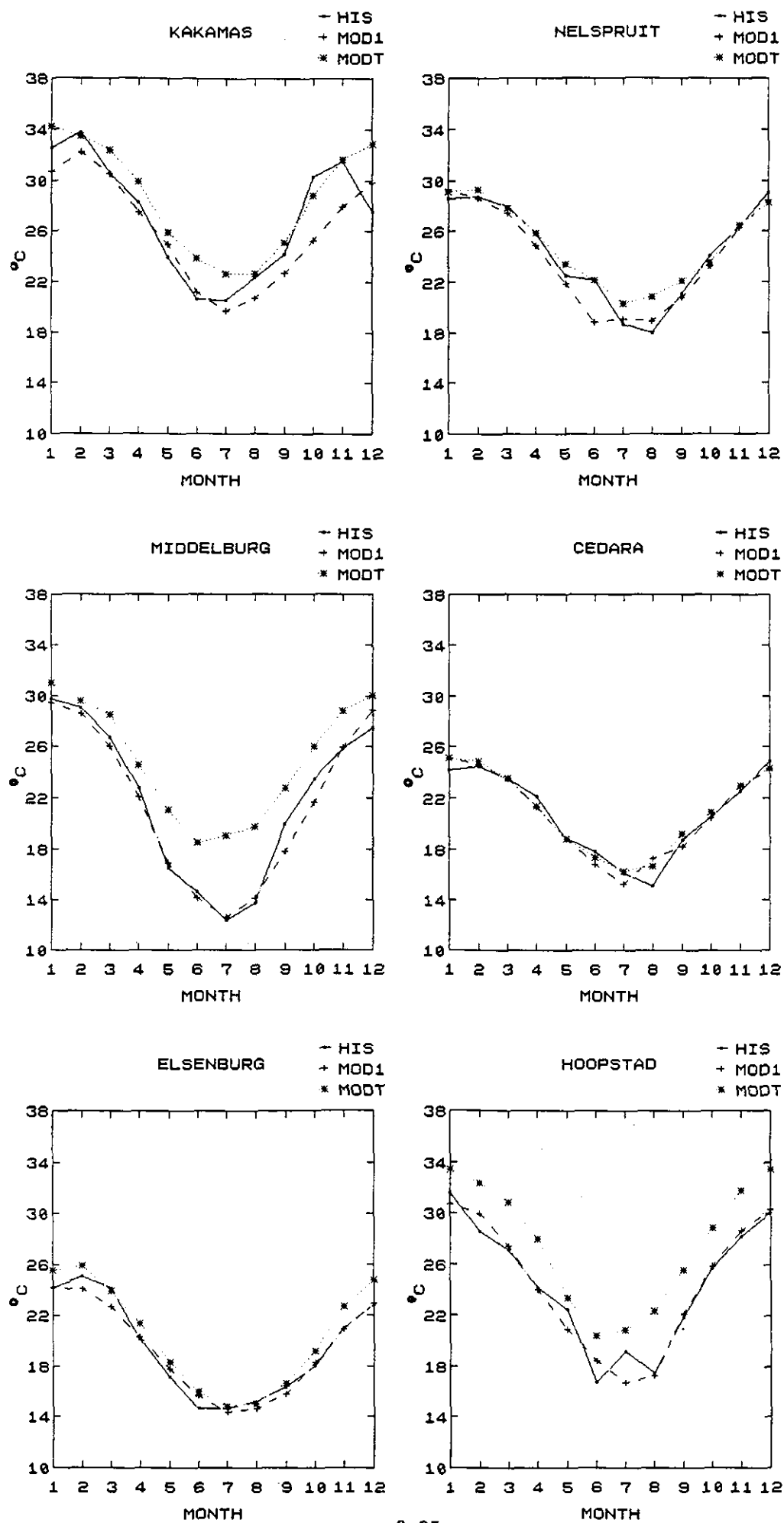
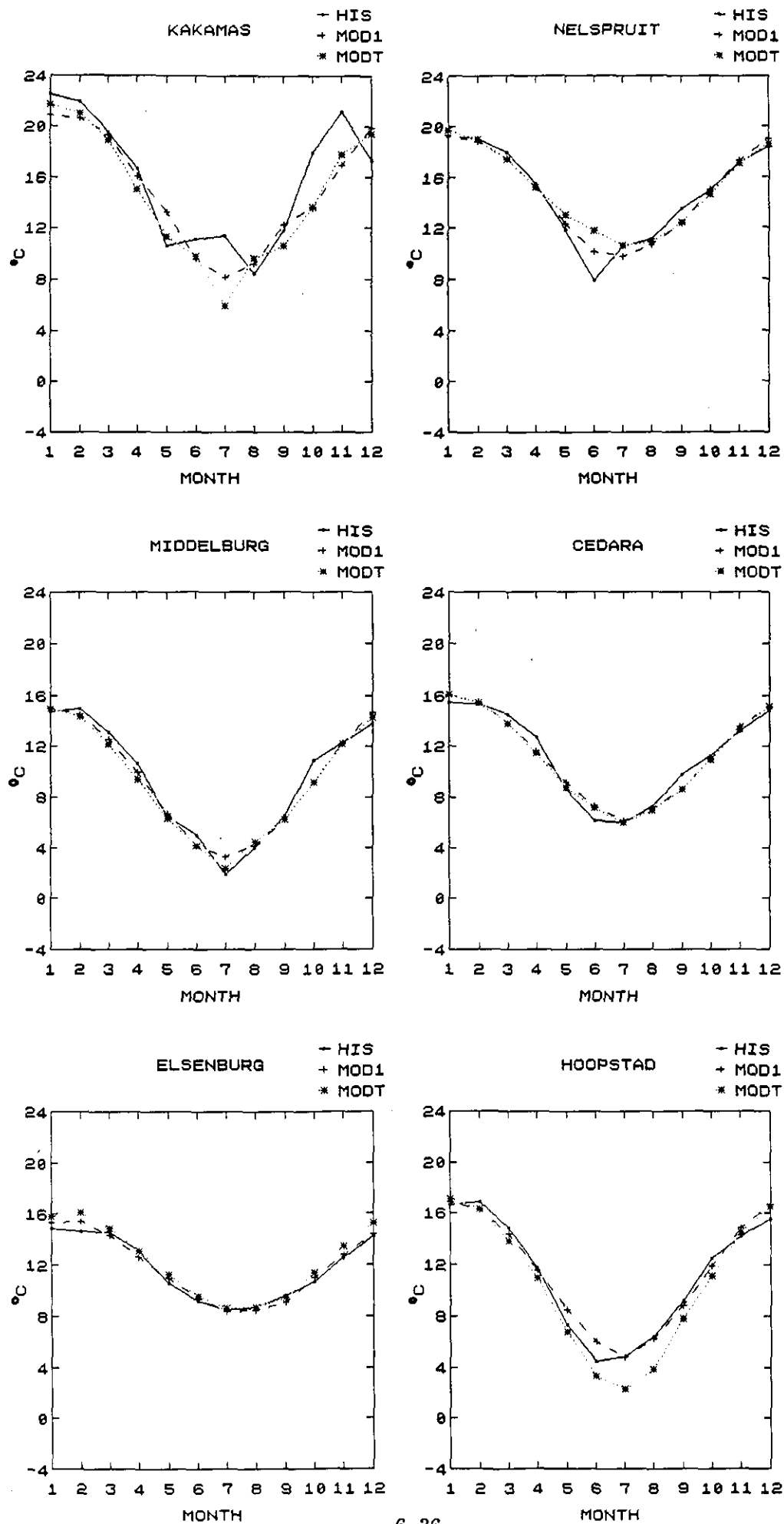


FIGURE 6.22 Monthly means for maximum temperature for wet days



**FIGURE 6.23** Monthly means for minimum temperature for wet days

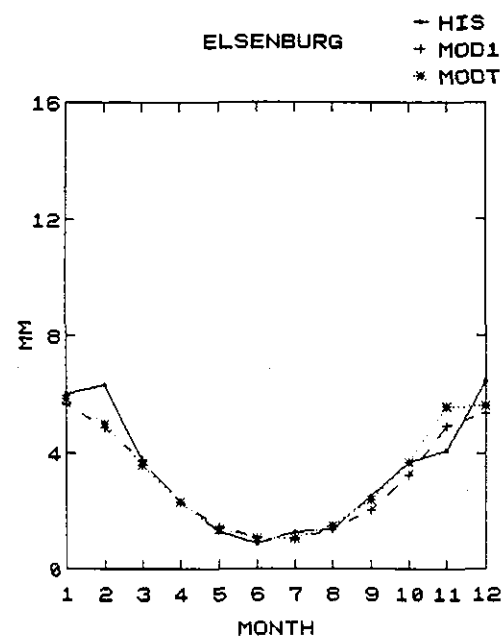
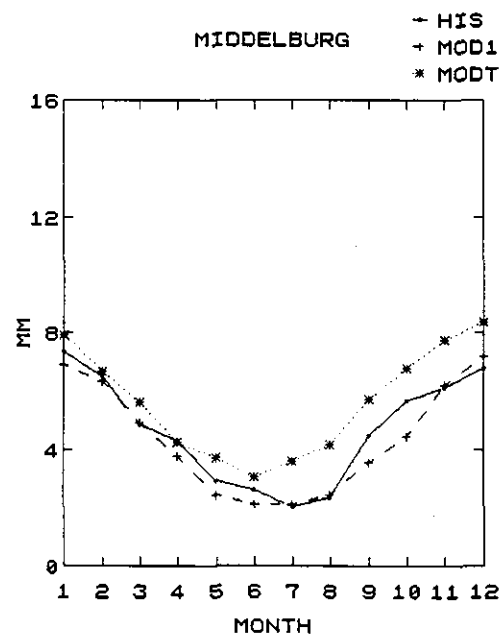
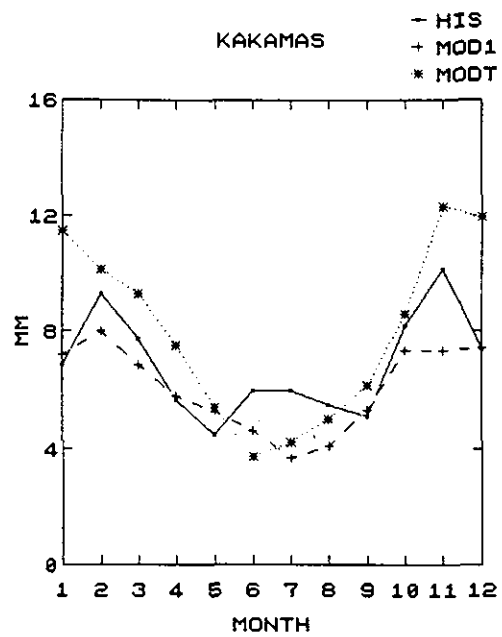
**FIGURE 6.24** Monthly means for evaporation for wet days

FIGURE 6.25 Monthly means for sunshine duration for wet days

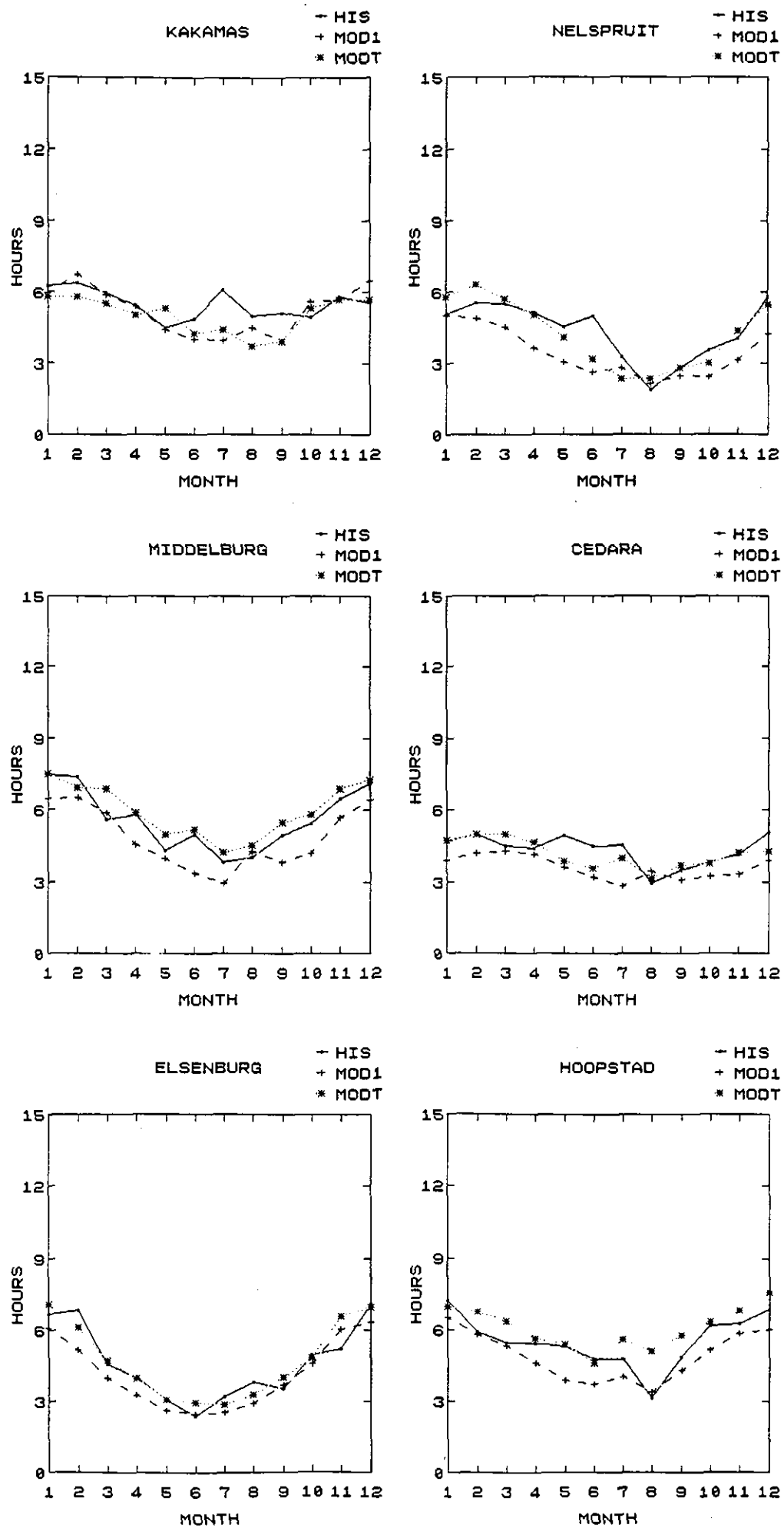
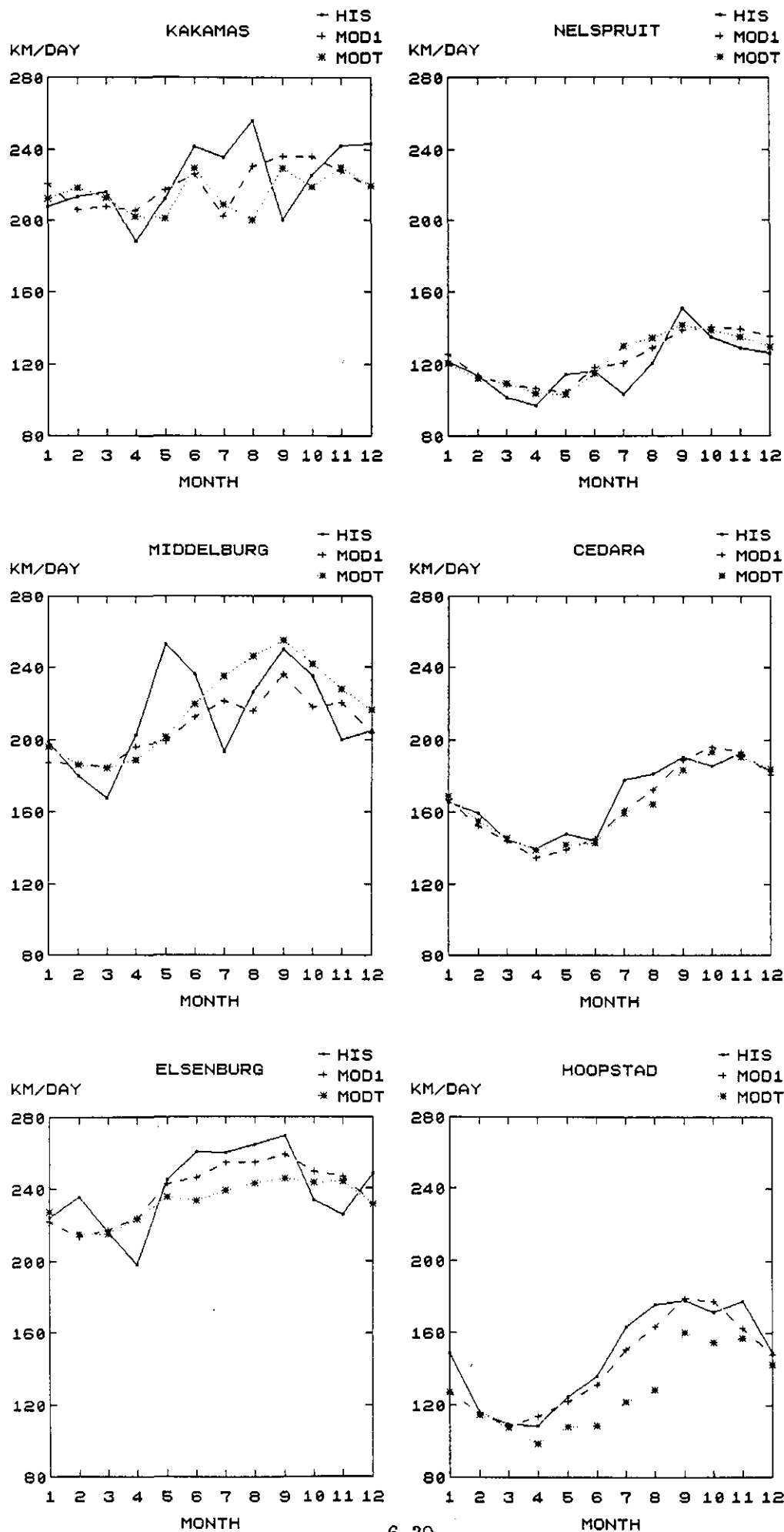


FIGURE 6.26 Monthly means for wind run for wet days



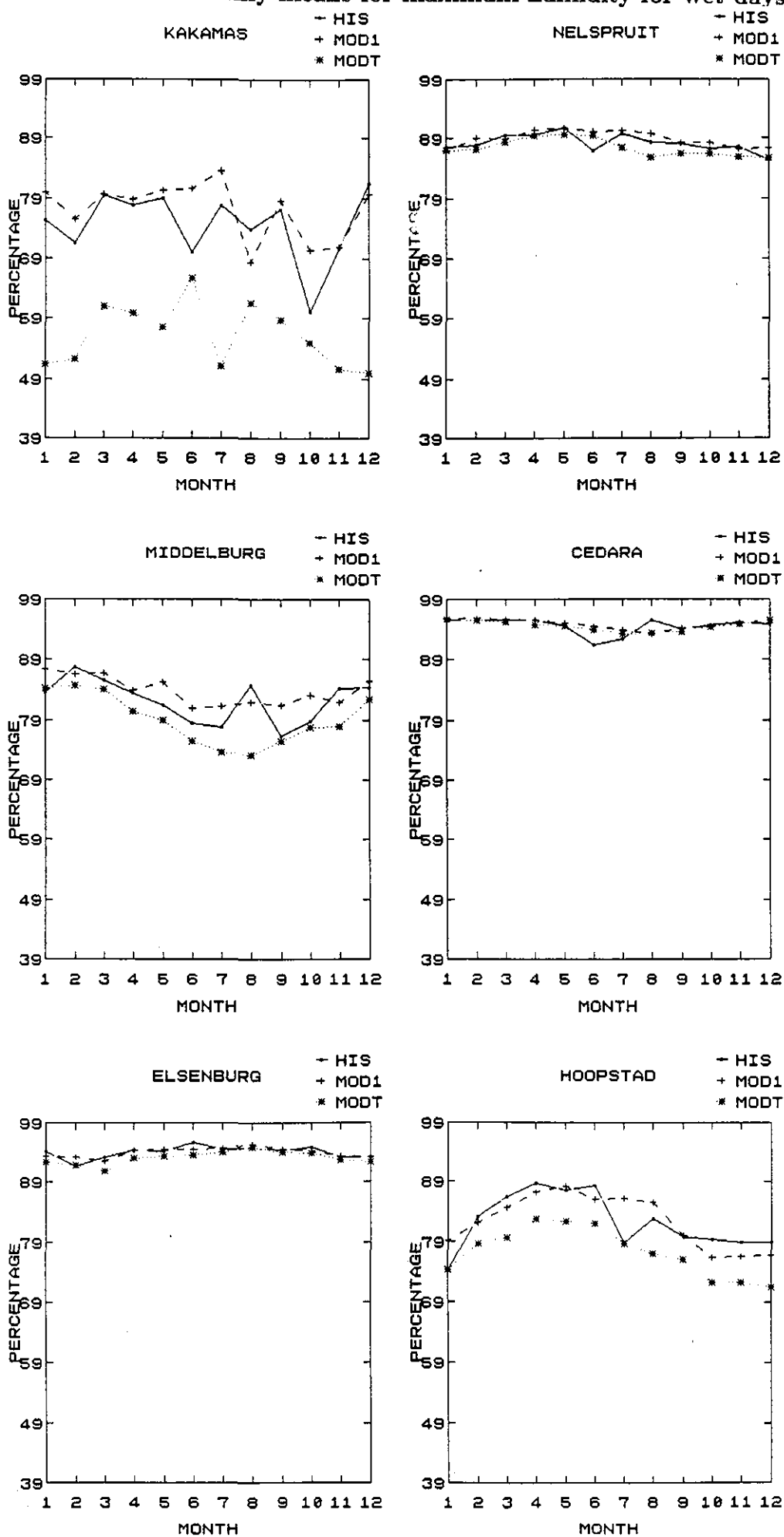
**FIGURE 6.27** Monthly means for maximum humidity for wet days

FIGURE 6.28 Monthly means for minimum humidity for wet days

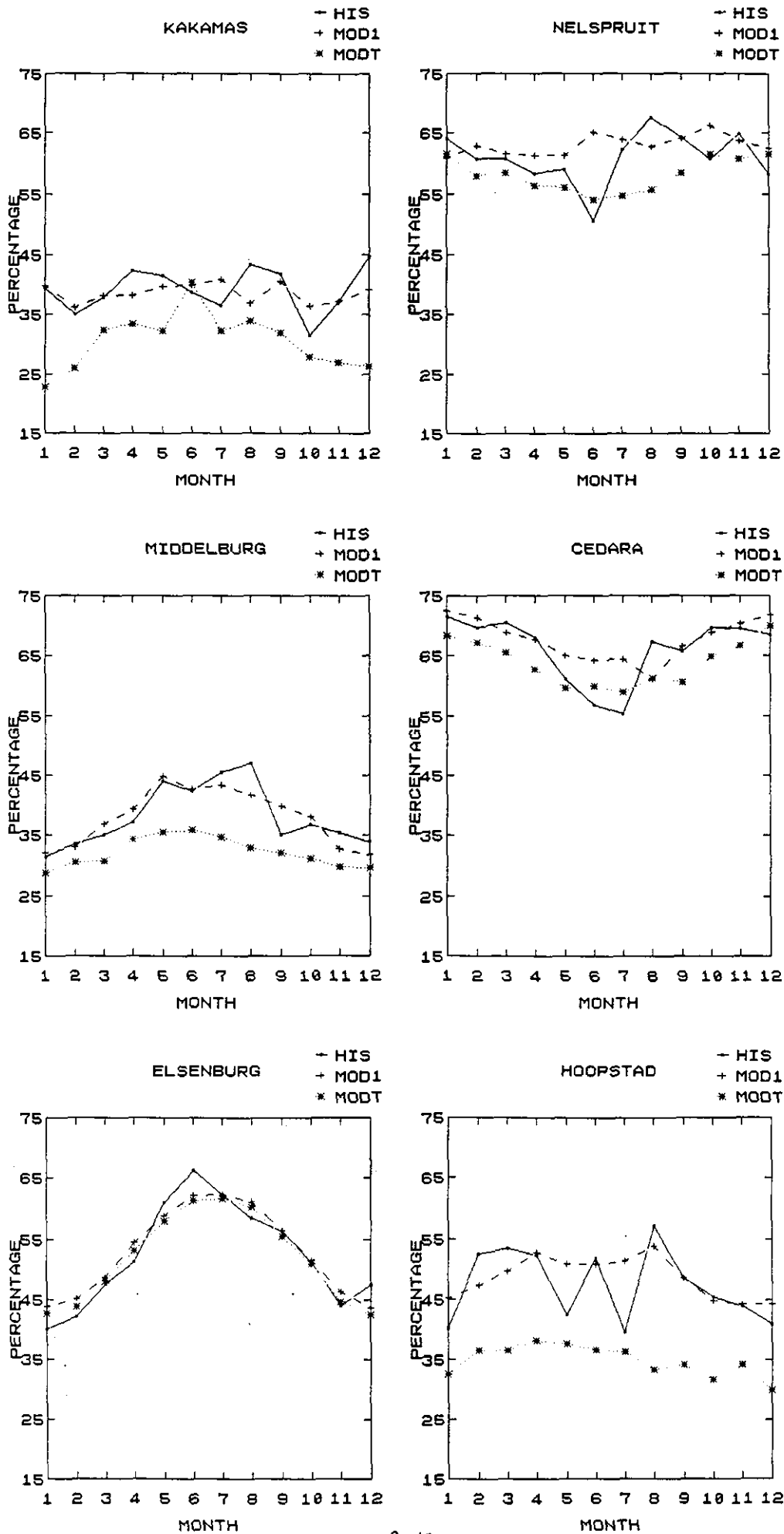


FIGURE 6.29 Monthly means for maximum temperature for dry days

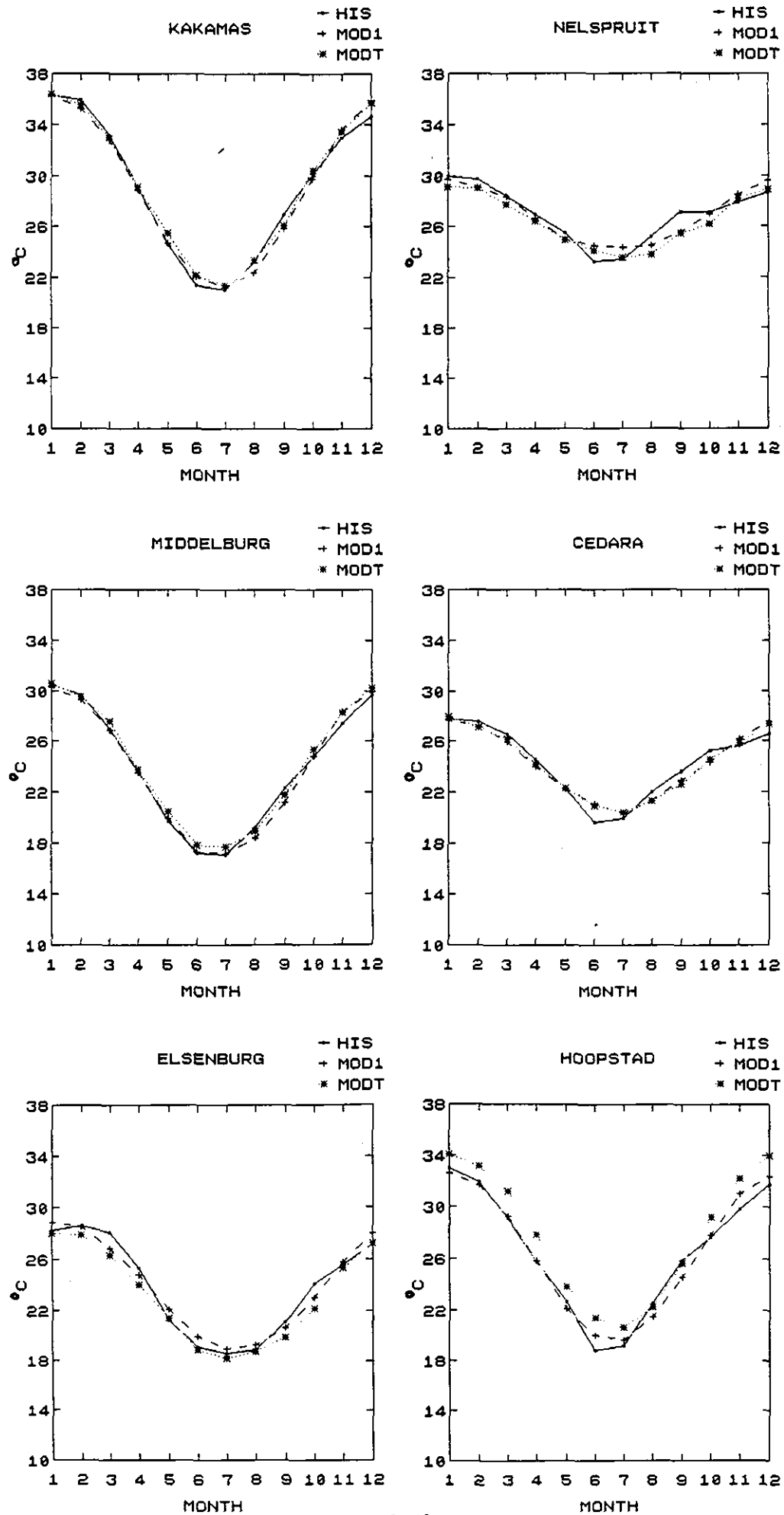
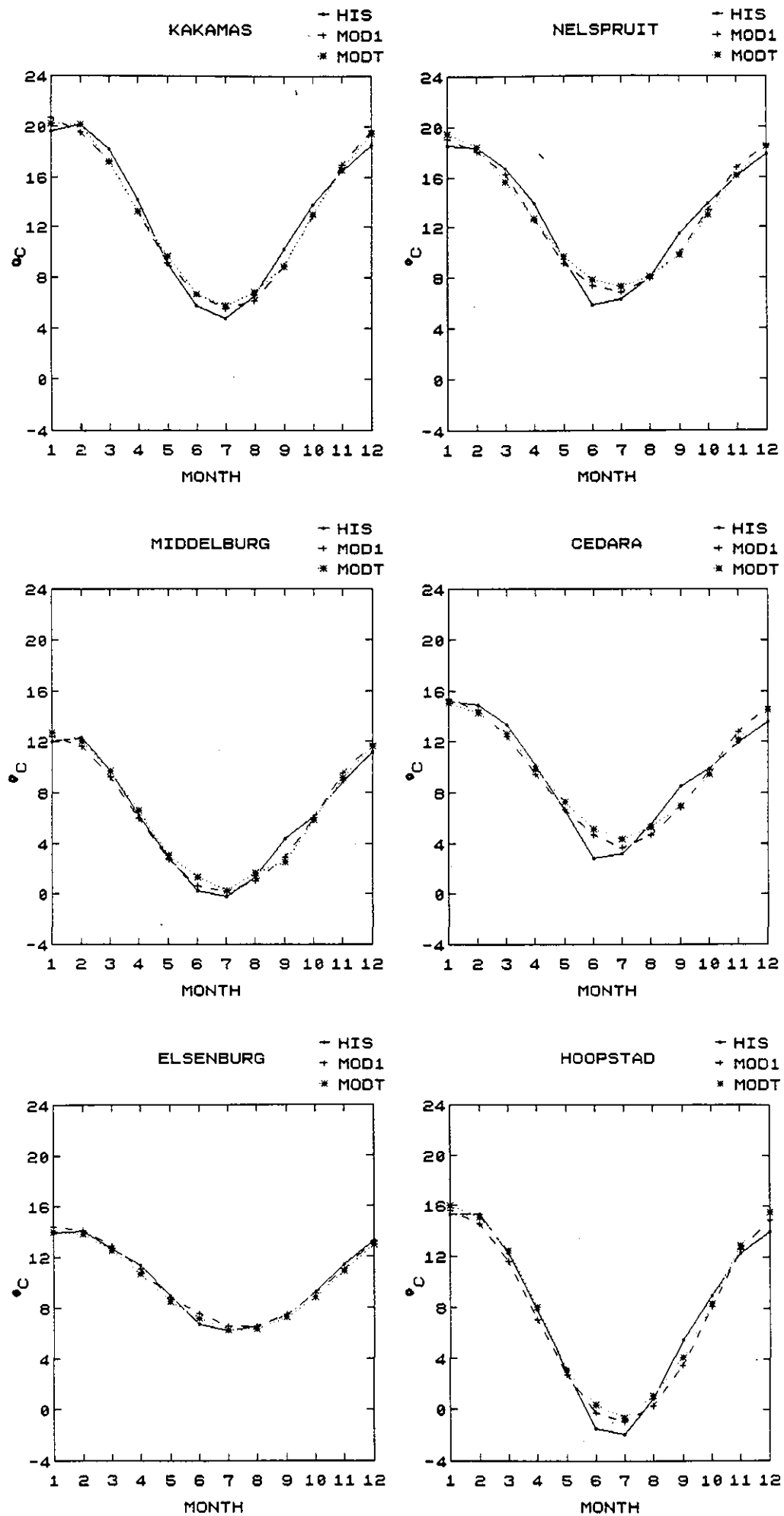


FIGURE 6.30 Monthly means for minimum temperature for dry days



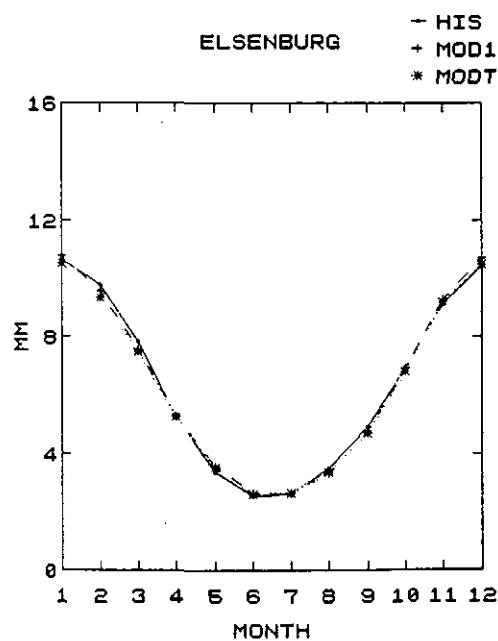
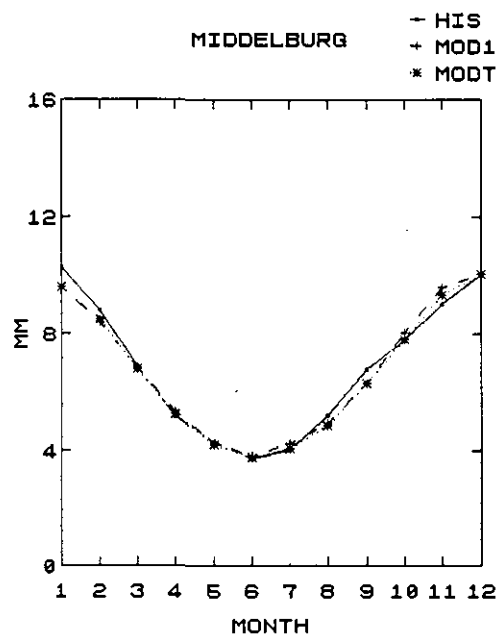
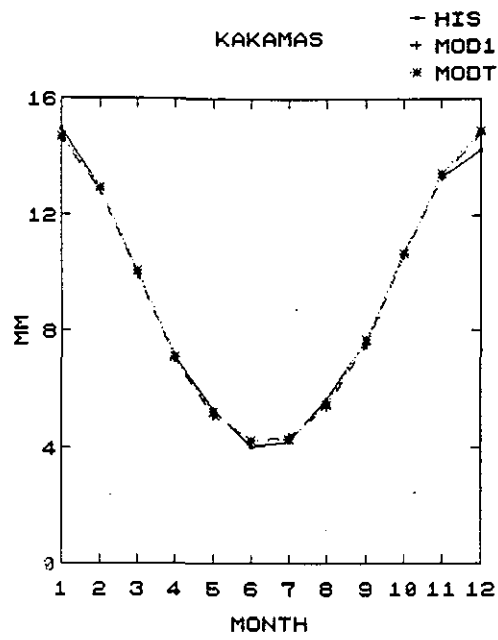
**FIGURE 6.31** Monthly means for evaporation for dry days

FIGURE 6.32 Monthly means for sunshine duration for dry days

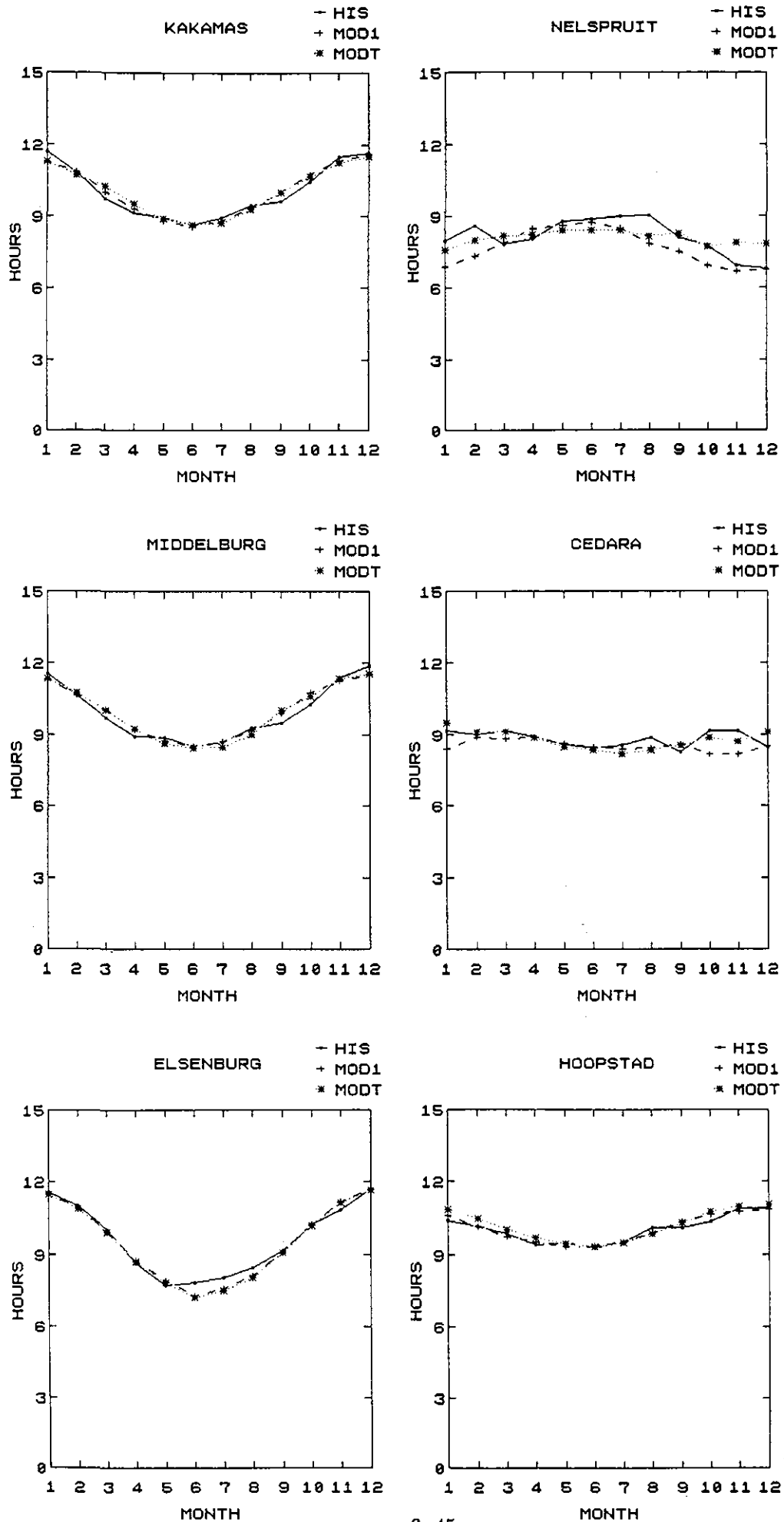


FIGURE 6.33 Monthly means for wind run for dry days

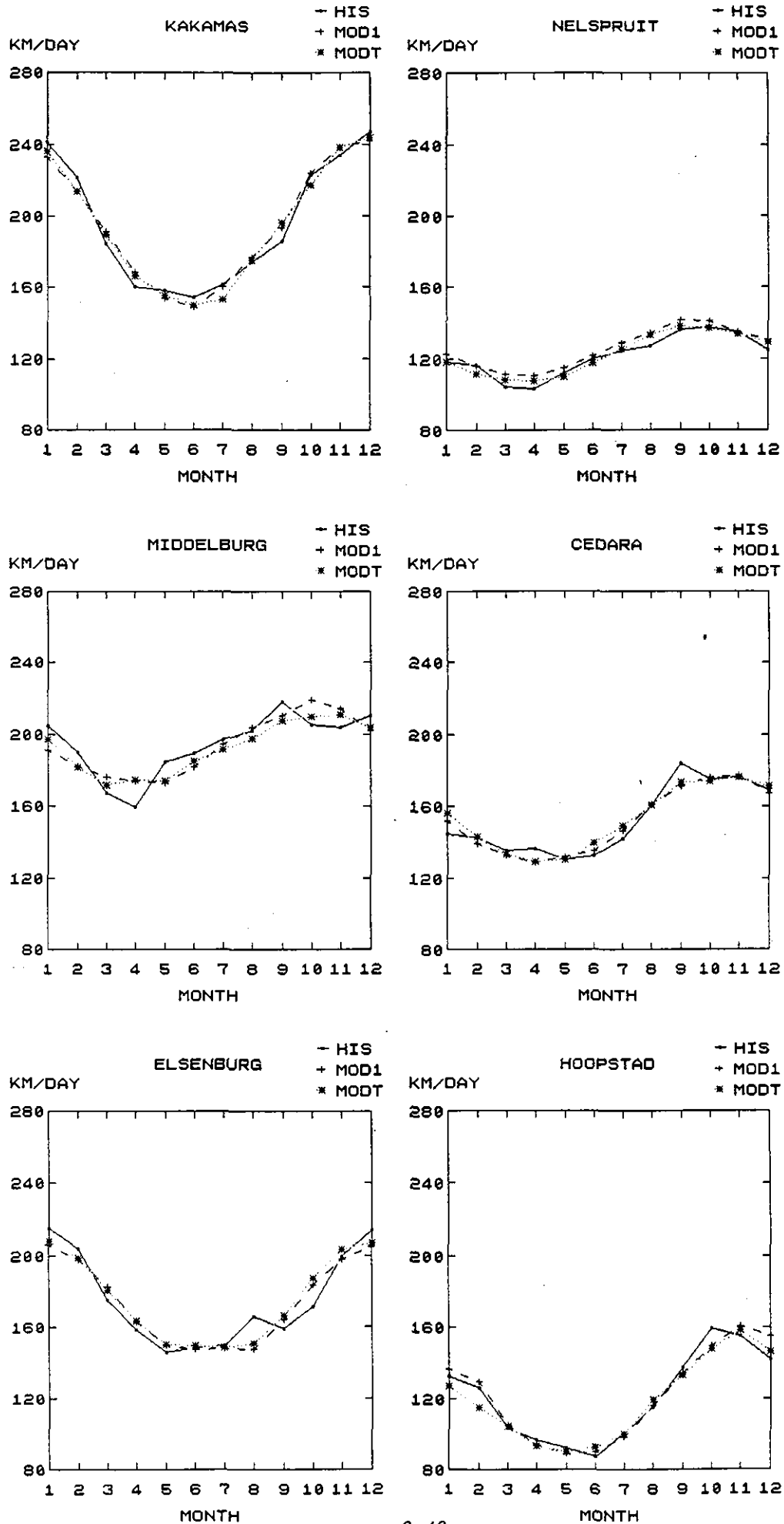
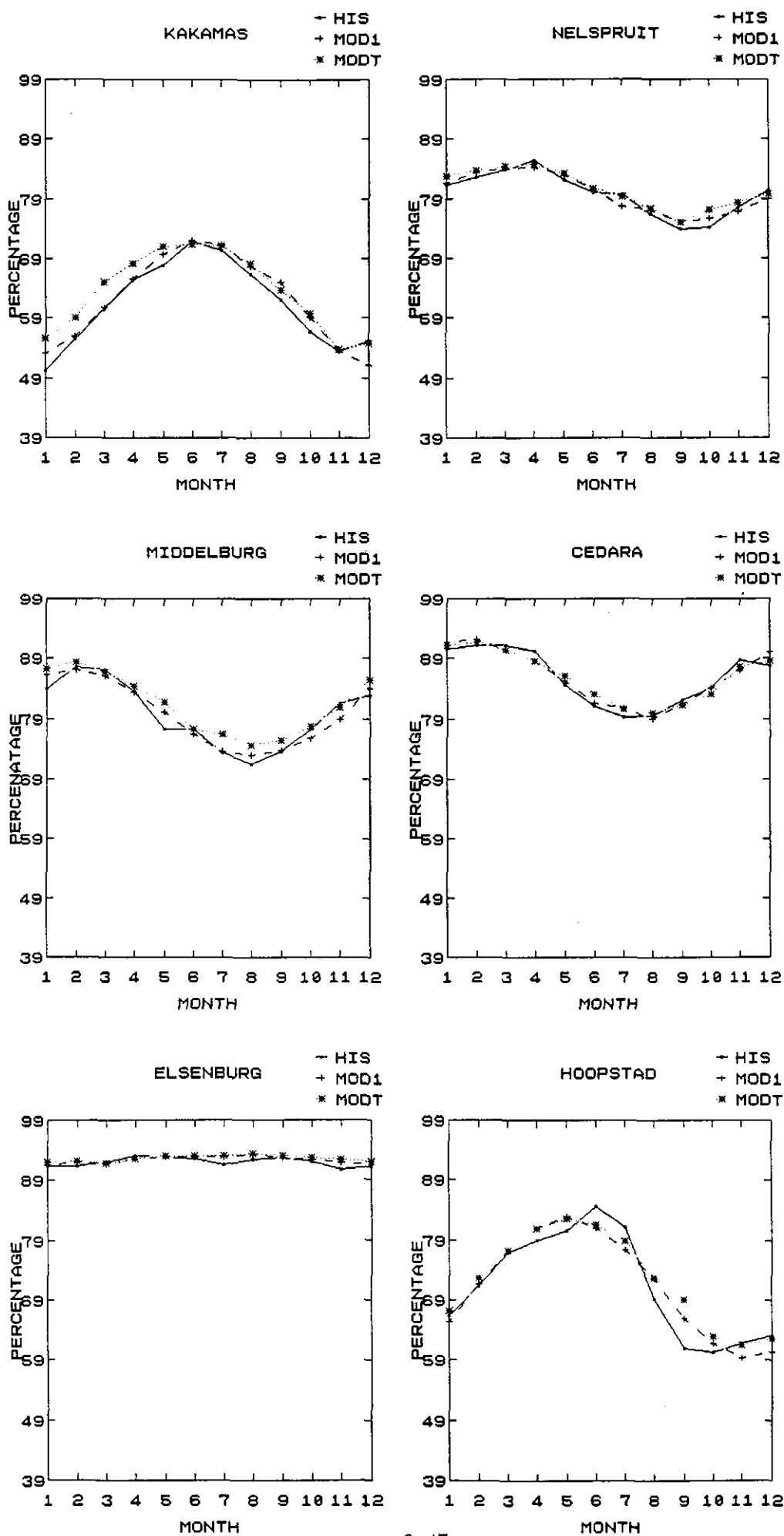
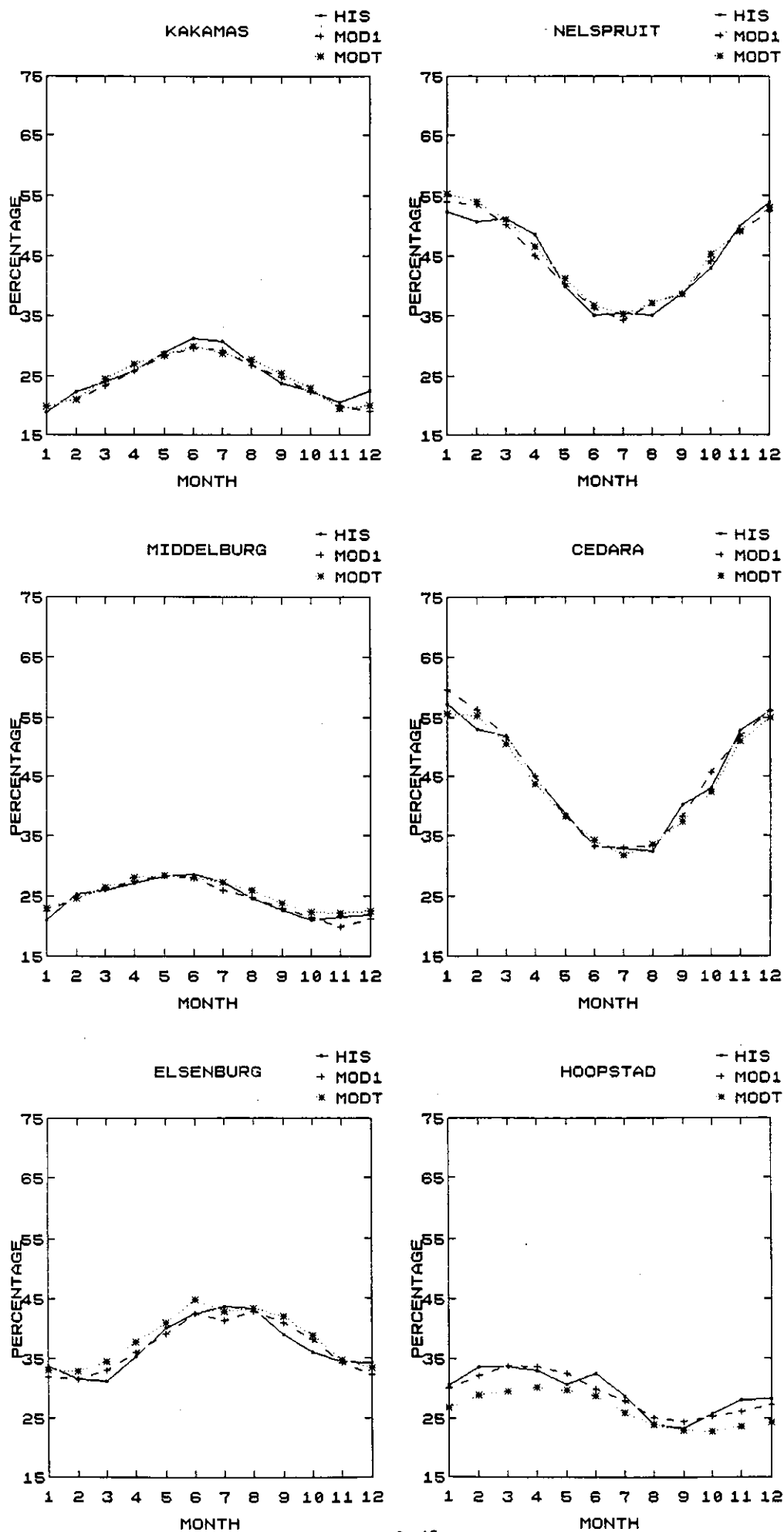


FIGURE 6.34 Monthly means for maximum humidity for dry days



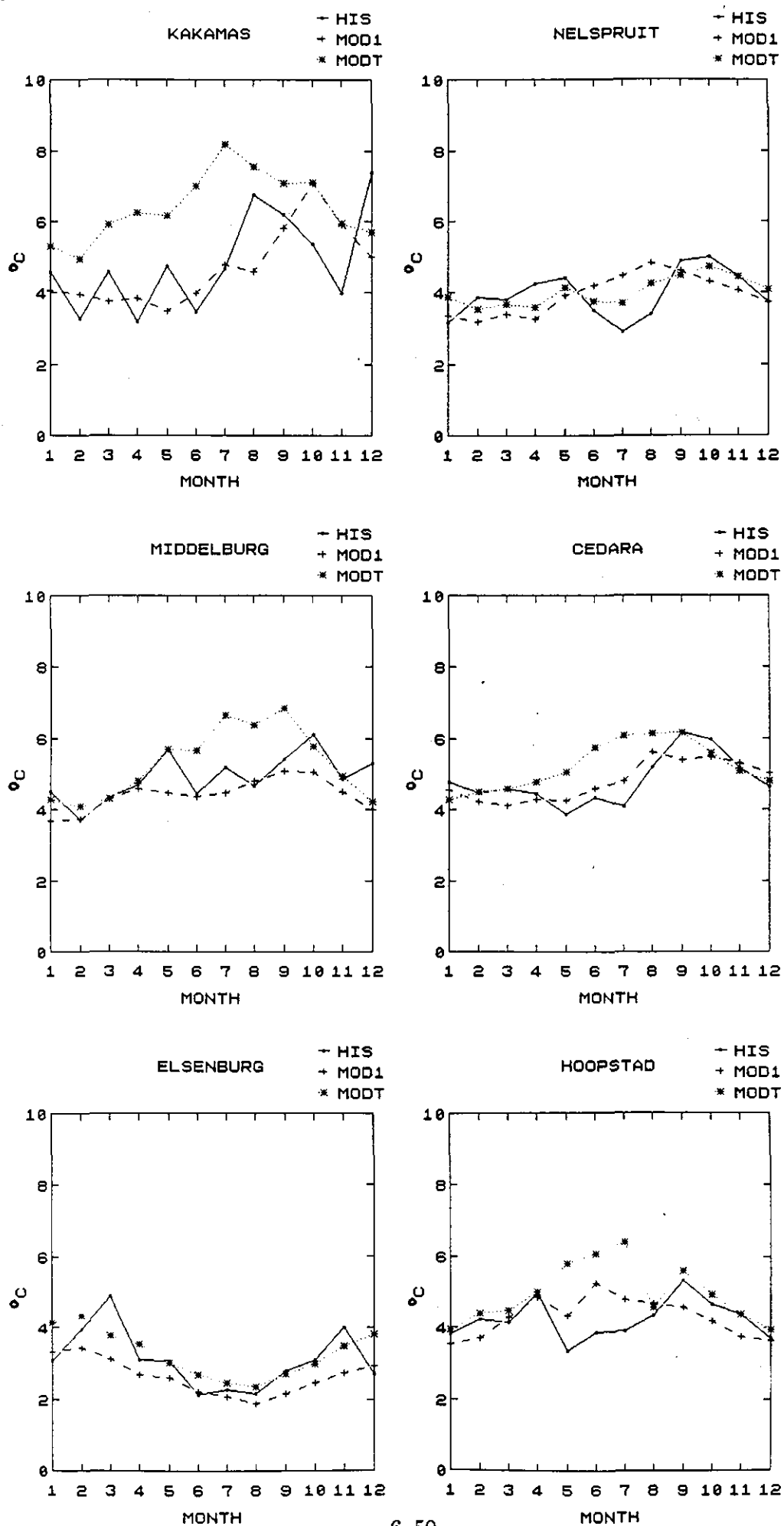
**FIGURE 6.35** Monthly means for minimum humidity for dry days

These figures show that monthly means are adequately preserved by both models, in particular, the models fit the monthly means very well on dry days. The same results are observed for the station Hoopstad for the variables maximum temperature and minimum humidity as when the sequences were treated as a whole. For the sequences of wet days, differences are observed between monthly means of the simulated sequences and the monthly means of the observed sequences, in particular for the variables wind run, maximum humidity and minimum humidity. As already mentioned there are relatively few observations of rainfall at these stations and therefore one does not expect the models to fit the historical records on wet days very accurately. This is supported by comparing the results obtained for the stations Kakamas and Cedara when the variables are conditioned on wet days. Kakamas is the station for which fewer rainfall days are observed, and Cedara is the station at which most rainfall days are observed, of the stations in this study. It is clear that both models preserve the monthly means for Cedara but do not perform as well for Kakamas.

**The plots also show that generally, Model 1 fits the data better than Model T.** This can be explained by observing that for Model 1 one is only separating the sequences into dry and wet days, while for Model T one separates the sequences into four parts, that is, into dry-dry, wet-wet, dry-wet and wet-dry sequences. Therefore the model parameters for Model T are estimated using very few observations especially when dealing with a wet sequence.

Figures 6.36 – 6.42 show the monthly standard deviations when the climate variables are conditioned on wet days. Figures 6.43 – 6.49 show the monthly standard deviations when the climate variables are conditioned on dry days. These plots show that both models have preserved monthly standard deviations when the climate variables are conditioned on dry days. Here again, very similar results are obtained to those when the sequences are taken as a whole. Generally, the models preserve the monthly standard deviations when the variables are conditioned on wet days. Some differences are observed between the simulated sequences and the historical records. These differences can be explained again for the reasons mentioned above. Where differences between simulated and historical sequences occur, it can be noted that generally Model T tends to overestimate the standard deviations, while Model 1 tends to underestimate them.

**FIGURE 6.36** Monthly standard deviations for maximum temperature for wet days



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**FIGURE 6.37** Monthly standard deviations for minimum temperature for wet days

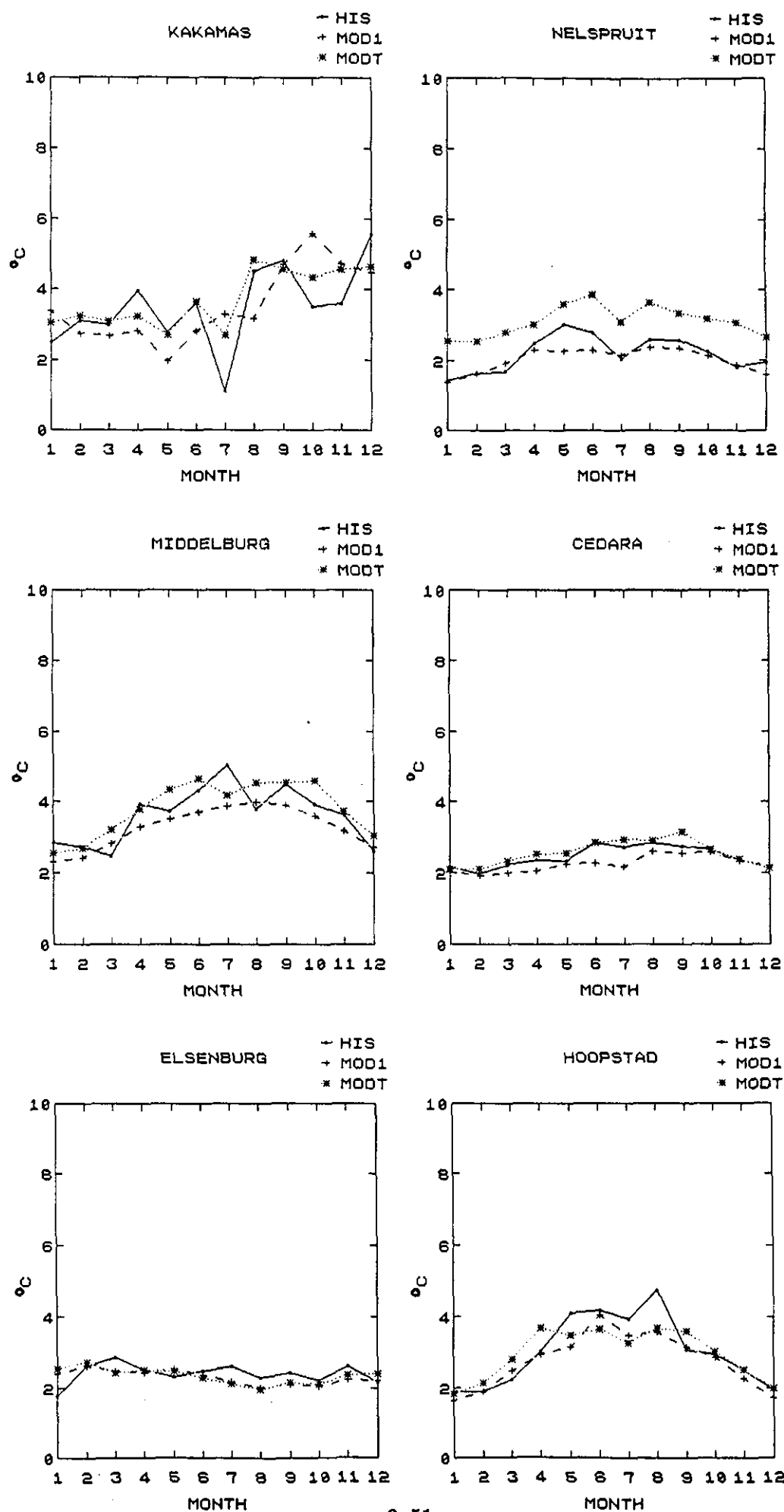


FIGURE 6.38 Monthly standard deviations for evaporation for wet days

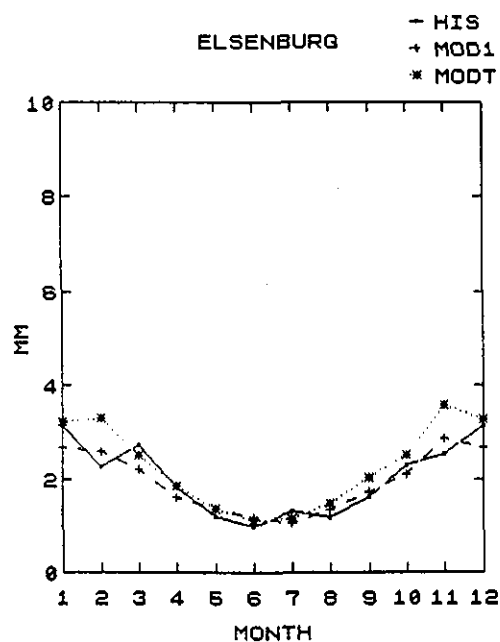
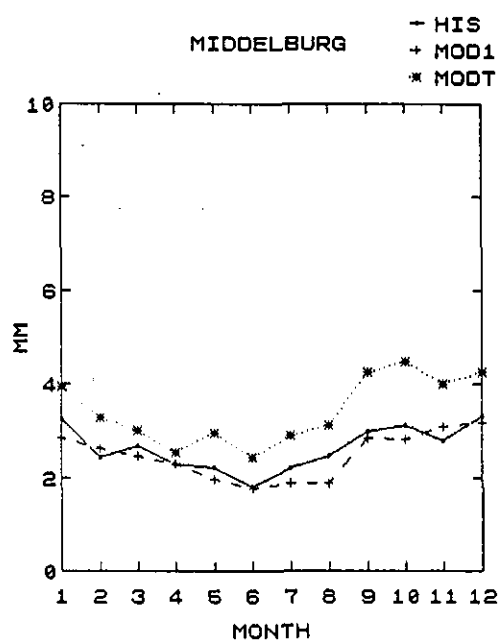
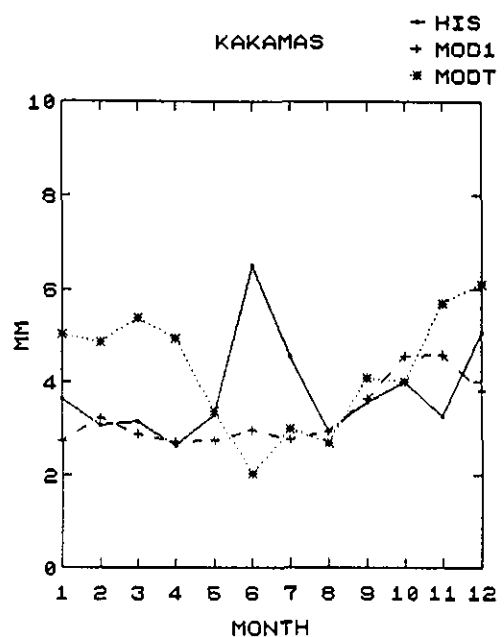


FIGURE 6.39 Monthly standard deviations for sunshine duration for wet days

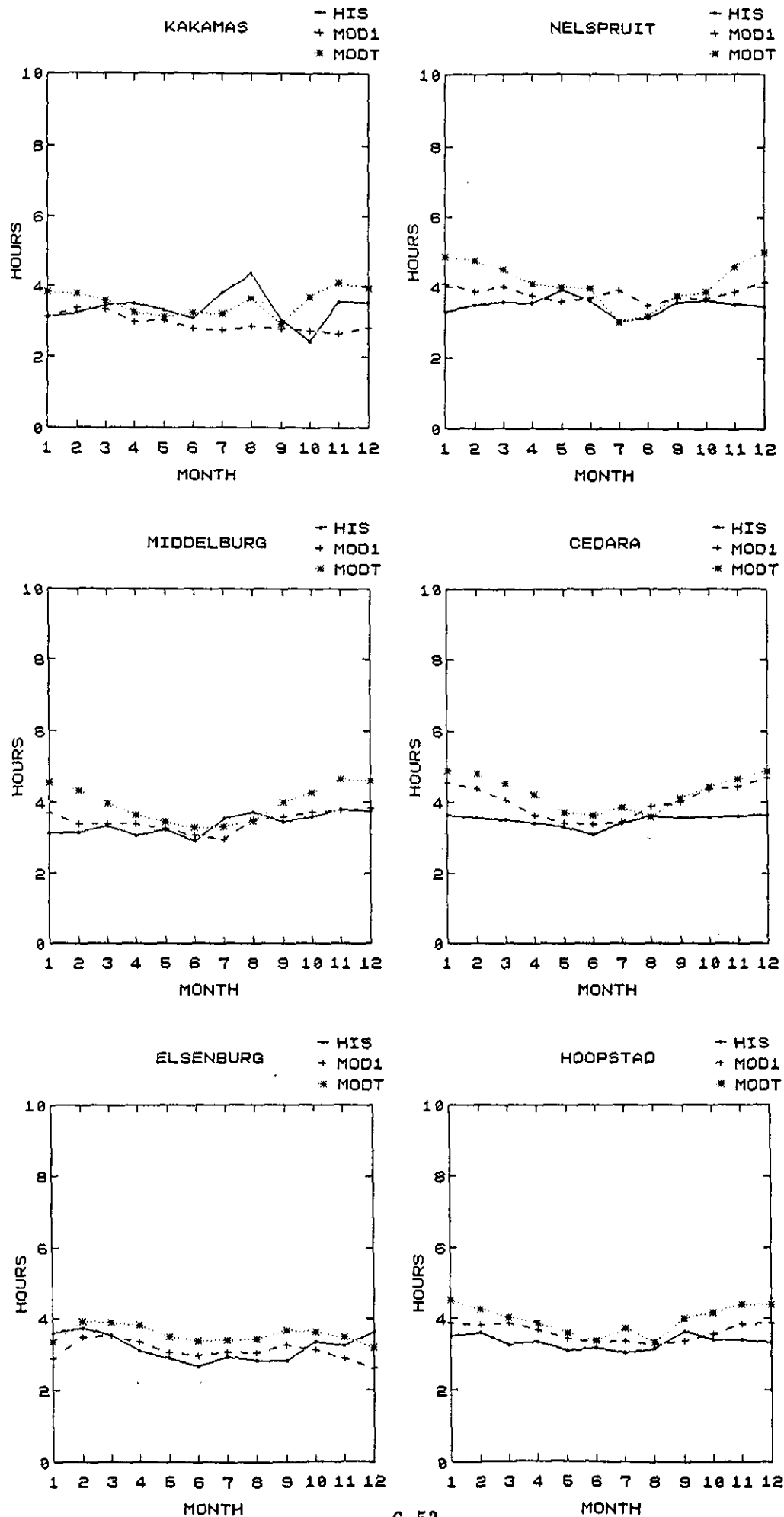


FIGURE 6.40 Monthly standard deviations for wind run for wet days

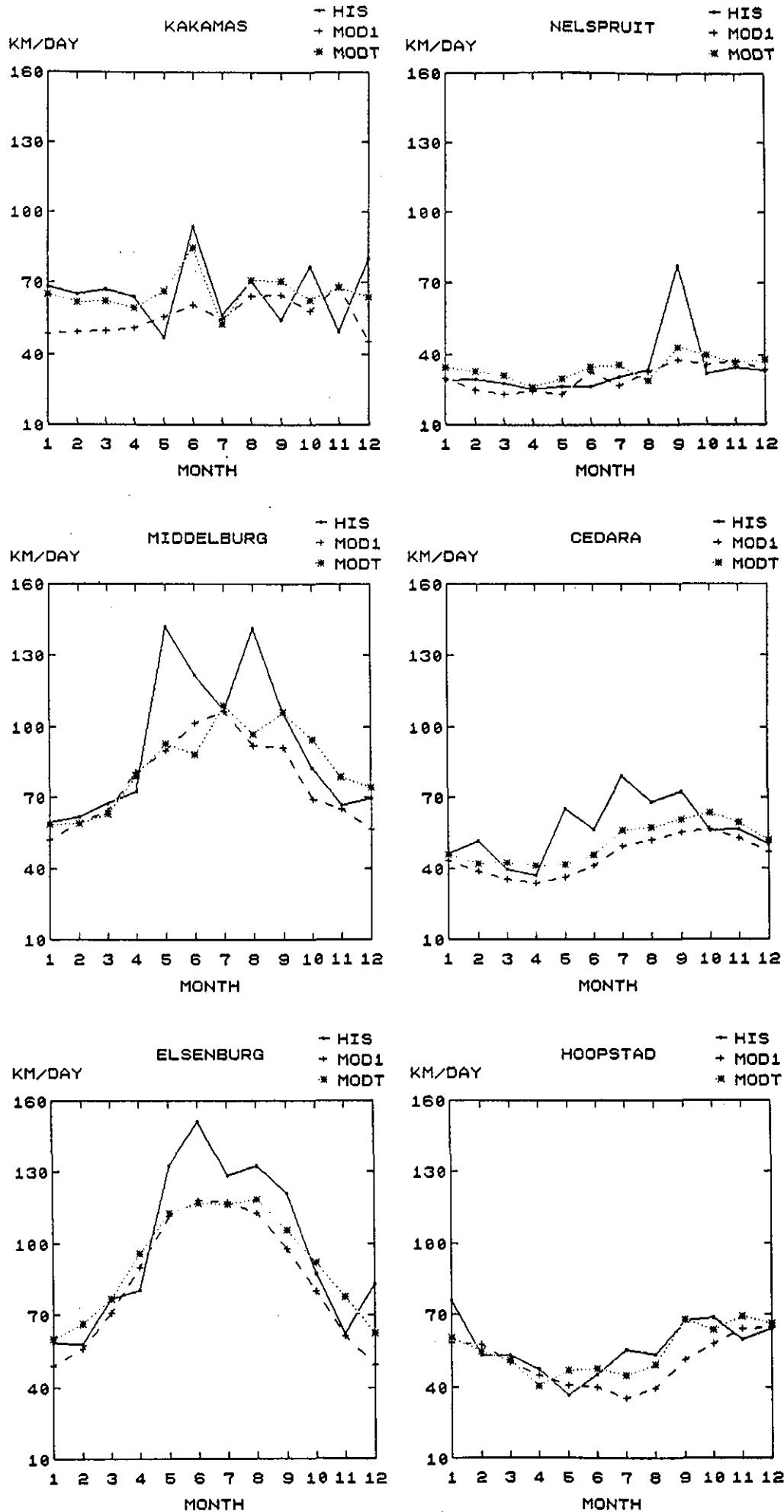


FIGURE 6.41 Monthly standard deviations for maximum humidity for wet days

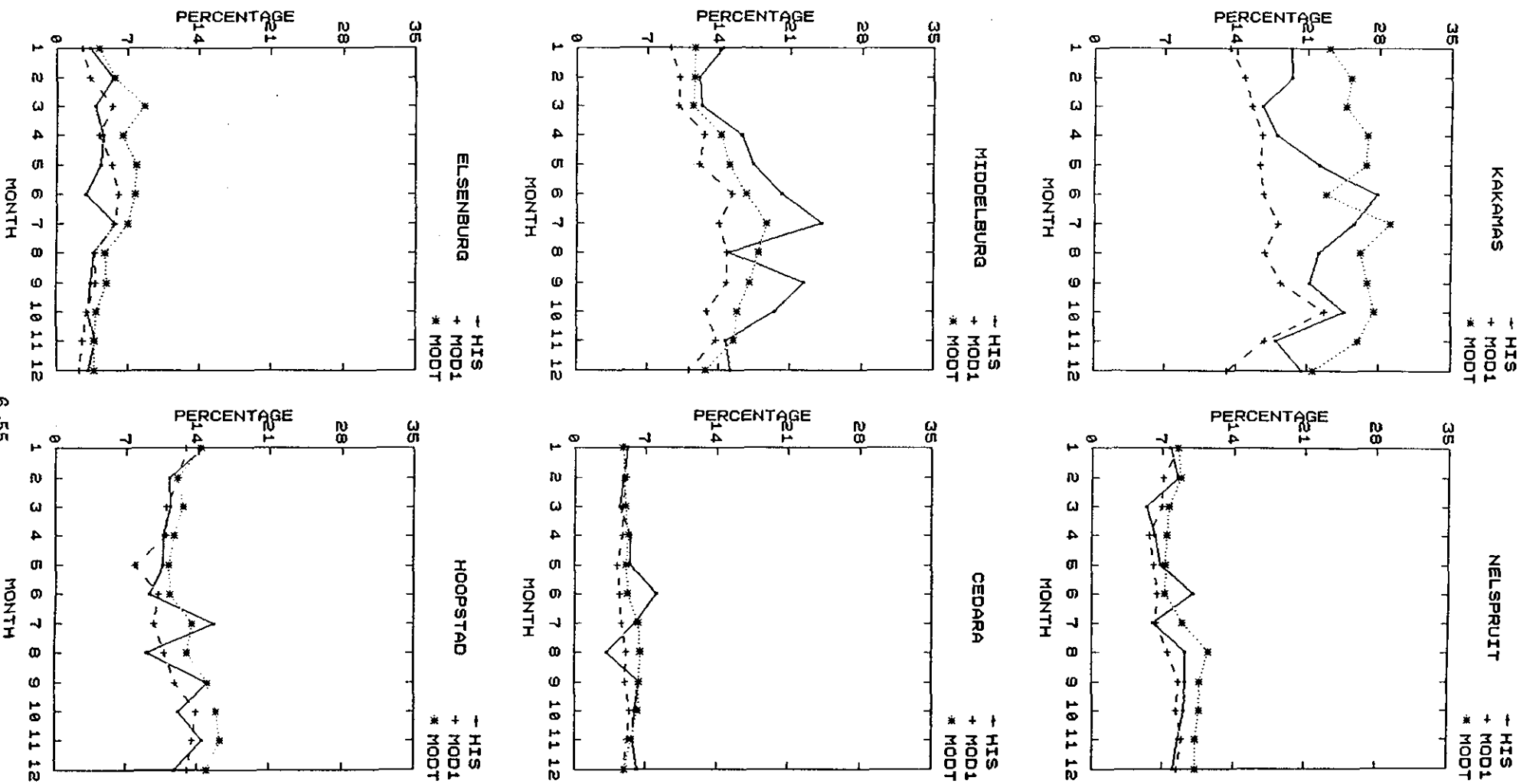


FIGURE 6.42 Monthly standard deviations for minimum humidity for wet days

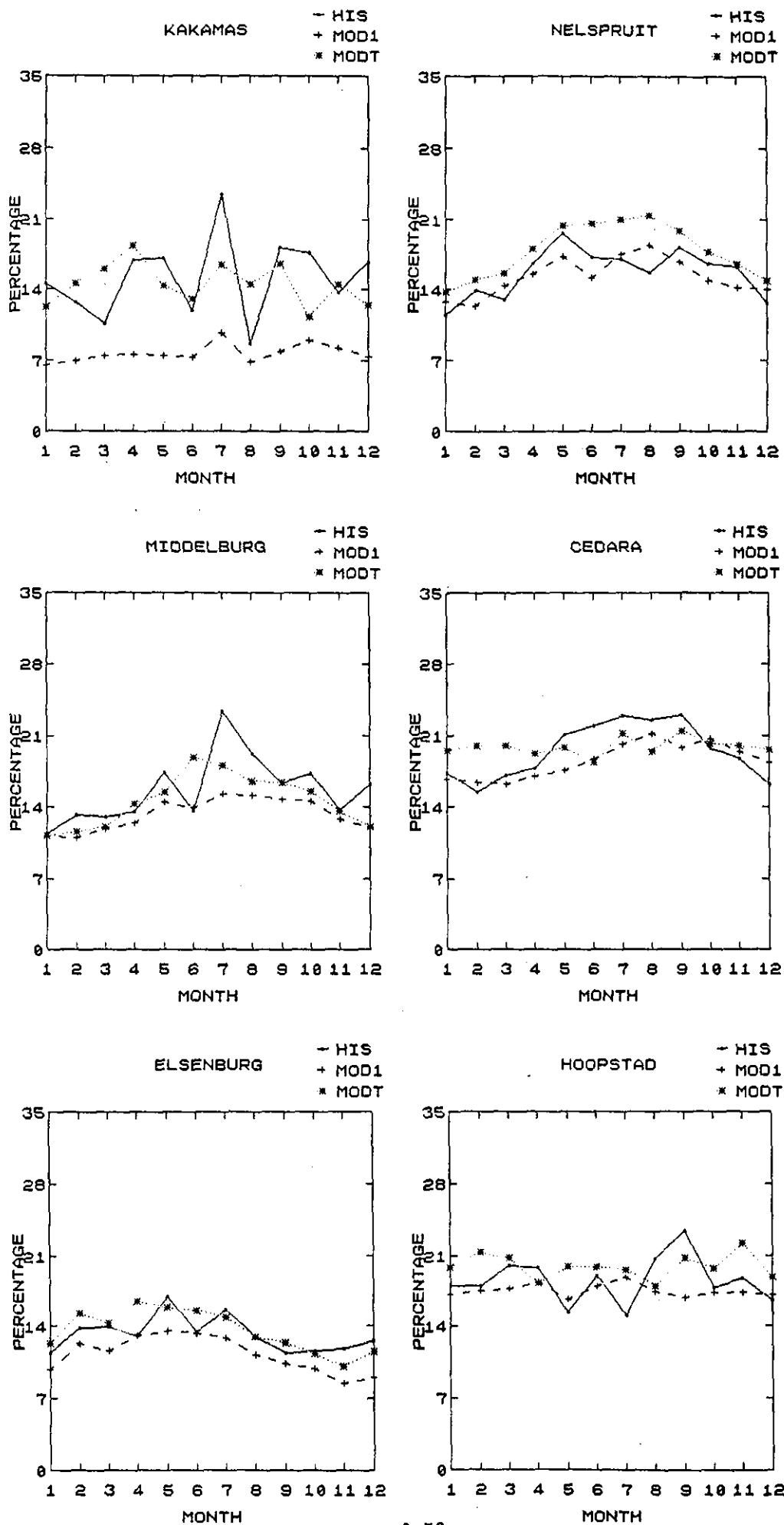
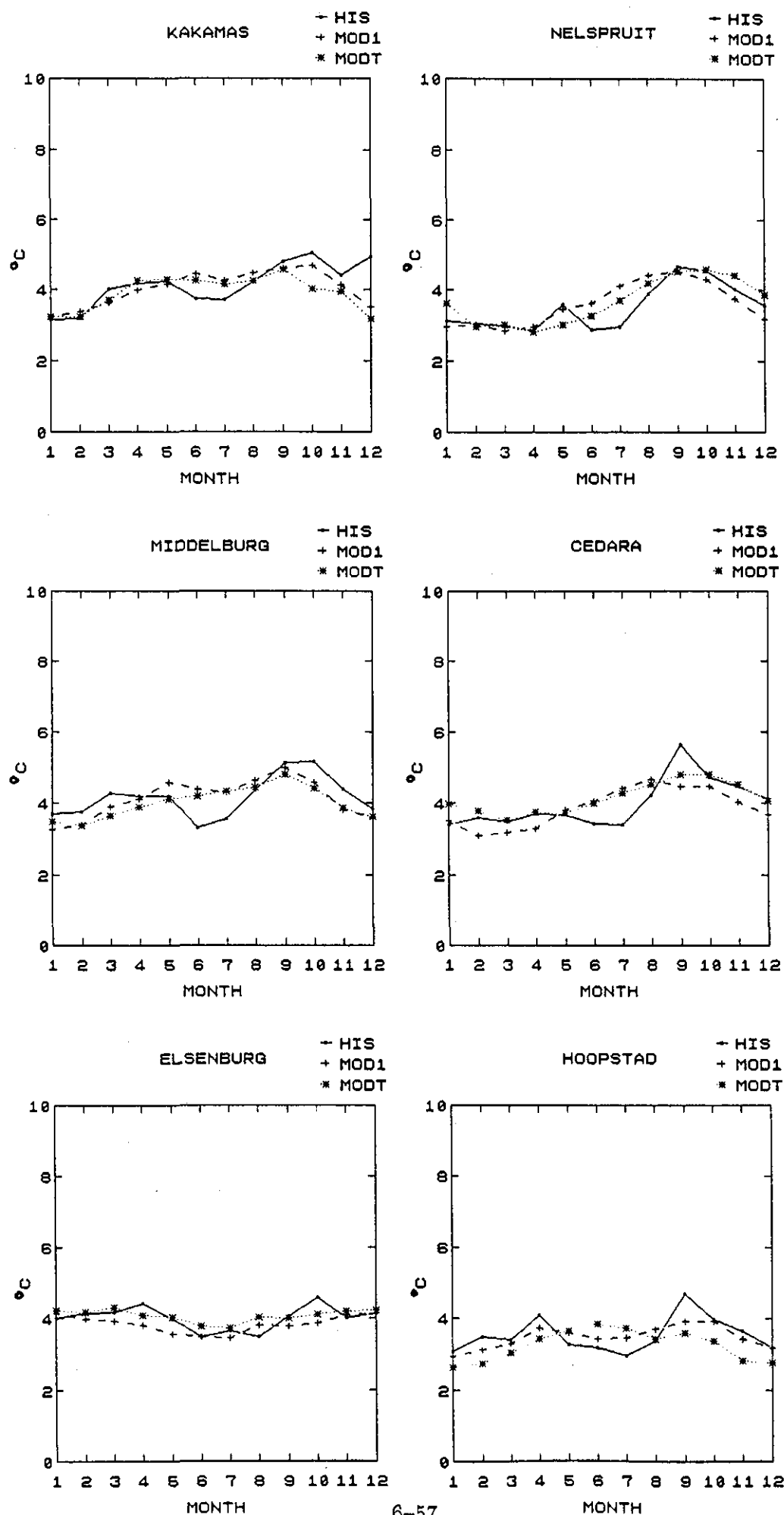


FIGURE 6.43 Monthly standard deviations for maximum temperature for dry days



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**FIGURE 6.44** Monthly standard deviations for minimum temperature for  
 dry days

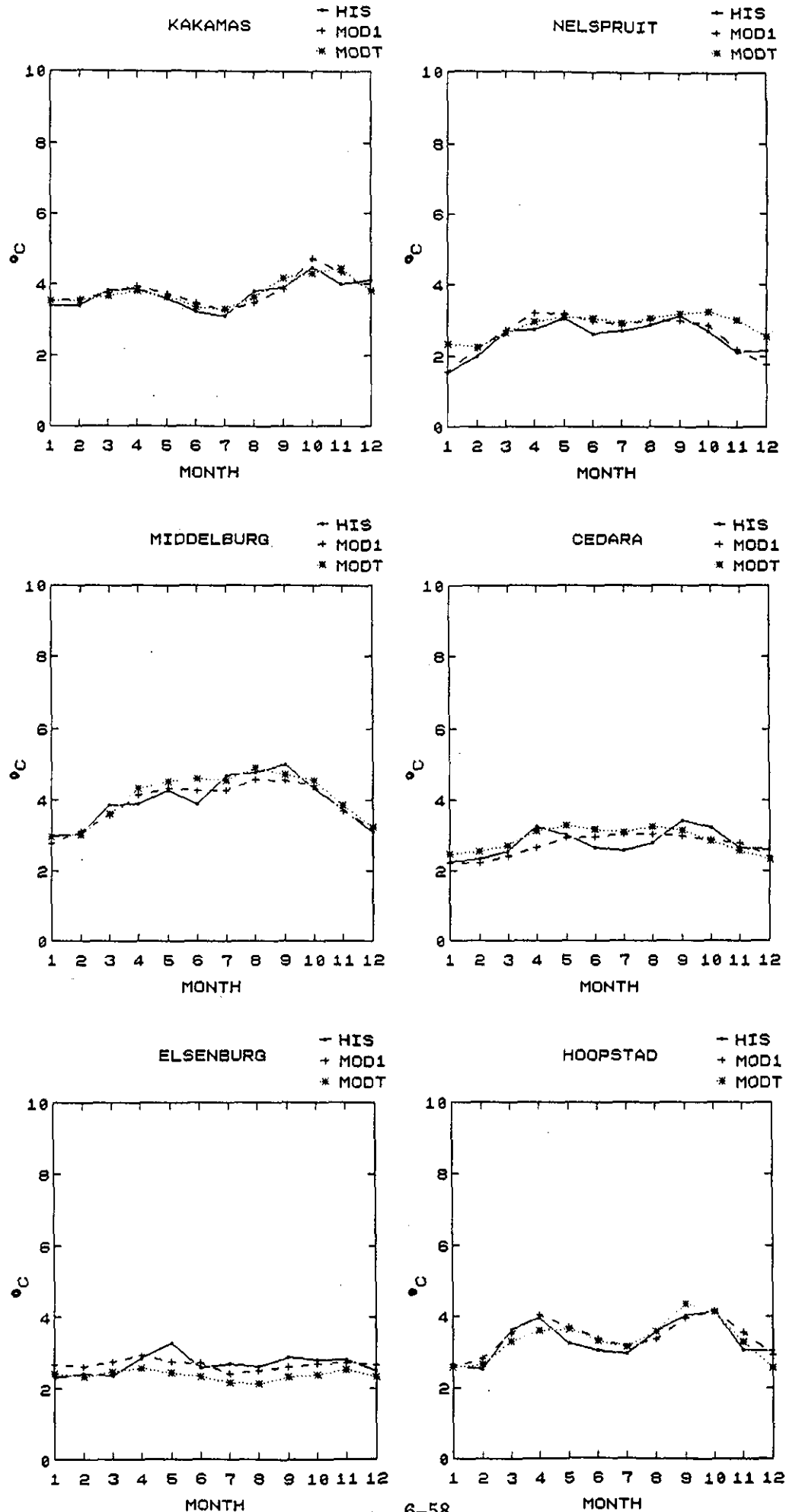


FIGURE 6.45 Monthly standard deviations for evaporation for dry days

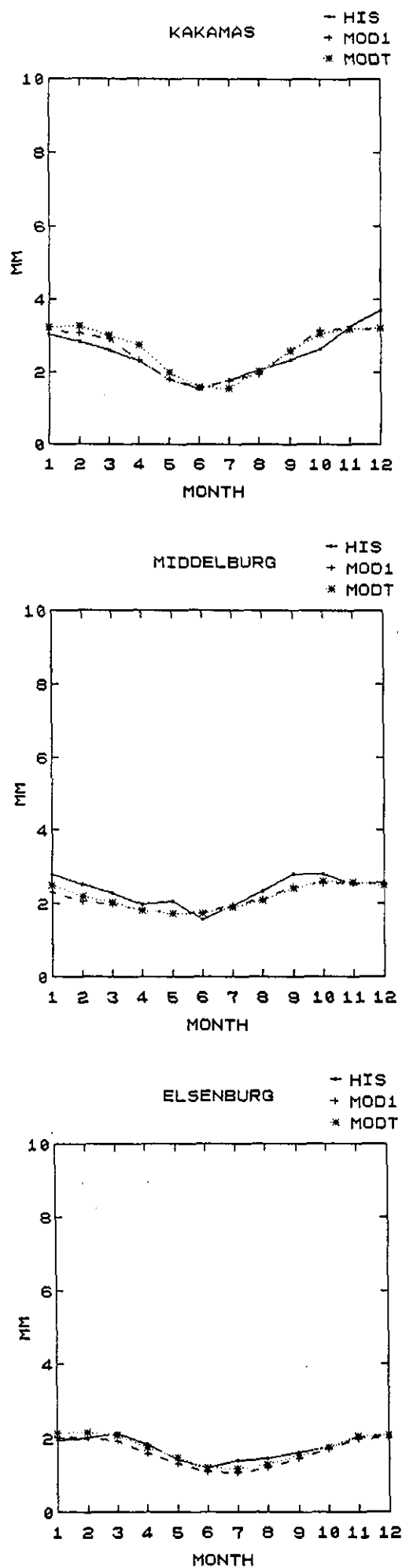


FIGURE 6.46 Monthly standard deviations for sunshine duration for dry days

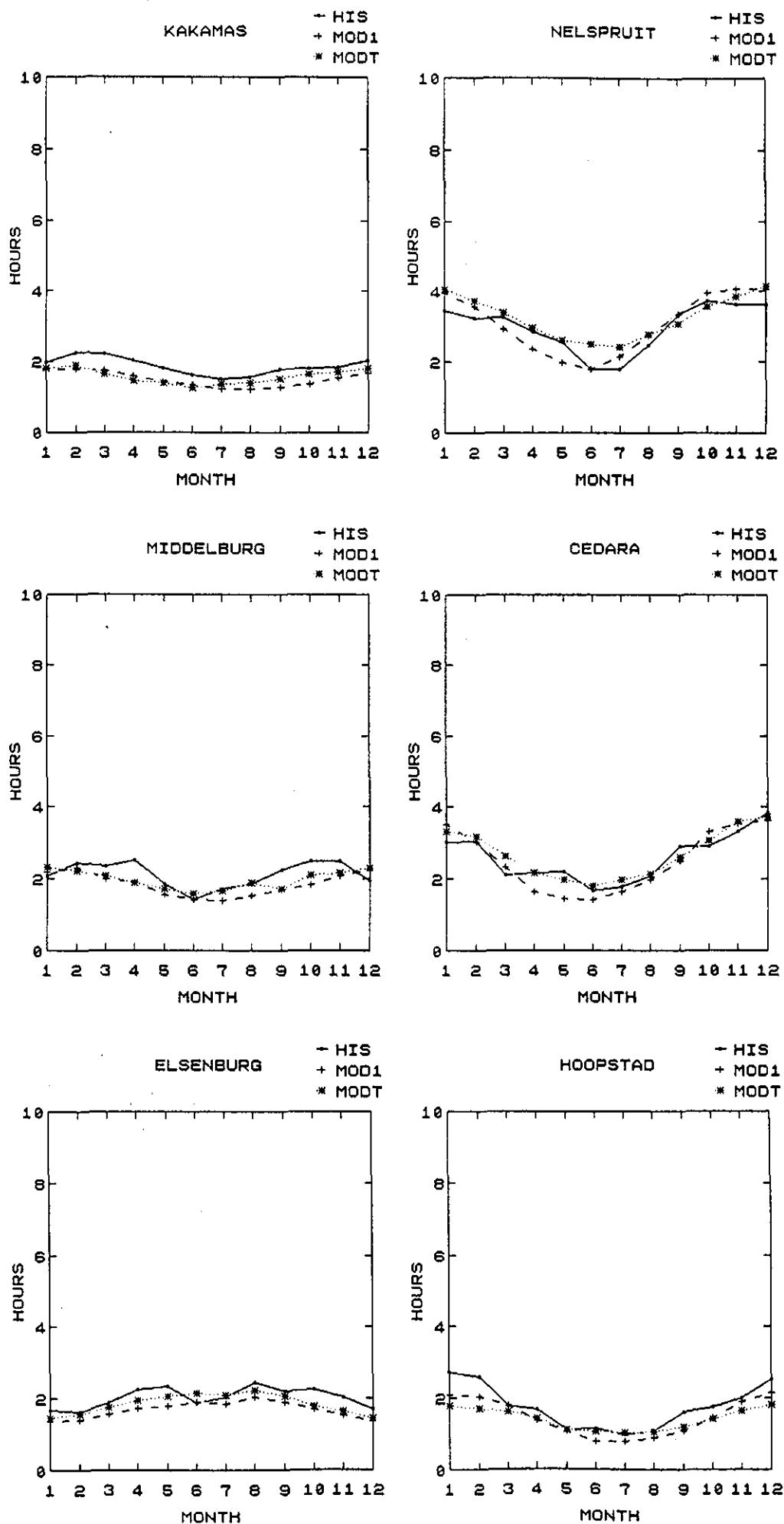


FIGURE 6.47 Monthly standard deviations for wind run for dry days

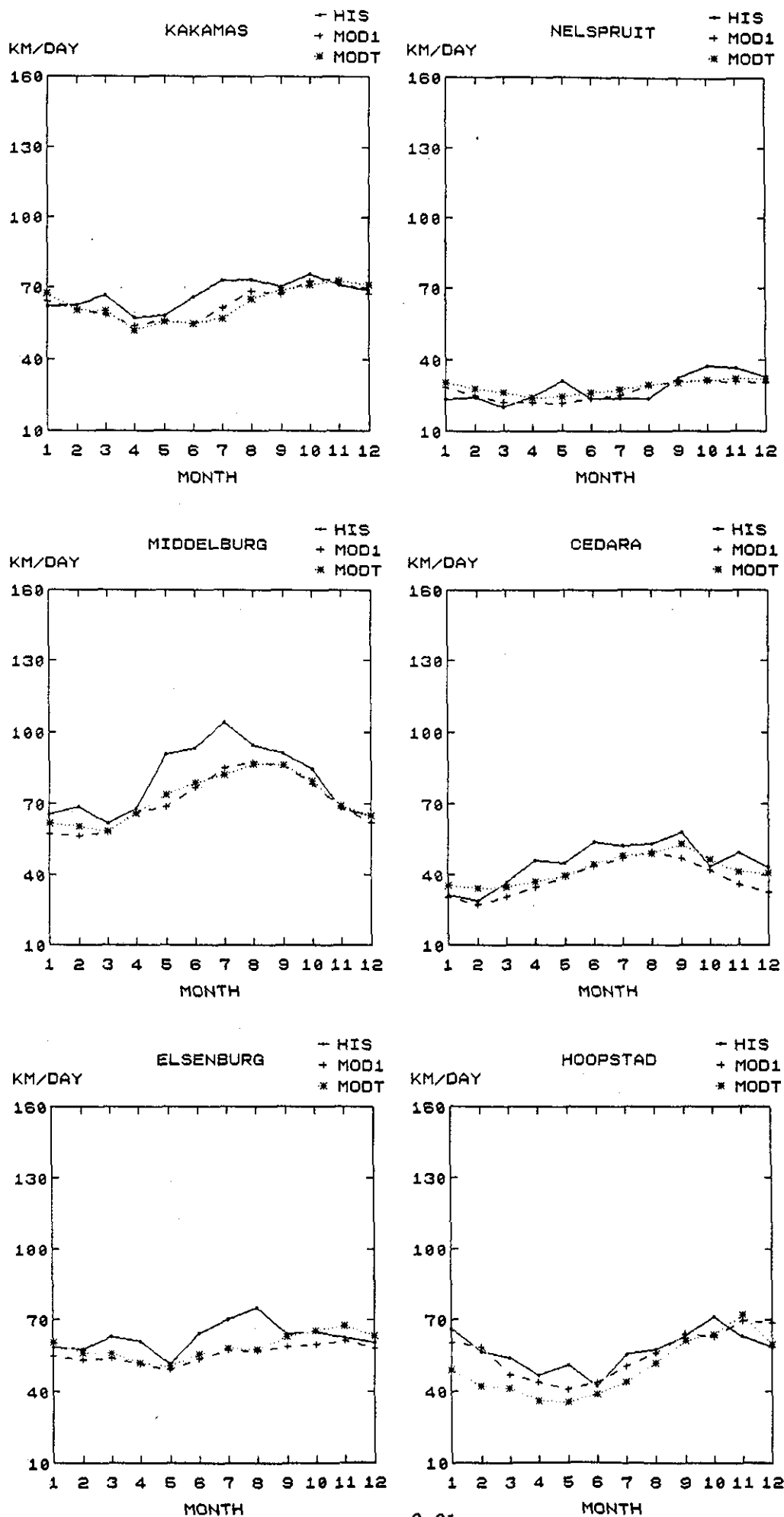


FIGURE 6.48 Monthly standard deviations for maximum humidity for dry days

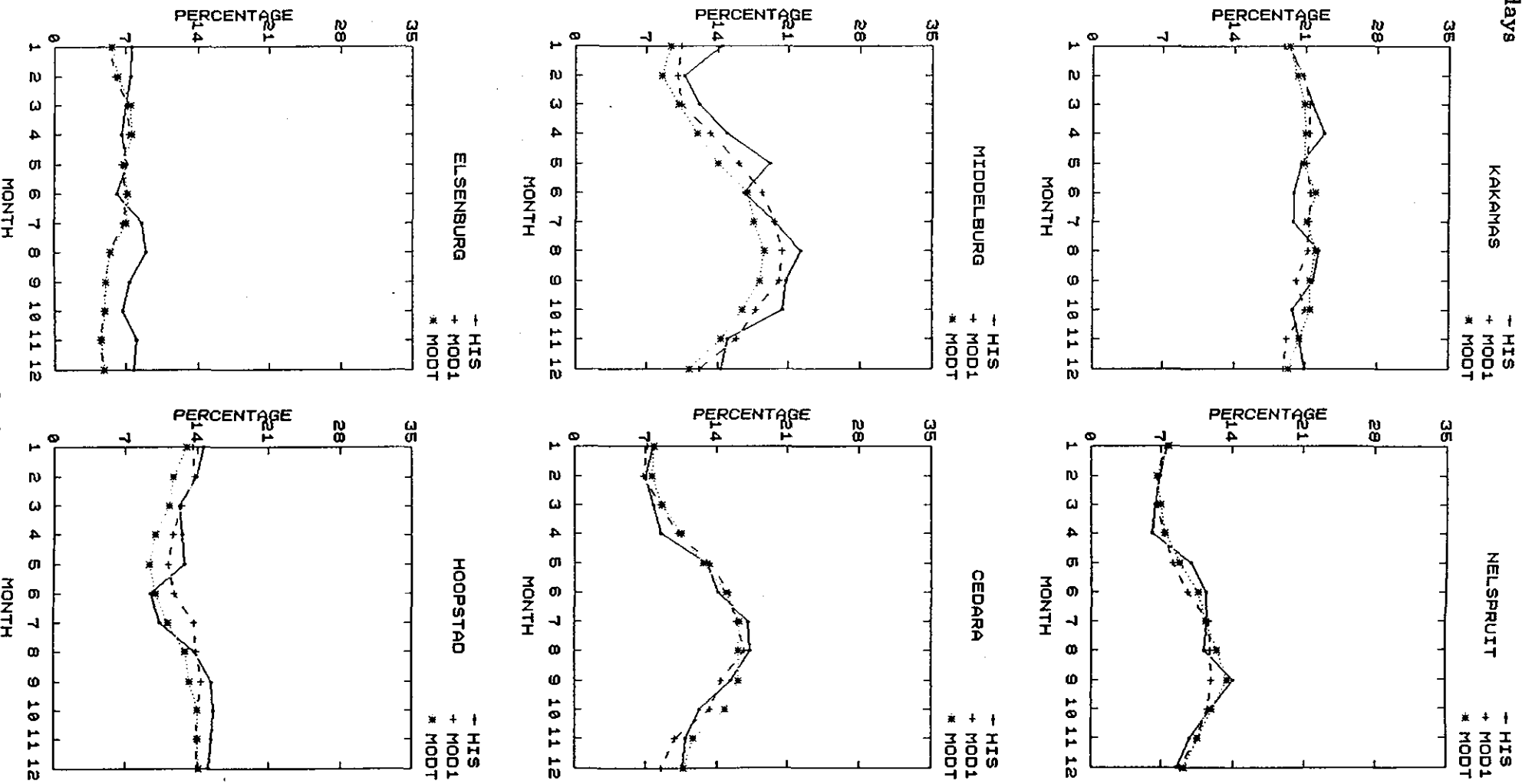
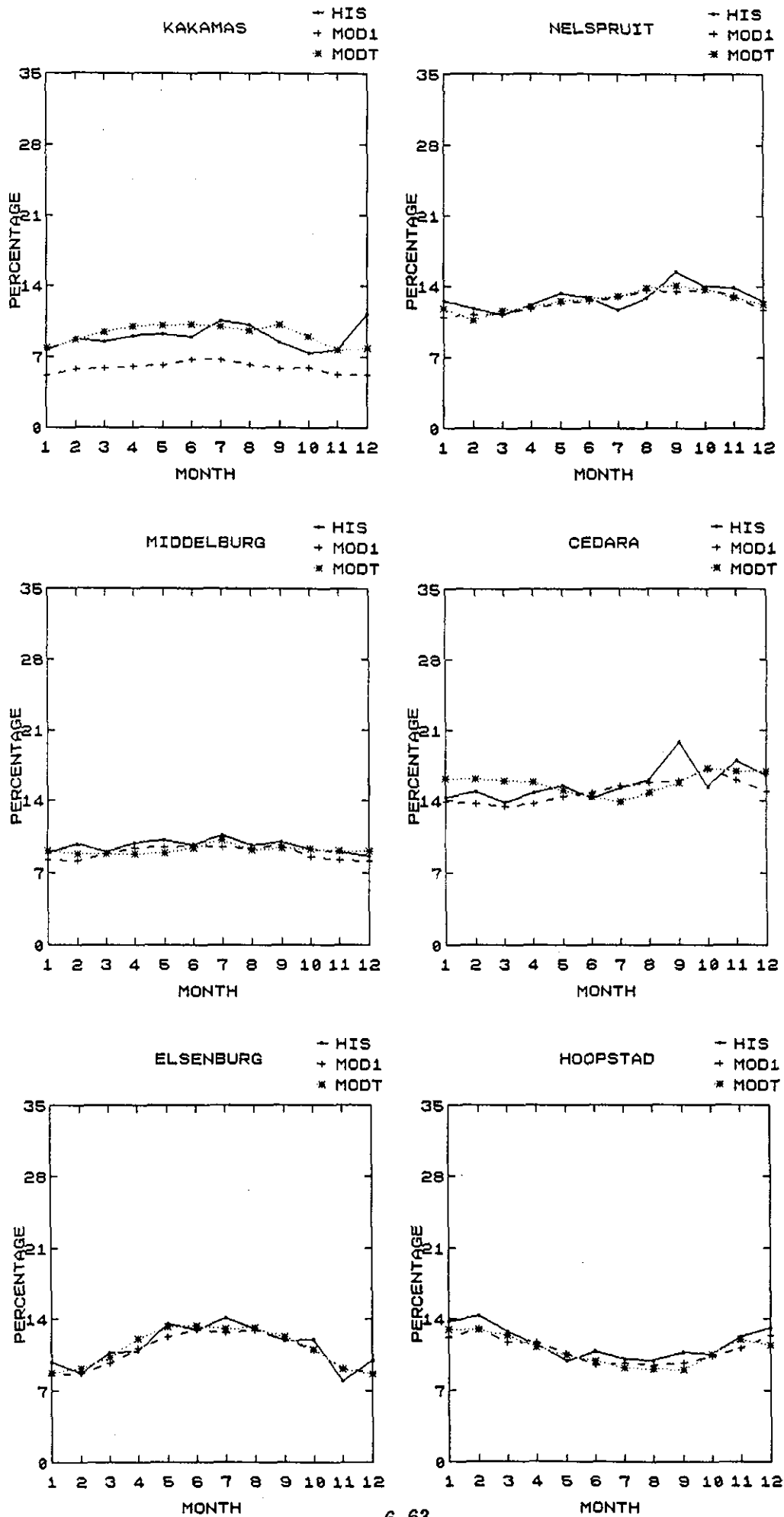


FIGURE 6.49 Monthly standard deviations for minimum humidity for dry days



**Autocorrelation property**

One of the properties exhibited by the historical record is that of autocorrelation, that is, within each climate variable there is a short-term persistence. For example, the temperature observed on a given day is statistically related to the temperature observed on the previous day.

To determine whether the models were successful in reproducing the autocorrelation structure that is present in the historical data, the autocorrelation coefficients (of up to a lag of four days) of each variable in the simulated sequence were compared to those of the historical record (Figures 6.50 – 6.57). From these comparisons it can be concluded that both models have described the autocorrelation property very well. Any differences observed between the simulated and historical sequences are mostly within 0.1 of the historical record. The variables that show these differences are generally sunshine, maximum humidity and minimum humidity. It must be noted here that the models assume an autoregressive process of order 1, that is, a lag of one day, and the bigger differences observed occur for lags of two or more days. Models with a higher autoregressive order might describe the autocorrelation structure of these variables, but this would mean increasing the complexity and the number of parameters in the models.

The autocorrelation coefficients of the simulated sequences were compared with those of the historical data, both for wet (Figures 6.57 – 6.63) and for dry sequences (Figures 6.64 – 6.70). The plots show that the autocorrelation structure in the simulated sequences closely resembles that of the observed data. Again the differences that are observed between the simulated and the historical sequences are mostly within 0.1 of the historical record.

**Cross-correlation property**

The cross-correlation coefficients for lag -1, 0 and 1 were used in the simulation technique. Therefore, it is necessary that the models should maintain this property. Figures 6.71 – 6.91 show the comparison of the historical and simulated cross-correlation coefficients for all climate variables. Generally, the models have successfully preserved the cross-correlation coefficients, in particular the lag 0 cross-correlation. The only exceptions to this are the cross-correlation coefficients of the simulated sequence of Model T, between the variables maximum and minimum humidity and the other climate variables, in particular for the stations Kakamas, Middelburg and Hoopstad. The cross-correlation of other variables

FIGURE 6.50 Autocorrelation coefficients for maximum temperature

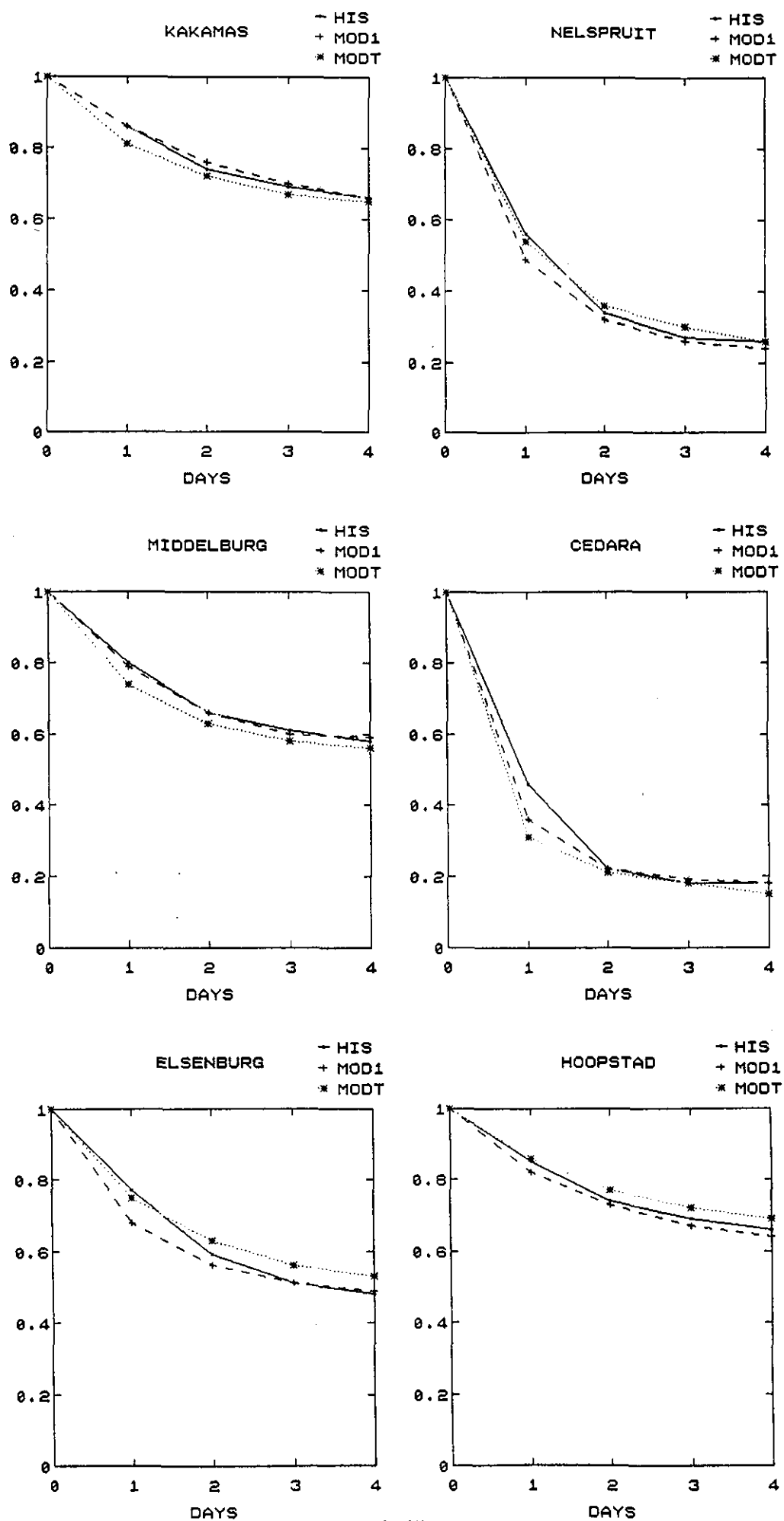


FIGURE 6.51 Autocorrelation coefficients for minimum temperature

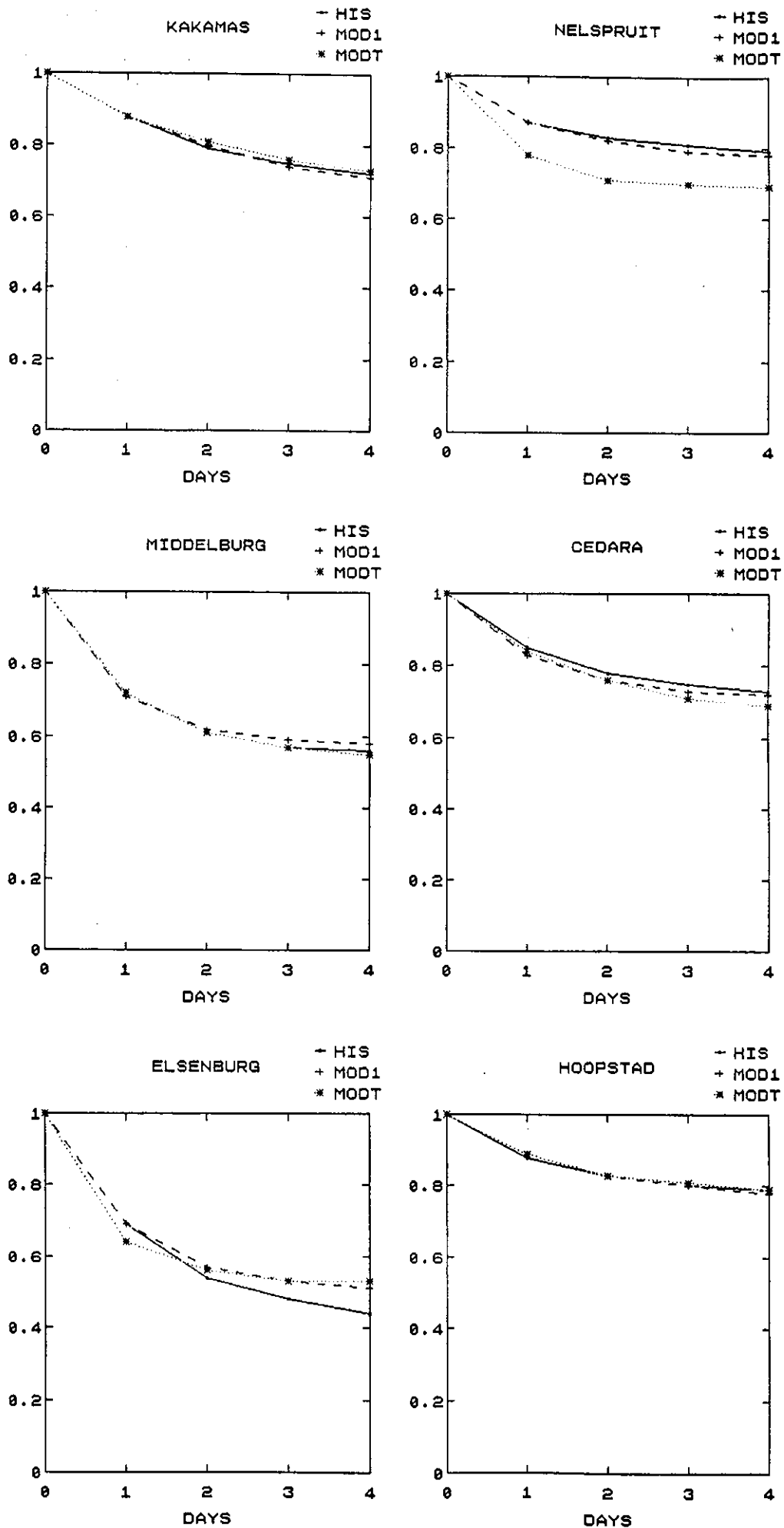


FIGURE 6.52 Autocorrelation coefficients for evaporation

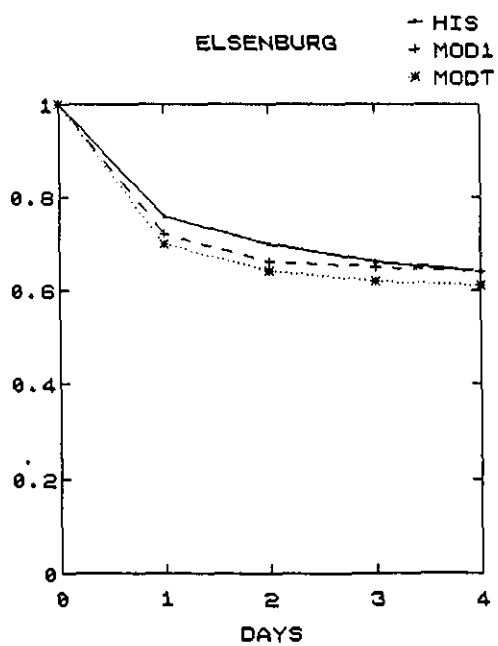
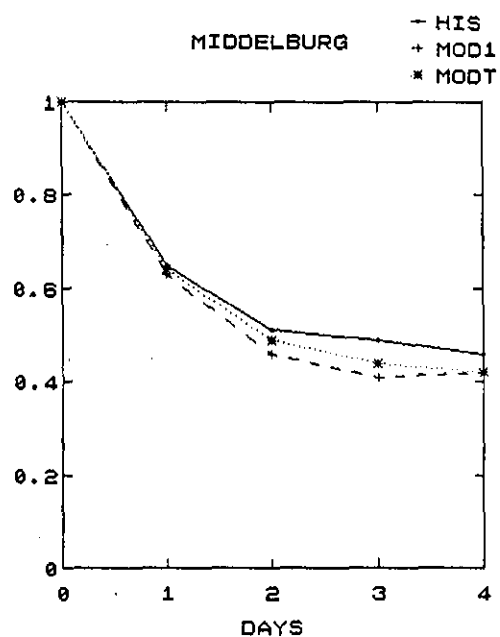
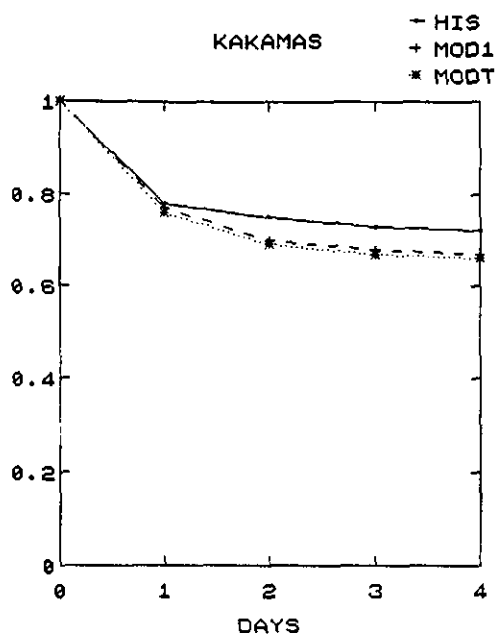
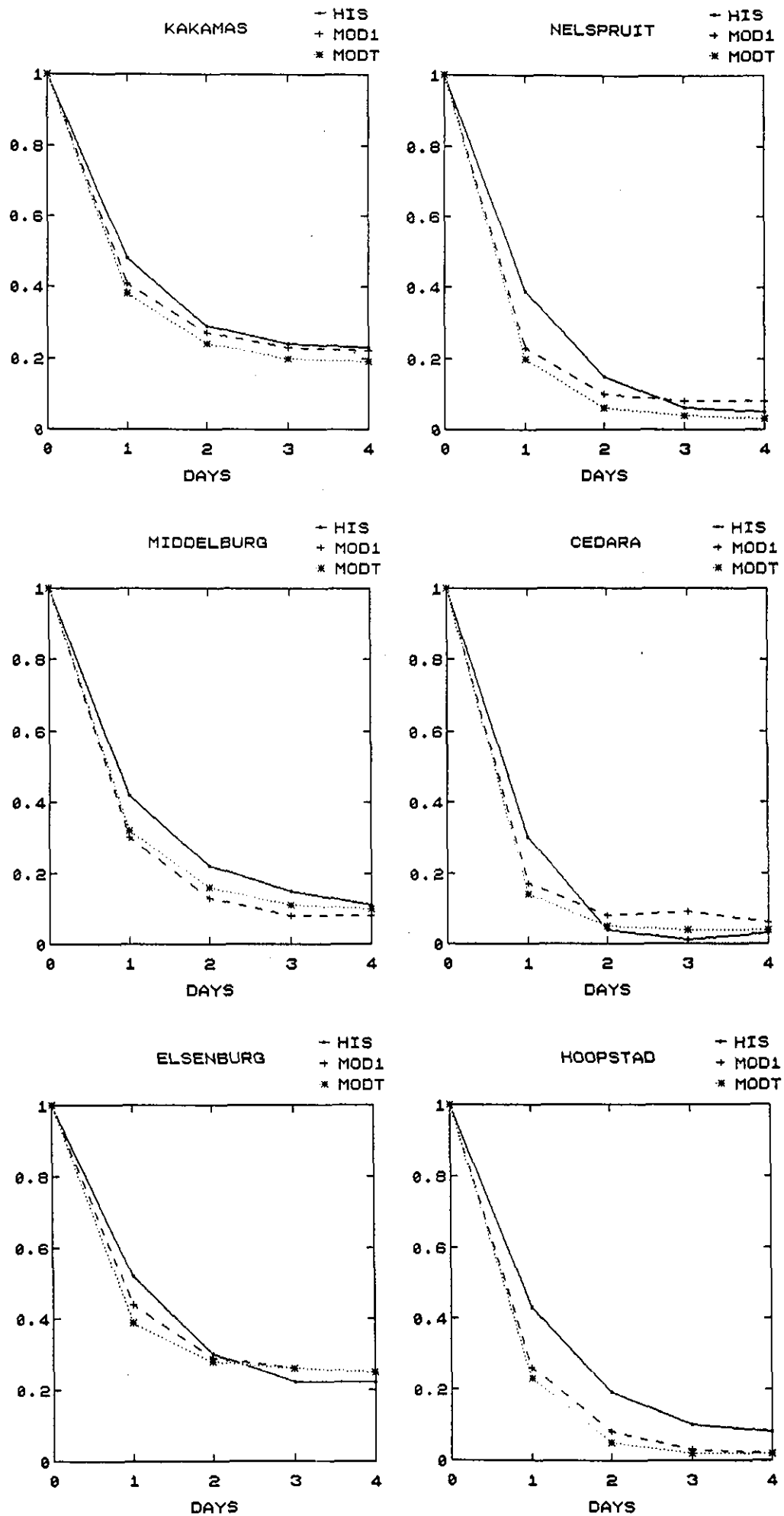


FIGURE 6.53 Autocorrelation coefficients for sunshine duration



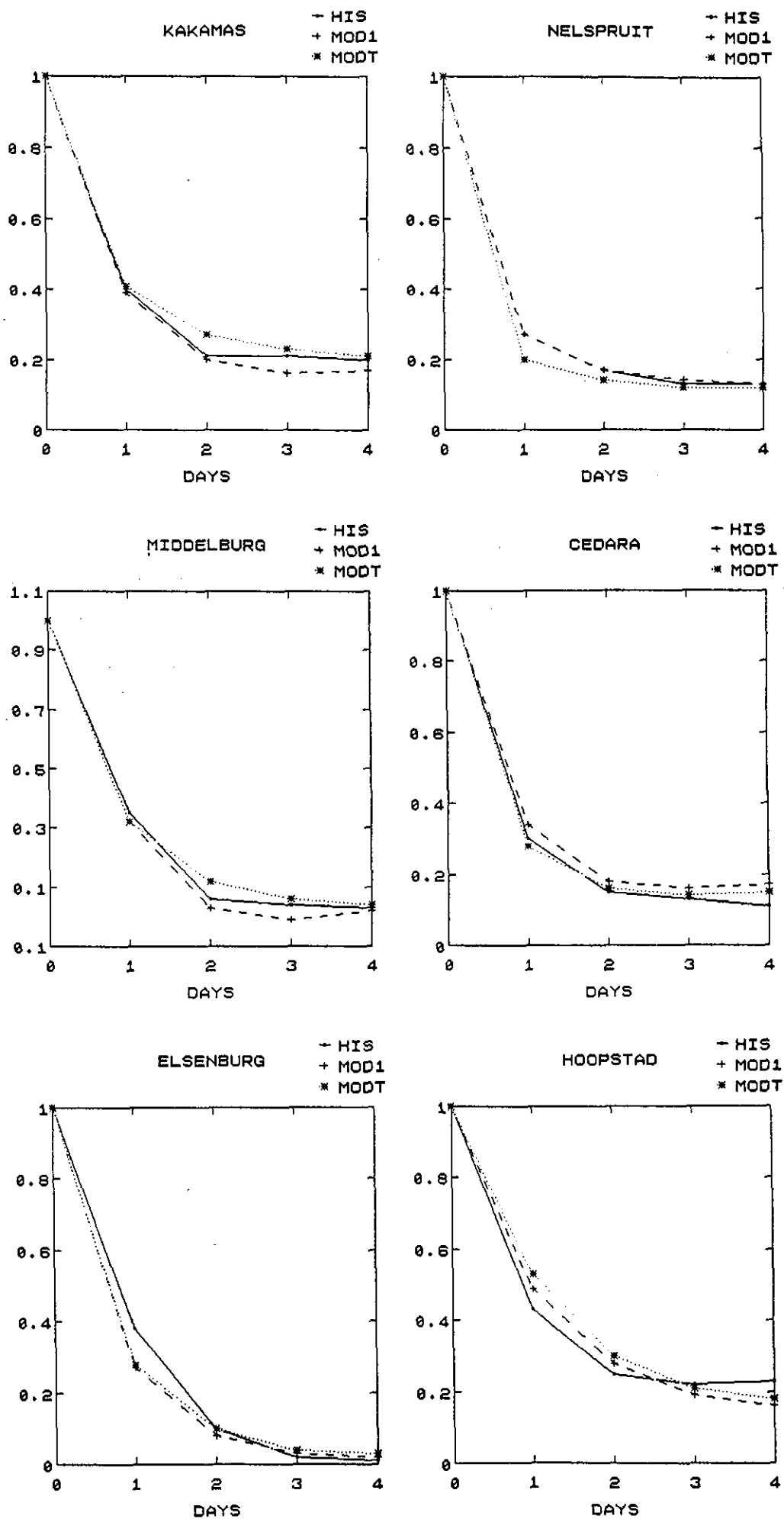


FIGURE 6.55 Autocorrelation coefficients for maximum humidity

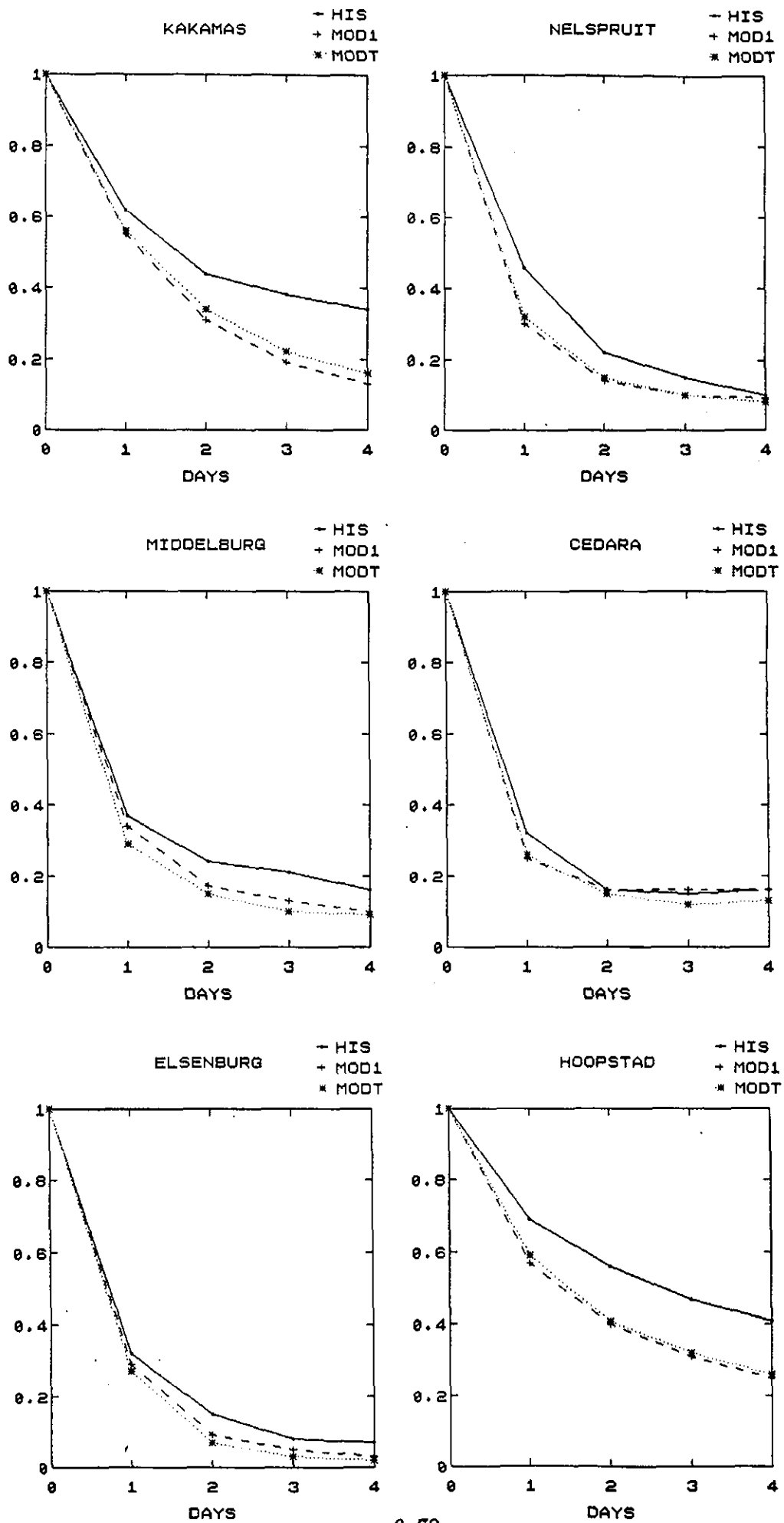


FIGURE 6.56 Autocorrelation coefficients for minimum humidity

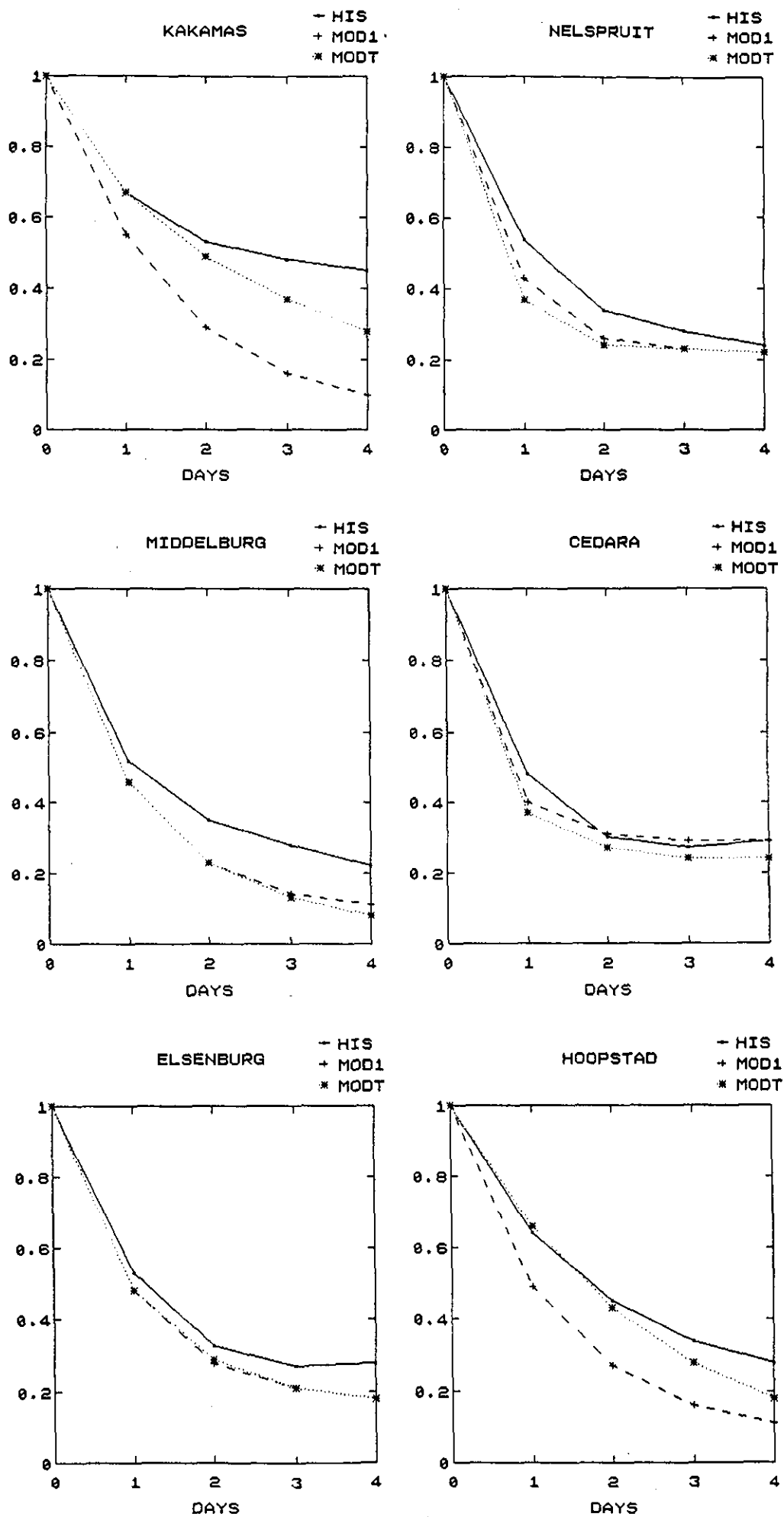
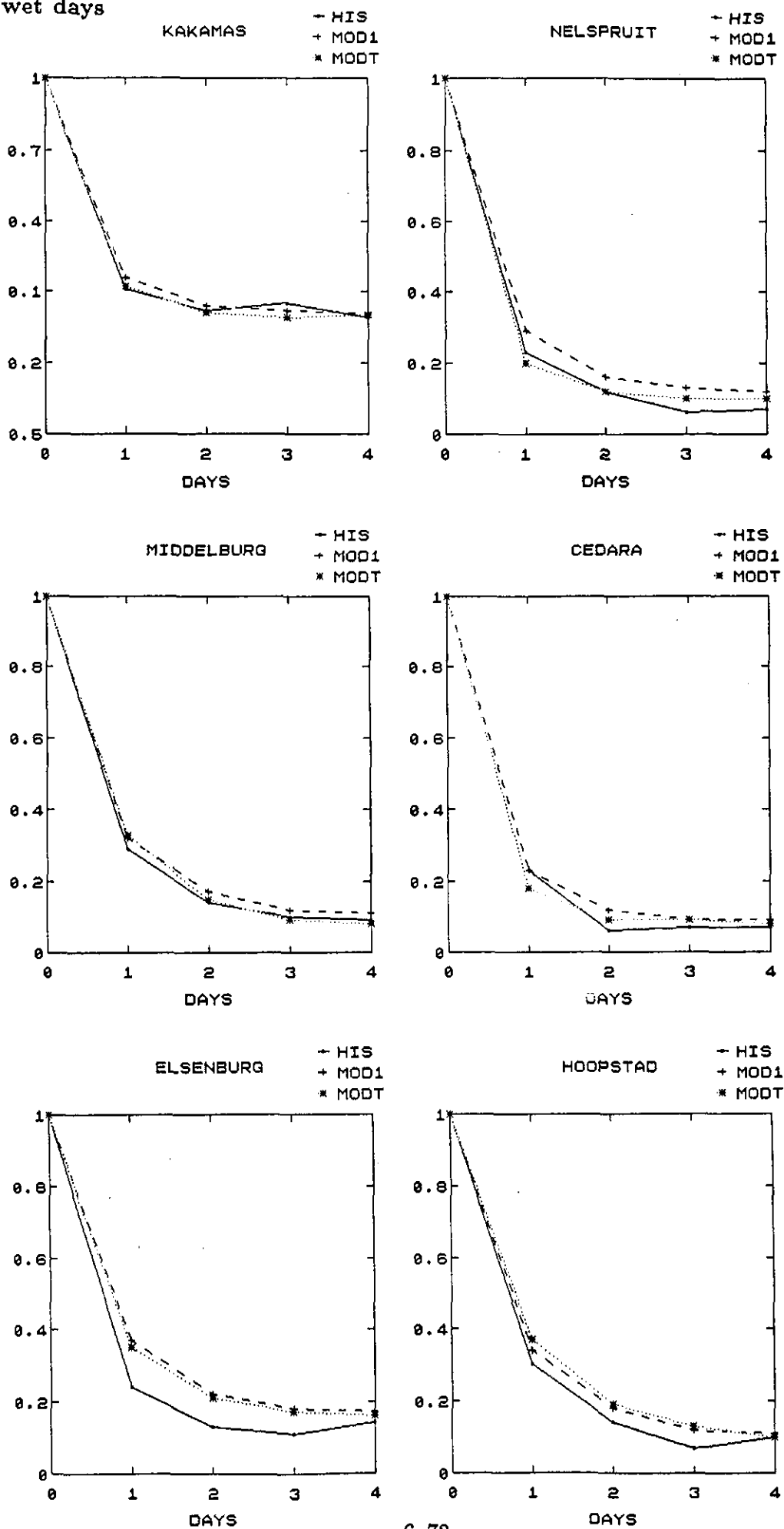


FIGURE 6.57 Autocorrelation coefficients for maximum temperature for wet days



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**FIGURE 6.58** Autocorrelation coefficients for minimum temperature for  
 wet days

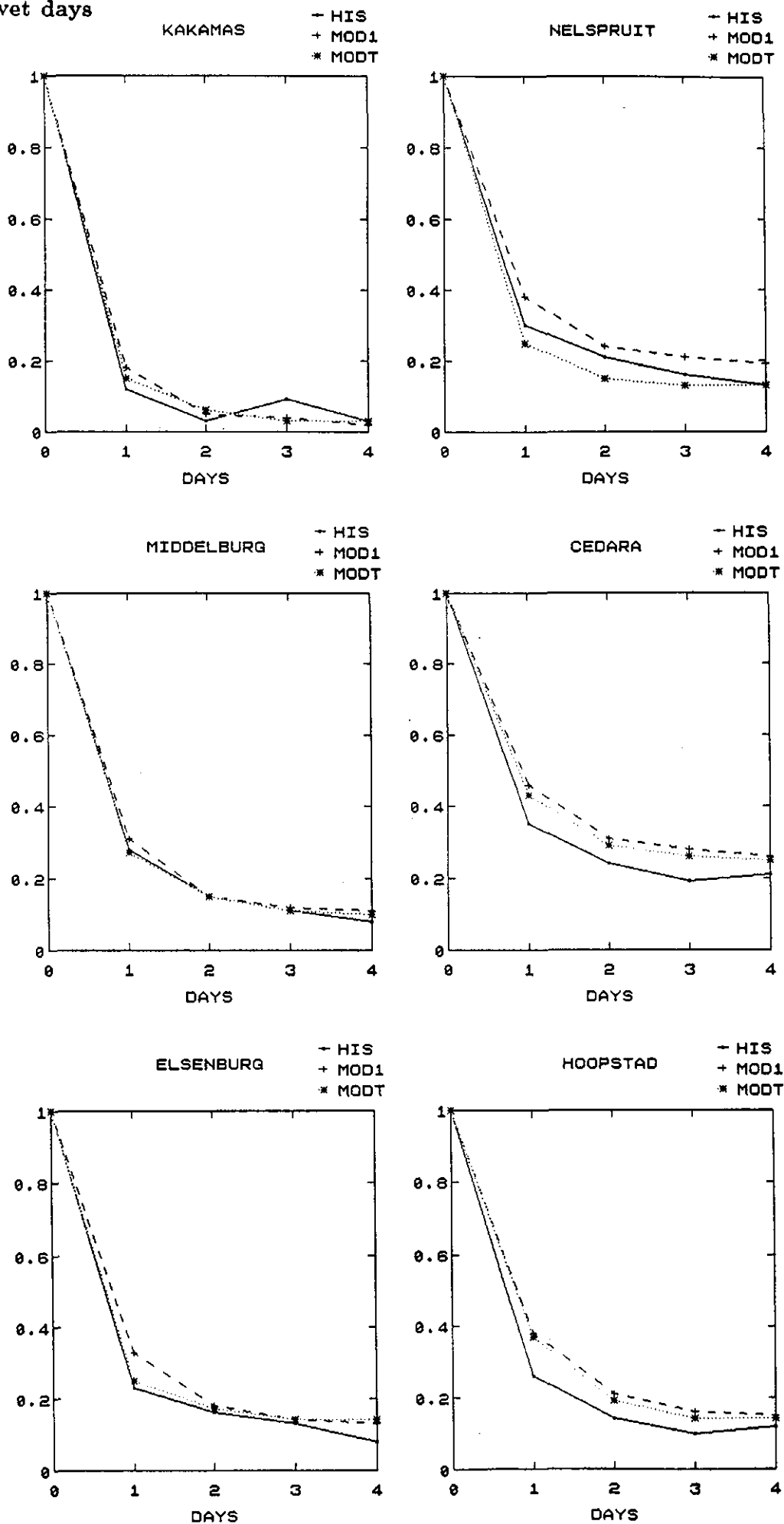


FIGURE 6.59 Autocorrelation coefficients for evaporation for wet days

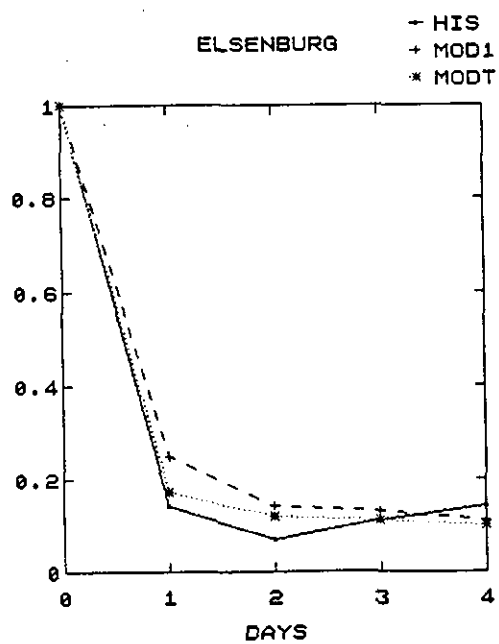
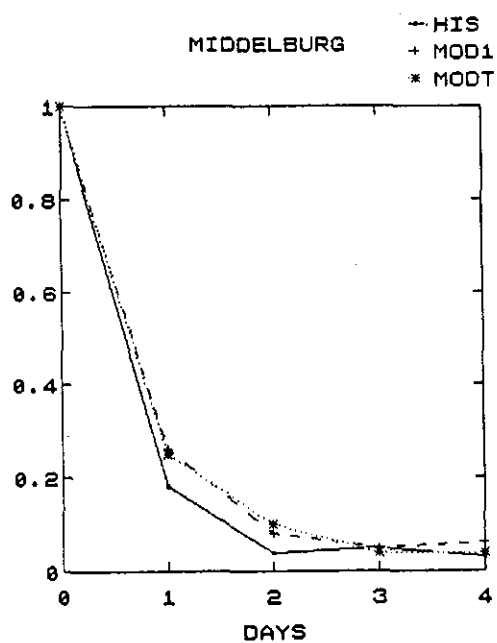
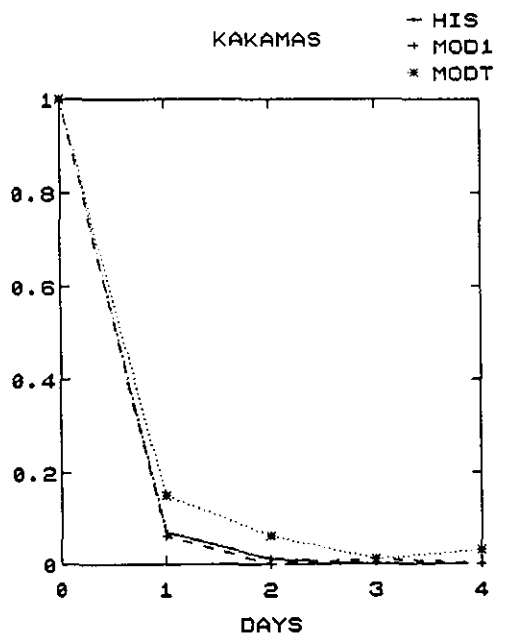


FIGURE 6.60 Autocorrelation coefficients for sunshine duration for wet days

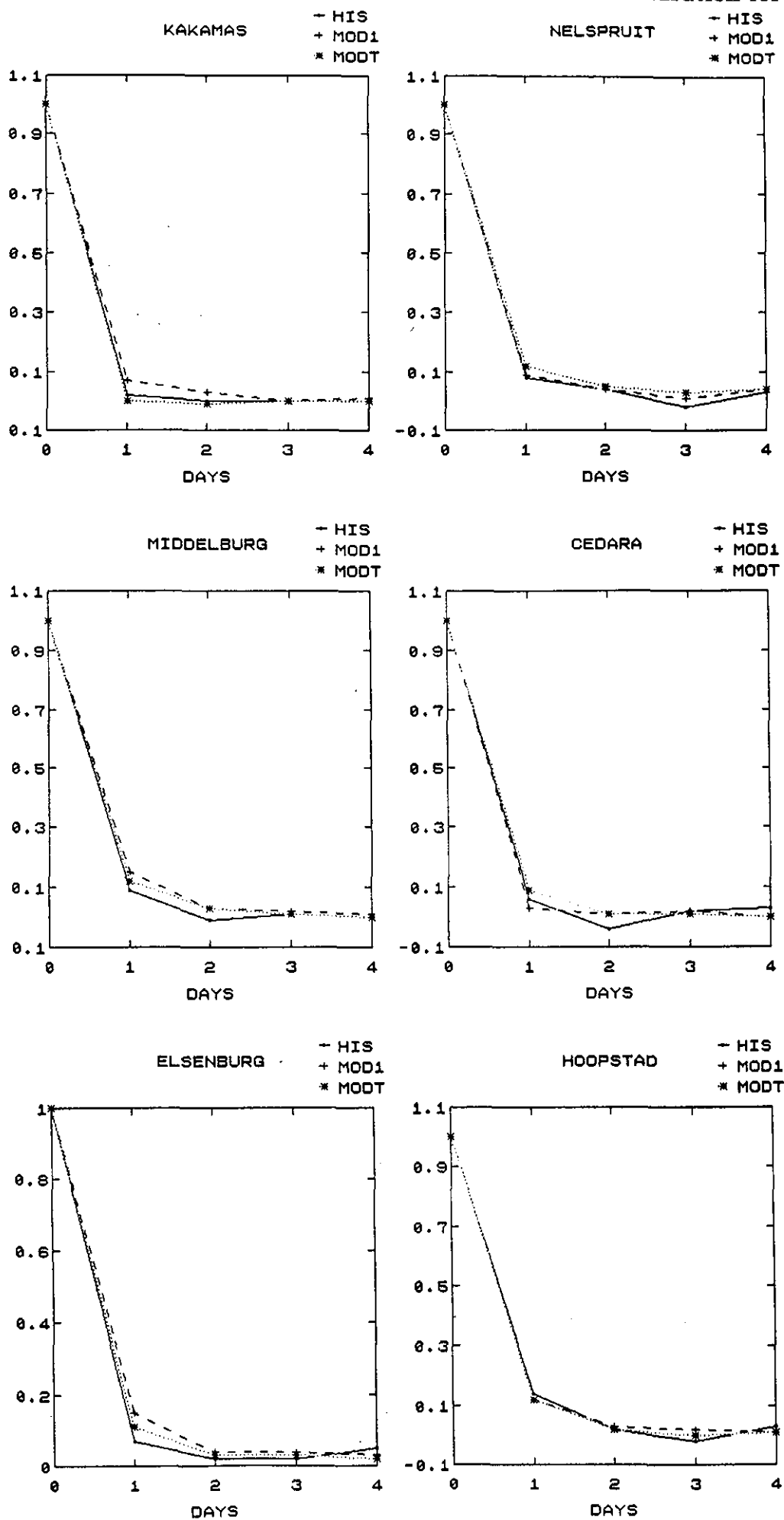


FIGURE 6.61 Autocorrelation coefficients for wind run for wet days

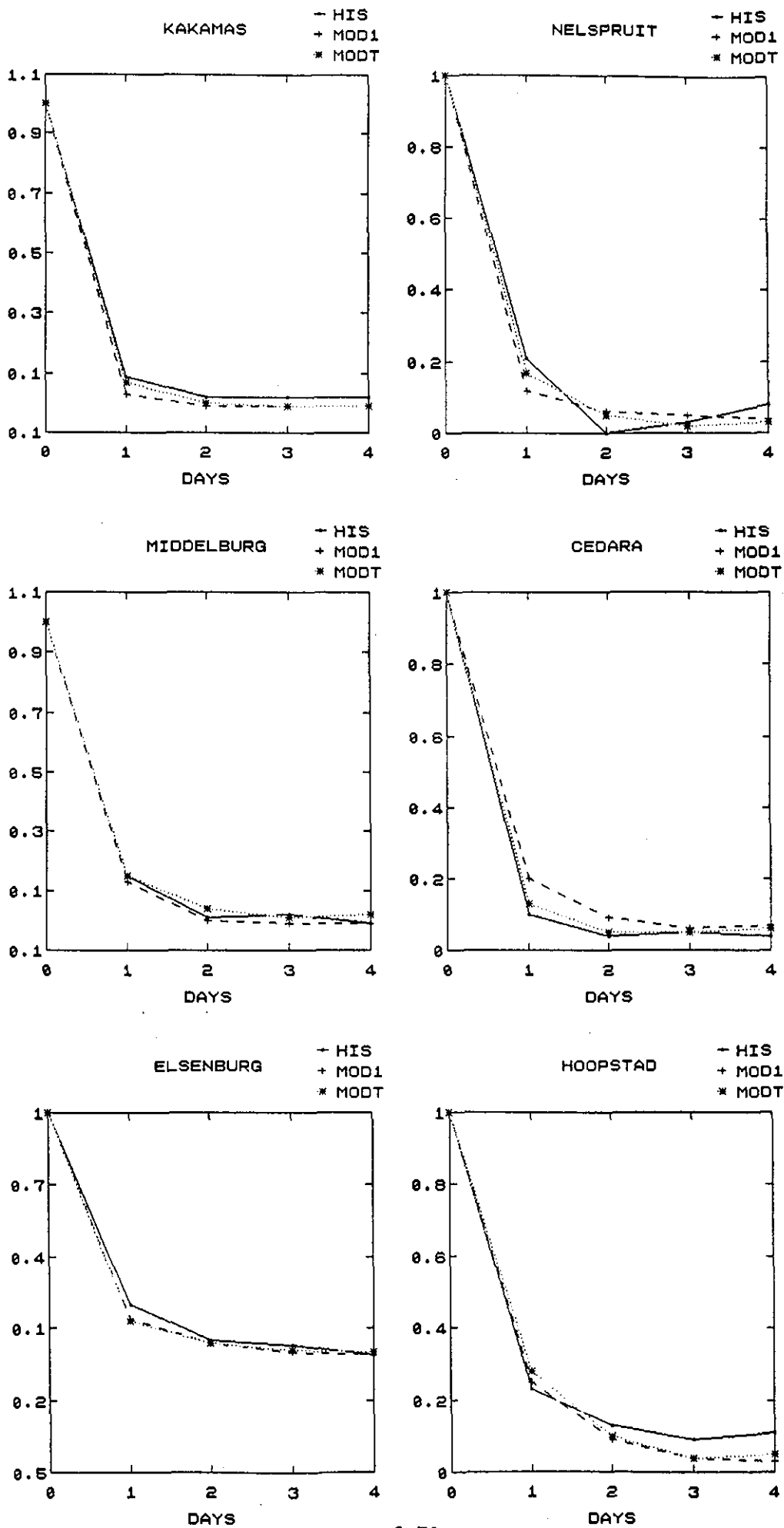


FIGURE 6.62 Autocorrelation coefficients for maximum humidity for wet days

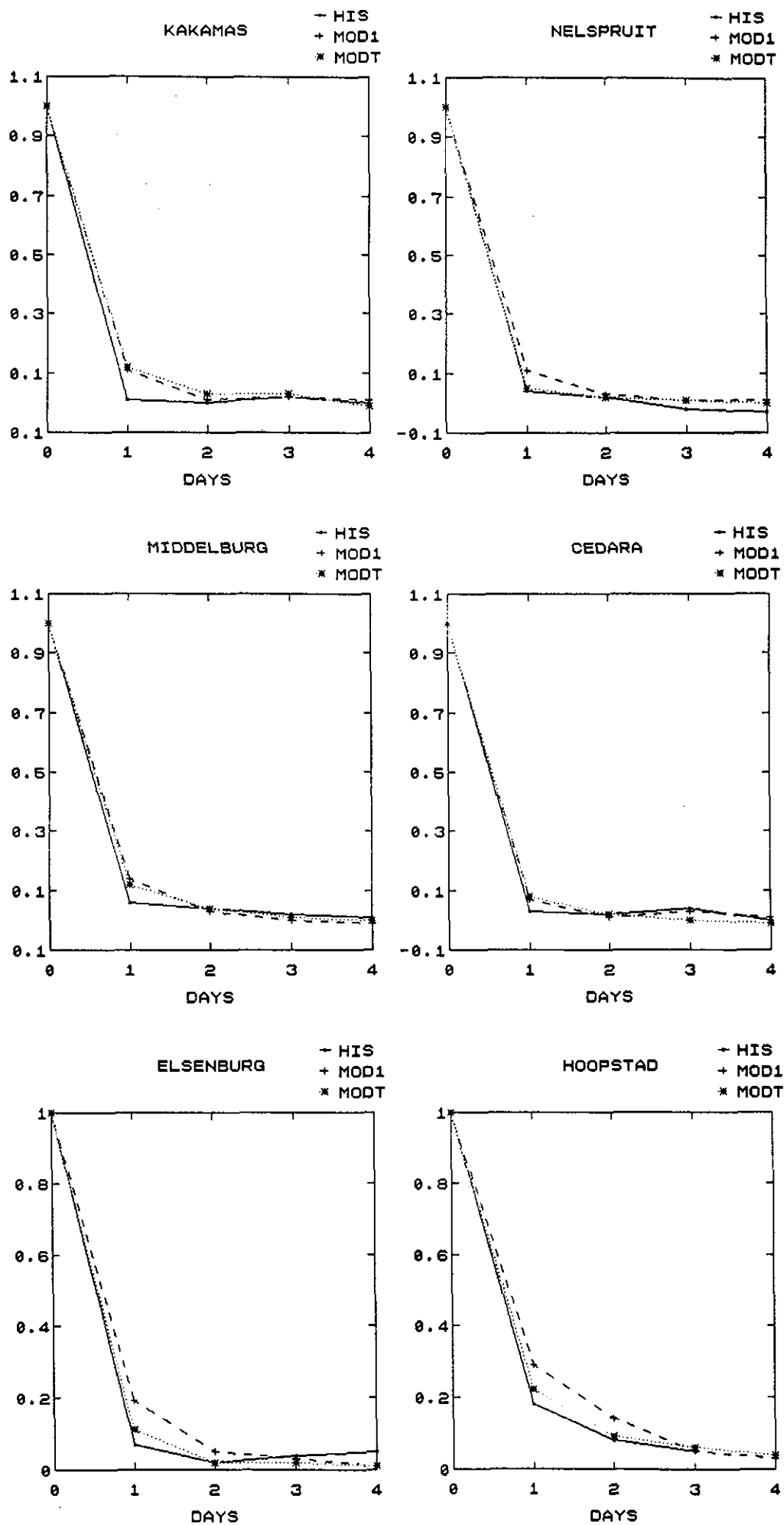


FIGURE 6.63 Autocorrelation coefficients for minimum humidity for wet days

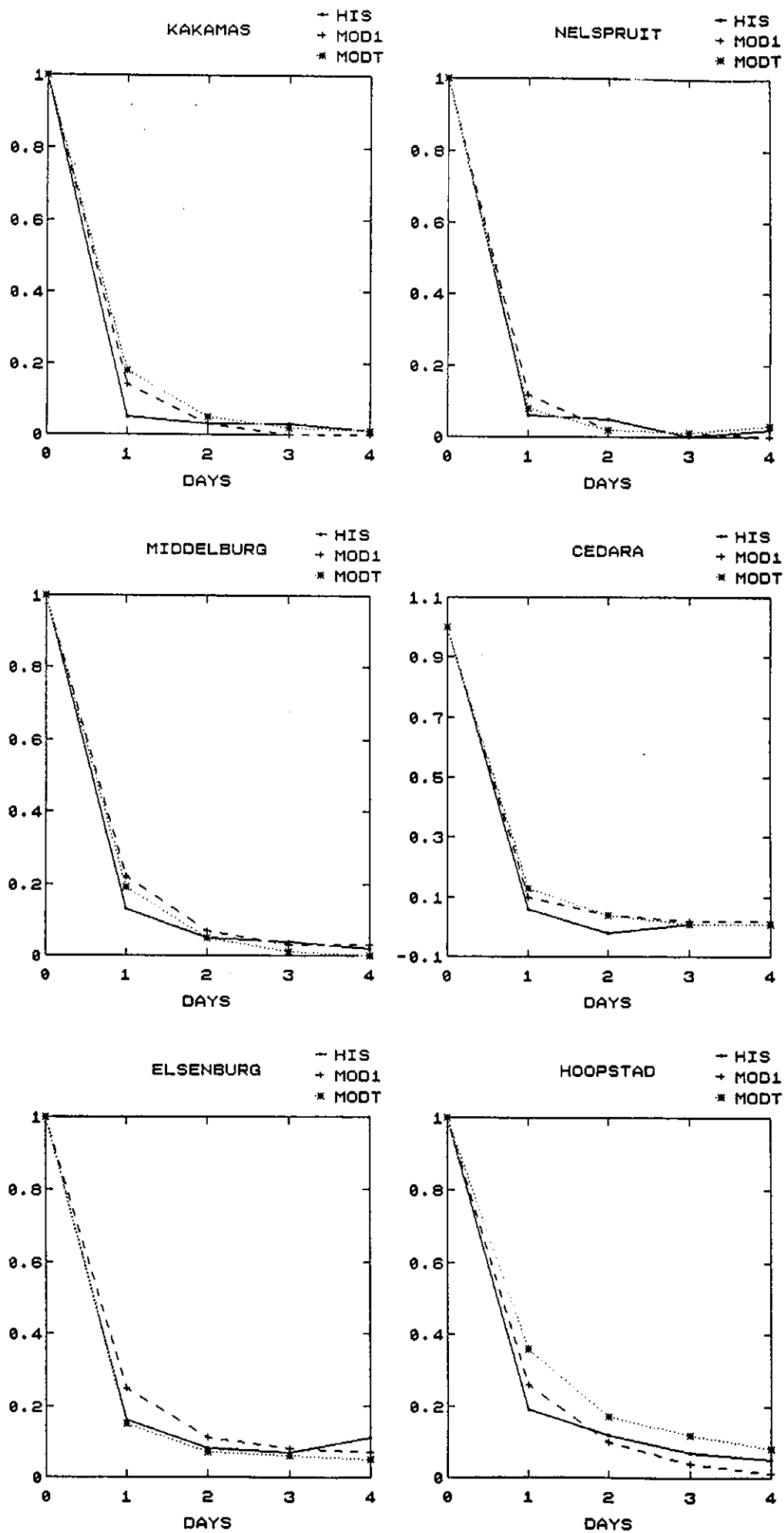


FIGURE 6.64 Autocorrelation coefficients for maximum temperature for dry days

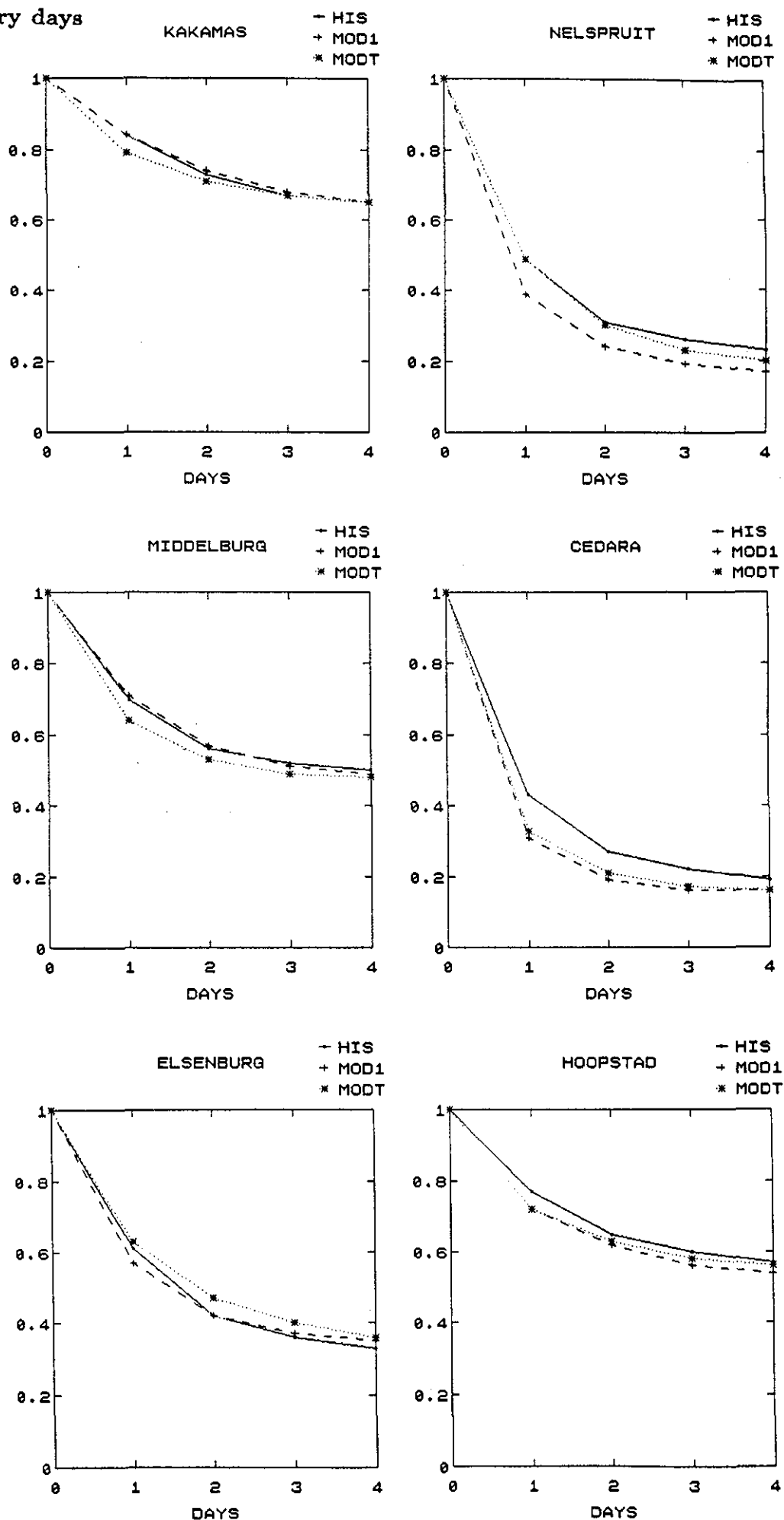


FIGURE 6.65 Autocorrelation coefficients for minimum temperature for dry days

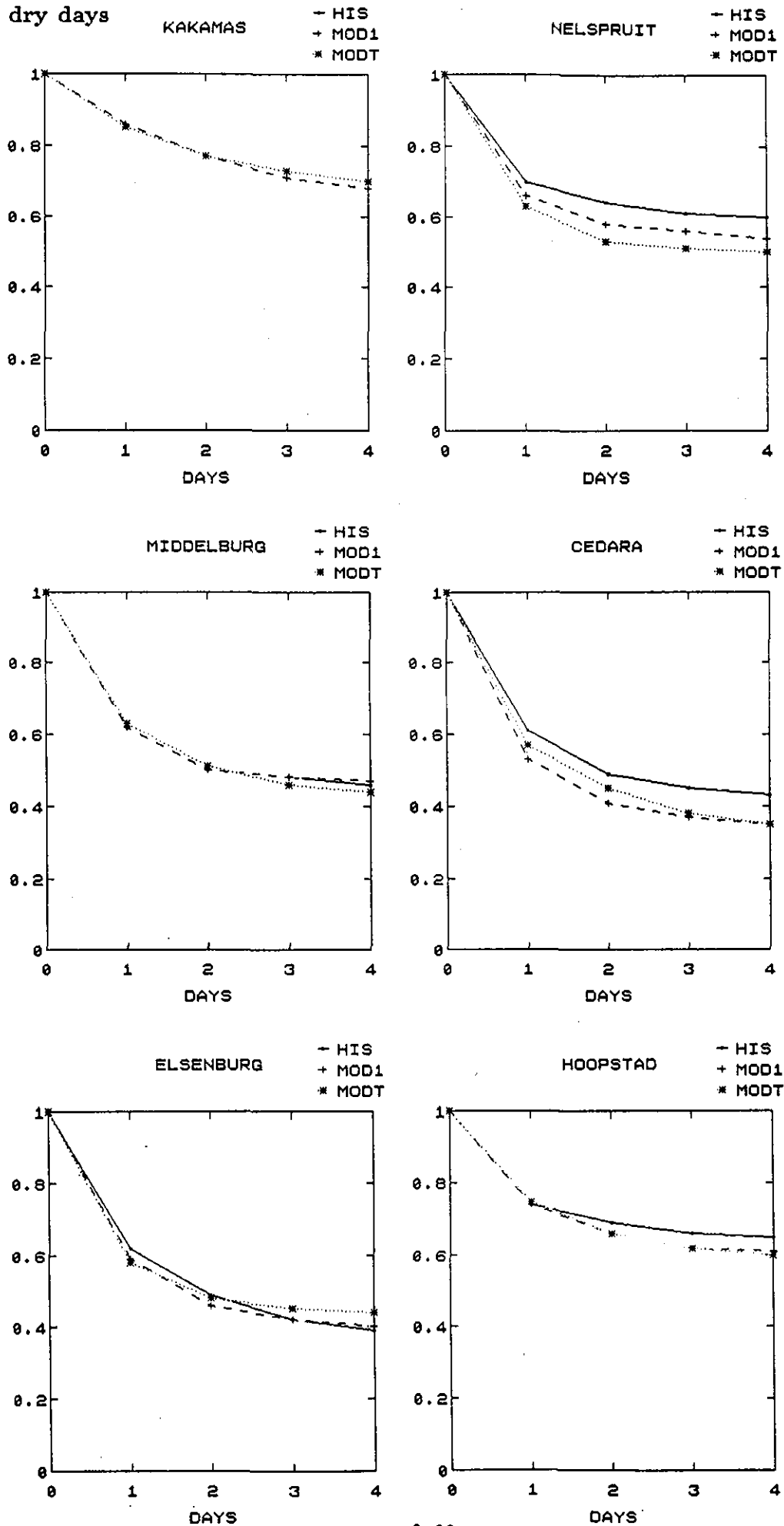


FIGURE 6.66 Autocorrelation coefficients for evaporation for dry days

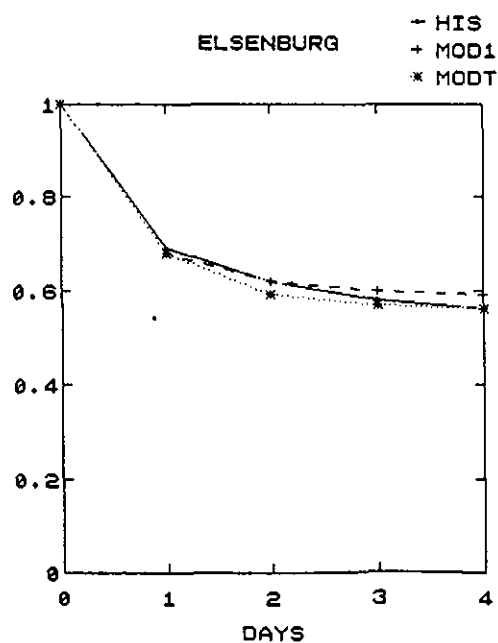
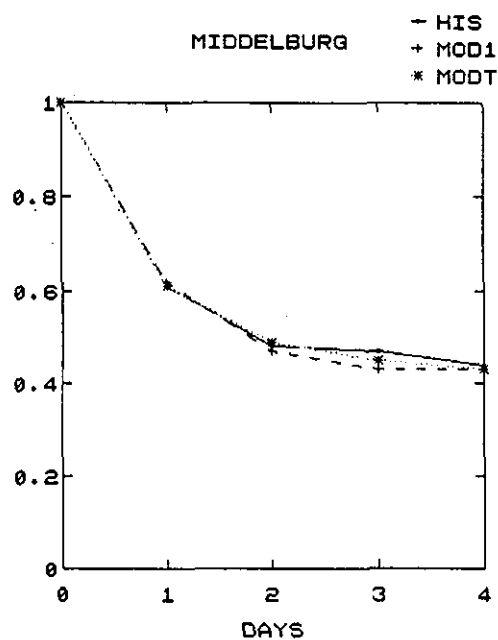
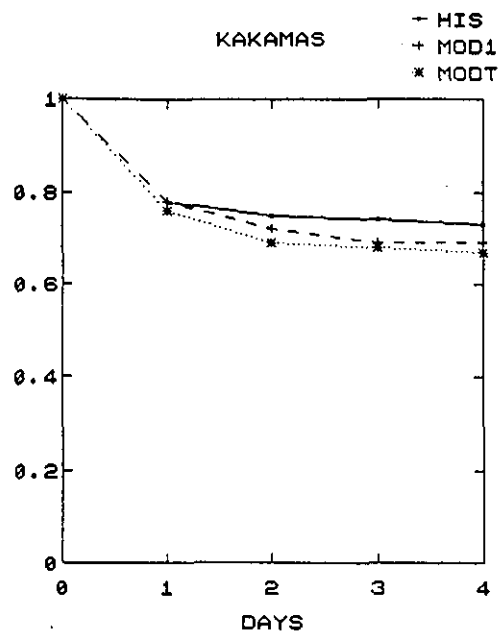


FIGURE 6.67 Autocorrelation coefficients for sunshine duration for dry days

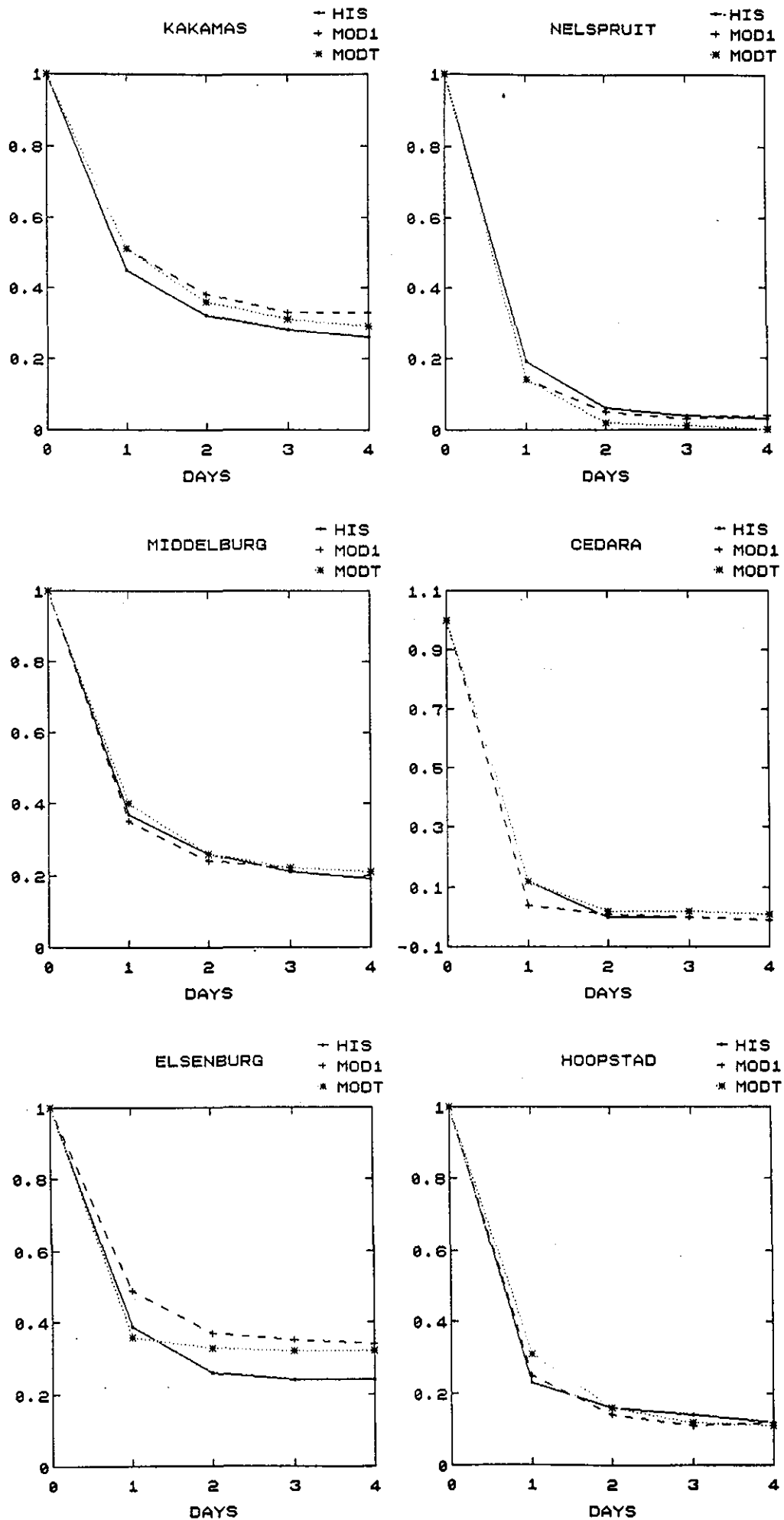


FIGURE 6.68 Autocorrelation coefficients for wind run for dry days

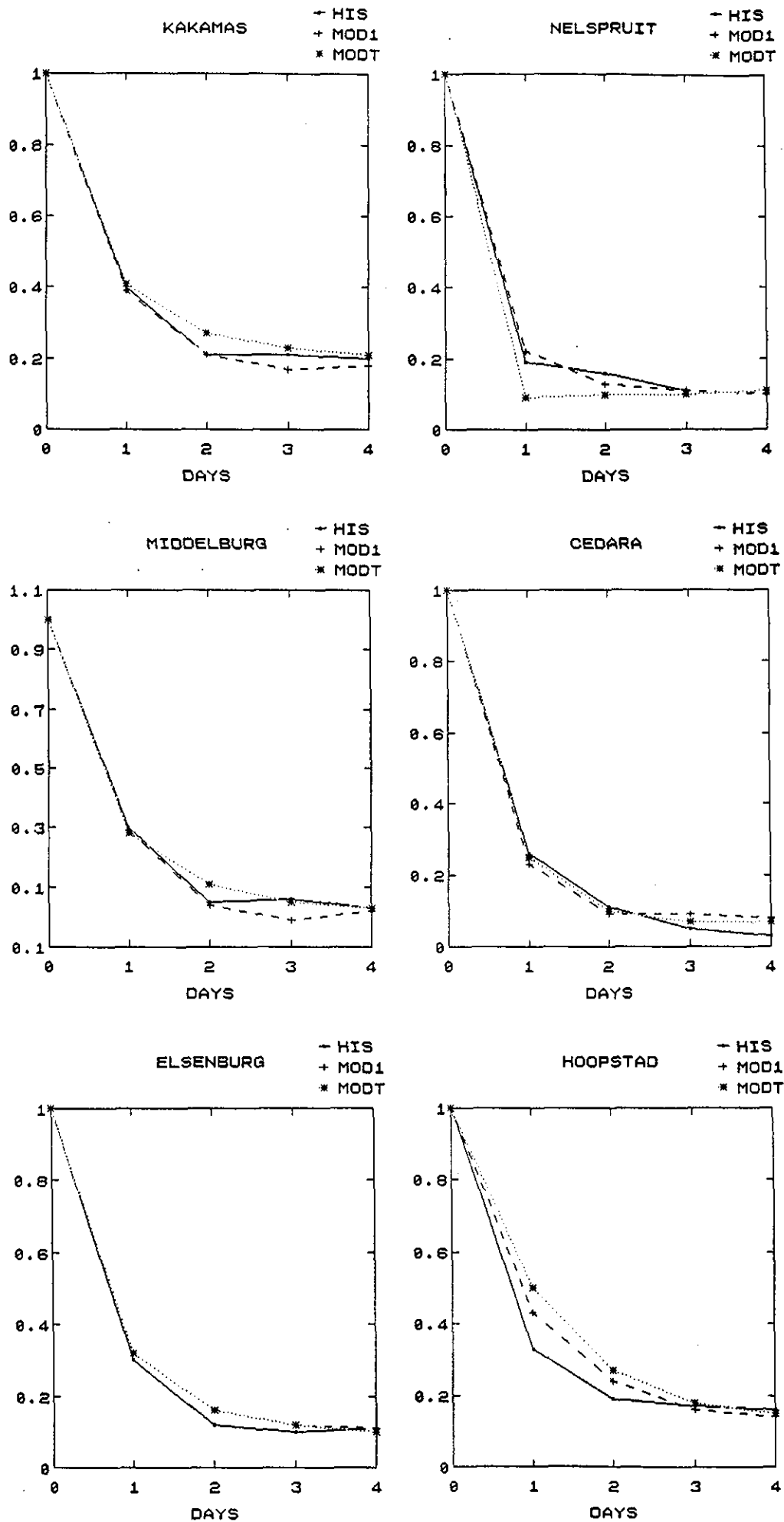


FIGURE 6.69 Autocorrelation coefficients for maximum humidity for dry days

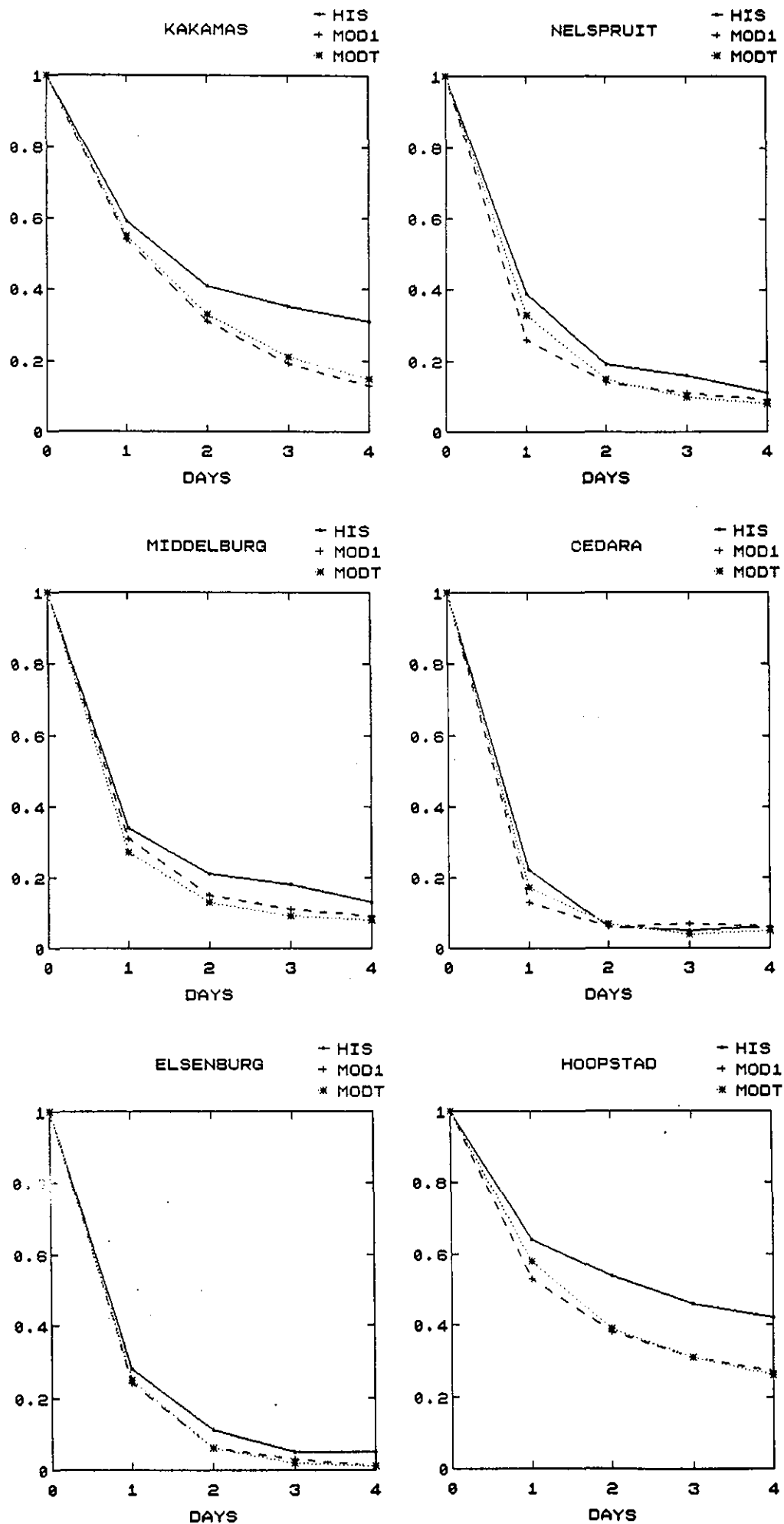
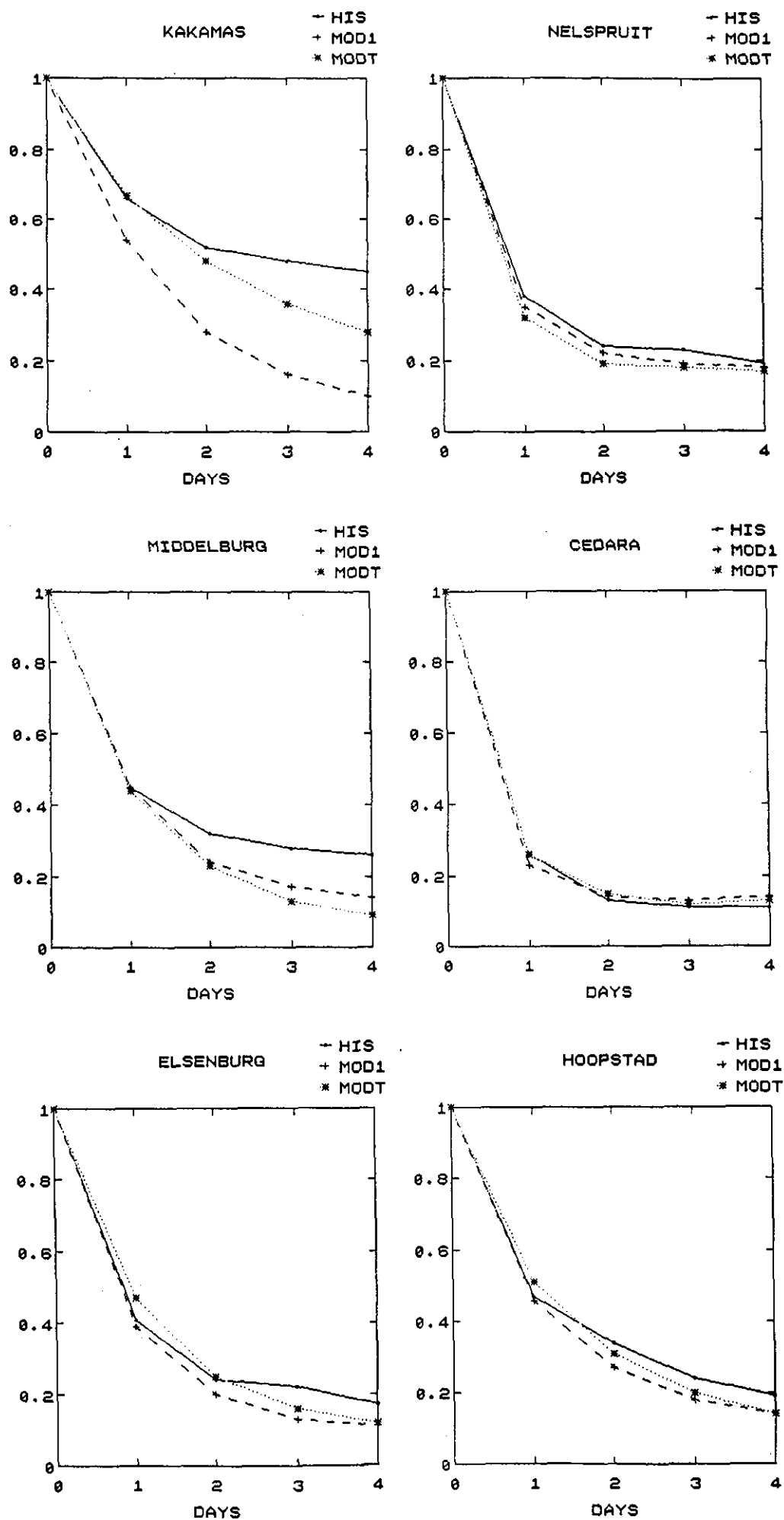


FIGURE 6.70 Autocorrelation coefficients for minimum humidity for dry days



CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.71** Cross-correlation coefficients for maximum temperature and  
 minimum temperature

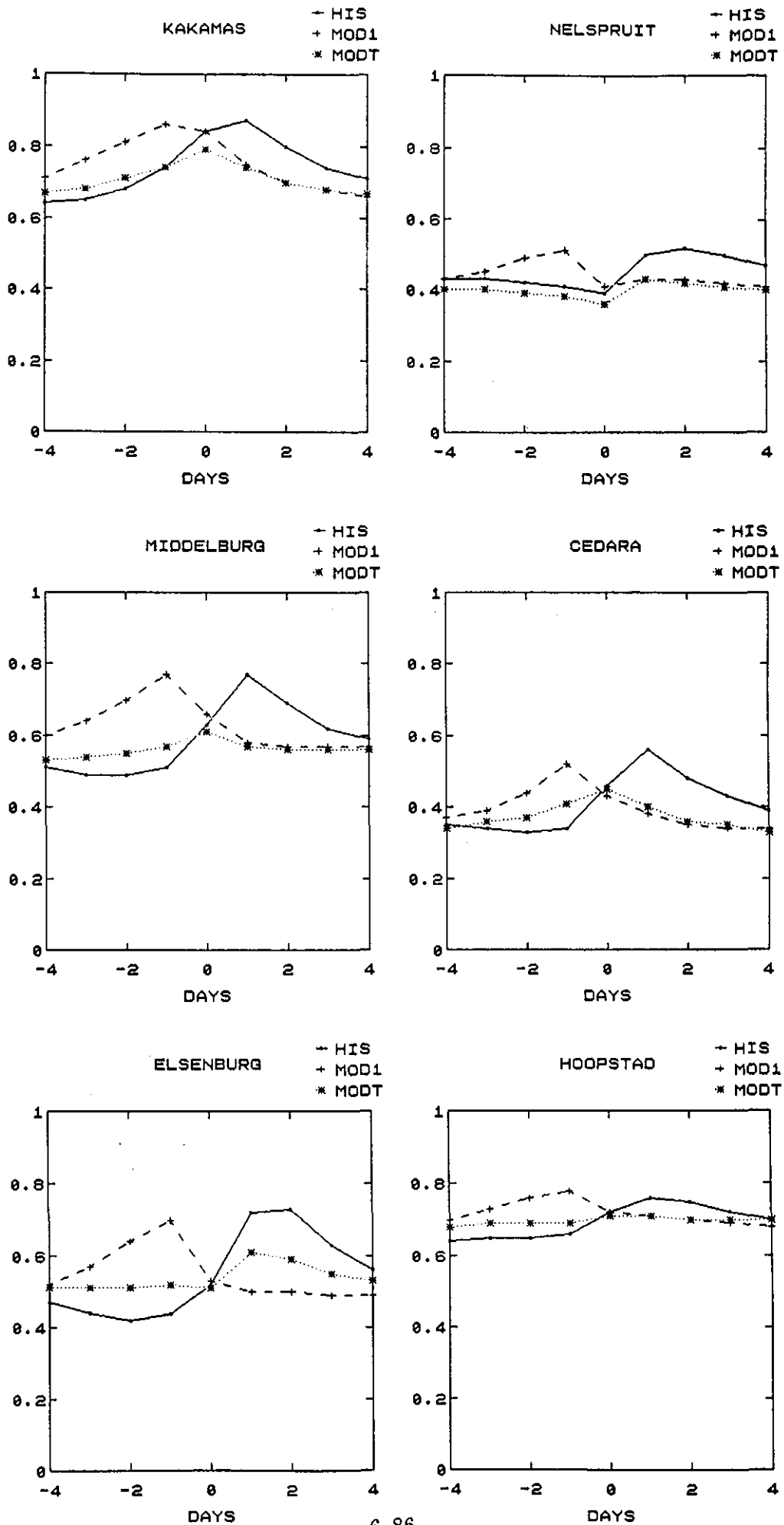
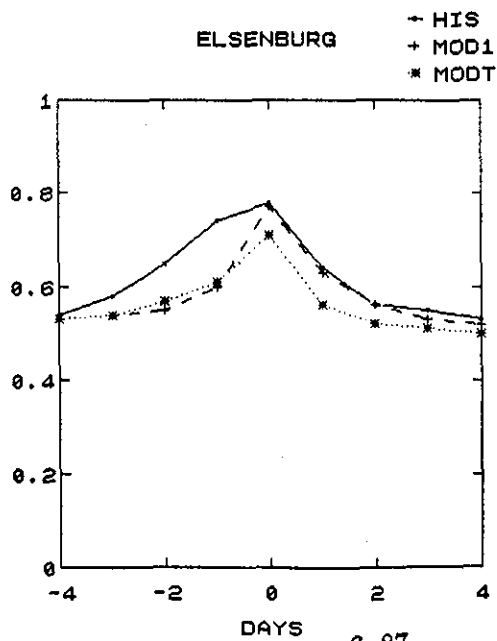
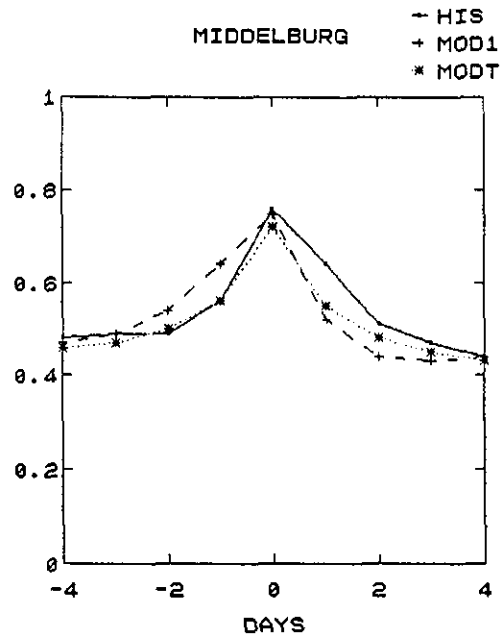
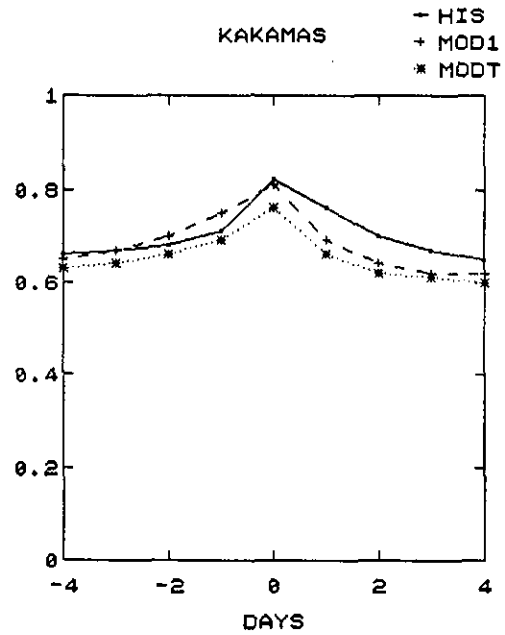
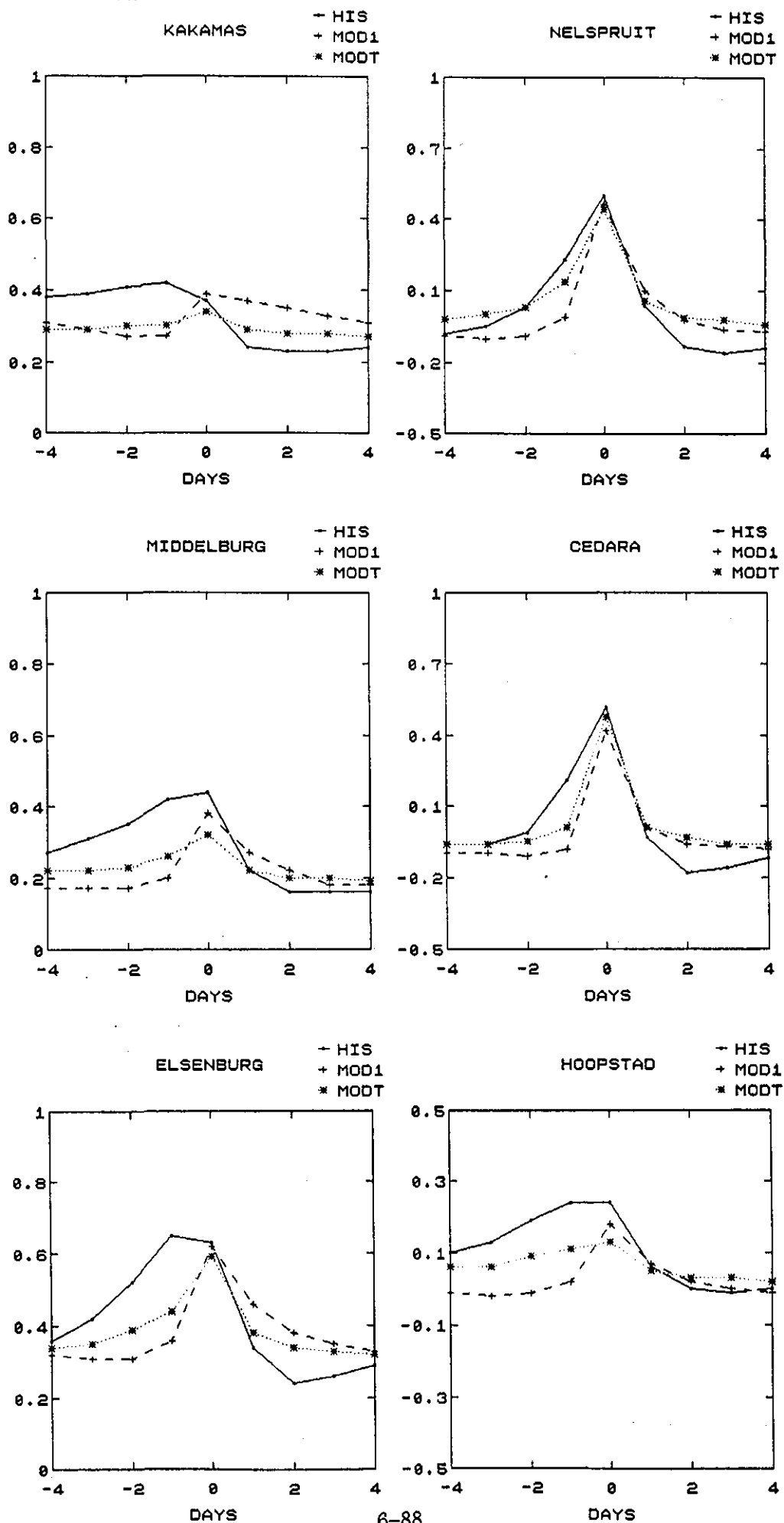


FIGURE 6.72 Cross-correlation coefficients for maximum temperature and evaporation



CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.73** Cross-correlation coefficients for maximum temperature and  
 sunshine duration



CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.74** Cross-correlation coefficients for maximum temperature and  
 wind run

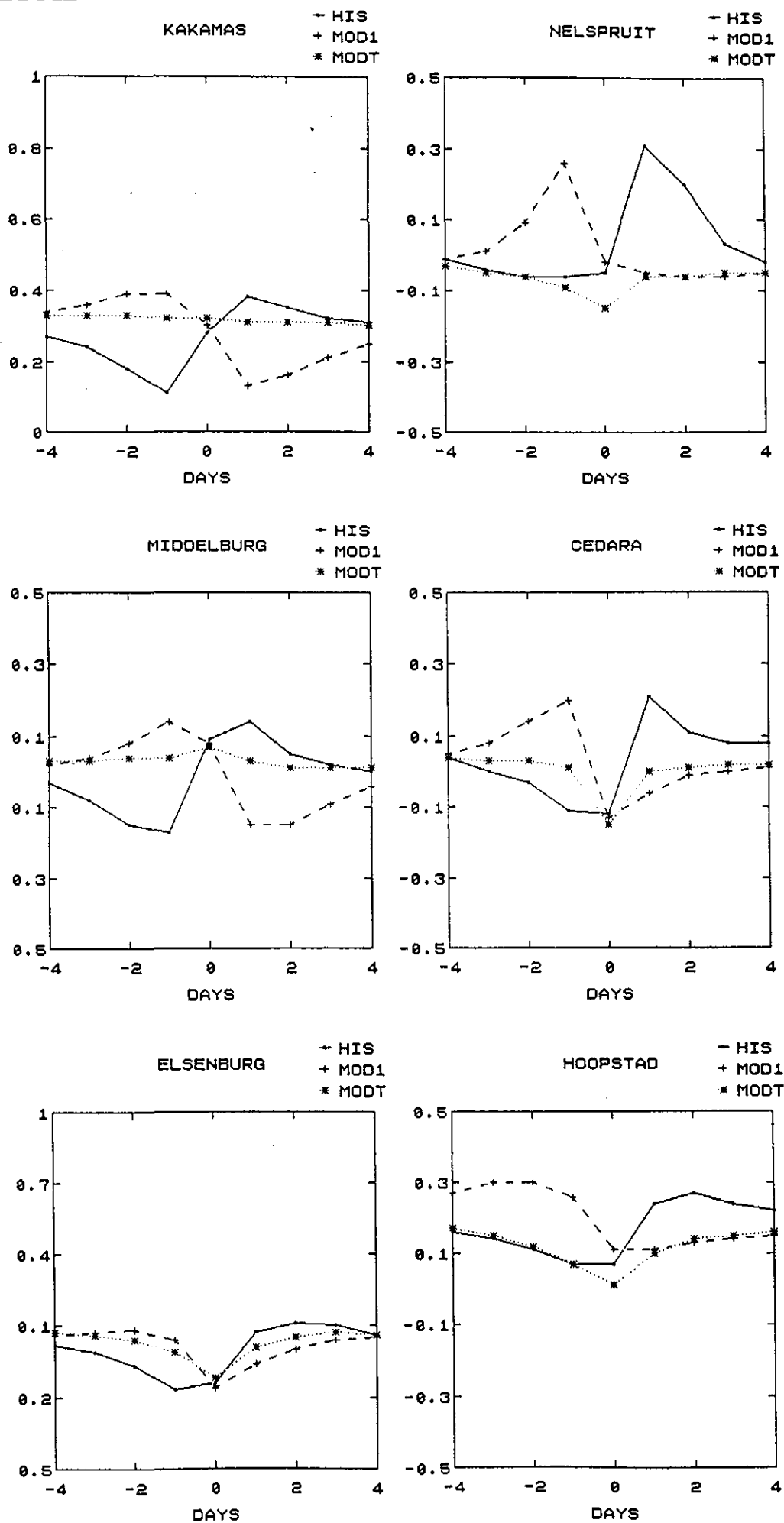
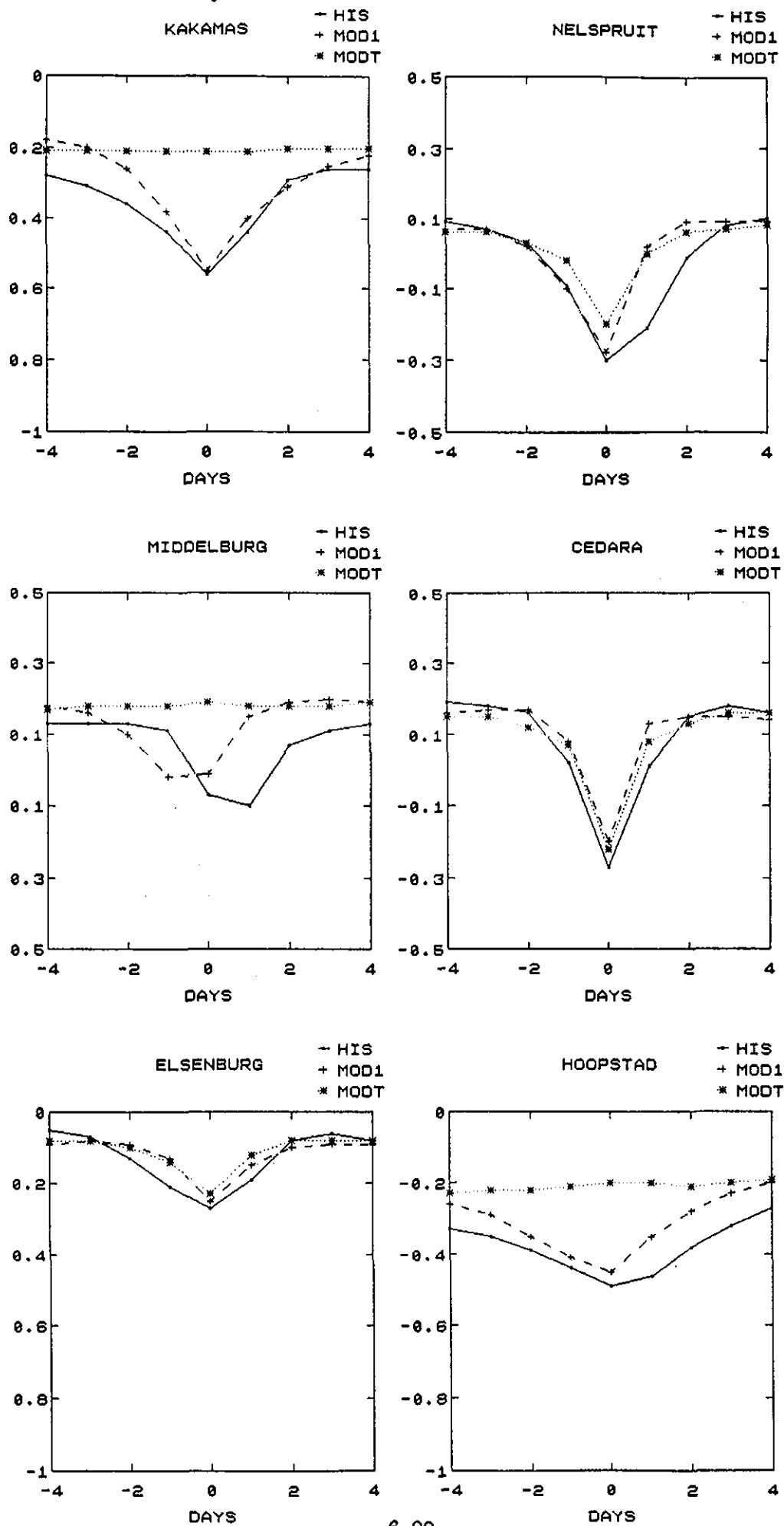
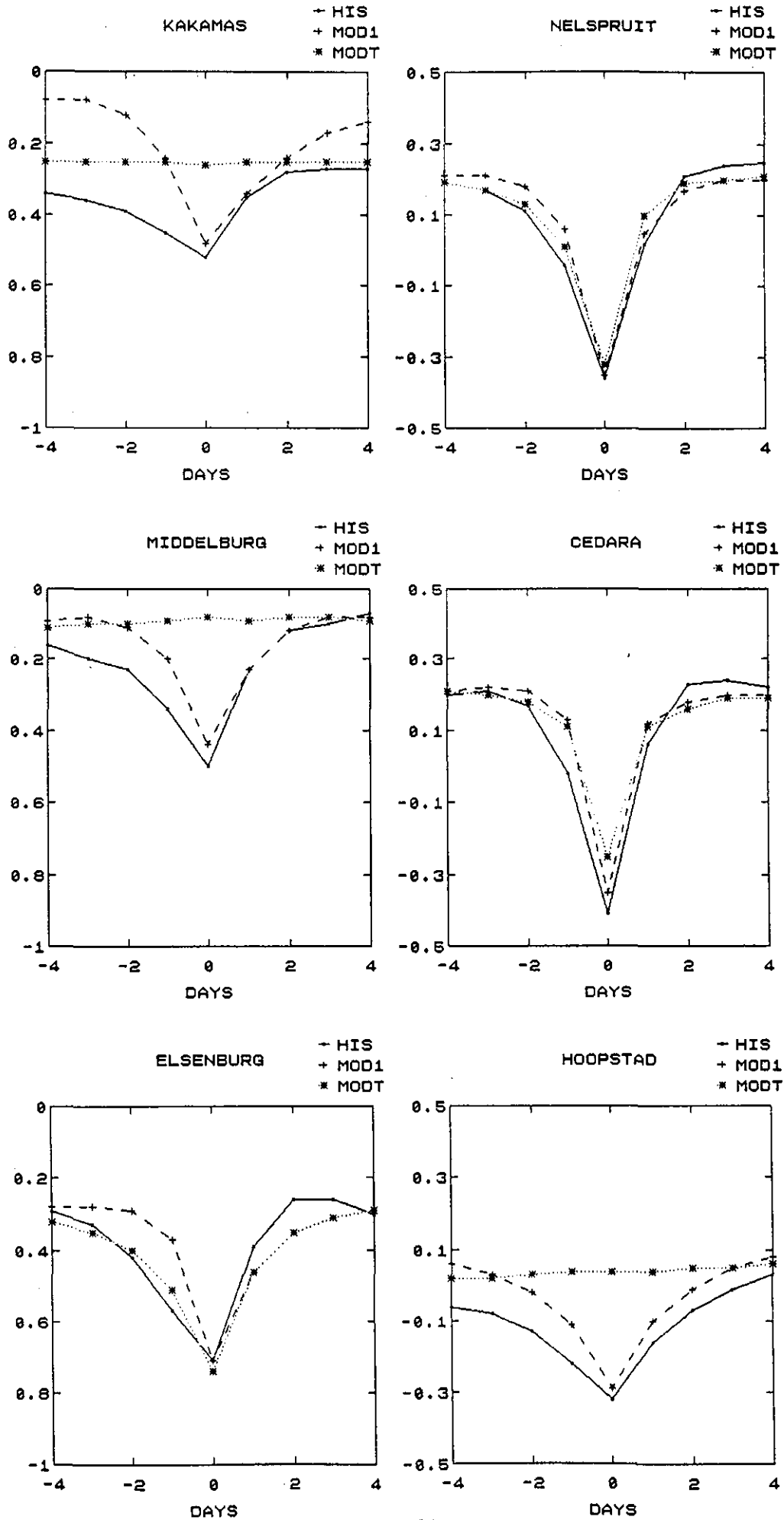


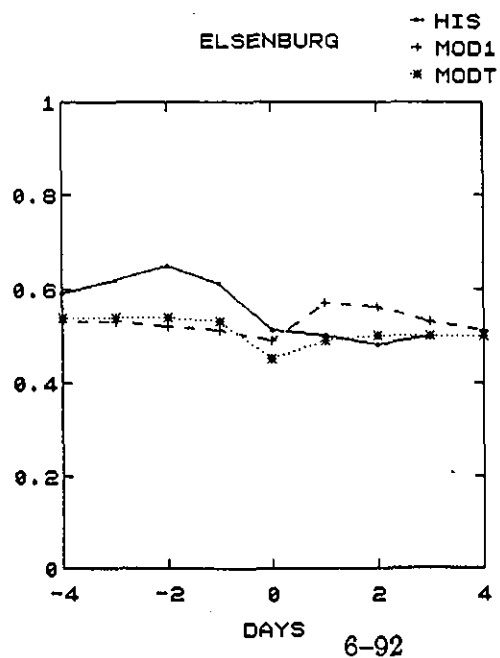
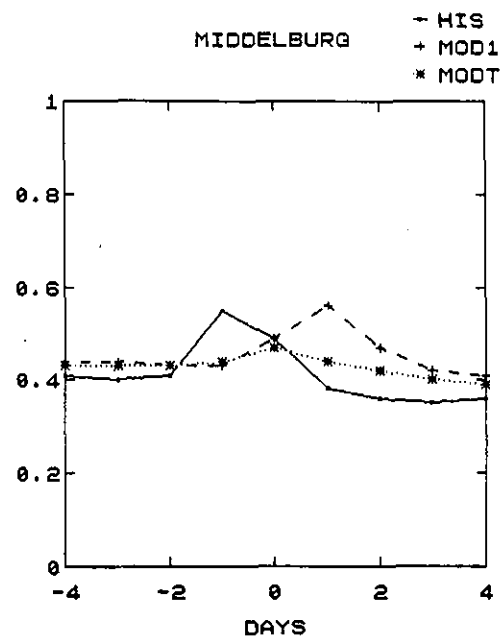
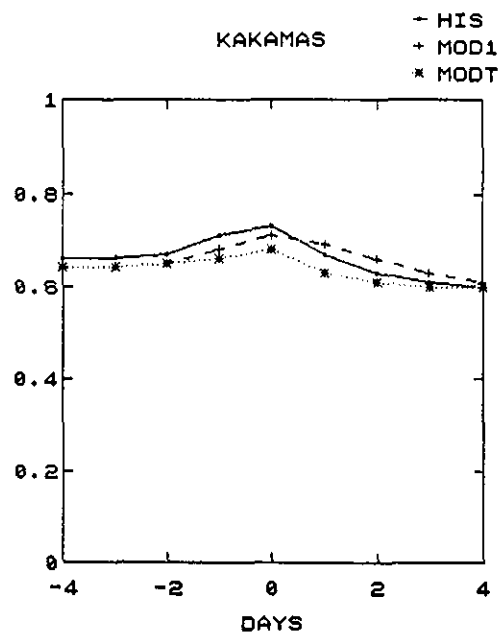
FIGURE 6.75 Cross-correlation coefficients for maximum temperature and maximum humidity



CHAPTER 6  
**FIGURE 6.76** Cross-correlation coefficients for maximum temperature and minimum humidity

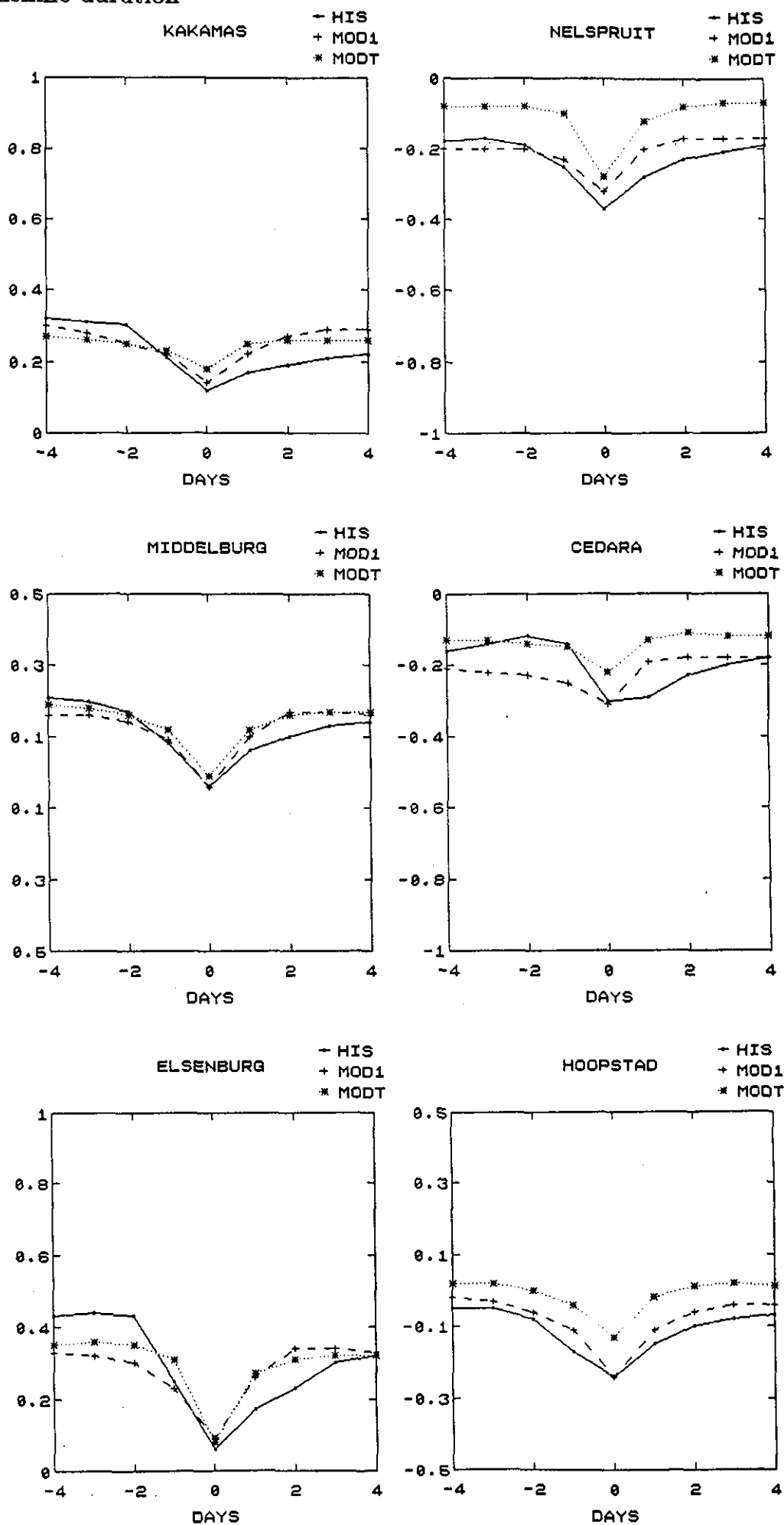
*Goodness of Fit*



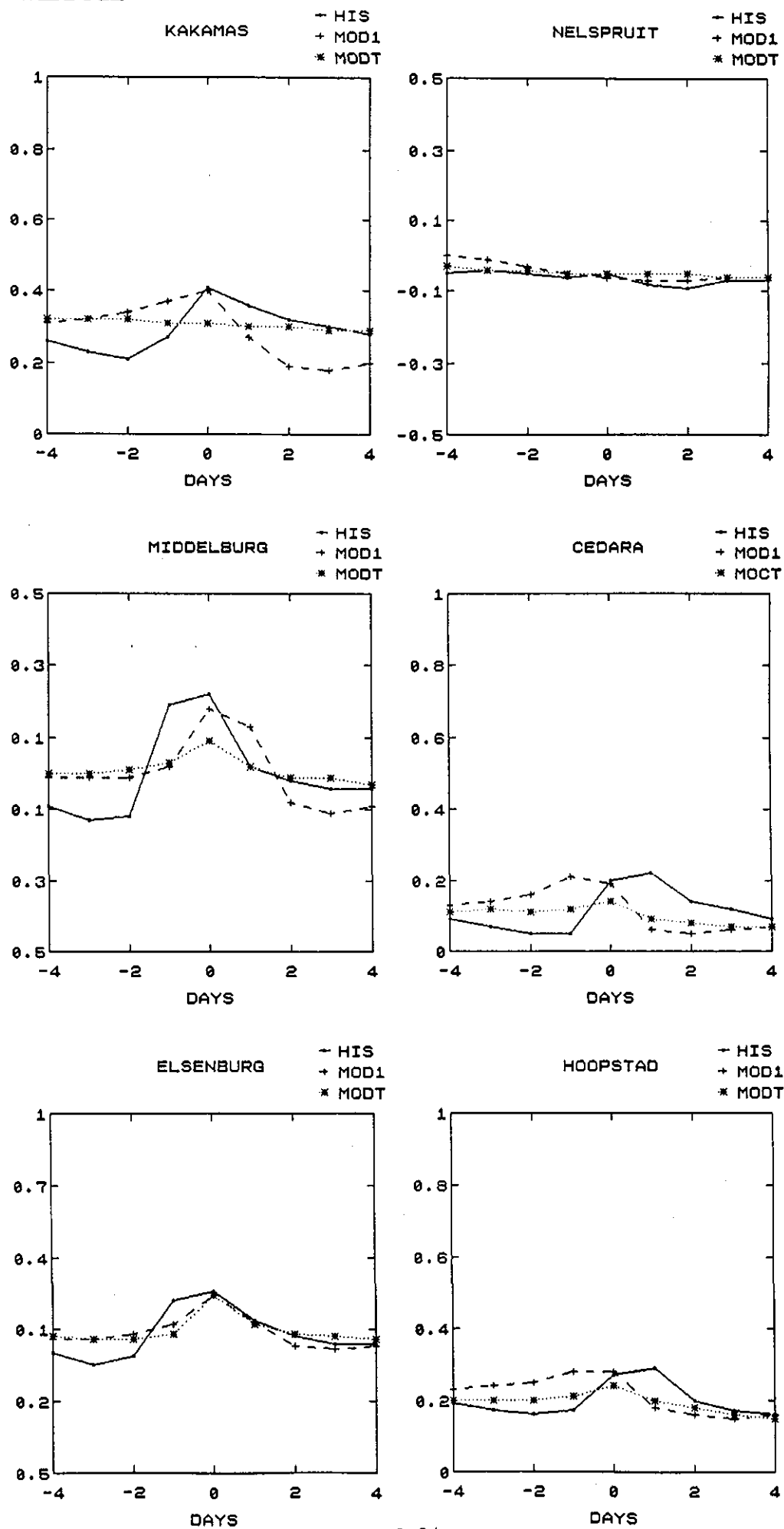


CHAPTER 6  
**FIGURE 6.78** Cross-correlation coefficients for minimum temperature and sunshine duration

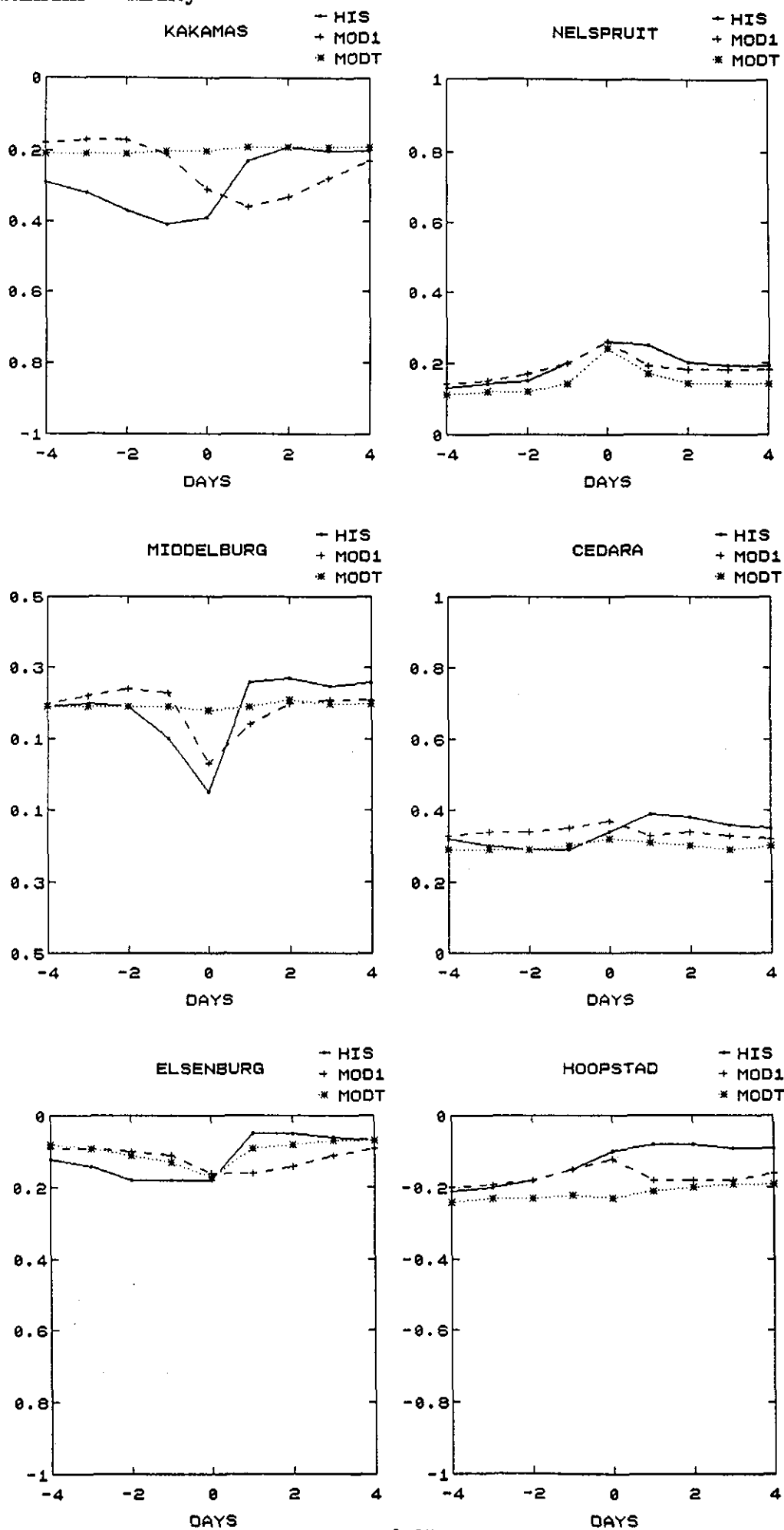
*Goodness of Fit*



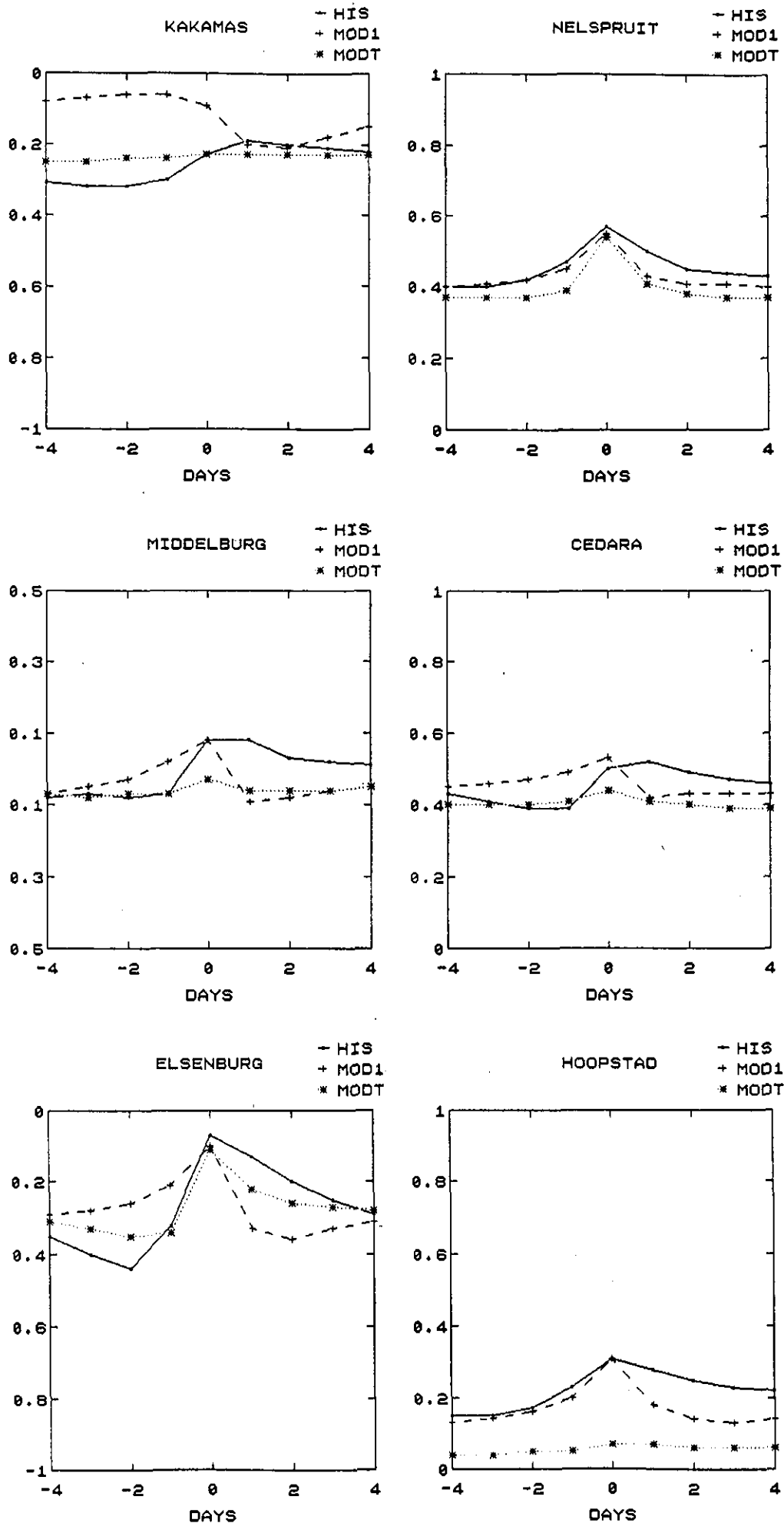
CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.79** Cross-correlation coefficients for minimum temperature and  
 wind run



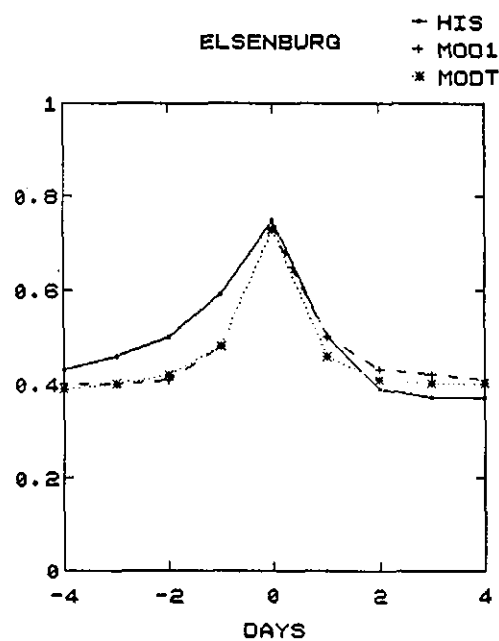
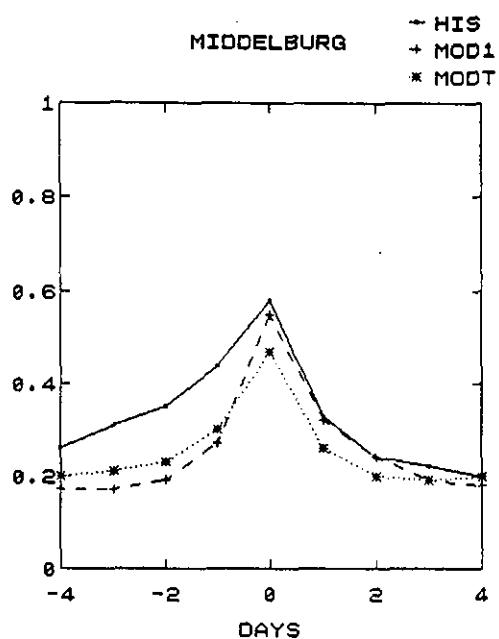
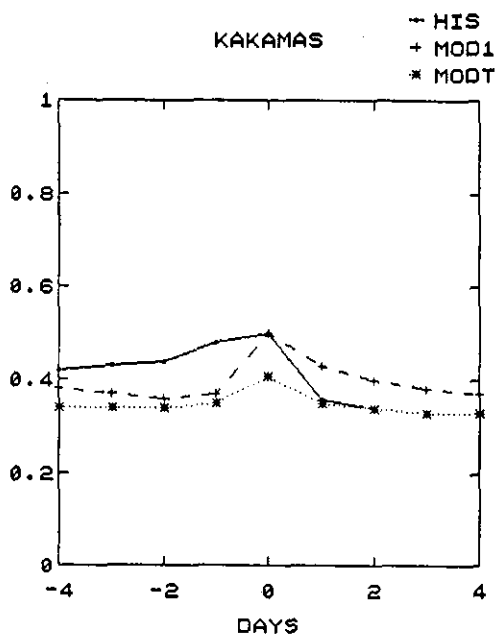
CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.80** Cross-correlation coefficients for minimum temperature and  
 maximum humidity



CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.81** Cross-correlation coefficients for minimum temperature and  
 minimum humidity



CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.82** Cross-correlation coefficients for evaporation and sunshine  
duration



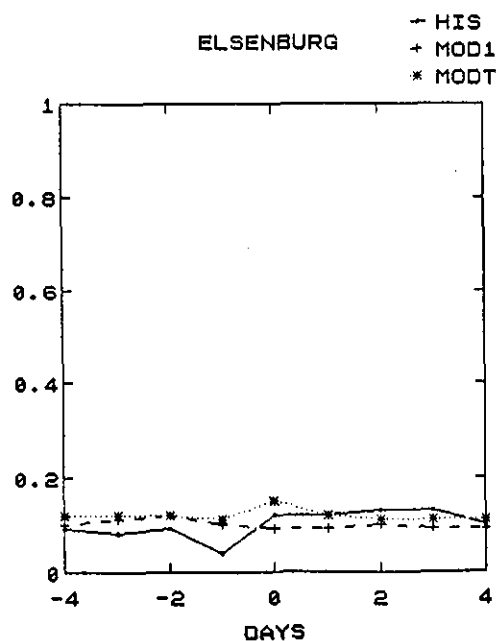
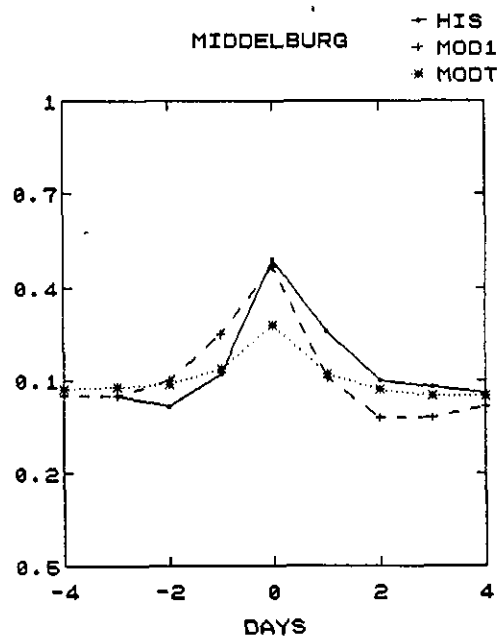
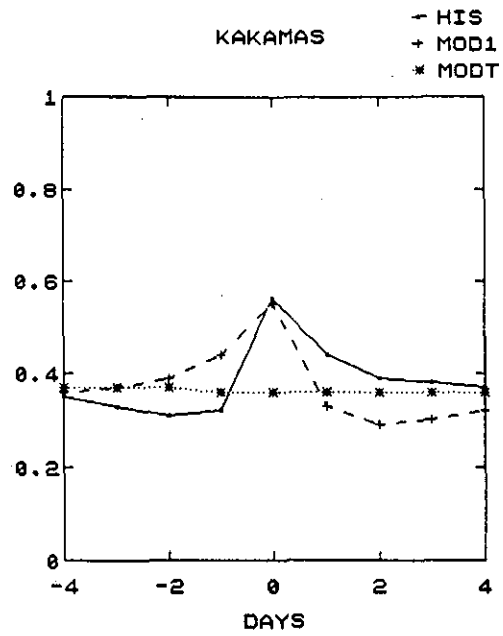
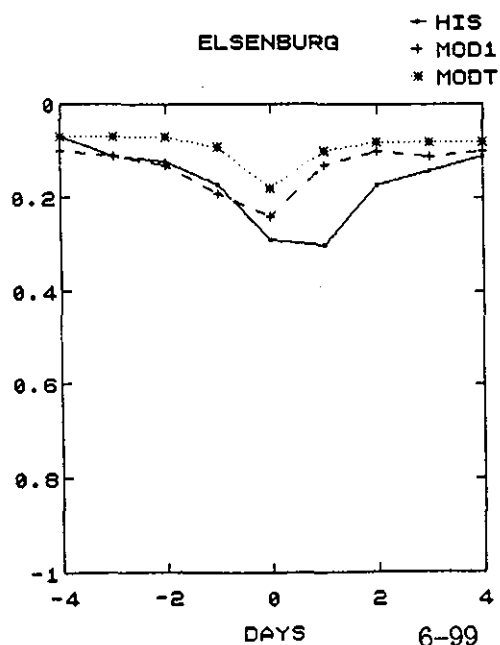
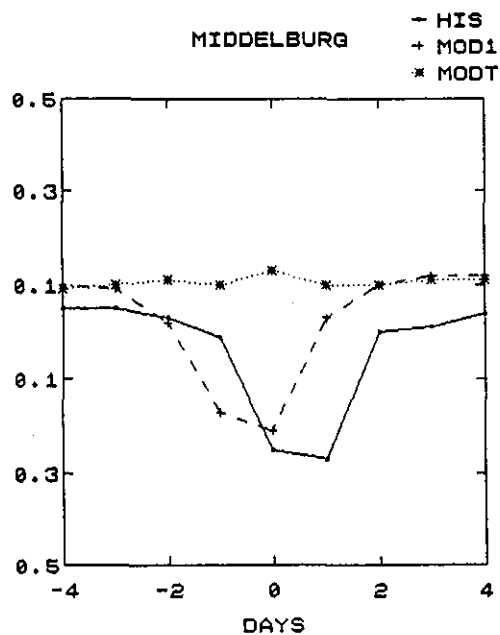
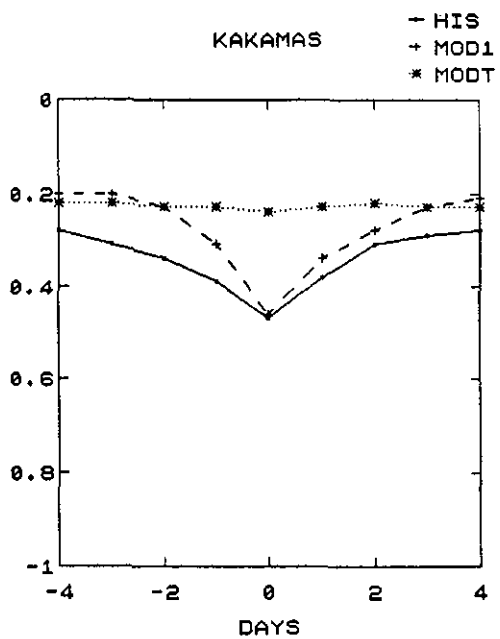
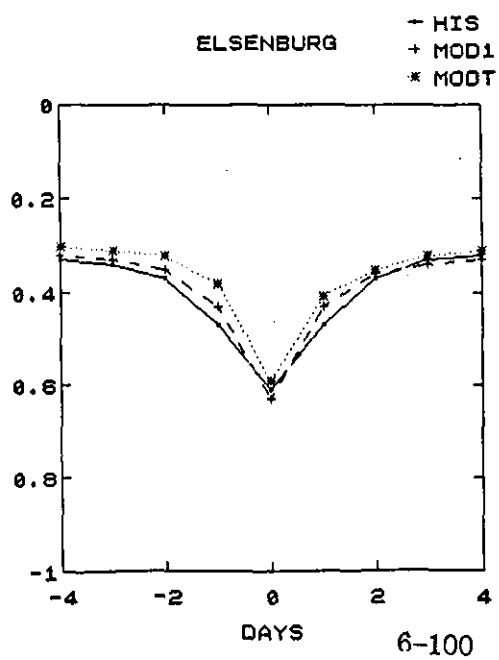
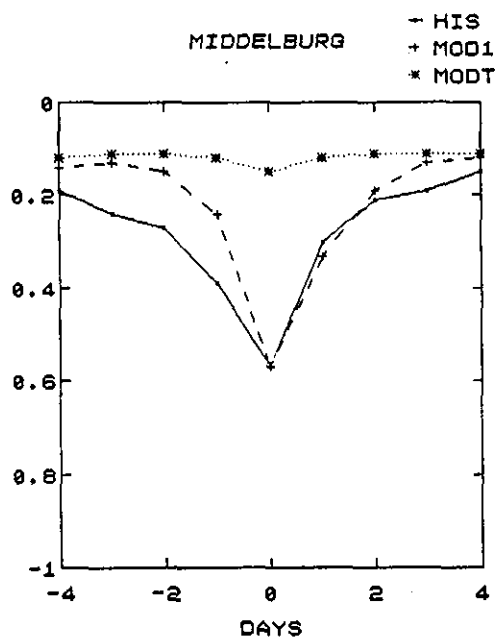
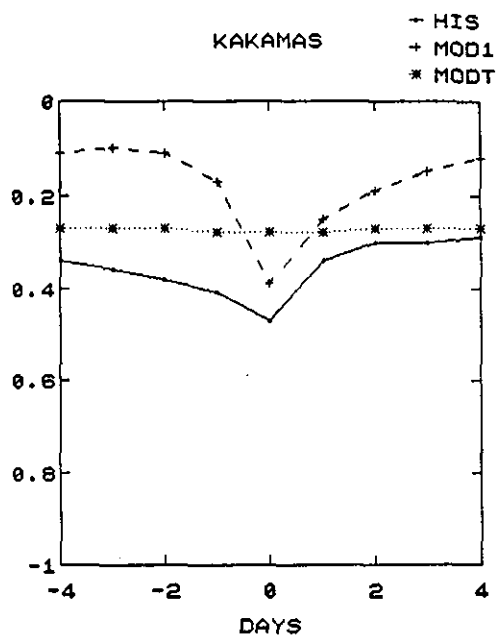


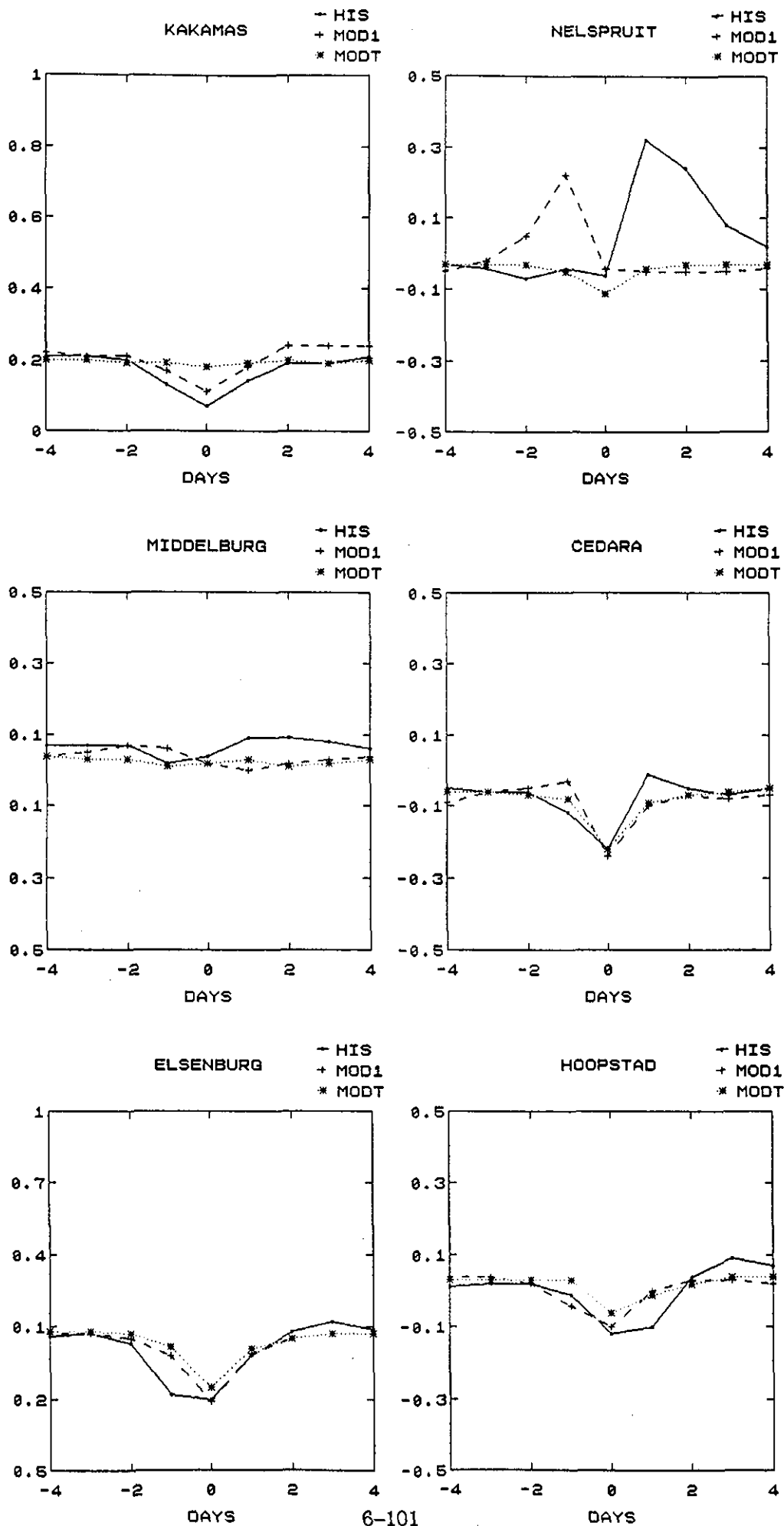
FIGURE 6.84 Cross-correlation coefficients for evaporation and maximum humidity



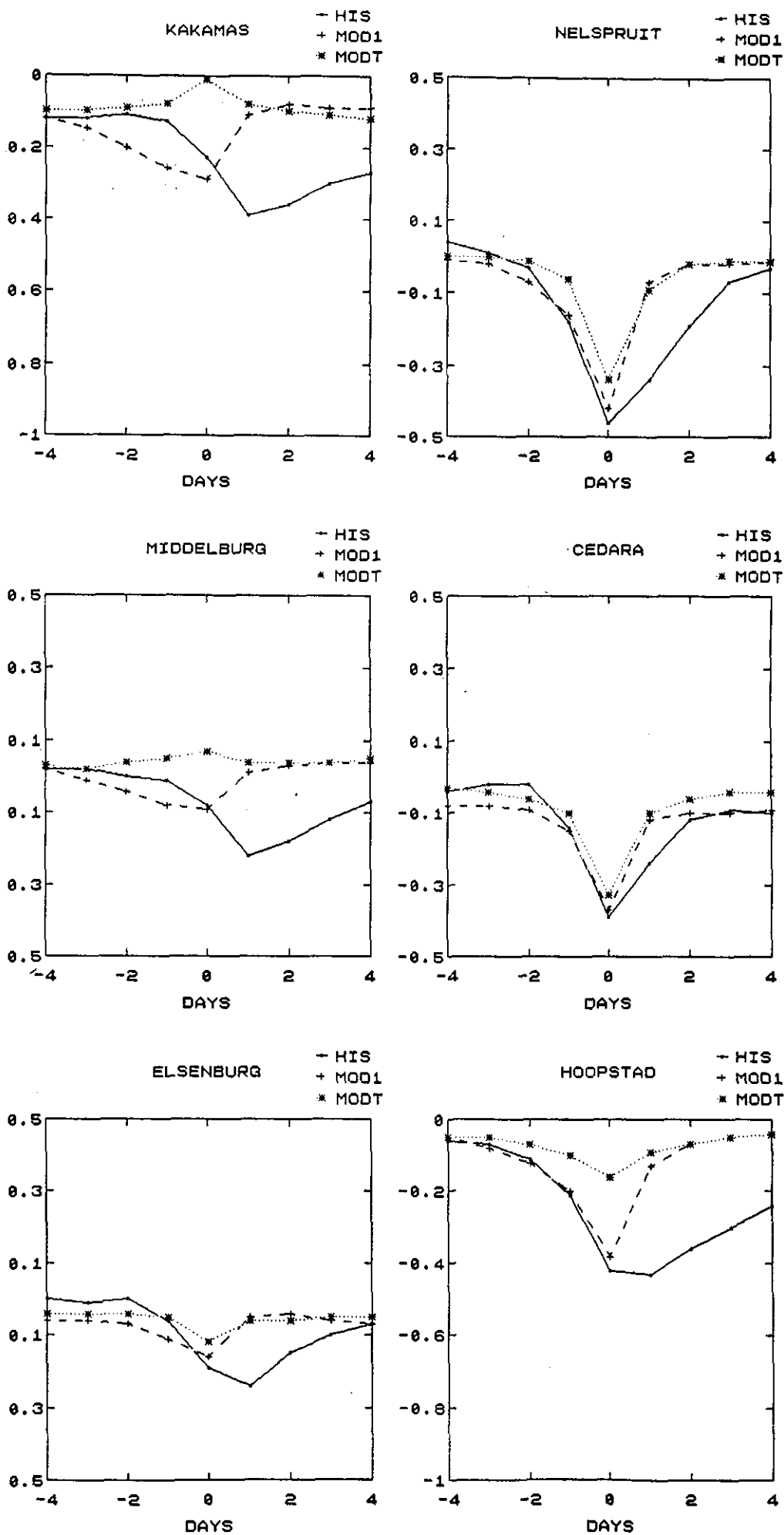
CHAPTER 6  
**FIGURE 6.85** Cross-correlation coefficients for evaporation and minimum humidity *Goodness of Fit*



run



CHAPTER 6 Goodness of Fit  
**FIGURE 6.87** Cross-correlation coefficients for sunshine duration and maxi-  
 mum humidity



CHAPTER 6 Goodness of Fit  
**FIGURE 6.88** Cross-correlation coefficients for sunshine duration and minimum  
humidity

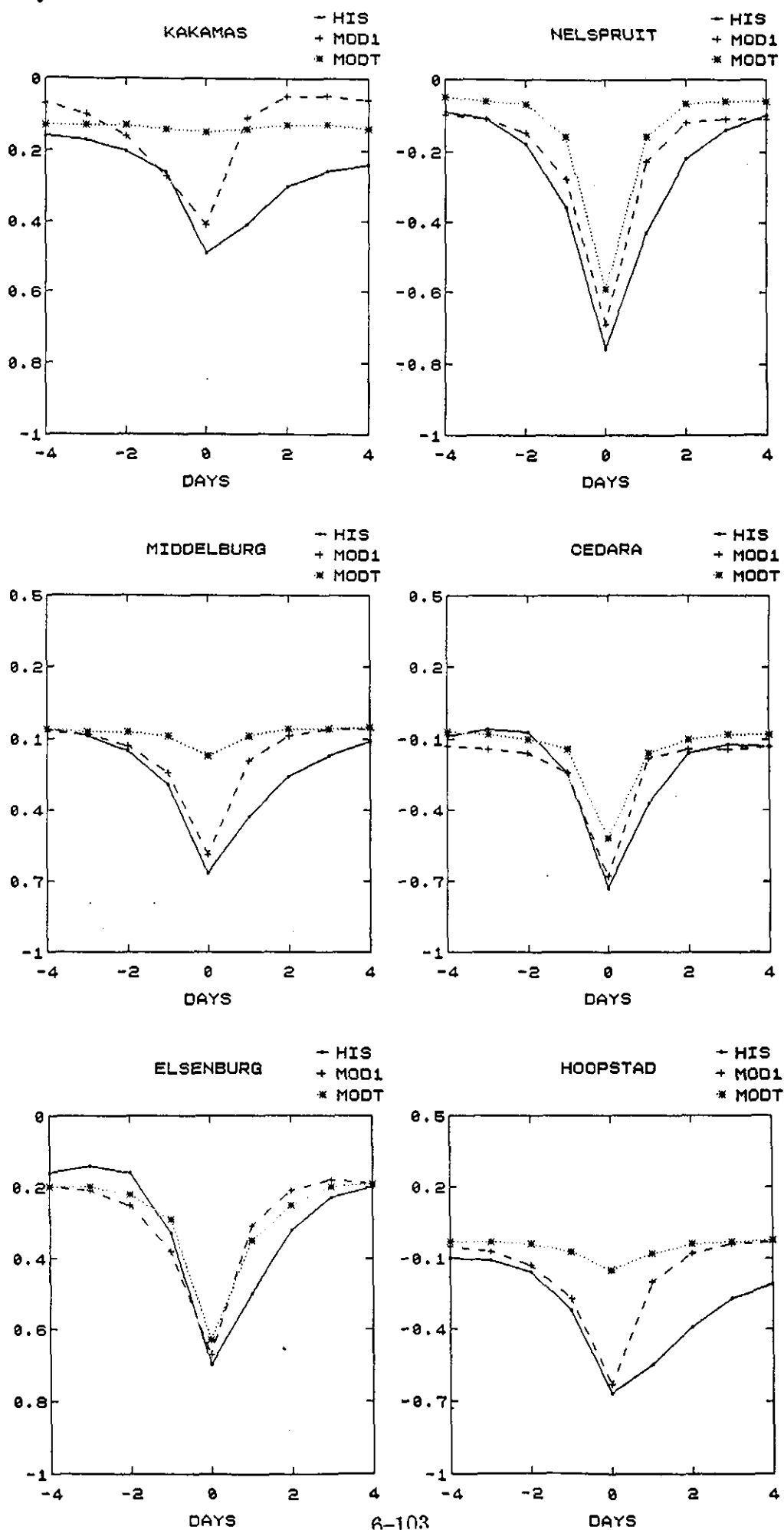
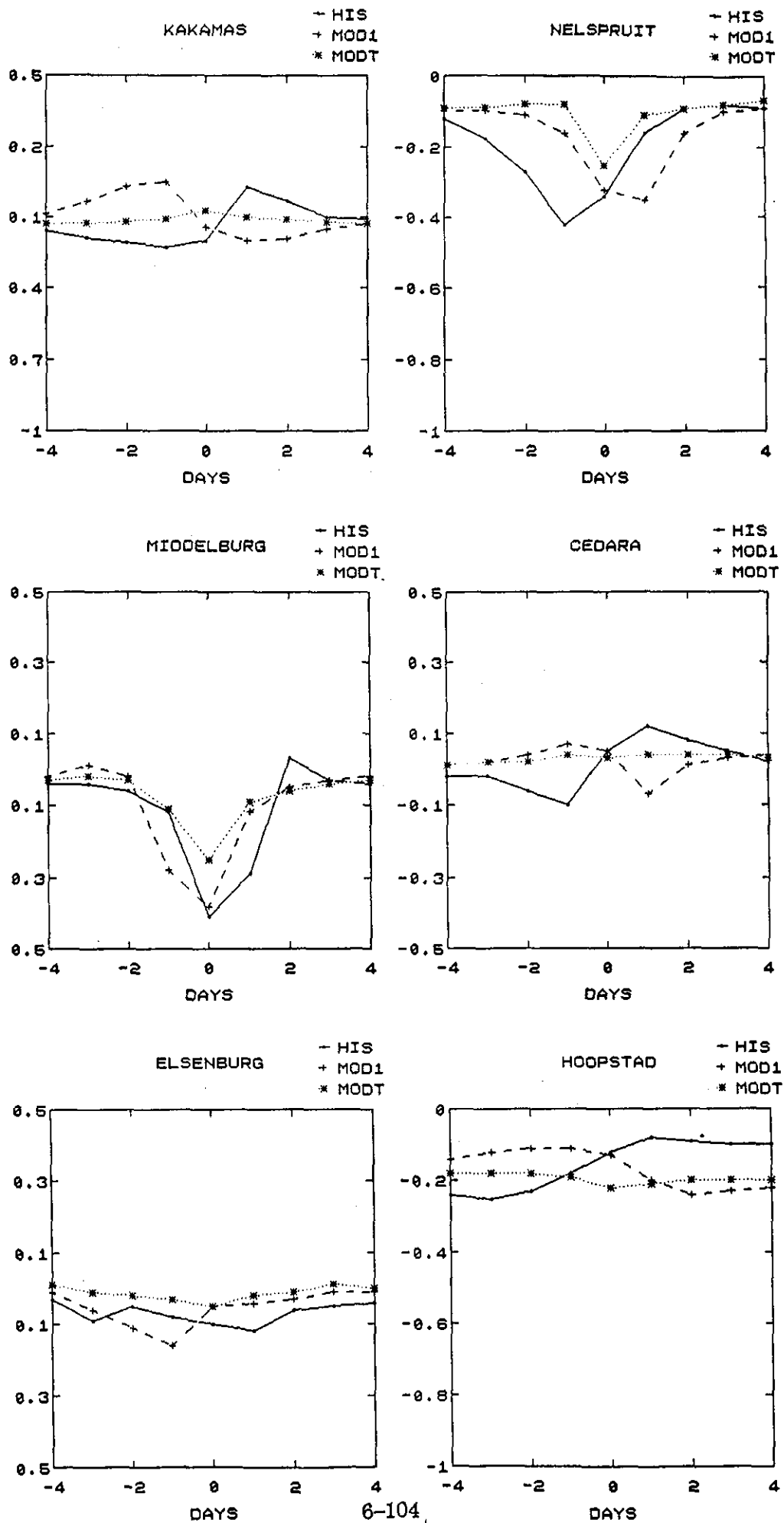


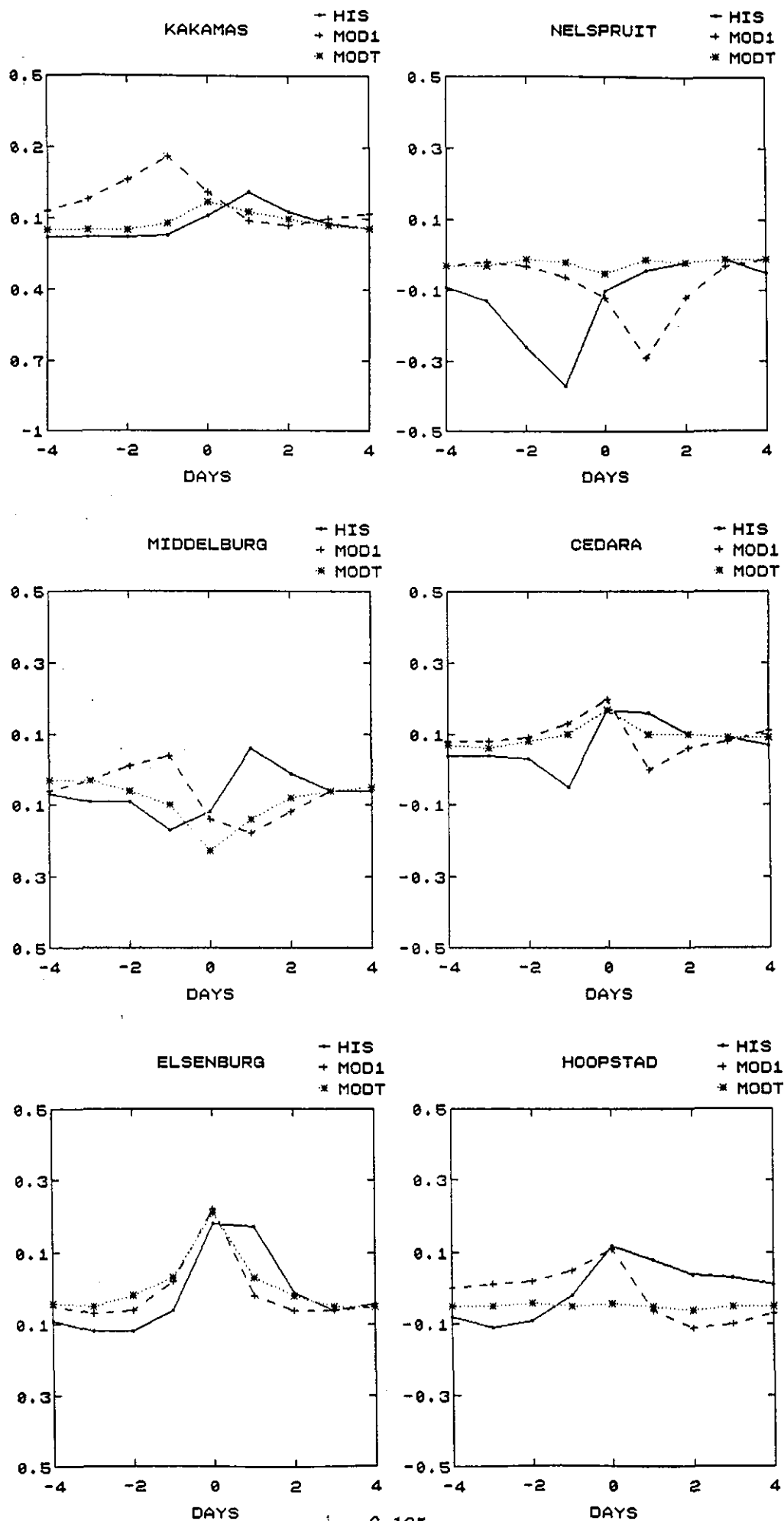
FIGURE 6.89 Cross-correlation coefficients for wind run and maximum humidity

*Goodness of Fit*

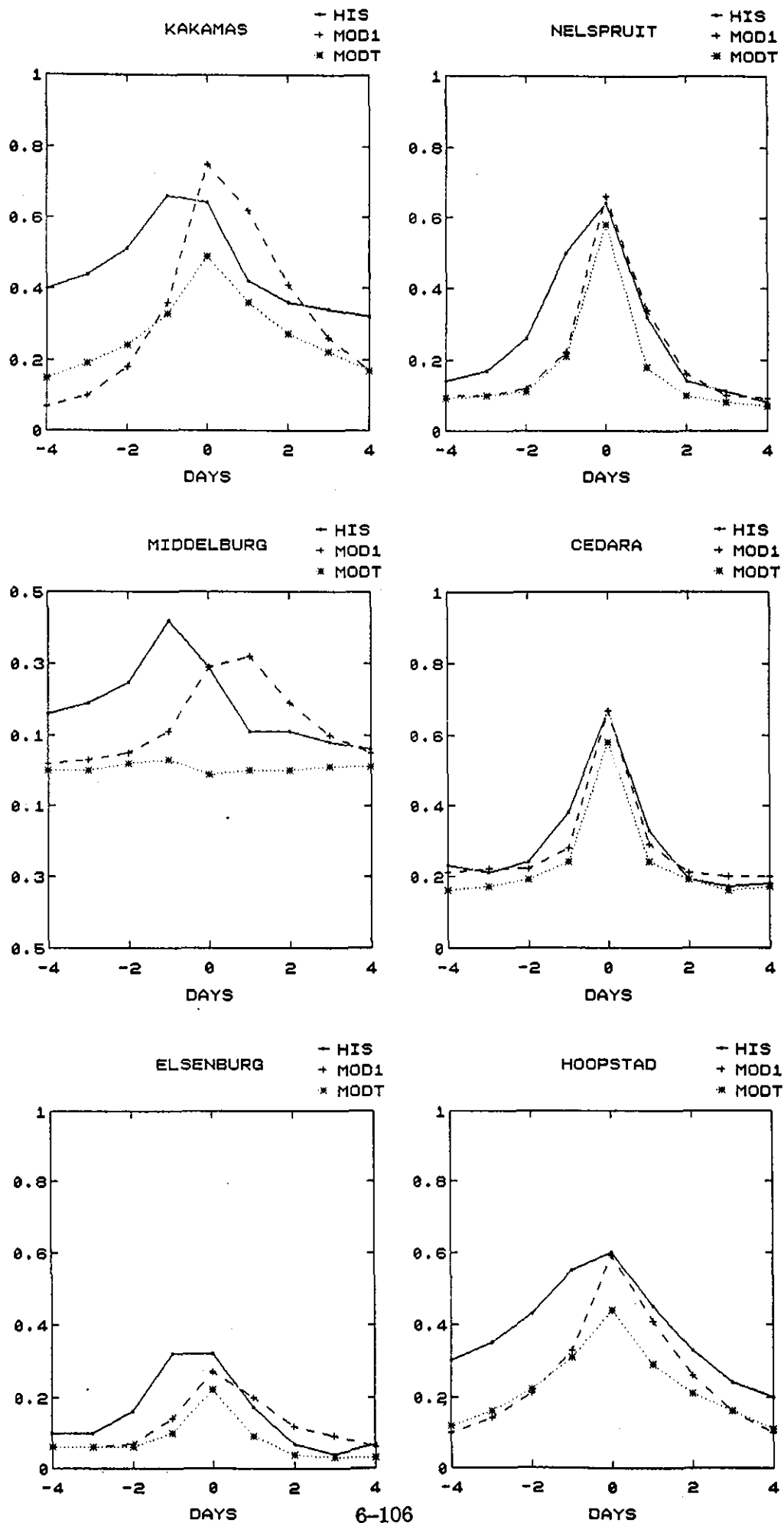


CHAPTER 6  
**FIGURE 6.90** Cross-correlation coefficients for wind run and minimum humidity

*Goodness of Fit*



CHAPTER 6 *Goodness of Fit*  
**FIGURE 6.91** Cross-correlation coefficients for maximum humidity and  
 minimum humidity



sometimes also differ from the historical cross-correlation coefficients for lags 1 and -1, but generally this difference is quite small.

### Summary

If one reflects on the complexity of the climate process and in particular on the large number of properties which the models are required to preserve, it can be reasonably concluded that the models perform remarkably well. All but a few of the relevant parameter functions and cross-correlations are preserved faithfully by the models. One other factor that one must keep in mind when evaluating the performance of the models is the quality and quantity of the historical record. For the model parameters to truly reflect the properties of climate variables there must be enough data records to have a representative sample of the climate variables. It is difficult to be specific about how long a record needs to be in order to be representative. Very roughly, and based on our experience, we would recommend the minimum of 20 years of record before one can feel confidence in the results.

There are a number of weaknesses displayed by the models. The most important are:

- (i) Model T does not retain the property of the monthly means in the variable minimum humidity for Hoopstad. It also does not preserve the monthly means on wet days for the variables maximum humidity and minimum humidity for most stations, in particular those that have relatively few rainfall observations.
- (ii) Model 1 does not retain the property of monthly standard deviations in the variable minimum humidity for Kakamas. It also does not preserve the monthly standard deviations on wet days for the variables maximum humidity and minimum humidity for some stations.
- (iii) Model T shows a weakness in maintaining the cross-correlation coefficients of maximum humidity and minimum humidity and the other variables in particular for the stations Kakamas, Middelburg and Hoopstad.

A choice of model at this point is not straightforward as the performance of the models is neither perfect nor totally without merit, but each model shows strengths and weaknesses. A criterion to base our preference on any particular model can be based on factors such as:

- (a) Implementation costs, that is derivation of theory for parameter estimation, complexity of model in terms of number of parameters needed and the computational simplicity aspect of the model.

(b) Preservation of climate variable properties by the model.

Model T is the more complex of the models in terms of computational difficulty. The model parameters are estimated iteratively and therefore more time consuming. It is also the most flexible of the models in that each variable is allowed to be modelled by a model that "best" describes it. In Model 1 all climate variables have the same structure.

Generally Model 1 appears to perform as well and sometimes better than Model T in describing some aspects of the climate variables. However, Model T cannot be simply dismissed as it does perform better than Model 1 in some aspects and one must bear in mind that in Model T the climate sequences are separated into four, while Model 1 only separates them into two. Thus fewer observations are available for the estimation of some of the parameters of Model T. An increase in the length of the historical record may therefore result in an improved performance by Model T.

## CHAPTER 7

### SUMMARY AND RECOMMENDATIONS

This chapter gives a brief summary of the study performed followed by the main research findings and finally by recommendations.

#### Summary

Five stochastic models to describe daily climate sequences of South Africa were considered. The climate variables included in these models are rainfall, maximum and minimum temperature, maximum and minimum humidity, evaporation (when records available), sunshine duration and wind run. Except for rainfall, which is an essential component of all the models, this list of variables may be either reduced or augmented. Thus the modelling procedure which has been developed is not restricted to this particular set of climate variables.

The models are required to preserve important properties exhibited by the daily climate sequences. These properties are seasonality, wet/dry day effect, autocorrelation, cross-correlation and boundedness. Suitable transformations need to be applied to the climate variables at the start of the modelling procedure to take care of the property of boundedness.

The technique employed was firstly to model rainfall using a first-order Markov chain with seasonal parameters to describe the occurrence of wet and dry days, while the Weibul distribution was used to describe the rainfall depth of wet days. The rainfall mean was allowed to vary seasonally. This model provides synthetic sequences of wet and dry days. Finally, the remaining climate variables were modelled according to the wet or dry status of each day.

The first model considered was proposed by Richardson (1981), where a stationary residual series is obtained by subtracting the seasonal mean and dividing by the standard deviation of each climate variable, each of the functions conditioned on the wet and dry status of the day. A weakly stationary process suggested by Matalas (1967) is used to model the residual series. It is assumed that the residual series is normally distributed and that the serial correlation of each variable can be described by a first-order autoregressive process.

Three models, referred to as Model 3, 4 and 5 (Model 2 was developed as a prototype

to the others) were developed to incorporate additional flexibility in the autocorrelation structure of Model 1. That is, the autocorrelation structure is allowed to depend on the wet and dry status of the day as well as that of the previous day.

The three new models which were developed form a compatible family of models of varying degrees of complexity. This feature leads to a number of advantages. It allows one to select relatively simple models for sites where historical records are short (as is presently the case at practically all sites in South Africa) and to change the selection to a more complex model when the records become sufficiently long. In addition it is possible to assemble the final multivariate model for a site from components from any of the three model types. Thus, for example, it is possible to apply Model 3 to wind run, Model 4 to minimum and maximum temperature and Model 5 to the remaining variables.

The problem of deciding which model or model combination is most appropriate for a particular site can be determined objectively. The Akaike Information Criterion was found to be suitable for this purpose and has been incorporated in the software package which was developed for this project.

Apart from the mathematical development of the new models, one of the most difficult obstacles that had to be overcome in the course of the project was that of controlling the range of extreme values generated by the models. In part this problem arises because some of the climate variables must fall within fixed boundaries and, in addition, some of these variables (for example maximum humidity) exhibit a high frequency of occurrence on or near their boundaries. Suitable transformations had to be found to ensure that the generated value would remain within their appropriate bounds. The results of our validation tests indicate that this type of difficulty can be successfully overcome.

A second problem that had to be solved related to the presence of gaps in the historical records. As well as the gaps that were present in the original record one has to add the gaps which are created by filtering out observations that are clearly incorrect (for example, that fall outside their permissible range). The number of missing values in the historical records used in this report ranged between 1% and 13% of the data. The serial correlation and cross-correlation structure of climate variables does not allow one to simply ignore missing values. Special methods had to be developed to deal with this problem. We found that a procedure based on the EM algorithm can be used to satisfactorily estimate the missing values thereby filling the gaps in the historical records.

The results of model validation for the rainfall model confirm the findings of Zucchini and Adamson (1984), namely that the assumptions regarding the characteristics of daily rainfall sequences, the rationale of model structure and the parameter estimation techniques are particularly successful in providing a model that can adequately reproduce the properties of daily rainfall sequences.

The requirement for an accurate simulation of the occurrence of wet and dry days is very important in the present study as these simulated sequences are used to determine the generation procedure to be adopted by the other climate variables. This component of the rainfall model was found to be very successful in preserving the characteristics of the occurrence of wet and dry days.

As already mentioned, the multivariate models for climate data were required to meet certain specifications. Namely, they had to preserve properties such as seasonality, wet/dry day effect, autocorrelation, cross-correlation as well as annual, monthly and daily properties, in particular the mean and the standard deviation functions. Tests of the multivariate models for climate data showed that the models were capable of representing almost all the characteristics exhibited by the historical data.

Whenever the models showed differences between the simulated and historical sequences, it was noted that it usually was for the variables wind run, maximum humidity and minimum humidity and mainly for the case where these sequences were conditioned on wet days. Further investigation revealed that the stations where these differences occurred were those for which relatively few rainfall records are observed.

When evaluating the performance of the climate models one must bear in mind the fact that the length of the historical records determines in a way the performance of the models. A relatively short historical record leads to three problems. Firstly, one is estimating a large number of parameters with very few data values thus decreasing the precision of the estimates. Secondly, because the models separate the sequences into wet and dry sequences, the effective record length for the conditioned estimates is further reduced, in particular for the wet sequence as rainfall events in some parts of South Africa are relatively rare. Thus long records of climate observations are needed to compensate for the lack of rainfall events. Thirdly, the fact that the record lengths of the stations in this study are quite small, combined with the fact that there are missing observations in the records means that the historical data might not wholly be representative of the long term climate for that

particular location.

The climate variables investigated in this study have a very complex joint distribution. Each variable exhibits a number of distinctive features and in addition the variables are interdependent. Any model which is to usefully describe climate sequences must preserve these properties. This study has shown that it is feasible to model climate on a daily basis and that there are at least four models which can be used to do so. Either Model 1 or a combination of Models 3, 4 and 5 can be used. A choice between Model 1 and Model T is not clear. Both models show some weaknesses and some merits. Model 1 does appear to perform better than Model T for those stations that have few rainfall observations, such as Kakamas. It is also less time consuming in parameter estimation as it does not estimate them iteratively. However, Model T (that is, a combination of Models 3, 4 and 5) does outperform Model 1 in some instances and we would expect that this will become increasingly the case as the historical records become longer.

The software which was developed in this project covers both the parameter estimation and the generating of artificial sequences for all the models that have been described in this report. The programs were coded in FORTRAN and make no use of licensed software. In addition they can be implemented on micro computers thereby making the methodology easily accessible to a wide range of users.

## **Recommendations**

### **Quality of historical records**

The main obstacle to the application of the techniques described in this report on a large scale is the lack of suitable historical records. This refers to both the quantity and the quality of available data. The records which were used for this report represent some of the best available in South Africa. Nevertheless, for the purpose of modelling daily climate, they are barely adequate. Although there is little that can be done to increase the length of records except to wait for more data to be collected, it should be possible to improve the quality of historical records. In particular it would be useful if some measure of the reliability of the observations were also recorded on a regular basis. As we have repeatedly pointed out in the body of the report, one of the problems which we encountered was that of identifying incorrect observations. This task would be considerably simplified if one had some index of reliability associated with (ideally) each recording or set of recordings.

**Transfer of technology**

For the methods developed in this project to realise their full potential it will be necessary to calibrate the models at many more sites. As was pointed out in the report, no special training is required to use the programs for generating climate sequences once the parameters of the model have been estimated. However, some training is required to use the program to prepare the data for estimation and to carry out the estimation for a new site. We estimate that, with instruction, it would require two to three weeks for a competent programmer to learn how to use the methodology.

We recommend that the Computing Centre for Water Research (CCWR) be approached to acquire the expertise to implement the estimation techniques and with the help of users, gradually build up a data base of estimates of the model parameters for as many sites as possible in South Africa. The CCWR already offer a similar data product, namely the parameter estimates of a daily rainfall model for 2550 sites in South Africa. These arose from a previous Water Research Commission project (Zucchini and Adamson (1984)). The CCWR also offer the artificial rainfall generating program which can be applied to any of these sites. Thus the programs developed in the course of this project constitute a logical extension of a service that the CCWR already offer.

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## APPENDIX A

### The choice of the Fourier approximation, $L$

The order of approximation of the Fourier representation of a function,  $\lambda(t)$ , is always taken to be an odd integer. This restriction is made partly for programming convenience and partly for the following reason:

If we rewrite the Fourier representation of  $\lambda(t, L)$  by its polar form, we get

$$\lambda(t, L) = \begin{cases} \alpha_0 + \sum_{i=1}^p \alpha_i \cos \left( \frac{2\pi i}{NT} ((t-1) - \beta_i) \right), & L \text{ odd} \\ \alpha_0 + \sum_{i=1}^p \alpha_i \cos \left( \frac{2\pi i}{NT} ((t-1) - \beta_i) \right) + \alpha_p \cos \frac{2\pi p(t-1)}{NT}, & L \text{ even} \end{cases}$$

where

$$\alpha_0 = \gamma_1$$

$$\alpha_i = (\gamma_{2i}^2 + \gamma_{2i+1}^2)^{\frac{1}{2}}, \quad i = 1, 2, \dots, p$$

$$\beta_i = \frac{NT}{2\pi i} \arctan \left( \frac{\gamma_{2i+1}}{\gamma_{2i}} \right), \quad i = 1, 2, \dots, p$$

and  $p$  is the integer part of  $\frac{L-1}{2}$ . The  $\alpha_i$  is called the amplitude and  $\beta_i$  is called the phase of the  $i$ th harmonic.

If  $L$  is even, then the highest harmonic does not have a phase parameter. Thus the quality of the fit of the model depends on the time origin selected. If  $L$  is odd we obtain the same degree of approximation for all time origins.

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## APPENDIX B

### Properties of the Fourier series approximation

We have used the Fourier representation of  $\lambda(t)$  as the basis for obtaining approximations. Other representations are also feasible, e.g. polynomials or rational functions. There are several reasons for selecting the Fourier representation rather than other possibilities. Firstly,  $\lambda(t)$  is known to be approximately sinusoidal in shape and consequently we can expect that even for small values of  $L$ , the approximation  $\lambda(t, L) \approx \lambda(t)$  will be reasonably accurate. Secondly,  $\lambda(t, L)$  is periodic, which is a property that  $\lambda(t)$  is known to have. Thirdly, the individual components in the representation are orthogonal, which is a convenient mathematical property.

## APPENDIX C

### The Cholesky decomposition

For  $A$  an  $(n \times n)$  symmetric, positive definite matrix, there exists a unique lower triangular matrix  $F$  with positive diagonal elements such that

$$A = FF^T.$$

This is known as the Cholesky decomposition. An algorithm to reduce a matrix  $A$  to its Cholesky decomposition is given below.

#### Notation

$f_{ij}$  is the  $ij$ th element of the matrix  $F$ .

$a_{ij}$  is the  $ij$ th element of the matrix  $A$ .

#### Algorithm

Step 1: Set  $f_{11} = \sqrt{a_{11}}$

Step 2: For  $j = 2, 3, \dots, n$

Set  $f_{j1} = \frac{a_{1j}}{f_{11}}$

Next  $j$ .

Step 3: For  $i = 2, 3, \dots, n-1$

$$\text{Set } f_{ii} = \sqrt{a_{ii} - \sum_{j=1}^{i-1} f_{ij}^2}$$

For  $j = i+1, i+2, \dots, n$ .

$$\text{Set } f_{ji} = \frac{a_{ij} - \sum_{k=1}^{i-1} f_{ik}f_{jk}}{f_{ii}}$$

Next  $j$ .

Next  $i$ .

Step 4:

$$\text{Set } f_{nn} = \sqrt{a_{nn} - \sum_{j=1}^{n-1} f_{nj}^2}$$

## APPENDIX D

A listing of the FORTRAN programs referred to in Chapter 5 are obtainable from the CCWR.  
The address is:

Computing Centre for Water Research

c/o University of Natal

P O Box 375

Pietermaritzburg

3200

Tel. (0331) 63320 ext. 177/178

Fax (0331) 61896

Step 5: End.

The elements of  $F$  above the main diagonal are defined to be zero. The above algorithm does not set them to zero, so if necessary the following step should be inserted immediately preceding "Next  $j$ " in Step 3:

Set  $f_{ij} = 0$ .

## APPENDIX E

### The EM algorithm

In any data record collected over a long period of time, one would expect to encounter gaps, where the number of gaps usually increases proportionally with the size of the data set.

Factors which contribute to the occurrence of these gaps may be, for example, loss of records, temporary absence of observers, breakdown of measuring devices or simply incorrect recordings noted. Whatever the reason for their occurrence, gaps in climate variables are problematic for the following reasons:

Firstly, the cross-correlation structure present in the multivariate time series will be destroyed if there are missing values present. Secondly, the autocorrelation structure breaks down when gaps occur and finally, the seasonal structure disappears if the data is not complete.

An effective way of dealing with incomplete data sets is to "fill" these gaps with data. A recent method known as the EM algorithm has been shown to work very satisfactorily when estimating missing values in rainfall data (Makhuvha, 1988). In fact, out of the several methods investigated, the EM algorithm was chosen as the most efficient method for estimation of missing rainfall records, and it performs at least as well as the other methods in terms of accuracy.

The theory and definition of the EM algorithm given here has been extracted from Makhuvha 1988. The same terminology has been adhered to, with only slight changes to suit it to the present problem.

Literature focuses attention on estimating model parameters in the presence of missing observations. However, we are interested in the missing values themselves. Thus the convergence criteria is based on the estimated missing values rather than on the successive parameter estimates.

### General description of the EM algorithm

The EM algorithm is a method which iteratively computes maximum likelihood estimates when some observations are missing. Let  $Z$  be a complete data set matrix of  $n$  observations on  $k$  climate variables, where  $k \geq 2$  and  $n \geq k + 2$ . We assume that the

data is generated by a model described by a density function  $f(Z|\phi)$  indexed by unknown parameter  $\phi$ . Given the model and parameter vector  $\phi$ ,  $f(Z|\phi)$  is a function of  $Z$ , that is, of the observations.

**Definition:** The likelihood function  $L(\phi|Z)$  is any function of  $\phi$  which is proportional to  $f(Z|\phi)$  when given the data value  $Z$ .

We denote the log-likelihood function by

$$\ell(\phi, Z) = \ln L(\phi, Z).$$

Let  $Z = (Z_{\text{obs}}, Z_{\text{mis}})$  where  $Z_{\text{obs}}$  denotes the observed values of  $Z$  and  $Z_{\text{mis}}$  denotes the missing values of  $Z$ . Write

$$Z_{\text{obs}} = (Z_{\text{obs},1}, Z_{\text{obs},2}, \dots, Z_{\text{obs},n})$$

where  $Z_{\text{obs},i}$  represents the set of climate variables having observation at  $i$ ,  $i = 1, 2, \dots, n$ .

Let  $f(Z|\phi) = f(Z_{\text{obs}}, Z_{\text{mis}}|\phi)$  denote the density function of the joint distribution of  $Z_{\text{obs}}$  and  $Z_{\text{mis}}$ . To obtain the marginal probability density of  $Z_{\text{obs}}$ , the missing data  $Z_{\text{mis}}$  is integrated out. That is,

$$f(Z_{\text{obs}}|\phi) = \int f(Z_{\text{obs}}, Z_{\text{mis}}|\phi) dZ_{\text{mis}}. \quad (1)$$

The likelihood function of  $\phi$  based on  $Z_{\text{obs}}$  is defined to be any function of  $\phi$  proportional to  $f(Z_{\text{obs}}|\phi)$ :

$$L(\phi, Z_{\text{obs}}) \propto f(Z_{\text{obs}}|\phi).$$

In situations where values are missing at random,  $L(\phi, Z_{\text{obs}})$  is called the true likelihood of  $\phi$  based on the observed data  $Z_{\text{obs}}$ . By making use of the complete data specification  $f(Z|\phi)$ , the EM algorithm is used to estimate the parameter  $\phi$  which maximizes  $f(Z_{\text{obs}}|\phi)$ . In other words, we try to maximize the likelihood function

$$L(\phi, Z_{\text{obs}}) = \int f(Z_{\text{obs}}, Z_{\text{mis}}|\phi) dZ_{\text{mis}} \quad (2)$$

with respect of  $\phi$ .

### Definition of the EM algorithm

The EM algorithm has a useful and simple interpretation when the complete data  $Z$  has a distribution from the regular exponential family defined by

$$f(Z|\phi) = \frac{b(Z) \exp(\phi t(Z)^T)}{a(\phi)} \quad (3)$$

where

$\phi$  denotes a  $(1 \times R)$  vector of parameters,

$t(Z)$  denotes a  $(1 \times R)$  vector of complete data sufficient statistics, and

$a$  and  $b$  are functions of  $\phi$  and  $Z$  respectively.

The parameterization of  $\phi$  in (3) is unique up to an arbitrary non-singular  $(R \times R)$  linear transformation, as is the corresponding choice of  $t(Z)$ .

We restrict our attention to only one class of the exponential type of distribution, namely, the Multivariate Normal distribution. We say that a distribution is Multivariate Normal if its density function is given by:

$$f(Z|\mu, \Sigma) = (2\pi)^{-k/2} |\Sigma|^{-\frac{1}{2}} \exp\left[-\frac{1}{2}(Z - \mu)^T \Sigma^{-1} (Z - \mu)\right] \quad (4)$$

where

$$\begin{aligned} Z^T &= (Z_1 \ Z_2 \ \dots \ Z_k), \\ \mu &= (\mu_1 \ \mu_2 \ \dots \ \mu_k), \\ \Sigma &= \begin{bmatrix} \sigma_1^2 & \dots & & \dots & \sigma_{1k} \\ \vdots & & \vdots & & \vdots \\ \vdots & & \sigma_{ij} & & \vdots \\ \sigma_{k1} & \dots & \vdots & \dots & \sigma_k^2 \end{bmatrix} \end{aligned} \quad (5)$$

where  $\sigma_{ij}$  is the covariance of the  $i$ th and  $j$ th component of  $Z$ .

Suppose we are dealing with more than one set of observations, that is, we have a matrix of  $n$  sets of observations such that

$$Z = \begin{bmatrix} Z_{11} & Z_{12} & \dots & Z_{1k} \\ Z_{21} & Z_{22} & \dots & Z_{2k} \\ \vdots & \vdots & \ddots & \vdots \\ Z_{n1} & Z_{n2} & \dots & Z_{nk} \end{bmatrix}. \quad (6)$$

The likelihood of the observations (6) is

$$L(\mu, \Sigma|Z) = (2\pi)^{-\frac{nk}{2}} |\Sigma|^{-\frac{n}{2}} \exp\left[-\frac{1}{2} \sum_{i=1}^n (Z_i - \mu)^T \Sigma^{-1} (Z_i - \mu)\right] \quad (7)$$

Using (7) we can find the sufficient statistics for the parameters.

$$\begin{aligned}
L(\mu, \Sigma | Z) &= (2\pi)^{\frac{-nk}{2}} |\Sigma|^{\frac{-n}{2}} \exp \left[ -\frac{1}{2} n \operatorname{tr}(\mu^T \mu \Sigma^{-1}) \right] \exp \left[ -\frac{1}{2} \sum_{i=1}^n \operatorname{tr}(Z_i Z_i Z_i^T) \begin{pmatrix} -2\Sigma^{-1} \\ \Sigma^{-1} \end{pmatrix} \right] \\
&= (2\pi)^{\frac{-nk}{2}} |\Sigma|^{\frac{-n}{2}} \exp \left[ -\frac{1}{2} n \operatorname{tr}(\mu^T \mu \Sigma^{-1}) \right] \\
&\quad \exp \left[ -\frac{1}{2} \sum_{i=1}^n \left( \sum_{j=1}^k \sum_{\ell=1}^k [Z_{ij} \mu_{\ell} \sigma_{j\ell} + Z_{ij} Z_{i\ell} \sigma_{j\ell}] \right) \right] \\
&= (2\pi)^{\frac{-nk}{2}} |\Sigma|^{\frac{-n}{2}} \exp \left[ -\frac{1}{2} n \operatorname{tr}(\mu^T \mu \Sigma^{-1}) \right] \\
&\quad \exp \left[ -\frac{1}{2} \sum_{i=1}^n (1_n \otimes Z_i)^T \begin{pmatrix} \mu_1 \sigma_1 \\ \vdots \\ \mu_k \sigma_k \end{pmatrix} - \frac{1}{2} \sum_{i=1}^n (Z_i \otimes Z_i)^T \sigma_i \right]
\end{aligned}$$

Therefore

$$t(Z) = \begin{pmatrix} \sum_{i=1}^n 1_n \otimes Z_i \\ \sum_{i=1}^n Z_i \otimes Z_i \end{pmatrix} \quad (8)$$

$$\phi = \begin{bmatrix} \mu_1 \sigma_1 \\ \vdots \\ \mu_k \sigma_k \\ -\frac{1}{2} \sigma_1 \\ \vdots \\ -\frac{1}{2} \sigma_k \end{bmatrix} \quad (9)$$

where

$\phi$  is a vector of parameters, and

$t(Z)$  is the sufficient statistics for  $\phi$  since it does not depend on any parameter.

Since the statistics  $t(Z)$  is sufficient for the parameter  $\phi$ , it therefore has all the relevant information contained in  $Z$  for inference about the parameter.

### The E step and the M step of EM.

Each iteration of the EM algorithm involves two steps which are called the expectation step (E step) and the maximization step (M step). The steps given below may be applied if equation (7) satisfies the conditions of it being a class of the exponential type of distribution.

Suppose that  $\phi^{(p)}$  denotes the current value of  $\phi$  after  $p$  cycles of the algorithm. The next cycle involves the following two steps:

E step: At the  $(p+1)$  cycle, the E step is the computation of the conditional expectation of the complete data sufficient statistics given:

- i) the observed data  $Z_{\text{obs}} = (Z_{\text{obs},1}, \dots, Z_{\text{obs},n})$  and
- ii) the estimated value of the parameter from the  $p$ th cycle.

That is, we compute

$$t^{(p)} = E[t(Z)|Z_{\text{obs}}, \phi^{(p)}]. \quad (10)$$

M step: At the  $(p+1)$  cycle, the M step is the maximization of the complete data likelihood function in which the complete data sufficient statistics  $t(Z)$  has been replaced by its conditional expectation obtained in the E step. We set the derivatives of the complete data likelihood function to zero and determine  $\phi^{(p+1)}$ , i.e. as the solution of the equation

$$E(t(Z)|\phi) = t^{(p)} \quad (11)$$

which defines the maximum likelihood estimator of  $\phi$  under the assumption that (7) is a class of the exponential family.

We now show how the E and M steps of the EM algorithm are obtained under the assumption that the distribution is multivariate normal.

If, at the  $p$ th iteration,  $\phi^{(p)}$  denotes the current estimates of the parameters, then the E step of the algorithm consists of calculating:

$$\begin{aligned} E\left(\sum_{i=1}^n Z_{ij}|Z_{\text{obs}}, \phi^{(p)}\right) &= \sum_{i=1}^n Z_{ij}^{(p)}, \quad j = 1, 2, \dots, k. \\ E\left(\sum_{i=1}^n Z_{ij}Z_{i\ell}|Z_{\text{obs}}, \phi^{(p)}\right) &= \sum_{i=1}^n Z_{ij}^{(p)}Z_{i\ell}^{(p)} + C_{j\ell}^{(p)}, \quad j, \ell = 1, \dots, k, \end{aligned}$$

where

$$\begin{aligned} Z_{ij}^{(p)} &= Z_{ij} && \text{if } Z_{ij} \text{ is observed} \\ &= E(Z_{ij}|Z_{\text{obs},i}, \phi^{(p)}) && \text{if } Z_{ij} \text{ is missing} \end{aligned}$$

and

$$\begin{aligned} C_{j\ell}^{(p)} &= 0 && \text{if } Z_{ij} \text{ or } Z_{i\ell} \text{ are observed} \\ &= \text{cov}(Z_{ij}, Z_{i\ell}|Z_{\text{obs},i}, \phi^{(p)}) && \text{if } Z_{ij} \text{ or } Z_{i\ell} \text{ are missing} \end{aligned}$$

Missing values  $Z_{ij}$  are therefore replaced by the conditional mean of  $Z_{ij}$  given the set of values  $Z_{\text{obs},i}$  observed for that observation.

Similarly, the maximization step (M step) is found from equation (8). The new estimates  $\phi^{(p+1)}$  of the parameters are estimated as follows:

$$\begin{aligned}\mu_j^{(p+1)} &= \frac{1}{n} \sum_{i=1}^n Z_{ij}^{(p)}, \quad j = 1, 2, \dots, k \\ \sigma_{j\ell}^{(p+1)} &= \frac{1}{n} E \left( \sum_{i=1}^n Z_{ij} Z_{i\ell} | Z_{\text{obs}} \right) - \mu_j^{(p+1)} \mu_\ell^{(p+1)} \\ &= \frac{1}{n} \sum_{i=1}^n \left[ (Z_{ij}^{(p)} - \mu_j^{(p+1)})(Z_{i\ell}^{(p)} - \mu_\ell^{(p+1)}) + C_{j\ell}^{(p)} \right], \quad j, \ell = 1, 2, \dots, k.\end{aligned}$$

### Estimation of missing values using the EM

The method described here estimates the missing data point, say  $y_\ell$ , by using all the records, i.e. estimated and real records. In this method all the observations are utilized after the initial estimation stage.

### Algorithm

Suppose we are considering a climate variables matrix  $Z$  of dimension  $n \times (k+1)$ . Partition the  $Z$  matrix into a vector of observations in the target variable,  $y$ , of dimension  $(n \times 1)$  and a matrix of observations in the control variables,  $X$ , of dimension  $(n \times k)$ . Note that any variable in the  $Z$  matrix can be regarded as the target variable, depending on which variable's missing values we are currently estimating.

Suppose we wish to estimate the missing value  $y_\ell$ :

Cycle 0

Step 1

Construct the vector  $y^*$  from  $y$  and the matrix  $X^*$  from the  $(n \times k)$  matrix  $X$ , by eliminating from both all the rows which contain one or more missing observations in either. For example, because  $y_\ell$  is one of the missing observations then the  $\ell$ th row in both  $y$  and  $X$  is eliminated. Suppose that  $y^*$  ends up with  $n^*$  entries, then  $X^*$  is an  $(n^* \times k)$  matrix. The vector  $y^*$  and matrix  $X^*$  should now contain no missing observations. Check that there is sufficient data to regress  $y^*$  on  $X^*$ . If there is not then some of the control variables will have to be removed and one must begin again.

Step 2

Calculate the least squares estimates of the regression parameters using the target variable vector  $y^*$  and the matrix of control variables  $X^*$ . That is find:

$$\hat{\beta}^{(0)} = (X^{*T} X^*)^{-1} X^{*T} y^*$$

and

$$\hat{\beta}_0^{(0)} = \bar{y}^* - \bar{X}^* \hat{\beta}^{(0)}$$

where

$$\bar{y}^* = \frac{1}{n^*} \sum_{i=1}^{n^*} y_i^*,$$

and

$$\bar{X}^* = \frac{1}{n^*} \sum_{i=1}^{n^*} x_{ij}^* \quad j = 1, 2, \dots, k$$

and where the superscript (0) represents the initial estimation cycle.

Estimate the missing record  $y_{\ell}$  using the regression model:

$$y_{\ell}^{(0)} = \beta_0^{(0)} + \sum_{j=1}^k x_{\ell j} \hat{\beta}_j^{(0)}$$

where  $x_{\ell j}$ , ( $j = 1, 2, \dots, k$ ) are the observed values from the control variables matrix  $X$ .

Step 3

After all the missing values in matrix  $Z$  have been estimated, create a new "data" matrix, say  $Z^{(0)}$  containing estimates obtained in place of missing values.

Cycle b

Step 1

Suppose that the new data matrix created in the previous cycle is  $Z^{(b-1)}$ , then partition  $Z^{(b-1)}$  into a matrix of control variables  $X^{(b-1)}$  and a vector of the target variable  $y^{(b-1)}$ .

Step 2

Calculate the least squares estimates by using the new target variable vector  $y^{(b-1)}$  and the matrix of control variables  $X^{(b-1)}$ , where superscript  $(b-1)$  represents the previous cycle. That is find:

$$\hat{\beta}^{(b)} = (X^{(b-1)T} X^{(b-1)})^{-1} X^{(b-1)T} y^{(b-1)}$$

and

$$\hat{\beta}_{(0)}^{(b)} = \bar{y}^{(b-1)} - \bar{X}^{(b-1)} \hat{\beta}^{(b)}$$

where

$$\bar{y}^{(b-1)} = \frac{1}{n} \sum_{i=1}^n y_i^{(b-1)},$$

and

$$\bar{X}^{(b-1)} = \frac{1}{n} \sum_{i=1}^n x_{ij}^{(b-1)}, \quad j = 1, 2, \dots, k.$$

Re-estimate the missing record  $y_\ell$  :

$$y_\ell^{(b)} = \hat{\beta}_{(0)}^{(b)} + \sum_{j=1}^k x_{\ell j}^{(b-1)} \hat{\beta}_j^{(b)}.$$

Let

$$\text{Conv}_\ell = \frac{y_\ell^{(b)} - y_\ell^{(b-1)}}{y_\ell^{(b)}}$$

where  $b$  represents the current iteration;  $b - 1$  represents the previous cycle.

Step 3

If all the missing values in the current variable have been estimated, then check for convergence by using the following criterion:

$$\text{Crit} = \sum_{j=1}^k \sum_{\ell=1}^{n-n_j^*} \text{conv}_{\ell j}$$

where  $n_j^*$  is the number of observed values in the current variable.

Step 4

If  $\text{Crit} \leq F, F$  a small number close to zero, then  $y_\ell$  is considered the required estimate of the missing value and the re-estimation is discontinued, otherwise repeat Steps 1, 2 and 3.

## APPENDIX D

### PROGRAMS

In this appendix, we give a listing of the FORTRAN programs which were developed to implement the methodology discussed in this report. The programs listed here are set for 7 climate variables, excluding rainfall, and for 12 years of historical records.

#### PROGRAM 1

```
-----
C      -----
C      ..... PROGRAM TO COMPUTE VECTORS REQUIRED FOR PARAMETER
C      ESTIMATION OF RAINFALL MODEL.
C      -----

C      NWW(T) = THE NUMBER OF TIMES IT WAS WET IN PERIOD T-1 AND
C      WET IN PERIOD T.

C      NDW(T) = THE NUMBER OF TIMES IT WAS DRY IN PERIOD T-1 AND
C      WET IN PERIOD T.

C      R(I,T) = THE Ith NON-ZERO RAINFALL DEPTH IN PERIOD T,
C      I = 1, 2, ....., NR(T); T = 1, 2, ....., NT.

C      NR(T) = THE NUMBER OF TIMES IT WAS WET (NON-ZERO RAIN) IN
C      PERIOD T.

C      NW(T) = THE NUMBER OF TIMES IT WAS WET IN PERIOD T-1 AND THERE
C      WAS AN OBSERVATION IN PERIOD T (i.e. THERE WAS NOT A
C      GAP IN PERIOD T).

C      ND(T) = THE NUMBER OF TIMES IT WAS DRY (ZERO RAIN) IN PERIOD
C      T-1 AND THERE WAS AN OBSERVATION IN PERIOD T.

C      FOR EACH T = 1,2,...,NT (WHERE NT = THE NUMBER OF PERIODS IN
C      THE YEAR):
C      THE ABOVE ARRAYS ARE REQUIRED BY THE ESTIMATION ALGORITHMS
C      AS FOLLOWS:-
C      i) NW( ) AND NWW( ) ARE REQUIRED TO ESTIMATE THE
C      PARAMETER FOR THE PROBABILITY THAT A WET PERIOD
C      FOLLOWS A WET PERIOD.
C      ii) ND( ) AND NDW( ) ARE REQUIRED TO ESTIMATE THE
C      PARAMETERS FOR THE PROBABILITY THAT A WET PERIOD
C      FOLLOWS A DRY PERIOD.
C      iii) NR( ) AND R( , ) ARE REQUIRED TO FIT THE PARAMETERS
C      OF THE MEAN RAINFALL RAIN IN A WET PERIOD AND THE
C      COEFFICIENT OF VARIATION.

C      THE OUTPUT OF THIS PROGRAM IS GIVEN IN TWO PARTS.  NAMELY, THE
C      FIRST PART:
C      T (T=1,2,...,NT), NW, NWW, ND, NDW, N, NR
C      THE SECOND PART:
C      R(I,T) (I=1,2,...,NR(T); T=1,2,...,NT).
```

```

C      MAIN PROGRAM
C      -----

C      ..... NT = THE NUMBER OF PERIODS IN THE YEAR (e.g. 365 FOR
C              DAILY DATA)
C      ..... NY = THE NUMBER OF YEARS OF DATA (INCLUDING THE MISSING
C              VALUES)
C      ..... NRT = THE MAXIMUM VALUE GIVEN FOR THE DIMENSION OF THE
C              ARRAY R(I,T)
C      ..... RAIN = THE ARRAY THAT CONTAINS THE DATA
C      ..... IND = AN INDICATOR OF THE STATUS OF THE PREVIOUS PERIOD
C              -1 => PREVIOUS OBSERVATION MISSING
C              0 => PREVIOUS PERIOD WAS DRY
C              1 => PREVIOUS PERIOD WAS WET

C      INTEGER          NT,NY,NRT,IND,T
C      PARAMETER        (NT=365)
C      PARAMETER        (NY=12)
C      PARAMETER        (NRT=50)
C      INTEGER          N(NT)
C      INTEGER          NR(NT)
C      INTEGER          NW(NT)
C      INTEGER          NWW(NT)
C      INTEGER          ND(NT)
C      INTEGER          NDW(NT)
C      REAL             R(NRT,NT)

15     FORMAT (7 (I4))
25     FORMAT (8 (F9.2))

      OPEN(UNIT=12,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN(UNIT=10,FILE='\\WATER\\DATA\\COUNTS.DAT',STATUS='UNKNOWN')
      OPEN(UNIT=20,FILE='\\WATER\\DATA\\RAIN.DAT',STATUS='UNKNOWN')

C      ..... THE REQUIRED VECTORS ARE COMPUTED.

      DO 10, I = 1, NT
          N(I) = 0
          NR(I) = 0
          NW(I) = 0
          NWW(I) = 0
          ND(I) = 0
          NDW(I) = 0
          DO 20, J = 1, NRT
              R(J,I) = 0
20         CONTINUE
10     CONTINUE
      IND = -1
      DO 30, J = 1, NY
          DO 40, I = 1, NT
              READ (12,*) RAIN
              IF (RAIN .EQ. 0) THEN

```

```

        N(I) = N(I) + 1
        IF (IND .EQ. 0) THEN
            ND(I) = ND(I) + 1
        ELSEIF (IND .EQ. 1) THEN
            NW(I) = NW(I) + 1
            IND = 0
        ELSEIF (IND .EQ. -1) THEN
            IND = 0
        ENDIF
    ELSEIF (RAIN .GT. 0) THEN
        NR(I) = NR(I) + 1
        R(NR(I),I) = RAIN
        IF (IND .EQ. 0) THEN
            NDW(I) = NDW(I) + 1
            IND = 1
        ELSEIF (IND .EQ. 1) THEN
            NWW(I) = NWW(I) + 1
        ELSEIF (IND .EQ. -1) THEN
            IND = 1
        ENDIF
    ELSEIF (RAIN .LT. 0) THEN
        IND = -1
    ENDIF
40    CONTINUE
30    CONTINUE
    DO 50, I = 1, NT
        N(I) = N(I) + NR(I)
        ND(I) = ND(I) + NDW(I)
        NW(I) = NW(I) + NWW(I)
50    CONTINUE

C      ..... THE VECTORS COMPUTED ARE WRITTEN OUT

    DO 60, I = 1, NT
        WRITE (10,15) I, NW(I), NWW(I), ND(I), NDW(I), N(I), NR(I)
60    CONTINUE
    DO 70, T = 1, NT
        WRITE (20,25) (R(I,T), I = 1, NR(T))
70    CONTINUE

    STOP
    END

```

PROGRAM 2

```

C -----
C ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR RAINFALL
C MODEL - PROBABILITIES OF WET & DRY SEQUENCES
C -----
C ..... THIS PROGRAM USES THE GENERIC NOTATION MM(T) & M(T)
C TO REPRESENT THE RELEVANT ARRAYS AS FOLLOWS:
C   i) WHEN WE ARE ESTIMATING THE PROBABILITY THAT A
C      WET PERIOD FOLLOWS A WET PERIOD, THEN
C      MM(T) = NW(T)
C      M(T) = NWW(T)
C   ii) WHEN WE ARE ESTIMATING THE PROBABILITY THAT A
C      WET PERIOD FOLLOWS A DRY PERIOD, THEN
C      MM(T) = ND(T)
C      M(T) = NDW(T)
C   iii) WHEN WE ARE ESTIMATING THE PROBABILITY THAT PERIOD
C        T IS WET, THEN
C      MM(T) = N(T)
C      M(T) = NR(T)
C
C NP = NUMBER OF PARAMETERS TO BE FITTED
C THETA(NP) = VECTOR OF PARAMETERS ESTIMATED
C AM(0:K) = CORRESPONDING AMPLITUDES, K = (NP-1)/2
C PH(K) = CORRESPONDING PHASES
C P(NT) = CURRENT ESTIMATES OF PROBABILITIES
C L(NT) = CURRENT ESTIMATES OF LOGITS
C DER(NT) = VECTOR OF 1ST PARTIAL DERIVATIVES
C DER2(NT,NT) = MATRIX OF 2ND PARTIAL DERIVATIVES
C PHI(NP,NT) = MATRIX OF FOURIER TERMS
C DELTA = CONVERGENCE CRITERION

```

```

INTEGER      NP,K,NT,ITER,T,MAXITER,NPMAX,KMAX
PARAMETER    (NPMAX=13)
PARAMETER    (KMAX=(NPMAX-1)/2)
PARAMETER    (NT=365)
REAL         MM(NT)
REAL         M(NT)
REAL         PI, LOGIT
PARAMETER    (PI=3.141593)
REAL         AM(0:KMAX)
REAL         PH(KMAX)
REAL         P(NT)
REAL         L(NT)
REAL         DER(NPMAX)
REAL         DER2(NPMAX,NPMAX)
REAL         THETA(NPMAX)
REAL         PHI(NT,NPMAX)

```

```

C 5  FORMAT (4X,2F4.0)  ----- PROB (W/W)
C 5  FORMAT (12X,FF4.0) ----- PROB (W/D)
C 5  FORMAT (20X,2F4.0) ----- PROB (W)

5  FORMAT (4X,2F4.0)
C 5  FORMAT (12X,2F4.0)
C 5  FORMAT (20X,2F4.0)
15  FORMAT (' EPS, MAXITER = ')
25  FORMAT (' ..... DID NOT CONVERGE')
35  FORMAT (/, ' .....', I3, ' ITERATION',/)
45  FORMAT (/, ' AMPLITUDE: ')
55  FORMAT (/, ' PHASE: ')
65  FORMAT (9F8.3)
75  FORMAT (' OPTIMAL PARAMETERS TO BE FITTED: ', I4)
85  FORMAT (' INITIAL ESTIMATES: ', F10.4)

OPEN (UNIT=4,FILE='CON')
OPEN (UNIT=9,FILE='LPT1')
OPEN (UNIT=10,FILE='\\WATER\\DATA\\COUNTS.DAT',STATUS='OLD')

C  ....... INPUT DATA

PRINT 15
READ (4,*) EPS, MAXITER

DO 10, T = 1, NT
    READ (10,5) MM(T), M(T)
10  CONTINUE

CALL TRIG (PHI,NPMAX,NT)

C  ....... DIFFERENT AMOUNT OF PARAMETERS FITTED AT A TIME.
C  ....... PROGRAM STOPS ONCE OPTIMAL NO. OF PARAMETERS ARE FITTED

CRITO = 10 ** 10
DO 300, NP = 1, NPMAX, 2
    WRITE (9,*) 'NO. OF PARAMETERS FITTED = ', NP

C  ....... COMPUTE INITIAL ESTIMATES OF THE PROBABILITIES
C  ....... AND LOGITS

DO 20, T = 1, NT
    IF (MM(T) .GT. 0) THEN
        P(T) = M(T) / MM(T)
    ELSE
        P(T) = -1
        GOTO 20
    ENDIF
    IF (M(T) .EQ. 0) THEN
        L(T) = -5
    ELSEIF (M(T) .EQ. MM(T)) THEN
        L(T) = 5
    ELSEIF ((M(T) .GT. 0.00001).AND.(M(T) .NE. MM(T))) THEN

```

```

                L(T) = LOG (P(T) / (1-P(T)))
            ENDIF
20      CONTINUE

      DO 30, I = 1, NP
        TO = 0
        T1 = 0
        DO 40, T = 1, NT
          IF (MM(T) .GT. 0.00001) THEN
            TO = TO + L(T) * PHI (T,I)
            T1 = T1 + PHI (T,I) ** 2
          ENDIF
40      CONTINUE
        THETA (I) = TO / T1
        WRITE (9,85) THETA (I)
30      CONTINUE

C      ..... ITERATIVE ESTIMATION OF PARAMETERS

      IC = 0
      DO 200, ITER = 1, MAXITER
        WRITE (9,35) ITER

C      ..... COMPUTE 1ST AND 2ND DERIVATIVES

        DO 50, I = 1, NP
          DER (I) = 0
          DO 60, J = 1, NP
            DER2(I,J) = 0
60      CONTINUE
50      CONTINUE
        DELTA = 0
        DO 70, T = 1, NT
          LOGIT = THETA(1)
          DO 80, I = 2, NP
            LOGIT = LOGIT + THETA(I) * PHI(T,I)
80      CONTINUE
          TO = EXP (LOGIT)
          PROB = TO / (1+TO)
          T1 = M(T) - MM(T) * PROB
          T2 = MM(T) * PROB / (1+TO)
          DELTA = DELTA + ABS (P(T) - PROB)
          P(T) = PROB
          DO 90, I = 1, NP
            DER(I) = DER(I) + T1 * PHI(T,I)
            DO 100, J = 1, I
              DER2(I,J) = DER2(I,J) - T2*PHI(T,I)*PHI(T,J)
100      CONTINUE
90      CONTINUE
70      CONTINUE
        DO 110, I = 1, NP
          DO 120, J = I+1, NP
            DER2(I,J) = DER2(J,I)
120      CONTINUE

```

```

110          CONTINUE

          CALL LINEAR (NP,MAX,NP,DER,DER2,THETA)

C          ..... TESTING FOR CONVERGENCE

          IF (DELTA .GT. EPS) THEN
            IC = 0
          ELSE
            IC = 1
          ENDIF
          IF (IC) 200, 200, 400
200          CONTINUE
          WRITE (9,25)

C          ..... TRANSFORMING PARAMETERS TO THEIR AMPLITUDE AND PHASE
C          REPRESENTATION

400          K = (NP-1) / 2
          CALL AMPHA (AM,PH,THETA,NP,MAX,K,PI,NT)

C          ..... MODEL SELECTION

          LLK = 0
          DO 210, T = 1, NT
            IF (MM(T) .GT. 0.000001) THEN
              LLK = LLK+M(T)*LOG(P(T))+(MM(T)-M(T))*LOG(1-P(T))
            ENDIF
210          CONTINUE
          CRIT = -LLK + NP
          IF (CRIT .LT. CRITO) THEN
            LO = NP
            CRITO = CRIT
          ELSE
            WRITE (9,75) LO
            STOP
          ENDIF
300          CONTINUE

          STOP
          END

```

# PROGRAM 3

```

C -----
C ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR THE
C           DISTRIBUTION OF RAINFALL DEPTHS.
C -----

C ..... THE FOLLOWING NOTATION IS USED:
C           NP = NUMBER OF PARAMETERS IN THE MEAN FUNCTION
C           R(,) = MATRIX OF RAINFALL DEPTHS OBSERVED
C           DER() = VECTOR OF 1ST PARTIAL DERIVATIVES
C           DER2() = MATRIX OF 2ND PARTIAL DERIVATIVES
C           Q() = VECTOR OF AVERAGE OBSERVED RAINFALL IN EACH PERIOD
C           THETA() = VECTOR OF PARAMETER ESTIMATES
C           AM() = CORRESPONDING AMPLITUDES
C           PH() = CORRESPONDING PHASES
C           F() = CURRENT ESTIMATE OF THE MEAN
C           SO() = OBSERVED DAILY STANDARD DEVIATIONS
C           SF() = FITTED DAILY STANDARD DEVIATIONS
C           DELTA = CONVERGENCE CRITERION

INTEGER      NP,NT,T,ITER,K,MAXITER
PARAMETER    (NP=3)
PARAMETER    (NT=365)
PARAMETER    (K=(NP-1)/2)
INTEGER      NR(NT)
REAL         R(SO,NT)
REAL         PI,DENOM,NUM
PARAMETER    (PI=3.141593)
REAL         COEFF,DELTA
REAL         DER(NP),SOLN(NP)
REAL         DER2(NP,NP)
REAL         PHI(NT,NP)
REAL         Q(NT),F(NT),SO(NT),SF(NT)
REAL         THETA(NP)
REAL         AM(0:K)
REAL         PH(K)

5  FORMAT (6(4X), I4)
15  FORMAT (14(I5))
25  FORMAT (' EPS, MAXITER = ')
35  FORMAT (' ..... DID NOT CONVERGE')
45  FORMAT (/, ' .....', I3, ' ITERATION',/)
55  FORMAT (/, ' AMPLITUDE: ')
65  FORMAT (/, ' PHASE: ')
75  FORMAT (9F8.3)
85  FORMAT (' OPTIMAL PARAMETERS TO BE FITTED: ', I4)
95  FORMAT (' INITIAL ESTIMATES: ', F10.4)
105 FORMAT (' COEFFICIENT OF VARIATION: ', F10.4)

OPEN (UNIT=4,FILE='CON')
OPEN (UNIT=9,FILE='LPT1')
OPEN (UNIT=10,FILE='\\WATER\\DATA\\RAIN.DAT',STATUS='OLD')

```

```

OPEN (UNIT=20,FILE='\\WATER\\DATA\\COUNTS.DAT',STATUS='OLD')

C      ..... INPUT DATA

      PRINT 25
      READ (4,*) EPS, MAXITER

      DO 10, T = 1, NT
        READ (20,5) NR(T)
10     CONTINUE
      DO 20, T = 1, NT
        READ (10,*) (R(I,T), I = 1, NR(T))
20     CONTINUE

      CALL TRIG (PHI,NP,NT)

C      ..... COMPUTE INITIAL ESTIMATES

      DO 30, T = 1, NT
        IF (NR(T) .GT. 0) THEN
          TERMO = 0
          DO 40, I = 1, NR(T)
            TERMO = TERMO + R(I,T)
40         CONTINUE
          Q(T) = TERMO / NR(T)
        ENDIF
30     CONTINUE
      DO 50, I = 1, NP
        TERMO = 0
        TERM1 = 0
        DO 60, T = 1, NT
          IF (NR(T) .NE. 0) THEN
            TERMO = TERMO + Q(T) * PHI(T,I)
            TERM1 = TERM1 + PHI(T,I) ** 2
          ENDIF
60         CONTINUE
          THETA (I) = TERMO / TERM1
          WRITE (9,95) THETA (I)
50     CONTINUE

C      ..... ITERATIVE PARAMETER ESTIMATION

      IC = 0
      DO 100, ITER = 1, MAXITER
        WRITE (9,45) ITER

C      ..... COMPUTE 1ST & 2ND PARTIAL DERIVATIVES

        DO 70, I = 1, NP
          DER(I) = 0
          DO 80, J = 1, I
            DER2(I,J) = 0
80         CONTINUE
70         CONTINUE

```

```

DO 90, T = 1, NT
  TERMO = THETA(1)
  DO 180, I = 2, NP
    TERMO = TERMO + THETA(I) * PHI(T,I)
180    CONTINUE
  F(T) = TERMO
90    CONTINUE
  DO 110, T = 1, NT
    IF (NR(T) .GT. 0) THEN
      DO 120, I = 1, NP
        DER(I) = DER(I) - NR(T) * (Q(T) - F(T)) * PHI(T,I)
        DO 130, J = 1, I
          DER2(I,J) = DER2(I,J) + NR(T) * PHI(T,I) * PHI(T,J)
130        CONTINUE
120      CONTINUE
    ENDIF
110    CONTINUE
    DO 140, I = 1, NP
      DO 150, J = I+1, NP
        DER2(I,J) = DER2(J,I)
150      CONTINUE
140    CONTINUE

    CALL LINEAR (NP,NP,DER,DER2,THETA)

C      ..... CONVERGENCE TEST

      DELTA = 0
      DO 170, I = 1, NP
        DELTA = DELTA + ABS (DER(I))
170      CONTINUE
      IF (DELTA .GT. EPS) THEN
        IC = 0
      ELSE
        IC = 1
      ENDIF
      IF (IC) 100,100,2
100    CONTINUE
      WRITE (9,35)

      2    DO 200, T = 1, NT
        TERMO = THETA(1)
        DO 220, I = 2, NP
          TERMO = TERMO + THETA(I) * PHI(T,I)
220        CONTINUE

C      ..... COMPUTE FITTED VALUES

      F(T) = TERMO
200    CONTINUE

C      ..... OUTPUT OBSERVED AND FITTED DAILY MEANS

      DO 300, T = 1, NT

```

```

          PRINT *, T, Q(T), F(T)
300      CONTINUE

C          ..... COMPUTE THE COEFFICIENT OF VARIATION

          DENOM = 0
          NUM = 0
          DO 510, T = 1, NT
              DO 520, I = 1, NR(T)
                  NUM = NUM + (R(I,T) - F(T)) ** 2
520          CONTINUE
              DENOM = DENOM + NR(T) * (F(T) ** 2)
510      CONTINUE
          COEFF = SQRT (NUM / DENOM)
          WRITE (9,105) COEFF

C          ..... COMPUTE THE AMPLITUDE AND PHASE REPRESENTATION

          CALL AMPHA (AM,PH,THETA,NP,K,K,PI,NT)

C          ..... COMPUTE THE FITTED AND OBSERVED STANDARD DEVIATIONS

          STOP
          END

```

PROGRAM 4

```

C -----
C ..... PROGRAM TO GENERATE RAINFALL SEQUENCES
C -----

      INTEGER          NT,NY,NP,PSTATE,STATE

C ..... PSTATE = PRESENT STATE OF DAY
C ..... STATE = PREVIOUS STATE OF DAY

      PARAMETER        (NT=365)
      PARAMETER        (NY=51)
      PARAMETER        (NP=3)

C ..... NT = £ OBSERVATIONS PER YEAR
C ..... NV = £ VARIABLES
C ..... NY = £ YEARS TO BE GENERATED
C ..... NP = £ PARAMETERS IN SEASONAL MODEL

      INTEGER          SEED
      REAL             RAIN
      REAL             GAM (2,NP)
      REAL             PHI (NP,0:NT)
      REAL             AMP (0:NP)
      REAL             PHASE (NP)

      COMMON            IDUM1,IDUM2

15      FORMAT (F9.2)
25      FORMAT (' GIVE 2 -VE Nos. TO INITIALIZE RANDOM GENERATOR',/)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\EST.DAT',STATUS='OLD')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\SIMU.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=22,FILE='CON')

C ..... COMPUTE THE FOURIER SERIES TERMS

      CALL COSSIN (PHI,NP,NT)
      DO 60, I = 1, NP
        PHI (I,0) = PHI (I,NT)
60      CONTINUE

      PI=3.14159

C ..... READING PARAMETER ESTIMATES

      READ (12, *) (GAM (1,J), J = 1, NP)
      READ (12, *) (GAM (2,J), J = 1, NP)
      READ (12, *) (AMP (I), I = 0, 1)
      READ (12, *) (PHASE (I), I = 1, 1)
      READ (12, *) CV

```

```

C      ..... INPUT SEED TO START RANDOM NUMBER GENERATOR.  MUST BE
C      NEGATIVE NUMBER.

PRINT 25
READ (22, *) SEED (1), SEED(2)

IDUM1 = SEED (1)
IDUM2 = SEED (2)

C      ..... COMPUTE PARAMETERS NEEDED FOR COMPUTATION OF RAINFALL
C      DEPTH

CALL CALBET (BETA,CV)
ALPH = 1 + 1 / BETA
GAMM = GAMMA (ALPH)
BI = 1 /BETA
W = 0.01721421

C      ..... SET INITIAL STATE OF DAY TO BE DRY
C      ..... SET INITIAL CLIMATE VALUE TO ZERO

C      ..... STATE = 1 ==> DRY
C      ..... STATE = 2 ==> WET

STATE = 1

DO 30, I = 1, NY
  DO 40, J = 1, NT

C      ..... GENERATE RAINFALL VALUE
C      -----

C      ..... COMPUTE PROBABILITY THAT A WET DAY FOLLOWS A WET DAY, OR
C      THE PROBABILITY THAT A WET DAY FOLLOWS A DRY DAY.

      CALL PIEST (NP,GAM,STATE,J,PHI,PI,NT)

C      ..... GENERATE A UNIFORM RANDOM NUMBER BETWEEN 0 AND 1.

      UNIFOR = URAN (IDUM)
      IF (UNIFOR .LT. PI) THEN
        PSTATE = 2
      ELSE
        PSTATE = 1
      ENDIF

C      ..... GENERATE RAINFALL DEPTH
C      -----

      CALL DEPTH3 (IDUM2,NP,RAIN,J,AMP,PHASE,GAMM,BI,W)

C      ..... OUTPUT GENERATED SEQUENCES

```

```

        IF (I .NE. 1) THEN
            WRITE (10,15) RAIN
        ENDIF

C      ..... UPDATE THE STATE OF THE PREVIOUS DAY

        IF (PSTATE .NE. STATE) THEN
            STATE = PSTATE
        ENDIF

40      CONTINUE
30      CONTINUE

        STOP
        END

```

PROGRAM 5

```

C -----
C ..... PROGRAM TO CONDITION DATA SET ACCORDING TO TH WET &
C          DRY STATUS OF THE DAY
C -----

      INTEGER          NY,NT
      PARAMETER        (NY=12)
      PARAMETER        (NT=365)
      INTEGER          SEQ(2,NY,NT)
      INTEGER          COUNT(2,NY)
      REAL             OBSN

15      FORMAT (14(I5))
25      FORMAT (I5)

      DO 20, J = 1, 2
          DO 40, I = 1, NY
              COUNT (J, I) = 0
40          CONTINUE
20      CONTINUE

      OPEN (UNIT=8,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')

      DO 10, J = 1, NY
          DO 50, I = 1, NT
              READ (8, *) OBSN
              IF (OBSN .EQ. 0) THEN
                  COUNT (1, J) = COUNT (1, J) + 1
                  SEQ (1, J, COUNT (1, J)) = I
              ELSEIF (OBSN .GT. 0) THEN
                  COUNT (2, J) = COUNT (2, J) + 1
                  SEQ (2, J, COUNT (2, J)) = I
              ENDIF
50          CONTINUE
10      CONTINUE

      OPEN (UNIT=10,FILE='\\WATER\\DATA\\SEQ2.DAT',STATUS='UNKNOWN')

      DO 60, I = 1, NY
          DO 30, J = 1, 2
              WRITE (10, 25) COUNT (J, I)
              WRITE (10, 15) (SEQ (J, I, K), K = 1, COUNT (J, I))
30          CONTINUE
60      CONTINUE

      STOP
      END

```

# PROGRAM 6

```

C -----
C ..... PROGRAM TO COMPUTE MEAN VECTOR
C -----

```

```

      INTEGER          NV,NY,NT,NPARM
      PARAMETER        (NV=7)
      PARAMETER        (NY=12)
      PARAMETER        (NT=365)
      PARAMETER        (NPARM=11)
      INTEGER          SEQ (2,NY,NT)
      INTEGER          COUNT (2,NY)
      INTEGER          DENOM (NT)
      REAL              MU (2,NT)
      REAL              PHI (NT,NPARM)
      REAL              THETA,OMEGA,PI
      PARAMETER        (PI = 3.14159265)
      DIMENSION        CLIMA[HUGE] (NY,NT)

```

```

5      FORMAT (11F6.2)
15     FORMAT (14 (I5))
25     FORMAT (I5)
35     FORMAT (18X,F9.2)
45     FORMAT (27X,F9.2)
55     FORMAT (36X,F9.2)
65     FORMAT (45X,F10.2)
75     FORMAT (55X,F10.2)
85     FORMAT (65X,F9.2)
95     FORMAT (9X,F9.2)

```

```

      OPEN (UNIT=30,FILE=' \WATER\DATA\MEAND.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=20,FILE=' \WATER\DATA\PHID.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=32,FILE=' \WATER\DATA\MEANW.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=22,FILE=' \WATER\DATA\PHIW.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=24,FILE=' \WATER\DATA\SEQ2.DAT',STATUS='OLD')
      OPEN (UNIT=18,FILE=' \WATER\DATA\CLIMA.DAT',STATUS='OLD')

```

```

C ..... INPUT SEQ OF DRY & WET DAYS

```

```

      DO 50, I = 1, NY
        DO 60, J = 1, 2
          READ (24, 25) COUNT (J, I)
          READ (24, 15) (SEQ (J, I, K), K = 1, COUNT (J, I))
60      CONTINUE
50     CONTINUE

      DO 30, K = 1, NV
        DO 10, I = 1, NY
          DO 20, J = 1, NT

```

```

C ..... INPUT ONE VARIABLE AT A TIME

```

```

        IF (K .EQ. 1) THEN
            READ (18,95) CLIMA (I, J)
        ELSEIF (K .EQ. 2) THEN
            READ (18,35) CLIMA (I, J)
        ELSEIF (K .EQ. 3) THEN
            READ (18,45) CLIMA (I, J)
        ELSEIF (K .EQ. 4) THEN
            READ (18,55) CLIMA (I, J)
        ELSEIF (K .EQ. 5) THEN
            READ (18,65) CLIMA (I, J)
        ELSEIF (K .EQ. 6) THEN
            READ (18,75) CLIMA (I, J)
        ELSEIF (K .EQ. 7) THEN
            READ (18,85) CLIMA (I, J)
        ENDIF

20          CONTINUE
10          CONTINUE

C          ..... COMPUTE MEAN VECTOR FOR WET & DRY DAYS

        DO 310, M = 1, 2
            DO 320, J = 1, NT
                DENOM (J) = 0
                MU (M,J) = -999.0
320          CONTINUE
                DO 330, I = 1, NY
                    DO 370, J = 1, COUNT (M,I)
                        L = SEQ (M, I, J)
                        IF (CLIMA (I,L) .NE. -999) THEN
                            IF (MU (M,L) .LE. -900) THEN
                                MU (M,L) = 0.0
                            ENDIF
                            MU (M,L) = MU (M,L) + CLIMA (I,L)
                            DENOM (L) = DENOM (L) + 1
                        ENDIF
                    CONTINUE
370          CONTINUE
330          CONTINUE
                DO 380, J = 1, NT
                    IF (MU (M,J) .NE. -999) THEN
                        MU (M,J) = MU (M,J) / DENOM (J)
                    ENDIF
380          CONTINUE
310          CONTINUE

        DO 130, M = 1, 2

        OMEGA = 2 * PI / NT
        KK = (NPARM - 1) / 2
        DO 510, T = 1, NT
            PHI (T,1) = 1
510          CONTINUE
            DO 520, J = 1, KK
                J1 = 2 * J

```

```

      J2 = J1 + 1
      THETA = OMEGA * J
      A = 2 * COS (THETA)
      PHI (1,J1) = 1
      PHI (2,J1) = A / 2
      PHI (1,J2) = 0
      PHI (2,J2) = SIN (THETA)
      DO 530, T = 3, NT
        PHI (T,J1) = A * PHI (T-1,J1) - PHI (T-2,J1)
        PHI (T,J2) = A * PHI (T-1,J2) - PHI (T-2,J2)
530      CONTINUE
520      CONTINUE

C      ..... SHRINK MEAN & FOURIER VECTOR BY OMITTING MISSING OBSNS.

      NC = 0
      DO 120, I = 1, NT
        IF (MU (M,I) .NE. -999) THEN
          MU (M,I-NC) = MU (M,I)
          DO 140, L = 1, NPARM
            PHI (I-NC,L) = PHI (I,L)
140          CONTINUE
          ELSE
            NC = NC + 1
          ENDIF
120      CONTINUE

C      ..... OUTPUT OF MEAN & FOURIER VECTORS FOR DRY & WET DAYS

      IF (M .EQ. 1) THEN
        PRINT *, 'NO. OBNS ON DRY DAYS: ', NT - NC
        DO 80, I = 1, NT - NC
          WRITE (30, *) MU (M,I)
          WRITE (20, 5) (PHI (I,L), L = 1, NPARM)
80        CONTINUE
        ELSE
          PRINT *, 'NO. OBNS ON WET DAYS: ', NT - NC
          DO 40, I = 1, NT - NC
            WRITE (32, *) MU (M,I)
            WRITE (22, 5) (PHI (I,L), L = 1, NPARM)
40          CONTINUE
        ENDIF
130      CONTINUE
        REWIND 18
30      CONTINUE

      STOP
      END

```

PROGRAM 7  
-----

```
C -----
C ..... PROGRAM TO ESTIMATE PARAMETERS FOR THE MEAN AND
C STANDARD DEVIATION FUNCTIONS
C --> NB THIS PROGRAM IS DESIGNED TO ESTIMATE
C PARAMETERS FOR A DRY SEQUENCE. IT CAN ALSO
C BE USED FOR WET SEQUENCES BY READING THE
C APPROPRIATE INPUT DATA FILES
C -----
```

```

      INTEGER          NT,NV,NPARM,PI
      PARAMETER        (NT=365)
      PARAMETER        (NV=7)
      PARAMETER        (NPARM=11)
      PARAMETER        (PI=3.141593)
      REAL              LLK
      REAL              MU (NT,1)
      REAL              MEAN (NT)
      REAL              PHI (NT,NPARM)
      REAL              TRSP (NPARM,NT)
      REAL              SOLN (NPARM,NPARM)
      REAL              RESULT (NPARM,NT)
      REAL              BETA (NPARM,1)

15      FORMAT (/,/ )
25      FORMAT (F10.3)
35      FORMAT (A13,I4,A11)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=6,FILE='CON')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\MEAND.DAT',STATUS='OLD')
      OPEN (UNIT=14,FILE='\\WATER\\DATA\\PHID.DAT',STATUS='OLD')

      WRITE (9,*) 'INITIAL ESTIMATES FOR MEAN (DRY) FUNCTION'
      WRITE (9,15)

      WRITE (6,*) ' MAXIMUM NUMBER OF PARAMETERS TO BE FITTED'
      READ (6,*) NPT

      DO 30, K = 1, NV
        CRITO = 10**10
        DO 20, I = 1, NT
          READ (12,*) MU (I,1)
          READ (14,*) (PHI (I,L), L =1, NPT)
20      CONTINUE

        DO 100, NP = 1, NPT, 2
          CALL TRNSP (PHI,NP,NT,TRSP,NPARM,NT)
          CALL XNP (TRSP,PHI,SOLN,NP,NT,NT,NP,NPARM,NT,NT,NPARM)
          CALL INVNP (NP,SOLN,NPARM)
          CALL XNP (SOLN,TRSP,RESULT,NP,NP,NP,NT,NPARM,NPARM,
&              NPARM,NT)

```

```

      CALL XNP (RESULT,MU,BETA,NP,NT,NT,1,NP,NT,NT,1)

C      ..... OUTPUT OF PARAMETER ESTIMATES

      WRITE (9,*) 'BETA ESTIMATES FOR VARIABLE: ', K
      DO 10, I = 1, NP
        WRITE (9,*) BETA (I,1)
10      CONTINUE

      DO 50, I = 1, NT
        MEAN (I) = 0.0
        DO 60, L = 1, NP
          MEAN (I) = MEAN (I) + BETA(L,1) * PHI(I,L)
60      CONTINUE
50      CONTINUE
      LLK = 0
      DO 40, I = 1, NT
        LLK = LLK + (MU(I,1) - MEAN (I))**2
40      CONTINUE
      LLK = -LLK/2 - (NT/2) * LOG (2*PI)
      CRIT = -LLK + NP
      WRITE (9,*) ' AKAIKE"S INFO CRITERION FOR VARIABLE ',K
      WRITE (9,*) CRIT
      WRITE (9,35) ' WHEN FITTING' ,NP, ' PARAMETERS'
      WRITE (9,15)
      IF (CRIT .LT. CRITO) THEN
        LL = NP
        CRITO = CRIT
      ELSE
        GOTO 200
      ENDIF
100      CONTINUE
200      WRITE (9,*) ' NUMBER OF PARAMETERS CHOSEN: ', LL
30      CONTINUE

      STOP
      END

```

PROGRAM 8

```

C -----
C ..... PROGRAM TO COMPUTE THE STANDARD DEVIATION VECTOR
C -----

```

```

      INTEGER          NV,NY,NT,NPARM,NP
      PARAMETER        (NV=7)
      PARAMETER        (NY=12)
      PARAMETER        (NT=365)
      PARAMETER        (NPARM=3)
      PARAMETER        (NP=11)
      INTEGER          SEQ (2,NY,NT)
      INTEGER          COUNT (2,NY)
      REAL              MU (2,NT)
      REAL              SIGMA (2,NT)
      REAL              PHI (NT,NP)
      REAL              ALPHA (NV,2,NPARM)
      DIMENSION        CLIMA[HUGE] (NY,NT)
      REAL              DENOM (NT)

```

```

5      FORMAT (11F6.2)
15     FORMAT (14(I5))
25     FORMAT (I5)
35     FORMAT (18X,F9.2)
45     FORMAT (27X,F9.2)
55     FORMAT (36X,F9.2)
65     FORMAT (45X,F10.2)
75     FORMAT (55X,F10.2)
85     FORMAT (65X,F9.2)
95     FORMAT (9X,F9.2)

```

```

      OPEN (UNIT=40,FILE='\\WATER\\DATA\\SIGMAD.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=20,FILE='\\WATER\\DATA\\PHD.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=42,FILE='\\WATER\\DATA\\SIGMAW.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=22,FILE='\\WATER\\DATA\\PHW.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=24,FILE='\\WATER\\DATA\\SEQ2.DAT',STATUS='OLD')
      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\EST-M.DAT',STATUS='OLD')

```

```

C ..... INPUT OF PARAMETER ESTIMATES FOR MEAN FUNCTION

```

```

      DO 170, K = 1, NV
        DO 40, M= 1, 2
          READ (10,*) (ALPHA (K,M,I), I = 1, NPARM)
40      CONTINUE
170    CONTINUE

```

```

C ..... INPUT SEQ OF DRY & WET DAYS

```

```

      DO 50, I = 1, NY
        DO 60, J = 1, 2
          READ (24,25) COUNT (J,I)

```

```

        READ (24,15) (SEQ (J,I,K), K = 1, COUNT (J,I))
60      CONTINUE
50      CONTINUE

DO 30, K = 1, NV
  DO 10, I = 1, NY
    DO 20, J = 1, NT

C      ..... INPUT ONE VARIABLE AT A TIME

        IF (K .EQ. 1) THEN
          READ (18,95) CLIMA (I, J)
        ELSEIF (K .EQ. 2) THEN
          READ (18,35) CLIMA (I, J)
        ELSEIF (K .EQ. 3) THEN
          READ (18,45) CLIMA (I, J)
        ELSEIF (K .EQ. 4) THEN
          READ (18,55) CLIMA (I, J)
        ELSEIF (K .EQ. 5) THEN
          READ (18,65) CLIMA (I, J)
        ELSEIF (K .EQ. 6) THEN
          READ (18,75) CLIMA (I, J)
        ELSEIF (K .EQ. 7) THEN
          READ (18,85) CLIMA (I, J)
        ENDIF

20      CONTINUE
10      CONTINUE
        CALL TRIG (PHI, NP, NT)

C      ..... GENERATE MEAN VECTOR

        CALL GMEAN (MU, PHI, NT, NPARM, ALPHA, K, NV)

C      ..... COMPUTE STD DEV VECTOR FOR WET & DRY DAYS

DO 330, M = 1, 2
  DO 310, I = 1, NT
    DENOM (I) = 0
    SIGMA (M, I) = -999.0
310    CONTINUE
    DO 320, J = 1, NY
      DO 370, L = 1, COUNT (M, J)
        L = SEQ (M, J, I)
        IF (CLIMA (J, L) .NE. -999) THEN
          IF (SIGMA (M, L) .EQ. -999) THEN
            SIGMA (M, L) = 0.0
          ENDIF
          SIGMA (M, L) = SIGMA (M, L) + (CLIMA (J, L) - MU (M, L)) ** 2
          DENOM (L) = DENOM (L) + 1
        ENDIF
      CONTINUE
    CONTINUE
    DO 380, I = 1, NT

```

```

                IF (SIGMA (M,I) .NE. -999) THEN
                    SIGMA (M,I) = SQRT(SIGMA (M,I) / DENOM (I))
                ENDIF
380             CONTINUE
330         CONTINUE

C         ..... SHRINK STD DEV VECTOR & FOURIER VECTOR BY OMITTING MISSING
C         OBSERVATIONS

                DO 130, M = 1, 2
                    NC = 0
                    DO 120, I = 1, NT
                        IF (SIGMA (M,I) .NE. -999) THEN
                            SIGMA (M,I-NC) = SIGMA (M,I)
                            DO 140, L = 1, NP
                                PHI (I-NC,L) = PHI (I,L)
140                             CONTINUE
                        ELSE
                            NC = NC + 1
                        ENDIF
120                 CONTINUE

C         ..... OUTPUT OF STDE DEV & FOURIER VECTORS FOR WET & DRY DAYS

                IF (M .EQ. 1) THEN
                    PRINT *, 'NO. OBNS ON DRY DAYS: ', NT - NC
                    DO 70, I = 1, NT - NC
                        WRITE (40, *) SIGMA (M,I)
                        WRITE (20, 5) (PHI (I,L), L = 1, NP)
70                     CONTINUE
                ELSE
                    PRINT *, 'NO. OBNS ON WET DAYS: ', NT - NC
                    DO 80, I = 1, NT - NC
                        WRITE (42, *) SIGMA (M,I)
                        WRITE (22, 5) (PHI (I,L), L = 1, NP)
80                     CONTINUE
                ENDIF
                CALL TRIG (PHI,NP,NT)
130             CONTINUE
                REWIND 18
30         CONTINUE

        STOP
        END

```

# PROGRAM 9

```

-----
C      -----
C      ..... PROGRAM TO STANDARDIZE RESIDUAL TIME SERIES
C      -----

      INTEGER          NV,NY,NT,NPARM
      PARAMETER        (NV=7)
      PARAMETER        (NY=12)
      PARAMETER        (NT=365)
      PARAMETER        (NPARM=3)
      REAL              CLIMA (0:NV)
      REAL              MU (2,NV,NT)
      REAL              SIGMA (2,NV,NT)
      REAL              PHI (NT,NPARM)
      REAL              PSI (NV,2,NPARM)
      REAL              ALPHA (NV,2,NPARM)

5      FORMAT (3F 10.6)
35     FORMAT (5F9.2,2F10.2,F9.2)

      OPEN (UNIT=12,FILE='\\WATER\\DATA\\CLIMAR.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\EST-M.DAT',STATUS='OLD')
      OPEN (UNIT=40,FILE='\\WATER\\DATA\\EST-S.DAT',STATUS='OLD')

C      ..... INPUT OF PARAMETER ESTIMATES FOR MEAN AND STANDARD
C      DEVIATION FUNCTION

      DO 10, K = 1, NV
        DO 20, M= 1, 2
          READ (10,*) (ALPHA (K,M,I), I = 1, NPARM)
20      CONTINUE
10     CONTINUE

      DO 60, K = 1, NV
        DO 70, M= 1, 2
          READ (40,*) (PSI (K,M,I), I = 1, NPARM)
70      CONTINUE
60     CONTINUE

      CALL TRIG (PHI,NPARM,NT)

C      ..... GENERATE MEAN AND STANDARD DEVIATION VECTORS

      CALL GAVSTD (MU,PHI,NT,NPARM,ALPHA,NV,PSI,SIGMA)

      DO 30, I = 1, NY
        DO 40, J = 1, NT
          READ (18,35) (CLIMA (K), K = 0, NV)
          IF (CLIMA(0) .EQ. 0) THEN
            M = 1
          ELSE

```

```

        M = 2
        ENDIF
        DO 50 , K = 1, NV
            IF (CLIMA (K) .NE. -999) THEN
                CLIMA(K)=(CLIMA(K)-MU(M,K,J))/SIGMA(M,K,J)
            ENDIF
50      CONTINUE

C      ..... OUTPUT OF STANDARDIZED TIME SERIES

        WRITE (12,35) (CLIMA (K), K = 0, NV)
40      CONTINUE
30      CONTINUE

        STOP
        END

```

# PROGRAM 10

```

C      -----
C      ..... PROGRAM TO COMPUTE CROSS-CORRELATION COEFFICIENTS
C      FOR LAGO AND LAG1.
C      -----

        INTEGER          NY, NT, NV
        PARAMETER        (NY=12)
        PARAMETER        (NT=365)
        PARAMETER        (NV=7)
        INTEGER          DENOM(0:1)
        REAL              CLIMA(NV,NY*NT)
        REAL              CROSS(0:1)
        REAL              AVEG(NV),DEV(NV)
        REAL              CLAGO(NV,NV)
        REAL              CLAG1(NV,NV)

25      FORMAT (9F8.4)
35      FORMAT (9X,4F9.2,2F10.2,F9.2)
45      FORMAT (7F9.3)

        OPEN (UNIT=9,FILE='LPT1')
        OPEN (UNIT=10,FILE='\\WATER\\DATA\\CLIMAR.DAT',STATUS='OLD')
        OPEN (UNIT=20,FILE='\\WATER\\DATA\\LAGO.DAT',STATUS='UNKNOWN')
        OPEN (UNIT=30,FILE='\\WATER\\DATA\\LAG1.DAT',STATUS='UNKNOWN')
        OPEN (UNIT=4,FILE='CON')

```

```

        NTIME = NY * NT
        DO 10, I = 1, NTIME
            READ (10,35) (CLIMA (K,I), K = 1, NV)
10        CONTINUE

        CALL AVSTD3 (CLIMA,AVEG,DEV,NY,NT,NV)

        DO 20, K = 1, NV
            DO 30, KK = 1, NV
                DO 40, I = 0, 1
                    CROSS (I) = 0.0
                    DENOM (I) = 0.0
40                CONTINUE

                DO 50, I = 0, 1
                    DO 60, J = 1, NTIME-I
                        IF ((CLIMA(K,J).GT.-900).AND.(CLIMA(KK,J+I)
&                        .GT.-900)) THEN
                            CROSS(I)=CROSS(I)+((CLIMA(K,J)-AVEG(K))*
&                            (CLIMA(KK,J+I)-AVEG(KK)))
                            DENOM(I) = DENOM(I) + 1
                        ENDIF
60                    CONTINUE
                    IF (DENOM(I).GT.0) THEN
                        CROSS(I)=(CROSS(I)/DENOM(I))/((DEV(K)*DEV(KK))
                    ENDIF
50                CONTINUE

                CLAGO(K,KK) = CROSS(0)
                CLAG1(K,KK) = CROSS(1)

30            CONTINUE
20        CONTINUE

        DO 90, K = 1, NV
            WRITE (20,45) (CLAGO(K,KK), KK = 1, NV)
            WRITE (30,45) (CLAG1(K,KK), KK = 1, NV)
90        CONTINUE

        STOP
        END

```

PROGRAM 11

```

C -----
C ..... PROGRAM TO COMPUTE THE MATRICES A & B FOR MODEL1
C -----

      INTEGER          NV
      PARAMETER        (NV=7)
      REAL              CLAGO(NV,NV)
      REAL              CLAG1(NV,NV)
      REAL              A(NV,NV)
      REAL              B(NV,NV)
      REAL              INV(NV,NV)
      REAL              TRSP(NV,NV)
      REAL              TERM(NV,NV)

15  FORMAT (7F9.3)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=20,FILE='\\WATER\\DATA\\LAGO.DAT',STATUS='OLD')
      OPEN (UNIT=30,FILE='\\WATER\\DATA\\LAG1.DAT',STATUS='OLD')
      OPEN (UNIT=40,FILE='\\WATER\\DATA\\A.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=50,FILE='\\WATER\\DATA\\B.DAT',STATUS='UNKNOWN')

      DO 10, K = 1, NV
        READ (20,15) (CLAGO (K,KK), KK = 1, NV)
10  CONTINUE
      DO 20, K = 1, NV
        READ (30,15) (CLAG1 (K,KK), KK = 1, NV)
20  CONTINUE

      CALL INVT (CLAGO,INV,NV)
      CALL MULT (CLAG1,INV,A,NV,NV,NV,NV)

      DO 30, K = 1, NV
        WRITE (40,15) (A(K,KK), KK = 1, NV)
30  CONTINUE

      CALL TRANSP (CLAG1,NV,NV,TRSP)
      CALL MULT (A,TRSP,TERM,NV,NV,NV,NV)
      CALL SUBTR (CLAGO,TERM,NV)
      CALL CHOLKY (B,TERM,NV)

      DO 40, K = 1, NV
        WRITE (50,15) (B(K,KK), KK = 1, NV)
40  CONTINUE

      STOP
      END

```

PROGRAM 12

```

C -----
C ..... PROGRAM TO GENERATE CLIMATE SEQUENCES ACCORDING TO
C      MODEL 1.
C -----

C ..... INTEGER VARIABLES

      INTEGER          NT,NV,NY,NP,PSTATE,STATE

C ..... PSTATE = PRESENT STATE OF DAY
C ..... STATE = PREVIOUS STATE OF DAY

C ..... PARAMETER STATEMENTS

      PARAMETER        (NT=365)
      PARAMETER        (NY=51)
      PARAMETER        (NV=7)
      PARAMETER        (NP=3)

C ..... NT = £ OBSERVATIONS PER YEAR
C ..... NV = £ VARIABLES
C ..... NY = £ YEARS TO BE GENERATED
C ..... NP = £ PARAMETERS IN SEASONAL MODEL

      INTEGER          SEED (9)
      REAL             RAIN
      REAL             GAM (2,NP)
      REAL             PHI (NP,0:NT)
      REAL             RAND (NV,1)
      REAL             SIGMA (2,NV,0:NT)
      REAL             MU (2,NV,0:NT)
      REAL             OBSN (NV,1)
      REAL             RES (NV,1)
      REAL             TEMP (NV)
      REAL             AMP (0:NP)
      REAL             PHASE (NP)
      REAL             A (NV,NV)
      REAL             B (NV,NV)
      REAL             C (NT)
      COMMON           IDUM1,IDUM2,IDUM3,IDUM4,IDUM5,IDUM6,IDUM7

15  FORMAT (5F9.2, 2F10.2, F9.2)
25  FORMAT (' GIVE 9 -VE Nos. TO INITIALIZE RANDOM GENERATOR',/)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\SIMU.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=22,FILE='CON')

C ..... COMPUTE THE FOURIER SERIES TERMS

```

```

        CALL COSSIN (PHI, NP, NT)
        DO 60, I = 1, NP
            PHI (I, 0) = PHI (I, NT)
60      CONTINUE

        PI=3.14159
        SMAX=135
        SMIN=110
        AVE=(SMAX+SMIN)/2
        AMPS=SMAX-SMIN
        DO 330, I = 1, NT
            C(I) = AVE+(AMPS/2)*COS((2*PI/NT)*(I+11))
330     CONTINUE

C       ..... READING PARAMETER ESTIMATES

        CALL DATA1 (GAM, MU, SIGMA, NP, NV, AMP, PHASE, CV, PHI, A, B, NT)

C       ..... INPUT SEEDS TO START RANDOM NUMBER GENERATOR.  MUST BE
C       NEGATIVE NUMBER.

        PRINT 25
        DO 50, II = 1, 9
            READ (22, *) SEED (II)
50      CONTINUE
        IDUM1 = SEED (1)
        IDUM2 = SEED (2)
        IDUM3 = SEED (3)
        IDUM4 = SEED (4)
        IDUM5 = SEED (5)
        IDUM6 = SEED (6)
        IDUM7 = SEED (7)
        IDUM8 = SEED (8)
        IDUM9 = SEED (9)

C       ..... COMPUTE PARAMETERS NEEDED FOR COMPUTATION OF RAINFALL
C       DEPTH

        CALL CALBET (BETA, CV)
        ALPH = 1 + 1 / BETA
        GAMM = GAMMA (ALPH)
        BI = 1 / BETA
        W = 0.01721421

C       ..... SET INITIAL STATE OF DAY TO BE DRY
C       ..... SET INITIAL CLIMATE VALUE TO IT'S MEAN AT TIME ZERO

C       ..... STATE = 1 ==> DRY
C       ..... STATE = 2 ==> WET

        STATE = 1
        DO 10, I = 1, NV
            OBSN (I, 1) = 0.0
10      CONTINUE

```

```

DO 30, I = 1, NY
  DO 40, J = 1, NT

C      ..... GENERATE RAINFALL VALUE
C      -----

C      ..... COMPUTE PROBABILITY THAT A WET DAY FOLLOWS A WET DAY, OR
C      THE PROBABILITY THAT A WET DAY FOLLOWS A DRY DAY.

          CALL PIEST (NP,GAM,STATE,J,PHI,PI,NT)

C      ..... GENERATE A UNIFORM RANDOM NUMBER BETWEEN 0 AND 1.

          UNIFOR = URAN8 (IDUM8)
          IF (UNIFOR .LT. PI) THEN
            PSTATE = 2
          ELSE
            PSTATE = 1
          ENDIF

C      ..... GENERATE A NORMAL RANDOM NUMBER

          CALL GRAND2 (RAND,NV)

C      ..... GENERATE CLIMATE SEQUENCES
C      -----

          CALL MODEL1 (RAND,NV,SIGMA,MU,J,OBSN,PSTATE,NT,A,B,RES)

C      ..... DETERMINE WHETHER IT RAINED AND SET RAIN VALUE

C      ..... RAIN = 0 ==> DID NOT RAIN
C      ..... RAIN = 1 ==> RAINED

          IF (PSTATE .EQ. 1) THEN
            RAIN = 0
          ELSE
            RAIN = 1
          ENDIF

C      ..... GENERATE RAINFALL DEPTH IF IT RAINED
C      -----

          IF (RAIN .EQ. 1) THEN
            CALL DEPTH3 (IDUM9,NP,RAIN,J,AMP,PHASE,GAMM,BI,W)
          ENDIF

C      ..... TRANSFORM VARIABLES TO THE ORIGINAL UNITS

          TEMP(2)=(230-100*EXP(RES(2,1)))/(EXP(RES(2,1))+1)
          TEMP(1)=(410+TEMP(2)*EXP(RES(1,1)))/(EXP(RES(1,1))+1)
          TEMP(3)=(C(J)-0.01-(0.01*EXP(RES(3,1))))/(EXP(RES(3,1))+1)
          TEMP(4)=(10000/(EXP(RES(4,1))+1))-0.01

```

```

        TEMP(5)=101/(EXP(RES(5,1))+1)
        TEMP(6)=TEMP(5)/(EXP(RES(6,1))+1)
C      ..... OUTPUT GENERATED SEQUENCES

        IF (I.NE.1) THEN
            WRITE (10,15) RAIN, (TEMP (K), K = 1, NV)
        ENDIF

C      ..... UPDATE THE STATE OF THE PREVIOUS DAY

        IF (PSTATE .NE. STATE) THEN
            STATE = PSTATE
        ENDIF

40      CONTINUE
30      CONTINUE

        STOP
        END

```

# PROGRAM 13

```

C      -----
C      ..... PROGRAM TO PREPARE DATA SETS OF POSSIBLE WET/DRY
C      SEQUENCES
C      -----

        INTEGER          NY,NT,PREV,RAIN
        PARAMETER        (NY=12)
        PARAMETER        (NT=365)
        INTEGER          SEQ (4,NY,NT)
        INTEGER          COUNT (4,NY)
        REAL             CLIMA

5      FORMAT (F9.2)
15     FORMAT (14(I5))
25     FORMAT (I5)

        PREV = 0
        DO 20, J = 1, 4
            DO 40, I = 1, NY
                COUNT(J,I) = 0
40      CONTINUE
20     CONTINUE

        OPEN (UNIT=8,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')

```

```

DO 10, J = 1, NY
  DO 50, I = 1, NT
    READ (8,5) CLIMA
    IF (CLIMA .EQ. 0) THEN
      RAIN = 0
    ELSEIF (CLIMA .GT. 0) THEN
      RAIN = 1
    ELSEIF (CLIMA .EQ. -999) THEN
      RAIN = 2
    ENDIF
    IF ((RAIN .NE. 2) .AND. (PREV .NE. 2)) THEN
      IF (RAIN .EQ. PREV) THEN
        IF (RAIN .EQ. 0) THEN
          COUNT(1,J) = COUNT(1,J) + 1
          SEQ(1,J,COUNT(1,J)) = I
        ELSE
          COUNT(2,J) = COUNT(2,J) + 1
          SEQ(2,J,COUNT(2,J)) = I
        ENDIF
      ELSE
        IF (RAIN .EQ. 0) THEN
          COUNT(4,J) = COUNT(4,J) + 1
          SEQ(4,J,COUNT(4,J)) = I
        ELSE
          COUNT(3,J) = COUNT(3,J) + 1
          SEQ(3,J,COUNT(3,J)) = I
        ENDIF
      ENDIF
    ENDIF
    PREV = RAIN
50    CONTINUE
10    CONTINUE

OPEN (UNIT=10,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='UNKNOWN')

DO 60, I = 1, NY
  DO 30, J = 1, 4
    WRITE (10, 25) COUNT(J,I)
    WRITE (10, 15) (SEQ(J,I,K), K = 1, COUNT(J,I))
30    CONTINUE
60    CONTINUE

STOP
END

```

PROGRAM 14

```

C -----
C ..... PROGRAM TO COMPUTE THE AUTOCORRELATION COEFFICIENT
C          CONDITIONED ON WET/DRY STATUS OF THE DAY
C -----

      INTEGER      NY,NV,NT,NPARM,P,PP,NRAU
      PARAMETER    (NRAU=4)
      PARAMETER    (NY=12)
      PARAMETER    (NV=7)
      PARAMETER    (NT=365)
      PARAMETER    (NPARM=3)
      INTEGER      COUNT (NRAU,NY)
      INTEGER      SEQ (NRAU,NY,NT)
      INTEGER      SUM (NRAU)
      REAL         CLIMA (NY,NT)
      REAL         MU (2,NT)
      REAL         PHI (NT,NPARM)
      REAL         ALPHA (NV,2,NPARM)
      REAL         RAU (NRAU)

      5  FORMAT (/)
      25  FORMAT (9X,F9.2)
      35  FORMAT (18X,F9.2)
      45  FORMAT (27X,F9.2)
      55  FORMAT (36X,F9.2)
      65  FORMAT (45X,F10.2)
      75  FORMAT (55X,F10.2)
      85  FORMAT (65X,F9.2)
      95  FORMAT (15)
      105  FORMAT (14I5)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\EST-M.DAT',STATUS='OLD')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='OLD')

C ..... INPUT SEQUENCE OF DRY/WET DAYS

      DO 140, J = 1, NY
        DO 150, I = 1, 4
          READ (12,95) COUNT (I,J)
          READ (12,105) (SEQ (I,J,K), K = 1, COUNT (I,J))
150      CONTINUE
140      CONTINUE

C ..... INPUT OF PARAMETER ESTIMATES FOR THE MEAN FUNCTION

      DO 170, K = 1, NV
        DO 90, M= 1, 2
          READ (10,*) (ALPHA (K,M,I), I = 1, NPARM)
90      CONTINUE

```

```

170    CONTINUE

      CALL TRIG (PHI,NPARM,NT)
      DO 220, JJ = 1, NRAU
        DO 230, I = 1, NY
          SUM (JJ) = SUM (JJ) + COUNT (JJ,I)
230      CONTINUE
220      CONTINUE

      DO 30, K = 1, NV
        DO 10, I = 1, NY
          DO 20, J = 1, NT

C      ..... INPUT ONE VARIABLE AT A TIME

          IF (K .EQ. 1) THEN
            READ (18,25) CLIMA (I,J)
          ELSEIF (K .EQ. 2) THEN
            READ (18,35) CLIMA (I,J)
          ELSEIF (K .EQ. 3) THEN
            READ (18,45) CLIMA (I,J)
          ELSEIF (K .EQ. 4) THEN
            READ (18,55) CLIMA (I,J)
          ELSEIF (K .EQ. 5) THEN
            READ (18,65) CLIMA (I,J)
          ELSEIF (K .EQ. 6) THEN
            READ (18,75) CLIMA (I,J)
          ELSEIF (K .EQ. 7) THEN
            READ (18,85) CLIMA (I,J)
          ENDIF
20      CONTINUE
10      CONTINUE

C      ..... GENERATE MEAN VECTOR

      CALL GMEAN (MU,PHI,NT,NPARM,ALPHA,K,NV)

C      ..... COMPUTE AUTOCORRELATION

      WRITE (9,*) 'INITIAL ESTIMATES FOR RAU OF VARIABLE: ', K

      DO 120, JJ = 1, 4

        NUM = 0
        DENOM = 0
        CNT = 0
        CNT2 = 0

        IF (JJ.EQ. 1) THEN
          M = 1
          L = 1
        ELSEIF (JJ .EQ. 2) THEN
          M = 2
          L = 2

```

```

ELSEIF (JJ .EQ. 3) THEN
    M = 2
    L = 1
ELSEIF (JJ .EQ. 4) THEN
    M = 1
    L = 2
ENDIF
DO 110, I = 1, NY
    DO 130, J = 1, COUNT (JJ,I)
        P = SEQ (JJ,I,J)
        N = 0
        IF ((P .EQ. 1) .AND. (I .EQ. 1)) THEN
            GOTO 130
        ENDIF
        IF ((P .EQ. 1) .AND. (I .GT. 1)) THEN
            N = 1
            PP = 365
        ELSE
            PP = SEQ (JJ,I,J) - 1
            N = 0
        ENDIF
        IF ((CLIMA(I,P).NE.-999).AND.(CLIMA(I-N,PP).NE.
&          -999)) THEN
&          NUM = NUM+(CLIMA(I,P)-MU(M,P))*(CLIMA(I-N,PP)-
&          MU(L,PP))
        ELSE
            CNT = CNT + 1
        ENDIF
        IF (CLIMA(I-N,PP).NE.-999) THEN
            DENOM = DENOM+(CLIMA(I-N,PP)-MU(L,PP))*2
        ELSE
            CNT2 = CNT2 + 1
        ENDIF
130      CONTINUE
110    CONTINUE

200    NUM = NUM/(SUM(JJ)-1-CNT)
        DENOM = DENOM/(SUM(JJ)-CNT2)
        RAU (JJ) = NUM / DENOM
        WRITE (9,*) RAU (JJ)
120    CONTINUE
        WRITE (9,5)
        REWIND 18
30    CONTINUE

STOP
END

```

PROGRAM 15

```

C -----
C ..... PROGRAM TO ESTIMATE INITIAL STANDARD DEVIATION FUNCTION
C      -- MEAN VECTOR GENERATED
C -----

      INTEGER          NY,NV,NT,NPARM,P,PP,NRAU
      PARAMETER        (NRAU=4)
      PARAMETER        (NY=12)
      PARAMETER        (NV=7)
      PARAMETER        (NT=365)
      PARAMETER        (NPARM=3)
      INTEGER          COUNT (NRAU,NY)
      INTEGER          SEQ (NRAU,NY,NT)
      INTEGER          SUM (NRAU)
      REAL             SIGMA (NRAU)
      REAL             CLIMA (NY,NT)
      REAL             MU (2,NT)
      REAL             PHI (NT,NPARM)
      REAL             ALPHA (NV,2,NPARM)
      REAL             RAU (NRAU,NV)

5      FORMAT (/)
15     FORMAT (F9.2)
25     FORMAT (9X,F9.2)
35     FORMAT (18X,F9.2)
45     FORMAT (27X,F9.2)
55     FORMAT (36X,F9.2)
65     FORMAT (45X,F10.2)
75     FORMAT (55X,F10.2)
85     FORMAT (65X,F9.2)
95     FORMAT (I5)
105    FORMAT (14I5)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\EST-M.DAT',STATUS='OLD')
      OPEN (UNIT=14,FILE='\\WATER\\DATA\\RAU-M.DAT',STATUS='OLD')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='OLD')

C ..... INPUT SEQUENCE OF DRY/WET DAYS

      DO 140, J = 1, NY
        DO 150, I = 1, 4
          READ (12,95) COUNT (I,J)
          READ (12,105) (SEQ (I,J,K), K = 1, COUNT (I,J))
150      CONTINUE
140     CONTINUE

C ..... INPUT OF PARAMETER ESTIMATES FOR THE MEAN FUNCTION

      DO 170, K = 1, NV

```

```

          DO 90, M= 1, 2
            READ (10,*) (ALPHA (K,M,I), I = 1, NPARM)
90        CONTINUE
170      CONTINUE

C      ..... INPUT OF PARAMETER ESTIMATES FOR RAU

          DO 470, K = 1, NV
            READ (14,*) (RAU (I,K), I = 1, NRAU)
470      CONTINUE

          CALL TRIG (PHI,NPARM,NT)
          DO 220, JJ = 1, 4
            DO 230, I = 1, NY
              SUM (JJ) = SUM (JJ) + COUNT (JJ,I)
230      CONTINUE
220      CONTINUE

          DO 30, K = 1, NV
            DO 10, I = 1, NY
              DO 20, J = 1, NT

C      ..... INPUT ONE VARIABLE AT A TIME

              IF (K .EQ. 1) THEN
                READ (18,25) CLIMA (I, J)
              ELSEIF (K .EQ. 2) THEN
                READ (18,35) CLIMA (I, J)
              ELSEIF (K .EQ. 3) THEN
                READ (18,45) CLIMA (I, J)
              ELSEIF (K .EQ. 4) THEN
                READ (18,55) CLIMA (I, J)
              ELSEIF (K .EQ. 5) THEN
                READ (18,65) CLIMA (I, J)
              ELSEIF (K .EQ. 6) THEN
                READ (18,75) CLIMA (I, J)
              ELSEIF (K .EQ. 7) THEN
                READ (18,85) CLIMA (I, J)
              ENDIF
20      CONTINUE
10      CONTINUE

C      ..... GENERATE MEAN VECTOR

          CALL GMEAN (MU,PHI,NT,NPARM,ALPHA,K,NV)

C      ..... COMPUTE STANDARD DEVIATIONS

          WRITE (9,*) 'INITIAL ESTIMATES FOR SIGMA OF VARIABLE: ', K
          DO 120, JJ = 1, 4
            IF (JJ.EQ. 1) THEN
              M = 1
              L = 1
            ELSEIF (JJ .EQ. 2) THEN

```

```

        M = 2
        L = 2
    ELSEIF (JJ .EQ. 3) THEN
        M = 2
        L = 1
    ELSEIF (JJ .EQ. 4) THEN
        M = 1
        L = 2
    ENDIF
    NUM = 0
    CNT = 0
    DO 110, I = 1, NY
        DO 130, J = 1, COUNT (JJ,I)
            P = SEQ (JJ,I,J)
            N = 0
            IF ((P .EQ. 1) .AND. (I .EQ. 1)) THEN
                CNT = CNT + 1
                GOTO 130
            ENDIF
            IF ((P .EQ. 1) .AND. (I .GT. 1)) THEN
                N = 1
                PP = 365
            ELSE
                PP = SEQ (JJ,I,J)-1
                N = 0
            ENDIF
            IF ((CLIMA(I,P).NE.-999).AND.(CLIMA(I-N,PP).NE.
&                -999)) THEN
&                NUM = NUM+(CLIMA(I,P)-MU(M,P)-RAU(JJ,K)*(CLIMA
&                (I-N,PP)-MU(L,PP)))*2
            ELSE
                CNT = CNT + 1
            ENDIF
130        CONTINUE
110    CONTINUE

200    SIGMA (JJ) = SQRT(NUM/(SUM(JJ)-1-CNT))
    WRITE (9,*) SIGMA (JJ)
120    CONTINUE
    WRITE (9,5)
    REWIND 18
30    CONTINUE

    STOP
    END

```

PROGRAM 16

```

-----
C      ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR MODEL 3
C      -----

      INTEGER      NV,NY,NT,NP,NPARM,NRAU,CONVG,T
      PARAMETER    (NV=7)
      PARAMETER    (NY=12)
      PARAMETER    (NT=365)
      PARAMETER    (NP=14)
      PARAMETER    (NPARM=3)
      PARAMETER    (NRAU=4)
      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      REAL         RESID (NV,NY,NT)
      REAL         MU (2,0:NT)
      REAL         LNLIKE,AKAIKE,PI
      PARAMETER    (PI=3.141593)
      REAL         CLIMA (NY,0:NT)
      REAL         ALPHA (2,NV,NPARM)
      REAL         SIGMA (NRAU,NV)
      REAL         DER (NP)
      REAL         DER2 (NP,NP)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)
      REAL         THETA (NP)
      REAL         A (NP,0:NP)

5      FORMAT (/)
15     FORMAT (' ESTIMATES OF MEAN FOR DRY DAYS:', 3F10.4)
25     FORMAT (' ESTIMATES OF MEAN FOR WET DAYS:', 3F10.4)
35     FORMAT (' ESTIMATES OF STANDARD DEVIATIONS:', 4F10.4)
55     FORMAT (' ESTIMATES OF AUTOCORRELATION:', 4F10.4)
65     FORMAT (7F10.4)
75     FORMAT (' AKAIKE"S CRITERION FOR VARIABLE:', I4, ' IS:', F30.4)
85     FORMAT (9X,F9.2)
95     FORMAT (18X,F9.2)
105    FORMAT (27X,F9.2)
115    FORMAT (36X,F9.2)
125    FORMAT (45X,F10.2)
135    FORMAT (55X,F10.2)
145    FORMAT (65X,F9.2)
155    FORMAT (I5)
165    FORMAT (14I5)

      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='OLD')
      OPEN (UNIT=14,FILE='\\WATER\\DATA\\RESI.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=6,FILE='CON')

      CALL COSSIN (PHI,NPARM,NT)

```

```

DO 10, I = 1, NPARM
  PHI (I,0) = PHI (I,NT)
10  CONTINUE

C      ..... INPUT OF TIME SERIES AND INITIAL PARAMETER ESTIMATES

CALL INTAL3 (EPS,MAXITER,ALPHA,SIGMA,RAU,NPARM,NV,NRAU)
DO 20, K = 1, NV
  CONVG = 1
  DO 130, I = 1, NY
    DO 140, J = 1, NT
      IF (K.EQ.1) THEN
        READ (18,85) CLIMA (I, J)
      ELSEIF (K.EQ.2) THEN
        READ (18,95) CLIMA (I, J)
      ELSEIF (K.EQ.3) THEN
        READ (18,105) CLIMA (I, J)
      ELSEIF (K.EQ.4) THEN
        READ (18,115) CLIMA (I, J)
      ELSEIF (K.EQ.5) THEN
        READ (18,125) CLIMA (I, J)
      ELSEIF (K.EQ.6) THEN
        READ (18,135) CLIMA (I, J)
      ELSEIF (K.EQ.7) THEN
        READ (18,145) CLIMA (I, J)
      ENDIF
140    CONTINUE
      IF (I.EQ.1) THEN
        CLIMA (I,0) = CLIMA (I,1) - 0.5
      ELSEIF (I.NE.1) THEN
        CLIMA(I,0) = CLIMA(I-1,NT)
      ENDIF
130    CONTINUE

    REWIND 18
    IF (K .EQ. 1) THEN
      DO 150, KK = 1, NY
        DO 160, I = 1, 4
          READ (12,155) COUNT (I,KK)
          READ (12,165) (SEQ (I,KK,J), J = 1, COUNT (I,KK))
160        CONTINUE
150      CONTINUE
    ENDIF

C      ..... ITERATIVE ESTIMATION OF PARAMETERS

      CALL NEWT3 (ALPHA,SIGMA,RAU,NPARM,MAXITER,NT,NY,CLIMA,SEQ,
&              COUNT,DER,DER2,PHI,EPS,NP,NV,K,A,THETA,NRAU,CONVG)

C      ..... OUTPUT OF FINAL PARAMETER ESTIMATES

      WRITE (9,5)
      WRITE (9,15) (ALPHA (1,K,L), L = 1, NPARM)
      WRITE (9,25) (ALPHA (2,K,L), L = 1, NPARM)

```

```

WRITE (9,35) (SIGMA (J,K), J = 1, NRAU)
WRITE (9,55) (RAU (J,K), J = 1, NRAU)
WRITE (9,5)

```

C       ..... COMPUTE RESIDUAL MATRIX

```

IF (CONVG.EQ.1) THEN
  DO 30, M = 1, 2
    DO 40, I = 0, NT
      MU (M,I) = 0.0
      DO 50, L = 1, NPARM
        MU(M,I) = MU(M,I)+ALPHA(M,K,L)*PHI(L,I)
50      CONTINUE
40      CONTINUE
30      CONTINUE

  DO 60, I = 1, NY
    DO 70, J = 1, NT
      RESID (K,I,J) = -999.00
70      CONTINUE
60      CONTINUE
      LNLIKE = 0
      TERM = 0

  DO 80, J = 1, 4
    IF (J .EQ. 1) THEN
      M = 1
      L = 1
    ELSEIF (J .EQ. 2) THEN
      M = 2
      L = 2
    ELSEIF (J .EQ. 3) THEN
      M = 2
      L = 1
    ELSEIF (J .EQ. 4) THEN
      M = 1
      L = 2
    ENDIF
    DO 90, I = 1, NY
      DO 100, KK = 1, COUNT (J,I)
        T = SEQ (J,I,KK)
        IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1)
&          .NE.-999)) THEN
&          RESID(K,I,T)=(CLIMA(I,T)-MU(M,T))/SIGMA(J,K)
&          -RAU(J,K)*((CLIMA(I,T-1)-MU(L,T-1))/
&          SIGMA(J,K))
          LNLIKE = LNLIKE + (RESID(K,I,T))**2
        ENDIF
        TERM = TERM + LOG(SIGMA(J,K))
100      CONTINUE
90      CONTINUE
80      CONTINUE

      LNLIKE = -((NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2

```

```

          AKAIKE = -2*LNLIKE+2*NP
          WRITE (9,75) K, AKAIKE
        ENDIF
20      CONTINUE

        DO 110, I = 1, NY
          DO 120, T = 1, NT
            WRITE (14,65) (RESID (K,I,T), K = 1, NV)
120          CONTINUE
110        CONTINUE

        STOP
        END

```

PROGRAM 17

```

C -----
C ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR MODEL 3
C          USING CONJUGATE GRADIENT METHODS IN MULTIDIMENSIONS
C -----

      INTEGER      NV,NY,NT,NP,NPARM,NRAU,T
      PARAMETER    (NV=7)
      PARAMETER    (NY=12)
      PARAMETER    (NT=365)
      PARAMETER    (NP=14)
      PARAMETER    (NPARM=3)
      PARAMETER    (NRAU=4)
      REAL         THETA (NP)
      REAL         LNLIKE,AKAIKE,PI
      PARAMETER    (PI=3.141593)
      REAL         MU (2,0:NT)
      REAL         RESID (NV,NY,NT)

      COMMON       K,ICOUNT(NRAU,NY),ISEQ(NRAU,NY,NT),CLIMA(NY,0:NT),
&                ALPHA(2,NV,NPARM),SIGMA(NRAU,NV),PHI(NPARM,0:NT),
&                RAU(NRAU,NV),ISCALE(3,NV)

      5  FORMAT (/)
      15  FORMAT (' ESTIMATES OF MEAN FOR DRY DAYS:', 3F10.4)
      25  FORMAT (' ESTIMATES OF MEAN FOR WET DAYS:', 3F10.4)
      35  FORMAT (' ESTIMATES OF STANDARD DEVIATIONS:', 4F10.4)
      55  FORMAT (' ESTIMATES OF AUTOCORRELATION:', 4F10.4)
      65  FORMAT (' PARAMETER ESTIMATES FOR VARIABLE:', I4)
      75  FORMAT (' CONVERGE ACHIEVED IN ', I4, ' ITERATIONS')
      115  FORMAT (3X, F5.0)
      125  FORMAT (9X, F9.2)
      135  FORMAT (18X, F9.2)
      145  FORMAT (27X, F9.2)
      155  FORMAT (36X, F9.2)
      165  FORMAT (45X, F10.2)
      175  FORMAT (55X, F10.2)
      185  FORMAT (65X, F9.2)
      195  FORMAT (I5)
      205  FORMAT (14I5)
      215  FORMAT (7F10.4)
      315  FORMAT (' AKAIKE"S CRITERION FOR VARIABLE:', I4, ' IS:', F30.4)

      OPEN (UNIT=14,FILE='\\WATER\\DATA\\RESIT.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='OLD')
      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=6,FILE='CON')

```

```

CALL COSSIN (PHI,NPARM,NT)
DO 10, I = 1, NPARM
    PHI (I,0) = PHI (I,NT)
10 CONTINUE

C      ..... INPUT OF TIME SERIES AND INITIAL PARAMETER ESTIMATES

CALL INT3 (ALPHA,SIGMA,RAU,NPARM,NV,NRAU,ISCALE)

PRINT *, 'WHICH VARIABLE TO BE ESTIMATED?'
READ (6,*) K

DO 20, I = 1, NY
    DO 30, J = 1, NT
        IF (K .EQ. 1) THEN
            READ (18,125) CLIMA (I, J)
        ELSEIF (K .EQ. 2) THEN
            READ (18,135) CLIMA (I, J)
        ELSEIF (K .EQ. 3) THEN
            READ (18,145) CLIMA (I, J)
        ELSEIF (K .EQ. 4) THEN
            READ (18,155) CLIMA (I, J)
        ELSEIF (K .EQ. 5) THEN
            READ (18,165) CLIMA (I, J)
        ELSEIF (K .EQ. 6) THEN
            READ (18,175) CLIMA (I, J)
        ELSEIF (K .EQ. 7) THEN
            READ (18,185) CLIMA (I, J)
        ENDIF
30 CONTINUE
        CLIMA (I,0) = CLIMA (I,1) - 0.5
20 CONTINUE

    DO 40, KK = 1, NY
        DO 50, I = 1, 4
            READ (12,195) ICOUNT (I,KK)
            READ (12,205) (ISEQ (I,KK,J), J = 1, ICOUNT (I,KK))
50 CONTINUE
40 CONTINUE

C      ..... ITERATIVE ESTIMATION OF PARAMETERS

WRITE (9,65) K

C      ..... TRANSFORMING THE SEPARATE PARAMETER ARRAYS INTO ONE
C      VECTOR

DO 60, J = 1, NPARM
    THETA (J) = ALPHA (1,K,J)*ISCALE(3,K)
    THETA (J+3) = ALPHA (2,K,J)*ISCALE(3,K)
60 CONTINUE
    DO 70, J = 1, NRAU
        THETA (J+6) = SIGMA (J,K)*ISCALE(2,K)
        THETA (J+10) = RAU (J,K)*ISCALE(1,K)

```

```

70    CONTINUE
      CALL POLRIB (THETA,NP,TOL,ITER,FMIN)

C      ..... UPDATE PARAMETER ESTIMATES

      DO 80 J = 1, NPARM
        ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
        ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
80    CONTINUE

      DO 90, J = 1, NRAU
        SIGMA (J,K) = THETA (J+6)/ISCALE(2,K)
        RAU (J,K) = THETA (J+10)/ISCALE(1,K)
90    CONTINUE

      WRITE (9,75) ITER

C      ..... OUTPUT OF FINAL PARAMETER ESTIMATES

      WRITE (9,5)
      WRITE (9,15) (ALPHA (1,K,L), L = 1, NPARM)
      WRITE (9,25) (ALPHA (2,K,L), L = 1, NPARM)
      WRITE (9,35) (SIGMA (J,K), J = 1, NRAU)
      WRITE (9,55) (RAU (J,K), J = 1, NRAU)
      WRITE (9,5)

C      ..... COMPUTE RESIDUAL MATRIX

      DO 100, M = 1, 2
        DO 120, I = 0, NT
          MU (M,I) = 0.0
          DO 130, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
130        CONTINUE
120      CONTINUE
100    CONTINUE

      DO 180, I = 1, NY
        DO 190, J = 1, NT
          RESID (K,I,J) = -999.00
190      CONTINUE
180    CONTINUE
      LNLIKE = 0
      TERM = 0

      DO 140, J = 1, 4
        IF (J .EQ. 1) THEN
          M = 1
          L = 1
        ELSEIF (J .EQ. 2) THEN
          M = 2

```

```

      L = 2
    ELSEIF (J .EQ. 3) THEN
      M = 2
      L = 1
    ELSEIF (J .EQ. 4) THEN
      M = 1
      L = 2
    ENDIF
    DO 150, I = 1, NY
      DO 160, KK = 1, ICOUNT (J,I)
        T = ISEQ (J,I,KK)
        IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
&          THEN
&            RESID(K,I,T)=(CLIMA(I,T)-MU(M,T))/SIGMA(J,K)-RAU(J,K)
&              *((CLIMA(I,T-1)-MU(L,T-1))/SIGMA(J,K))
            LNLIKE = LNLIKE + (RESID(K,I,T))**2
          ENDIF
          TERM = TERM + LOG(SIGMA(J,K))
160      CONTINUE
150    CONTINUE
140  CONTINUE

    LNLIKE = -((NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2
    AKAIKE = -2*LNLIKE+2*NP

    WRITE (9,315) K, AKAIKE

    DO 200, I = 1, NY
      DO 170, T = 1, NT
        WRITE (14,215) (RESID (K,I,T), K = 1, NV)
170      CONTINUE
200    CONTINUE

    STOP
    END

```

PROGRAM 18

```

C -----
C ..... PROGRAM TO COMPUTE THE AUTOCORRELATION COEFFICIENT
C      -- FOR UNCONDITIONED DATA SET
C -----

      INTEGER          NY,NV,NT,NPARM
      PARAMETER        (NY=12)
      PARAMETER        (NV=7)
      PARAMETER        (NT=365)
      PARAMETER        (NPARM=3)
      REAL              RAIN (NY*NT)
      REAL              CLIMA (NY*NT)
      REAL              MU (2,NT)
      REAL              PHI (NT,NPARM)
      REAL              ALPHA (NV,2,NPARM)

      5  FORMAT (/)
      15 FORMAT (F9.2)
      25 FORMAT (9X,F9.2)
      35 FORMAT (18X,F9.2)
      45 FORMAT (27X,F9.2)
      55 FORMAT (36X,F9.2)
      65 FORMAT (45X,F10.2)
      75 FORMAT (55X,F10.2)
      85 FORMAT (65X,F9.2)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\EST-M.DAT',STATUS='OLD')

C ..... INPUT OF RAINFALL DATA

      DO 40, J = 1, NY*NT
      READ (18, 15) RAIN(J)
      40  CONTINUE
      REWIND 18

C ..... INPUT OF PARAMETER ESTIMATES FOR THE MEAN FUNCTION

      DO 170, K = 1, NV
      DO 90, M= 1, 2
      READ (10,*) (ALPHA (K,M,I), I = 1, NPARM)
      90  CONTINUE
      170 CONTINUE

      CALL TRIG (PHI,NPARM,NT)
      DO 30, K = 1, NV
      DO 10, I = 1, NY*NT

C ..... INPUT ONE VARIABLE AT A TIME

```

```

        IF (K .EQ. 1) THEN
            READ (18,25) CLIMA (I)
        ELSEIF (K .EQ. 2) THEN
            READ (18,35) CLIMA (I)
        ELSEIF (K .EQ. 3) THEN
            READ (18,45) CLIMA (I)
        ELSEIF (K .EQ. 4) THEN
            READ (18,55) CLIMA (I)
        ELSEIF (K .EQ. 5) THEN
            READ (18,65) CLIMA (I)
        ELSEIF (K .EQ. 6) THEN
            READ (18,75) CLIMA (I)
        ELSEIF (K .EQ. 7) THEN
            READ (18,85) CLIMA (I)
        ENDIF
10      CONTINUE

C      ..... GENERATE MEAN VECTOR

        CALL GMEAN (MU,PHI,NT,NPARM,ALPHA,K,NV)

C      ..... COMPUTE AUTOCORRELATION

        CNT = 0
        NUM = 0
        DENOM = 0
        COUNT = 0
        COUNT2 = 0
        DO 20, J = 2, NY*NT
            IF (RAIN(J).EQ.0) THEN
                M = 1
            ELSE
                M = 2
            ENDIF
            IF (J.GT.NT*(CNT+1)) THEN
                CNT = CNT + 1
            ENDIF
            I = J-NT*CNT
            IF (I .EQ. 1) THEN
                II = 365
            ELSE
                II = I-1
            ENDIF
            IF (RAIN (J-1).EQ.0) THEN
                L = 1
            ELSE
                L = 2
            ENDIF
            IF ((CLIMA(J).NE.-999).AND.(CLIMA(J-1).NE.-999)) THEN
                NUM = NUM+(CLIMA(J)-MU(M,I))*(CLIMA(J-1)-MU(L,II))
            ELSE
                COUNT = COUNT + 1
            ENDIF
            IF (CLIMA(J-1).NE.-999) THEN

```

```

        DENOM = DENOM+(CLIMA(J-1)-MU(L,II))**2
    ELSE
        COUNT2 = COUNT2 + 1
    ENDIF

20    CONTINUE

    IF (RAIN (NY*NT).EQ.0) THEN
        L = 1
    ELSE
        L = 2
    ENDIF

    IF (CLIMA(NY*NT).NE.-999) THEN
        DENOM = DENOM+(CLIMA(NY*NT)-MU(L,NT))**2
    ELSE
        COUNT2 = COUNT2 + 1
    ENDIF

    NUM = NUM/(NY*NT-1-COUNT)
    DENOM = DENOM/(NY*NT-COUNT2)
    RAU = NUM / DENOM
    WRITE (9,*) 'INITIAL ESTIMATE FOR RAU OF VARIABLE: ', K
    WRITE (9,*) RAU
    WRITE (9,5)
    REWIND 18

30    CONTINUE

    STOP
    END

```

PROGRAM 19

```

C -----
C ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR MODEL 4
C -----

      INTEGER      NV,NY,NT,NP,NPARM,NRAU,CONVG,T
      PARAMETER    (NV=7)
      PARAMETER    (NY=12)
      PARAMETER    (NT=365)
      PARAMETER    (NP=13)
      PARAMETER    (NPARM=3)
      PARAMETER    (NRAU=1)
      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      REAL         CLIMA (NY,0:NT)
      REAL         ALPHA (2,NV,NPARM)
      REAL         PSI (2,NV,NPARM)
      REAL         DER (NP)
      REAL         DER2 (NP,NP)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)
      REAL         THETA (NP)
      REAL         A (NP,0:NP)

5      FORMAT (/)
15     FORMAT (' ESTIMATES OF MEAN FOR DRY DAYS:', 3F10.4)
25     FORMAT (' ESTIMATES OF MEAN FOR WET DAYS:', 3F10.4)
35     FORMAT (' ESTIMATES OF VAR FOR DRY DAYS:', 3F10.4)
45     FORMAT (' ESTIMATES OF VAR FOR WET DAYS:', 3F10.4)
55     FORMAT (' ESTIMATE OF AUTOCORRELATION:', F10.4)
65     FORMAT (9X,F9.2)
75     FORMAT (18X,F9.2)
85     FORMAT (27X,F9.2)
95     FORMAT (36X,F9.2)
105    FORMAT (45X,F10.2)
115    FORMAT (55X,F10.2)
125    FORMAT (65X,F9.2)
135    FORMAT (I5)
145    FORMAT (14I5)

      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='OLD')
      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=6,FILE='CON')

      CALL COSSIN (PHI,NPARM,NT)
      DO 10, I = 1, NPARM
        PHI (I,0) = PHI (I,NT)
10     CONTINUE

C ..... INPUT OF TIME SERIES AND INITIAL PARAMETER ESTIMATES

```

CALL INITIAL (EPS,MAXITER,ALPHA,PSI,RAU,NPARM,NV,NRAU)

DO 20, K = 1, NV

CONVG = 1

DO 30, I = 1, NY

DO 40, J = 1, NT

IF (K.EQ.1) THEN

READ (18,65) CLIMA (I, J)

ELSEIF (K.EQ.2) THEN

READ (18,75) CLIMA (I, J)

ELSEIF (K.EQ.3) THEN

READ (18,85) CLIMA (I, J)

ELSEIF (K.EQ.4) THEN

READ (18,95) CLIMA (I, J)

ELSEIF (K.EQ.5) THEN

READ (18,105) CLIMA (I, J)

ELSEIF (K.EQ.6) THEN

READ (18,115) CLIMA (I, J)

ELSEIF (K.EQ.7) THEN

READ (18,125) CLIMA (I, J)

ENDIF

40 CONTINUE

IF (I.EQ.1) THEN

CLIMA (I,0) = CLIMA (I,1) - 0.5

ELSEIF (I.NE.1) THEN

CLIMA(I,0) = CLIMA(I-1,NT)

ENDIF

30 CONTINUE

REWIND 18

IF (K .EQ. 1) THEN

DO 50, KK = 1, NY

DO 60, I = 1, 4

READ (12,135) COUNT (I, KK)

READ (12,145) (SEQ (I, KK, J), J = 1, COUNT (I, KK))

60 CONTINUE

50 CONTINUE

ENDIF

C ..... ITERATIVE ESTIMATION OF PARAMETERS

CALL NEWT4 (ALPHA,PSI,RAU,NPARM,MAXITER,NT,NY,CLIMA,SEQ,  
& COUNT,DER,DER2,PHI,EPS,NP,NV,K,A,THETA,NRAU,CONVG)

C ..... OUTPUT OF FINAL PARAMETER ESTIMATES

WRITE (9,5)

WRITE (9,15) (ALPHA (1,K,L), L = 1, NPARM)

WRITE (9,25) (ALPHA (2,K,L), L = 1, NPARM)

WRITE (9,35) (PSI (1,K,L), L = 1, NPARM)

WRITE (9,45) (PSI (2,K,L), L = 1, NPARM)

WRITE (9,55) (RAU (J,K), J = 1, NRAU)

WRITE (9,5)

```

C          ..... COMPUTE RESIDUAL MATRIX

          IF ((CONVG.EQ.1).OR.(K.EQ.7)) THEN
&              CALL M4RES (RAU,ALPHA,PSI,PHI,COUNT,SEQ,CLIMA,NT,NY,
20              NPARM,NV,K,NRAU,NP,CONVG)
          ENDIF
          CONTINUE

          STOP
          END

```

PROGRAM 20

```

C          -----
C          ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR MODEL 4
C          USING CONJUGATE GRADIENT METHODS IN MULTIDIMENSIONS
C          -----

          INTEGER          NV,NY,NT,NP,NPARM,NRAU,T
          PARAMETER        (NV=7)
          PARAMETER        (NY=12)
          PARAMETER        (NT=365)
          PARAMETER        (NP=13)
          PARAMETER        (NPARM=3)
          PARAMETER        (NRAU=1)
          REAL              THETA (NP)
          REAL              AKAIKE,LNLIKE,PI
          PARAMETER        (PI=3.141593)
          REAL              MU (2,0:NT)
          REAL              SIGMA (2,0:NT)
          REAL              RESID (NV,NY,NT)

          COMMON            K,ICOUNT(4,NY),ISEQ(4,NY,NT),CLIMA(NY,0:NT),
&              ALPHA(2,NV,NPARM),PSI(2,NV,NPARM),PHI(NPARM,0:NT),
&              RAU(NRAU,NV),ISCALE(3,NV)

```

```

5      FORMAT (/)
15     FORMAT (' ESTIMATES OF MEAN FOR DRY DAYS:', 3F10.4)
25     FORMAT (' ESTIMATES OF MEAN FOR WET DAYS:', 3F10.4)
35     FORMAT (' ESTIMATES OF VAR FOR DRY DAYS:', 3F10.4)
45     FORMAT (' ESTIMATES OF VAR FOR WET DAYS:', 3F10.4)
55     FORMAT (' ESTIMATE OF AUTOCORRELATION:', F10.4)
65     FORMAT (' PARAMETER ESTIMATES FOR VARIABLE: ', I4)
75     FORMAT (' CONVERGE ACHIEVED IN ', I4, ' ITERATIONS')
115    FORMAT (3X, F5.0)
205    FORMAT (9X, F9.2)
305    FORMAT (18X, F9.2)
405    FORMAT (27X, F9.2)
505    FORMAT (36X, F9.2)
605    FORMAT (45X, F10.2)
705    FORMAT (55X, F10.2)
805    FORMAT (65X, F9.2)
905    FORMAT (I5)
105    FORMAT (14I5)
125    FORMAT (7F10.4)
135    FORMAT (' AKAIKE"S CRITERION FOR VARIABLE:', I4, ' IS:', F30.4)

OPEN (UNIT=18, FILE='\\WATER\\DATA\\CLIMA.DAT', STATUS='OLD')
OPEN (UNIT=12, FILE='\\WATER\\DATA\\SEQ.DAT', STATUS='OLD')
OPEN (UNIT=9, FILE='LPT1')
OPEN (UNIT=6, FILE='CON')
OPEN (UNIT=14, FILE='\\WATER\\DATA\\RESIT.DAT', STATUS='UNKNOWN')

TOL = 0.0000000001
CALL COSSIN (PHI, NPARM, NT)
DO 30, I = 1, NPARM
    PHI (I, 0) = PHI (I, NT)
30  CONTINUE

C      ..... INPUT OF TIME SERIES AND INITIAL PARAMETER ESTIMATES

CALL INT4 (ALPHA, PSI, RAU, NPARM, NV, NRAU, ISCALE)

PRINT *, 'WHICH VARIABLE TO BE ESTIMATED?'
READ (6, *) K

DO 10, I = 1, NY
    DO 220, J = 1, NT
        IF (K .EQ. 1) THEN
            READ (18, 205) CLIMA (I, J)
        ELSEIF (K .EQ. 2) THEN
            READ (18, 305) CLIMA (I, J)
        ELSEIF (K .EQ. 3) THEN
            READ (18, 405) CLIMA (I, J)
        ELSEIF (K .EQ. 4) THEN
            READ (18, 505) CLIMA (I, J)
        ELSEIF (K .EQ. 5) THEN
            READ (18, 605) CLIMA (I, J)
        ELSEIF (K .EQ. 6) THEN
            READ (18, 705) CLIMA (I, J)
        ELSEIF (K .EQ. 7) THEN
            READ (18, 805) CLIMA (I, J)
        ENDIF
220    CONTINUE
        CLIMA (I, 0) = CLIMA (I, 1) - 0.5
10    CONTINUE

```

```

REWIND 18
DO 440, KK = 1, NY
    DO 330, I = 1, 4
        READ (12,905) ICOUNT (I, KK)
        READ (12,105) (ISEQ (I, KK, J), J = 1, ICOUNT (I, KK))
330    CONTINUE
440    CONTINUE

C      ..... ITERATIVE ESTIMATION OF PARAMETERS

WRITE (9,65) K

C      ..... TRANSFORMING THE SEPARATE PARAMETER ARRAYS INTO ONE
C      VECTOR

DO 20, J = 1, NPARM
    THETA (J) = ALPHA (1, K, J) * ISCALE (3, K)
    THETA (J+3) = ALPHA (2, K, J) * ISCALE (3, K)
    THETA (J+6) = PSI (1, K, J) * ISCALE (2, K)
    THETA (J+9) = PSI (2, K, J) * ISCALE (2, K)
20    CONTINUE
DO 70, J = 1, NRAU
    THETA (J+12) = RAU (J, K) * ISCALE (1, K)
70    CONTINUE

CALL POLRIB (THETA, NP, TOL, ITER, FMIN)

C      ..... UPDATE PARAMETER ESTIMATES

DO 40 J = 1, NPARM
    ALPHA (1, K, J) = THETA (J) / ISCALE (3, K)
    ALPHA (2, K, J) = THETA (J+3) / ISCALE (3, K)
    PSI (1, K, J) = THETA (J+6) / ISCALE (2, K)
    PSI (2, K, J) = THETA (J+9) / ISCALE (2, K)
40    CONTINUE

DO 80, J = 1, NRAU
    RAU (J, K) = THETA (J+12) / ISCALE (1, K)
80    CONTINUE

WRITE (9,75) ITER

C      ..... OUTPUT OF FINAL PARAMETER ESTIMATES

WRITE (9,5)
WRITE (9,15) (ALPHA (1, K, L), L = 1, NPARM)
WRITE (9,25) (ALPHA (2, K, L), L = 1, NPARM)
WRITE (9,35) (PSI (1, K, L), L = 1, NPARM)
WRITE (9,45) (PSI (2, K, L), L = 1, NPARM)
WRITE (9,55) (RAU (J, K), J = 1, NRAU)
WRITE (9,5)

```

C           ..... COMPUTE RESIDUAL MATRIX

```

DO 50, M = 1, 2
  DO 60, I = 0, NT
    MU (M,I) = 0.0
    SIGMA (M,I) = 0.0
    DO 90, L = 1, NPARM
      MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
      SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
90    CONTINUE
60    CONTINUE
50    CONTINUE

DO 100, I = 1, NY
  DO 110, J = 1, NT
    RESID (K,I,J) = -999.00
110    CONTINUE
100    CONTINUE
    LNLIKE = 0
    TERM = 0
DO 120, J = 1, 4
  IF (J .EQ. 1) THEN
    M = 1
    L = 1
  ELSEIF (J .EQ. 2) THEN
    M = 2
    L = 2
  ELSEIF (J .EQ. 3) THEN
    M = 2
    L = 1
  ELSEIF (J .EQ. 4) THEN
    M = 1
    L = 2
  ENDIF
DO 130, I = 1, NY
  DO 140, KK = 1, ICOUNT (J,I)
    T = ISEQ (J,I,KK)
    IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
      THEN
      RESID(K,I,T) = (CLIMA(I,T)-MU(M,T))/SIGMA(M,T)-RAU(1,K)
      & *((CLIMA(I,T-1)-MU(L,T-1))/SIGMA(L,T-1))
      LNLIKE = LNLIKE + (RESID(K,I,T))**2
      ENDIF
    TERM = TERM + LOG(SIGMA(M,T))
140    CONTINUE
130    CONTINUE
120    CONTINUE

LNLIKE = -((NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2
AKAIKE = -2*LNLIKE+2*NP

WRITE (9,135) K, AKAIKE

DO 150, I = 1, NY
  DO 160, T = 1, NT
    WRITE (14,125) (RESID (K,I,T), K = 1, NV)
160    CONTINUE
150    CONTINUE

STOP
END

```

PROGRAM 21

```

C -----
C ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR MODEL 5
C -----

```

```

C PROGRAM EST-M5
C -----

```

```

INTEGER      NV,NY,NT,NP,NPARM,NRAU,CONVG,T
PARAMETER    (NV=7)
PARAMETER    (NY=12)
PARAMETER    (NT=365)
PARAMETER    (NP=16)
PARAMETER    (NPARM=3)
PARAMETER    (NRAU=4)
INTEGER      COUNT (4,NY)
INTEGER      SEQ (4,NY,NT)
REAL         CLIMA (NY,0:NT)
REAL         ALPHA (2,NV,NPARM)
REAL         PSI (2,NV,NPARM)
REAL         DER (NP)
REAL         DER2 (NP,NP)
REAL         PHI (NPARM,0:NT)
REAL         RAU (NRAU,NV)
REAL         THETA (NP)
REAL         A (NP,0:NP)

```

```

5  FORMAT (/)
15  FORMAT (' ESTIMATES OF MEAN FOR DRY DAYS:', 3F10.4)
25  FORMAT (' ESTIMATES OF MEAN FOR WET DAYS:', 3F10.4)
35  FORMAT (' ESTIMATES OF VAR FOR DRY DAYS:', 3F10.4)
45  FORMAT (' ESTIMATES OF VAR FOR WET DAYS:', 3F10.4)
55  FORMAT (' ESTIMATE OF AUTOCORRELATION:', 4F10.4)
75  FORMAT (9X,F9.2)
85  FORMAT (18X,F9.2)
95  FORMAT (27X,F9.2)
105 FORMAT (36X,F9.2)
115 FORMAT (45X,F10.2)
125 FORMAT (55X,F10.2)
135 FORMAT (65X,F9.2)
145 FORMAT (I5)
155 FORMAT (14I5)

```

```

OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
OPEN (UNIT=12,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='OLD')
OPEN (UNIT=9,FILE='LPT1')

```

```

CALL COSSIN (PHI,NPARM,NT)
DO 10, I = 1, NPARM
    PHI (I,0) = PHI (I,NT)

```

```

10    CONTINUE

C      ..... INPUT OF TIME SERIES AND INITIAL PARAMETER ESTIMATES

CALL INITIAL (EPS,MAXITER,ALPHA,PSI,RAU,NPARM,NV,NRAU)
DO 20, K = 1, NV
  CONVG = 1
  DO 30, I = 1, NY
    DO 40, J = 1, NT
      IF (K.EQ.1) THEN
        READ (18,75) CLIMA (I, J)
      ELSEIF (K.EQ.2) THEN
        READ (18,85) CLIMA (I, J)
      ELSEIF (K.EQ.3) THEN
        READ (18,95) CLIMA (I, J)
      ELSEIF (K.EQ.4) THEN
        READ (18,105) CLIMA (I, J)
      ELSEIF (K.EQ.5) THEN
        READ (18,115) CLIMA (I, J)
      ELSEIF (K.EQ.6) THEN
        READ (18,125) CLIMA (I, J)
      ELSEIF (K.EQ.7) THEN
        READ (18,135) CLIMA (I, J)
      ENDIF
40    CONTINUE
    IF (I.EQ.1) THEN
      CLIMA (I,0) = CLIMA (I,1) - 0.5
    ELSEIF (I.NE.1) THEN
      CLIMA(I,0) = CLIMA(I-1,NT)
    ENDIF
30  CONTINUE

  REWIND 18
  IF (K .EQ. 1) THEN
    DO 50, KK = 1, NY
      DO 60, I = 1, 4
        READ (12,145) COUNT (I,KK)
        READ (12,155) (SEQ (I,KK,J), J = 1, COUNT (I,KK))
60      CONTINUE
50    CONTINUE
  ENDIF

C      ..... ITERATIVE ESTIMATION OF PARAMETERS

      CALL NEWT5 (ALPHA,PSI,RAU,NPARM,MAXITER,NT,NY,CLIMA,SEQ,
&              COUNT,DER,DER2,PHI,EPS,NP,NV,K,A,THETA,NRAU,CONVG)

C      ..... OUTPUT OF FINAL PARAMETER ESTIMATES

      WRITE (9,5)
      WRITE (9,15) (ALPHA (1,K,L), L = 1, NPARM)
      WRITE (9,25) (ALPHA (2,K,L), L = 1, NPARM)
      WRITE (9,35) (PSI (1,K,L), L = 1, NPARM)
      WRITE (9,45) (PSI (2,K,L), L = 1, NPARM)

```

```

        WRITE (9,55) (RAU (J,K), J = 1, NRAU)
        WRITE (9,5)
C      ..... COMPUTE RESIDUAL MATRIX
        IF ((CONVG.EQ.1).OR.(K.EQ.7)) THEN
          CALL M5RES (RAU,ALPHA,PSI,PHI,COUNT,SEQ,CLIMA,NT,NY,
&                  NPARM,NV,K,NRAU,NP,CONVG)
        ENDIF
20    CONTINUE

        STOP
        END

```

PROGRAM 22

```

C -----
C ..... PROGRAM TO COMPUTE PARAMETER ESTIMATES FOR MODEL 5
C USING CONJUGATE GRADIENT METHODS IN MULTIDIMENSIONS
C -----

      INTEGER      NV,NY,NT,NP,NPARM,NRAU,T
      PARAMETER    (NV=7)
      PARAMETER    (NY=12)
      PARAMETER    (NT=365)
      PARAMETER    (NP=16)
      PARAMETER    (NPARM=3)
      PARAMETER    (NRAU=4)
      REAL         THETA (NP)
      REAL         LNLIKE,AKAIKE,PI
      PARAMETER    (PI=3.141593)
      REAL         MU (2,0:NT)
      REAL         SIGMA (2,0:NT)
      REAL         RESID (NV,NY,NT)

      COMMON       K,ICOUNT(4,NY),ISEQ(4,NY,NT),CLIMA(NY,0:NT),
&                ALPHA(2,NV,NPARM),PSI(2,NV,NPARM),PHI(NPARM,0:NT),
&                RAU(NRAU,NV),ISCALE(3,NV)

      5  FORMAT (/)
      15  FORMAT (' ESTIMATES OF MEAN FOR DRY DAYS:', 3F10.4)
      25  FORMAT (' ESTIMATES OF MEAN FOR WET DAYS:', 3F10.4)
      35  FORMAT (' ESTIMATES OF VAR FOR DRY DAYS:', 3F10.4)
      45  FORMAT (' ESTIMATES OF VAR FOR WET DAYS:', 3F10.4)
      55  FORMAT (' ESTIMATES OF AUTOCORRELATION:', 4F10.4)
      65  FORMAT (' PARAMETER ESTIMATES FOR VARIABLE: ', I4)
      75  FORMAT (' CONVERGE ACHIEVED IN ', I4, ' ITERATIONS')
      105 FORMAT (3X, F5.0)
      205 FORMAT (9X, F9.2)
      305  FORMAT (18X, F9.2)
      405  FORMAT (27X, F9.2)
      505  FORMAT (36X, F9.2)
      605  FORMAT (45X, F10.2)
      705  FORMAT (55X, F10.2)
      805  FORMAT (65X, F9.2)
      905  FORMAT (I5)
      115  FORMAT (14I5)
      315  FORMAT (7F10.4)
      415  FORMAT (' AKAIKE"S CRITERION FOR VARIABLE:', I4, ' IS:', F10.4)

      OPEN (UNIT=14,FILE='\\WATER\\DATA\\RESIT.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=18,FILE='\\WATER\\DATA\\CLIMA.DAT',STATUS='OLD')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\SEQ.DAT',STATUS='OLD')
      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=6,FILE='CON')

```

```

TOL = 0.0000000001

CALL COSSIN (PHI,NPARM,NT)
DO 10, I = 1, NPARM
    PHI (I,0) = PHI (I,NT)
10 CONTINUE

C      ..... INPUT OF TIME SERIES AND INITIAL PARAMETER ESTIMATES

CALL INT4 (ALPHA,PSI,RAU,NPARM,NV,NRAU,ISCALE)

PRINT *, 'WHICH VARIABLE TO BE ESTIMATED?'
READ (6,*) K

DO 20, I = 1, NY
    DO 30, J = 1, NT
        IF (K .EQ. 1) THEN
            READ (18,205) CLIMA (I, J)
        ELSEIF (K .EQ. 2) THEN
            READ (18,305) CLIMA (I, J)
        ELSEIF (K .EQ. 3) THEN
            READ (18,405) CLIMA (I, J)
        ELSEIF (K .EQ. 4) THEN
            READ (18,505) CLIMA (I, J)
        ELSEIF (K .EQ. 5) THEN
            READ (18,605) CLIMA (I, J)
        ELSEIF (K .EQ. 6) THEN
            READ (18,705) CLIMA (I, J)
        ELSEIF (K .EQ. 7) THEN
            READ (18,805) CLIMA (I, J)
        ENDIF
30    CONTINUE
        CLIMA (I,0) = CLIMA (I,1) - 0.5
20    CONTINUE

REWIND 18
DO 40, KK = 1, NY
    DO 50, I = 1, NRAU
        READ (12,905) ICOUNT (I,KK)
        READ (12,115) (ISEQ (I,KK,J), J = 1, ICOUNT (I,KK))
50    CONTINUE
40    CONTINUE

C      ..... ITERATIVE ESTIMATION OF PARAMETERS

WRITE (9,65) K

C      ..... TRANSFORMING THE SEPARATE PARAMETER ARRAYS INTO ONE
C      VECTOR

DO 60, J = 1, NPARM
    THETA (J) = ALPHA (1,K,J)*ISCALE(3,K)

```

```

        THETA (J+3) = ALPHA (2,K,J)*ISCALE(3,K)
        THETA (J+6) = PSI (1,K,J)*ISCALE(2,K)
        THETA (J+9) = PSI (2,K,J)*ISCALE(2,K)
60    CONTINUE

        DO 70, J = 1, NRAU
            THETA (J+12) = RAU (J,K)*ISCALE(1,K)
70    CONTINUE

        CALL POLRIB (THETA,NP,TOL,ITER,FMIN)

C      ..... UPDATE PARAMETER ESTIMATES

        DO 80 J = 1, NPARM
            ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
            ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
            PSI (1,K,J) = THETA (J+6)/ISCALE(2,K)
            PSI (2,K,J) = THETA (J+9)/ISCALE(2,K)
80    CONTINUE

        DO 90, J = 1, NRAU
            RAU (J,K) = THETA (J+12)/ISCALE(1,K)
90    CONTINUE

        WRITE (9,75) ITER

C      ..... OUTPUT OF FINAL PARAMETER ESTIMATES

        WRITE (9,5)
        WRITE (9,15) (ALPHA (1,K,L), L = 1, NPARM)
        WRITE (9,25) (ALPHA (2,K,L), L = 1, NPARM)
        WRITE (9,35) (PSI (1,K,L), L = 1, NPARM)
        WRITE (9,45) (PSI (2,K,L), L = 1, NPARM)
        WRITE (9,55) (RAU (J,K), J = 1, NRAU)
        WRITE (9,5)

C      ..... COMPUTE RESIDUAL MATRIX

        DO 100, M = 1, 2
            DO 120, I = 0, NT
                MU (M,I) = 0.0
                SIGMA (M,I) = 0.0
                DO 130, L = 1, NPARM
                    MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
                    SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
130                CONTINUE
120            CONTINUE
100        CONTINUE

        DO 140, I = 1, NY

```

```

DO 150, J = 1, NT
  RESID (K,I,J) = -999.00
150  CONTINUE
140  CONTINUE
  LNLIKE = 0
  TERM = 0

DO 160, J = 1, 4
  IF (J .EQ. 1) THEN
    M = 1
    L = 1
  ELSEIF (J .EQ. 2) THEN
    M = 2
    L = 2
  ELSEIF (J .EQ. 3) THEN
    M = 2
    L = 1
  ELSEIF (J .EQ. 4) THEN
    M = 1
    L = 2
  ENDIF

DO 170, I = 1, NY
  DO 180, KK = 1, ICOUNT (J,I)
    T = ISEQ (J,I,KK)
    IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
      & THEN
      RESID(K,I,T) = (CLIMA(I,T)-MU(M,T))/SIGMA(M,T)-
      & RAU(J,K)*((CLIMA(I,T-1)-MU(L,T-1))/
      & SIGMA(L,T-1))
      LNLIKE = LNLIKE + (RESID(K,I,T))**2
      ENDIF
      TERM = TERM + LOG(SIGMA(M,T))
180  CONTINUE
170  CONTINUE
160  CONTINUE
  LNLIKE = -((NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2
  AKAIKE = -2*LNLIKE+2*NP
  WRITE (9,415) K, AKAIKE

DO 190, I = 1, NY
  DO 200, T = 1, NT
    WRITE (14,315) (RESID (K,I,T), K = 1, NV)
200  CONTINUE
190  CONTINUE

STOP
END

```

PROGRAM 23

```

C -----
C ..... PROGRAM TO RECORD TIME PERIODS FOR WHICH A MISSING
C OBSERVATION OCCURS
C -----

      INTEGER          NV, TIME, BOUND
      PARAMETER        (NV=7)
      PARAMETER        (TIME=4380)
      PARAMETER        (BOUND=500)
      INTEGER          SEQMISS (NV, BOUND)
      INTEGER          COUNT (NV)
      REAL             CLIMA (NV)

15  FORMAT (14(I5))
25  FORMAT (I5)

      OPEN (UNIT=10, FILE=' \WATER\DATA\RESIDU.DAT', STATUS='OLD')
      OPEN (UNIT=8, FILE=' \WATER\DATA\SEQM.DAT', STATUS='UNKNOWN')
      DO 20, K = 1, NV
        COUNT (K) = 0
20  CONTINUE
      DO 10, J = 1, TIME
        READ (10, *) (CLIMA (K), K = 1, NV)
        DO 30, K = 1, NV
          IF (CLIMA (K) .LT. -900) THEN
            COUNT (K) = COUNT (K) + 1
            SEQMISS (K, COUNT (K)) = J
          ENDIF
30  CONTINUE
10  CONTINUE
      DO 60, K = 1, NV
        WRITE (8, 25) COUNT (K)
        WRITE (8, 15) (SEQMISS (K, I), I = 1, COUNT (K))
60  CONTINUE

      STOP
      END

```

# PROGRAM 24

```

C -----
C ..... A PROGRAM TO PATCH THE MISSING OBSERVATIONS IN A
C          GIVEN DATA SET USING THE EM-ALGORITHM
C -----
C
C In this program there are missing observations in almost all the
C the variables.
C
C The variables are read as one big matrix which consists of a
C column of the dependent variable - which should always be the
C first column, and the of the columns being the matrix of the
C independent variables.
C
C Each row of data represents an observation and where a "-999" is
C encountered, that would be representing a missing observation.
C
C The data is stored in a matrix called the Z-matrix, and that is
C subdivided into :
C       Y-matrix = A matrix of the dependent variable
C       X-matrix = A matrix of the independent variables
C
C The maximum dimensions of the matrices are:
C       Dependent variable      : 1
C       Independent variables   : 25
C       Observations            : 100
C
C Note that only one Y-variable can be patched at a time, and
C therefore we can only have one Y-variable at a time.  If
C there are missing observations in more than one variable, then
C it is therefore necessary to swop the variable s columns so
C that the variable which needs to be patched is always in the
C first column of the matrix.
C
C Note again that most of the routines which are in this program
C were copied from the programs written by Dr Ross Sparks.
C
C ..... VARIABLES DECLARATION
C
C INTEGER NOBS, NSTAT, IV, DV
C PARAMETER(NOBS=4380)
C PARAMETER(NSTAT=7)
C PARAMETER(IV=12)
C PARAMETER(DV=1)
C PARAMETER(NI=500)
C
C ..... NOBS = Number of all the records
C          NSTAT = Number of all the stations i.e. target & control
C          IV = Number of control stations
C          DV = Number of target stations
C
C REAL Z(NOBS,NSTAT)

```

```

REAL TMAT (NOBS,NSTAT)
REAL ZCEN(NOBS,NSTAT), PATCH(NOBS)
REAL MEANZ(NSTAT,DV), MEANZZ(DV,NSTAT)
REAL ZTZ(NSTAT,NSTAT)
REAL TEMP1(NSTAT), TEMPO, TEMP2
REAL MEAN1(NSTAT), MEAN2(NSTAT)
REAL BHAT(NSTAT,DV), BETA(7,NI), CONV(DV,DV)
INTEGER COUNT (NSTAT)
INTEGER SEQMISS (NSTAT,500)
INTEGER ROW, COL, ROUND, NROW
INTEGER NROUND

1  FORMAT(9F8.0)
2  FORMAT(20F6.0)

OPEN (UNIT=9,FILE='LPT1')
OPEN (UNIT=10,FILE='\\WATER\\DATA\\RESIDU.DAT',STATUS='OLD')
OPEN (UNIT=20,FILE='\\WATER\\DATA\\SEQM.DAT',STATUS='OLD')

C      ..... This DO-LOOP reads a matrix of all the rainfall
C      stations and all the observations in a MATRIX Z.

DO 10 ROW = 1, NOBS

    READ(10,*) (Z(ROW,COL), COL = 1, NSTAT)

10  CONTINUE

C      ..... This DO-LOOP reads a vector of the amount of missing
C      values for each variable in MATRIX COUNT and a matrix of
C      the specific times when missing values occur for each of
C      the variables in a MATRIX SEQMISS.

DO 20, K = 1, NSTAT
    READ (20,*) COUNT (K)
    READ (20,*) (SEQMISS (K,I), I = 1, COUNT (K))
20  CONTINUE

C      ..... FIND THE MEANS OF THE DIFFERENT COLS, I.E. FIND THE MEAN
C      OF ALL THE OBSERVATIONS IN COL 1, COL2, ETC.

DO 110 COL = 1, NSTAT
    MEANZ(COL,1) = 0.0
    MEAN1(COL) = 0.0
    DO 100 ROW = 1, NOBS
        IF (Z(ROW,COL) .NE. -999) THEN
            MEAN1(COL) = MEAN1(COL) + Z(ROW,COL)
        ENDIF
100    CONTINUE
    MEANZ(COL,1) = MEAN1(COL) / (NOBS - COUNT (COL))
110  CONTINUE

C      ..... SUBSTITUTE THE MISSING OBSERVATIONS BY THE
C      CALCULATED MEANS

```

```

      DO 130 COL = 1, NSTAT
        DO 120 K = 1, COUNT (COL)
          ROW = SEQMISS (COL,K)
          Z(ROW,COL) = MEANZ(COL,1)
120      CONTINUE
130      CONTINUE

      NROUND = 1

12121 CALL CNTRAL(ZCEN,NOBS,NSTAT,Z,NOBS,NSTAT,NOBS,NSTAT)
      CALL TMULT(ZCEN,NOBS,NSTAT,ZTZ,NSTAT,NSTAT,NSTAT,NOBS,NSTAT)
      CALL INV(ZTZ,NSTAT,NSTAT)

      DO 167 COL = 1, NSTAT
        MEAN2(COL) = 0.0
        DO 163 K = 1, COUNT (COL)
          ROW = SEQMISS (COL,K)
          MEAN2(COL) = MEAN2(COL) + Z(ROW,COL)
163      CONTINUE
        MEANZZ(1,COL) = (MEAN1(COL) + MEAN2(COL)) / NOBS
167      CONTINUE

      ROUND = 1

13131 DO 810 ROW = 1, NSTAT
        TEMP1(ROW) = (-1.0) * ZTZ(ROW,ROUND) / ZTZ(ROUND,ROUND)
810      CONTINUE

      TEMP2 = 0.0
      DO 830 ROW = 1, NSTAT
        IF (ROW .NE. ROUND) THEN
          TEMP2 = TEMP2 + MEANZZ(1,ROW) * TEMP1(ROW)
        ENDIF
830      CONTINUE

      TEMPO = MEANZZ(1,ROUND) - TEMP2

      TEMP1(ROUND) = TEMPO

      DO 450 ROW = 1, NSTAT
        BETA(ROW,NROUND) = TEMP1(ROW)
450      CONTINUE

      IF (NROUND .GT. 1) THEN

        DO 460 ROW = 1, NSTAT
          BHAT(ROW,1) = BETA(ROW,NROUND) - BETA(ROW,NROUND-1)
460      CONTINUE

        CALL TMULT(BHAT,NSTAT,DV,CONV,DV,DV,DV,NSTAT,DV)

```

```

ENDIF
C      ..... PATCH THE MISSING OBSERVATIONS

DO 210 ROW = 1, NOBS
    PATCH (ROW) = Z (ROW,ROUND)
210  CONTINUE
DO 200 K = 1, COUNT (ROUND)
    ROW = SEQMISS (ROUND,K)
    PATCH(ROW) = 0.0
    DO 192 COL = 1, NSTAT
        IF (COL .EQ. ROUND) THEN
            GO TO 192
        ENDIF
        PATCH(ROW) = PATCH(ROW) + Z(ROW,COL) * TEMP1(COL)
192  CONTINUE
        PATCH(ROW) = TEMP1(ROUND) + PATCH(ROW)
200  CONTINUE

DO 220, ROW = 1, NOBS
    TMAT (ROW,ROUND) = PATCH (ROW)
220  CONTINUE

IF (NROUND .GT. 1) THEN

    IF (CONV(1,1) .LT. 0.0000001) THEN

        WRITE (9,*) 'VALUES PATCHED AFTER ', NROUND, ' ALTERATIONS.'
        CALL PPMAT (BETA,NSTAT,NROUND,NSTAT, NROUND)
        CALL PMAT (TMAT,NOBS,NSTAT,NOBS,NSTAT)
        CALL PPMAT (TEMP1,NSTAT,DV,NSTAT,DV)
        GO TO 998

    ENDIF
ENDIF

ROUND = ROUND + 1

IF (ROUND .GT. NSTAT) THEN
    NROUND = NROUND + 1
    CALL COPY (TMAT,NOBS,NSTAT,Z,NOBS,NSTAT,NOBS,NSTAT)
    IF (NROUND .GT. NI) THEN
        WRITE (9,*) 'NO CONVERGENCE AFTER ', NI, ' ITERATIONS.'
        GOTO 998
    ENDIF
    GO TO 12121
ENDIF

GO TO 13131

998  STOP
END

```

PROGRAM 25

```

C -----
C ..... PROGRAM TO ESTIMATE THE CORRELATION MATRIX AND
C THE VECTOR OF VARIANCES.
C -----

      INTEGER          NT,NV,NV
      PARAMETER        (NT=4380)
      PARAMETER        (NV=7)
      REAL              TERM (5)
      REAL              CORR (NV,NV)
      REAL              RES (NV, NT)
      REAL              VARI (NV)

5      FORMAT (7 F10.4)
15     FORMAT (/, ' THE CORRELATION MATRIX: ')
25     FORMAT (' THE VARIANCE OF EACH VARIABLE: ')

      OPEN (UNIT=10,FILE='\\WATER\\DATA\\RESI.DAT',STATUS='OLD')
      OPEN (UNIT=12,FILE='\\WATER\\DATA\\CORR.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=9,FILE='LPT1')

      DO 20, I = 1, NT
        READ (10, *) (RES (K, I), K = 1, NV)
20     CONTINUE
      DO 40, I = 1, NV
        CORR (I, I) = 1
40     CONTINUE
      DO 60, K = 1, NV
        DO 70, J = K+1, NV
          DO 120, II = 1, 5
            TERM (II) = 0
120    CONTINUE
          DO 80, I = 1, NT
            TERM (1) = TERM (1) + RES (K, I) * RES (J, I)
            TERM (2) = TERM (2) + RES (K, I)
            TERM (3) = TERM (3) + RES (J, I)
            TERM (4) = TERM (4) + RES (K, I) ** 2
            TERM (5) = TERM (5) + RES (J, I) ** 2
80    CONTINUE
            TERM (1) = TERM (1) / NT
            TERM (4) = SQRT ((TERM (4) / NT) - (TERM (2) / NT) ** 2)
            VARI (K) = TERM (4) ** 2
            TERM (5) = SQRT ((TERM (5) / NT) - (TERM (3) / NT) ** 2)
            TERM (2) = TERM (2) * TERM (3) / NT ** 2
            TERM (1) = TERM (1) - TERM (2)
            TERM (4) = TERM (4) * TERM (5)
            CORR (K, J) = TERM (1) / TERM (4)
70    CONTINUE
60    CONTINUE
      TERM (2) = 0
      TERM (4) = 0

```

```

DO 10, I = 1, NT
    TERM (2) = TERM (2) + RES (NV, I)
    TERM (4) = TERM (4) + RES (NV, I) ** 2
10  CONTINUE
    VARI (NV) = (TERM (4) / NT) - (TERM (2) / NT) ** 2
    WRITE (9, 25)
    WRITE (9, 5) (VARI (K), K = 1, NV)
    WRITE (9, 15)
    DO 110, K = 1, NV
        WRITE (9, 5) (CORR (K, J), J = K, NV)
        WRITE (12, 5) (CORR (K, J), J = K, NV)
110  CONTINUE

    STOP
    END

```

PROGRAM 26

```

C -----
C ..... PROGRAM TO GENERATE CLIMATE SEQUENCES USING MODEL T
C -----

      INTEGER          NT,NV,NV3,NV4,NV5,NY,NP,PSTATE,STATE,A,NRAU

C ..... PSTATE = PRESENT STATE OF DAY
C ..... STATE = PREVIOUS STATE OF DAY

      PARAMETER        (NT=365)
      PARAMETER        (NY=51)
      PARAMETER        (NV=7)
      PARAMETER        (NV3=4)
      PARAMETER        (NV4=1)
      PARAMETER        (NV5=2)
      PARAMETER        (NP=3)
      PARAMETER        (NRAU=4)

C ..... NT = £ OBSERVATIONS PER YEAR
C ..... NV = £ VARIABLES
C ..... NY = £ YEARS TO BE GENERATED
C ..... NP = £ PARAMETERS IN SEASONAL MODEL

      INTEGER          SEED (9)
      REAL              RAIN
      REAL              GAM (2,NP)
      REAL              PHI (NP,0:NT)
      REAL              RAU3 (NRAU,NV3)
      REAL              RAU4 (NV4)
      REAL              RAU5 (NRAU,NV5)
      REAL              DECOMP (NV,NV)
      REAL              RAND (1,NV)
      REAL              SIGMA3 (NRAU,NV3)
      REAL              SIGMA4 (2,NV4,0:NT)
      REAL              SIGMA5 (2,NV5,0:NT)
      REAL              MU (2,NV,0:NT)
      REAL              OBSN (NV),TEMP(NV)
      REAL              AMP (0:NP)
      REAL              PHASE (NP)
      REAL              CORR (NV,NV)
      REAL              C (NT)

      COMMON            IDUM1,IDUM2,IDUM3,IDUM4,IDUM5,IDUM6,IDUM7

15  FORMAT (4F9.2, 2F10.2, F9.2)
25  FORMAT (' GIVE 9 -VE Nos. TO INITIALIZE RANDOM GENERATOR',/)

      OPEN (UNIT=9,FILE='LPT1')
      OPEN (UNIT=10,FILE='\\WATER\\DATA\\SIMU.DAT',STATUS='UNKNOWN')
      OPEN (UNIT=22,FILE='CON')

```

```

C      ..... COMPUTE THE FOURIER SERIES TERMS

      CALL COSSIN (PHI,NP,NT)
      DO 60, I = 1, NP
        PHI (I,0) = PHI (I,NT)
60     CONTINUE

      PI=3.14159
      SMAX=135
      SMIN=110
      AVE=(SMAX+SMIN)/2
      AMPS=SMAX-SMIN

      DO 330, I = 1, NT
        C(I) = AVE+(AMPS/2)*COS((2*PI/NT)*(I+1))
330    CONTINUE

C      ..... READING PARAMETER ESTIMATES

      CALL DATA (GAM,RAU3,RAU4,RAU5,MU,SIGMA3,SIGMA4,SIGMA5,NP,NV,AMP,
&              PHASE,CV,PHI,CORR,NT,NRAU,NV3,NV4,NV5)

C      ..... COMPUTE THE CHOLESKY DECOMPOSITION OF THE CORRELATION
C      MATRIX.  INPUT MATRIX HERE AS WELL.

      CALL CHOLESKY (DECOMP,CORR,NV)

C      ..... TRANSPOSE COVARIANCE MATRIX

      CALL GTRANP (DECOMP,NV)

C      ..... INPUT SEEDS TO START RANDOM NUMBER GENERATOR.  MUST BE
C      NEGATIVE NUMBER.

      PRINT 25
      DO 50, II = 1, 9
        READ (22, *) SEED (II)
50     CONTINUE

      IDUM1 = SEED (1)
      IDUM2 = SEED (2)
      IDUM3 = SEED (3)
      IDUM4 = SEED (4)
      IDUM5 = SEED (5)
      IDUM6 = SEED (6)
      IDUM7 = SEED (7)
      IDUM8 = SEED (8)
      IDUM9 = SEED (9)

C      ..... COMPUTE PARAMETERS NEEDED FOR COMPUTATION OF RAINFALL
C      DEPTH

      CALL CALBET (BETA,CV)
      ALPH = 1 + 1 / BETA

```

```

GAMM = GAMMA (ALPH)
BI = 1 /BETA
W = 0.01721421

C      ..... SET INITIAL STATE OF DAY TO BE DRY
C      ..... SET INITIAL CLIMATE VALUE TO ZERO

C      ..... STATE = 1 ==> DRY
C      ..... STATE = 2 ==> WET

STATE = 1
DO 10, I = 1, NV
    OBSN (I) = MU (STATE,I,0)
10  CONTINUE
DO 30, I = 1, NY
    DO 40, J = 1, NT

C      ..... GENERATE RAINFALL VALUE
C      -----

C      ..... COMPUTE PROBABILITY THAT A WET DAY FOLLOWS A WET DAY, OR
C      THE PROBABILITY THAT A WET DAY FOLLOWS A DRY DAY.

        CALL PIEST (NP,GAM,STATE,J,PHI,PI,NT)

C      ..... GENERATE A UNIFORM RANDOM NUMBER BETWEEN 0 AND 1.

        UNIFOR = URAN8 (IDUMB)
        IF (UNIFOR .LT. PI) THEN
            PSTATE = 2
        ELSE
            PSTATE = 1
        ENDIF

C      ..... GENERATE A NORMAL RANDOM NUMBER

        CALL GAUSS (DECOMP,RAND)

C      ..... GENERATE CLIMATE SEQUENCES
C      -----

        DO 80, K = 1, NV
            IF ((K.EQ.1).OR.(K.EQ.4).OR.(K.EQ.6)) THEN
                CALL MOD3 (RAND,STATE,NV3,NV,SIGMA3,MU,RAU3,K,
                    J,OBSN,PSTATE,NT,NRAU)
            ELSEIF ((K.EQ.2).OR.(K.EQ.5)) THEN
                CALL MOD5 (RAND,STATE,NV5,NV,SIGMA5,MU,RAU5,K,
                    J,OBSN,PSTATE,NT,NRAU)
            ELSEIF ((K.EQ.3)) THEN
                CALL MOD4 (RAND,STATE,NV4,NV,SIGMA4,MU,RAU4,K,
                    J,OBSN,PSTATE,NT)
            ENDIF
60      CONTINUE
80

```

```

C      ..... DETERMINE WHETHER IT RAINED AND SET RAIN VALUE
C
C      ..... RAIN = 0 ==> DID NOT RAIN
C      ..... RAIN = 1 ==> RAINED
C
C      IF (PSTATE .EQ. 1) THEN
C          RAIN = 0
C      ELSE
C          RAIN = 1
C      ENDIF
C
C      ..... GENERATE RAINFALL DEPTH IF IT RAINED
C      -----
C
C      IF (RAIN .EQ. 1) THEN
C          CALL DEPTH3 (IDUM9,NP,RAIN,J,AMP,PHASE,GAMM,BI,W)
C      ENDIF
C
C      ..... TRANSFORM VARIABLES TO THE ORIGINAL FORM
C
C      TEMP(2)=(230-100*EXP(OBSN(2)))/(EXP(OBSN(2))+1)
C      TEMP(1)=(410+TEMP(2)*EXP(OBSN(1)))/(EXP(OBSN(1))+1)
C      TEMP(3)=(C(J)-0.01-(0.01*EXP(OBSN(3))))/(EXP(OBSN(3))+1)
C      TEMP(4)=(10000/(EXP(OBSN(4))+1))-0.01
C      TEMP(6)=100/(EXP(OBSN(6))+1)
C      TEMP(5)=(101+TEMP(6)*EXP(OBSN(5)))/(EXP(OBSN(5))+1)
C
C      ..... OUTPUT GENERATED SEQUENCES
C
C      IF (I .NE. 1) THEN
C          WRITE (10,15) RAIN, (TEMP (K), K = 1, NV)
C      ENDIF
C
C      ..... UPDATE THE STATE OF THE PREVIOUS DAY
C
C      IF (PSTATE .NE. STATE) THEN
C          STATE = PSTATE
C      ENDIF
C
40      CONTINUE
30      CONTINUE
C
C      STOP
C      END

```

# SUBROUTINES

```

-----
..... THIS SUBROUTINE ITERATIVELY ESTIMATES THE MODEL
PARAMETERS BY THE NEWTON-RAPHSON METHOD FOR M3.
-----

```

```

SUBROUTINE NEWT3 (ALPHA,SIGMA,RAU,NP,NP,MAXITER,NT,NY,CLIMA,SEQ,
& COUNT,DER,DER2,PHI,EPS,NP,NV,K,A,THETA,NRAU,CONVG)
-----

```

```

INTEGER      COUNT (4,NY)
INTEGER      SEQ (4,NY,NT)
REAL         A (NP,0:NP)
REAL         SIGMA (NRAU,NV)
REAL         ALPHA (2,NV,NP,NP)
REAL         PHI (NP,NP,0:NT)
REAL         DER (NP)
REAL         DER2 (NP,NP)
REAL         CLIMA (NY,0:NT)
REAL         THETA (NP)
REAL         RAU (NRAU,NV)

```

```

15  FORMAT (' THE SUCCESSIVE THETA VALUES FOR VARIABLE: ', I4)
25  FORMAT (' .... DID NOT CONVERGE')
35  FORMAT (/, ' .... ', I3, ' ITERATION', /)

```

```

OPEN (UNIT=9,FILE='LPT1')

```

```

IC = 0
WRITE (9,15) K

```

```

..... TRANSFORMING THE SEPARATE PARAMETER ARRAYS INTO ONE
VECTOR

```

```

DO 20, J = 1, NP,NP
  THETA (J) = ALPHA (1,K,J)
  THETA (J+3) = ALPHA (2,K,J)
20  CONTINUE
DO 70, J = 1, NRAU
  THETA (J+6) = SIGMA (J,K)
  THETA (J+10) = RAU (J,K)
70  CONTINUE

```

```

..... ITERATIVE PARAMETER ESTIMATION

```

```

DO 10, ITER = 1, MAXITER

```

```

..... VECTOR OF 1ST DERIVATIVES AND MATRIX OF 2ND DERIVATIVES
IS COMPUTED

```

```

& CALL M3DERV (NP,NP,NT,ALPHA,SIGMA,RAU,CLIMA,SEQ,COUNT,
  DER,DER2,PHI,NP,NV,K,NRAU)

```

```

        DO 40, KK = 1, NP
          DO 50, J = KK, NP
            DER2 (J, KK) = DER2 (KK, J)
50      CONTINUE
40      CONTINUE
        PRINT 35, ITER

C      ..... NEW PARAMETER ESTIMATES ARE COMPUTED

        CALL NEWPARM (NP, DER, DER2, THETA, EPS, IC, A)

C      ..... UPDATE PARAMETER ESTIMATES

        DO 30 J = 1, NPARM
          ALPHA (1, K, J) = THETA (J)
          ALPHA (2, K, J) = THETA (J+3)
30      CONTINUE

        DO 80, J = 1, NRAU
          SIGMA (J, K) = THETA (J+6)
          RAU (J, K) = THETA (J+10)
80      CONTINUE

C      ..... TEST FOR CONVERGENCE

        IF (IC) 10, 10, 60
10      CONTINUE
        WRITE (9, 25)
        CONVG = 0

60      RETURN
        END

C      SUBROUTINE CALBET (BETA, CV)
      -----

      REAL          NUM, DENOM

      C2 = CV ** 2
      C3 = CV ** 3
      NUM = 339.5410 + 148.4445*CV + 192.7492*C2 + 22.4401*C3
      DENOM = 1 + 257.1162*CV + 287.8362*C2 + 157.2230*C3
      BETA = NUM / DENOM

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE SOLVES A SYSTEM OF EQUATIONS AND
C AND EXTRACTS NEW PARAMETER ESTIMATES
C -----

SUBROUTINE NEWPARM (NP,DER,DER2,THETA,EPS,IC,A)
-----

REAL          A (NP,0:NP)
REAL          DER (NP)
REAL          DER2 (NP,NP)
REAL          THETA (NP)

15  FORMAT (' MATRIX IS SINGULAR')
25  FORMAT (' NEW PARAMETER ESTIMATES: ', F10.4)

OPEN (UNIT=9,FILE='LPT1')

C THIS SETS UP THE A MATRIX WHICH IS USED IN SOLVING THE SYSTEM
C OF EQUATIONS

DO 10, I = 1, NP
  A (I,0) = DER (I)
  DO 20, J = 1, NP
    A (I,J) = DER2 (I,J)
20  CONTINUE
10  CONTINUE

C THIS SOLVES THE SYSTEM OF EQUATIONS
C THE DIFFERENCE BETWEEN THE VALUE OF THETA (Q) IN THIS
C ITERATION AND IN THE PREVIOUS ITERATION ARE STORED IN A (Q,0)

DO 30, I1 = 1, NP
  I2 = I1
  T1 = 0
  DO 40, I3 = I1, NP
    IF (ABS (A (I3,I1)) .GT. (ABS (T1))) THEN
      I2 = I3
      T1 = A (I3,I1)
    ENDIF
40  CONTINUE
    IF (T1 .EQ. 0) THEN
      WRITE (9,15)
      STOP
    ENDIF
    IF (I2 .NE. I1) THEN
      DO 50, IO = 0, NP
        TEMP = A (I1,IO)
        A (I1,IO) = A (I2,IO)
        A (I2,IO) = TEMP
50  CONTINUE
      ENDIF
      T2 = 1 / (A (I1,I1))

```

```

      NQ = NP
      DO 60, I4 = 0, NQ
        A (I1,I4) = A (I1,I4) * T2
60    CONTINUE
      DO 70, I3 = 1, NP
        IF (I1 .NE. I3) THEN
          T2 = A (I3,I1)
          A (I3,0) = A (I3,0)-A(I1,0)* T2
          DO 80, I0 = I1, NP
            A(I3,I0) = A(I3,I0) - A(I1,I0) * T2
80          CONTINUE
        ENDIF
      CONTINUE
70    CONTINUE
30    CONTINUE

C      ..... CONVERGENCE TEST

      CRIT = 0
      DO 205, I = 1, NP
        CRIT = CRIT + ABS(A(I,0))
205    CONTINUE
      IF (CRIT .GT. EPS) THEN
        IC = 0
      ELSE
        IC = 1
      ENDIF

C      ..... THIS EXTRACTS THE NEW PARAMETER VALUES

      DO 90, I = 1, NP
        THETA (I) = THETA (I) - A (I,0)
C      WRITE (9,25) THETA (I)
        PRINT 25, THETA (I)
90    CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C AND THE MATRIX OF SECOND DERIVATIVES FOR MODEL3.
C -----

C SUBROUTINE M3DERV (NPARM,NY,NT,ALPHA,SIGMA,RAU,CLIMA,SEQ,COUNT,
C & DER,DER2,PHI,NP,NV,K,NRAU)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      REAL         CLIMA (NY,0:NT)
      REAL         MU (2,0:365)
      REAL         SIGMA (NRAU,NV)
      REAL         DER (NP)
      REAL         DER2 (NP,NP)
      REAL         ALPHA (2,NV,NPARM)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

      DO 10, M = 1, 2
        DO 30, I = 0, NT
          MU (M,I) = 0.0
          DO 40, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
40          CONTINUE
30          CONTINUE
10          CONTINUE

      DO 80, I = 1, NP
        DER (I) = 0.0
        DO 90, J = 1, NP
          DER2 (I,J) = 0.0
90          CONTINUE
80          CONTINUE

      CALL M3DER1 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER,
&                NV,K,NRAU)
      CALL M3DER2 (NY,NT,NP,NPARM,COUNT,SEQ,SIGMA,RAU,PHI,DER2,NV,K,
&                NRAU)
      CALL M3DER3 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER2,
&                NV,K,NRAU)
      CALL M3DER4 (NY,NT,NP,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,DER2,NV,K,NRAU)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C FOR MODEL 3.
C -----

```

```

C SUBROUTINE M3DER1 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,
C & RAU,PHI,DER,NV,K,NRAU)
C -----

```

```

      INTEGER      COUNT (NRAU,NY)
      INTEGER      SEQ (NRAU,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER (NP)
      REAL         MU (2,0:NT)
      REAL         SIGMA (NRAU,NV)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

```

```

      DO 850, LL = 1, NPARM
        DO 870, M = 1, 2
          IF (M .EQ. 1) THEN
            N = 2
            NN = 1
            J = 3
            KK = 4
          ELSEIF (M .EQ. 2) THEN
            N = 1
            NN = 2
            J =
            KK = 3
          ENDIF

```

```

C ..... THE VARIABLE DER1 COMPUTES THE DERIVATIVE FOR THE MEAN
C FUNCTION

```

```

      DER1 = 0
      DO 10, IY = 1, NY
        DO 330, T = 1, COUNT (M,IY)
          P = SEQ (M,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
& .NE.-999)) THEN
            MIDDLE = ((CLIMA(IY,P)-MU(M,P))/SIGMA(M,K)-RAU(M,K)*
& ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,K)))
            DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(M,K)+RAU(M,K)*
& PHI(LL,P-1)/SIGMA(M,K))
          ENDIF
330      CONTINUE

      DO 350, T = 1, COUNT (J,IY)
        P = SEQ (J,IY,T)

```

```

      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
      MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(J,K)-RAU(J,K)*
&      ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(J,K)))
      DER1 = DER1+MIDDLE*(RAU(J,K)*PHI(LL,P-1)/SIGMA
&      (J,K))
      ENDIF
350      CONTINUE

```

```

      DO 360, T = 1, COUNT (KK,IY)
      P = SEQ (KK,IY,T)
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
      MIDDLE=((CLIMA(IY,P)-MU(NN,P))/SIGMA(KK,K)-RAU(KK,K)
&      *((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(KK,K)))
      DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(KK,K))
      ENDIF
360      CONTINUE
10      CONTINUE

```

```

      IF (M.EQ. 1) THEN
      DER (LL) = -DER1
      ELSEIF (M.EQ. 2) THEN
      DER(LL+3) = -DER1
      ENDIF

```

```

870      CONTINUE
850      CONTINUE

```

```

C      ..... THE DERIVATIVE WITH RESPECT TO THE AUTOCORRELATION
C      COEFFICIENT IS COMPUTED AS WELL AS THE DERIVATIVE
C      W.R.T. THE STANDARD DEVIATIONS

```

```

      DO 20, IY = 1, NY
      DO 700, T = 1, COUNT (1,IY)
      P = SEQ (1,IY,T)
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
      PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,K)
      PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,K)
      MIDDLE = PART1-RAU(1,K)*PART2
      DER(11) = DER(11)+MIDDLE*PART2
      PART1 = -((CLIMA(IY,P)-MU(1,P))/SIGMA(1,K)**2)
      PART2 = ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,K)**2)
      DER(7)=DER(7)-MIDDLE*(PART1+RAU(1,K)*PART2)-1/SIGMA(1,K)
      ENDIF
700      CONTINUE

```

```

      DO 701, T = 1, COUNT (2,IY)
      P = SEQ (2,IY,T)
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
      PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,K)
      PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,K)

```

```

        MIDDLE = PART1-RAU(2,K)*PART2
        DER(12) = DER(12)+MIDDLE*PART2
        PART1 = -((CLIMA(IY,P)-MU(2,P))/SIGMA(2,K)**2)
        PART2 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,K)**2)
        DER(8)=DER(8)-MIDDLE*(PART1+RAU(2,K)*PART2)-1/SIGMA(2,K)
    ENDIF
701    CONTINUE

    DO 702, T = 1, COUNT (3,IY)
        P = SEQ (3,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&        .NE.-999)) THEN
            PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(3,K)
            PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(3,K)
            MIDDLE = PART1-RAU(3,K)*PART2
            DER(13) = DER(13)+MIDDLE*PART2
            PART1 = -((CLIMA(IY,P)-MU(2,P))/SIGMA(3,K)**2)
            PART2 = ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(3,K)**2)
            DER(9)=DER(9)-MIDDLE*(PART1+RAU(3,K)*PART2)-1/SIGMA(3,K)
        ENDIF
702    CONTINUE

    DO 703, T = 1, COUNT (4,IY)
        P = SEQ (4,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&        .NE.-999)) THEN
            PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(4,K)
            PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(4,K)
            MIDDLE = PART1-RAU(4,K)*PART2
            DER(14) = DER(14)+MIDDLE*PART2
            PART1 = -((CLIMA(IY,P)-MU(1,P))/SIGMA(4,K)**2)
            PART2 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(4,K)**2)
            DER(10)=DER(10)-MIDDLE*(PART1+RAU(4,K)*PART2)-1/SIGMA(4,K)
        ENDIF
703    CONTINUE
20    CONTINUE

    RETURN
    END

```

```

C      -----
C      ..... SUBROUTINE TO SUBTRACT TWO MATRICES
C      -----

C      SUBROUTINE SUBTR (CLAGO,TERM,NV)
C      -----

      REAL          CLAGO (NV,NV)
      REAL          TERM (NV,NV)

      DO 10, I = 1, NV
        DO 20, J = 1, NV
          TERM (I,J) = CLAGO (I,J) - TERM (I,J)
20      CONTINUE
10     CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE FOLLOWING 2ND DERIVATIVES:
C      ALPHAJD-ALPHAID, ALPHAJD-ALPHAIW, ALPHAIW-ALPHAJW FOR
C      MODEL 3.
C -----

```

```

C SUBROUTINE M3DER2 (NY,NT,NP,NPARM,COUNT,SEQ,SIGMA,RAU,PHI,
C -----
C      &      DER2,NV,K,NRAU)
C -----

```

```

      INTEGER      COUNT (NRAU,NY)
      INTEGER      SEQ (NRAU,NY,NT)
      INTEGER      T,P
      REAL         MIDDLE
      REAL         DER2 (NP,NP)
      REAL         SIGMA (NRAU,NV)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

```

```

      OPEN (UNIT=9,FILE='LPT1')

```

```

      DO 10, LL = 1, NPARM
        DO 20, LLL = 1, NPARM
          DO 30, M = 1, 2
            IF (M .EQ. 1) THEN
              N = 2
              NN = 1
              J = 3
              KK = 4
            ELSEIF (M .EQ. 2) THEN
              N = 1
              NN = 2
              J = 4
              KK = 3
            ENDIF

```

```

C ..... THE VARIABLE DER COMPUTES THE 2ND DERIVATIVES FOR
C      ALPHAD-ALPHAD AND ALPHAW-ALPHAW WHILE DER3 COMPUTES
C      ALPHAD-ALPHAW

```

```

      DER = 0
      DER3 = 0
      DO 40, IY = 1, NY
        DO 50, T = 1, COUNT (M,IY)
          P = SEQ (M,IY,T)
          PART = (-PHI(LL,P)/SIGMA(M,K))+RAU(M,K)*PHI(LL,P-1)
          &      /SIGMA(M,K)
          PART2 = (-PHI(LLL,P)/SIGMA(M,K))+RAU(M,K)*PHI
          &      (LLL,P-1)/SIGMA(M,K)
          DER = DER+PART*PART2
50      CONTINUE

```

```

DO 60, T = 1, COUNT (J,IY)
  P = SEQ (J,IY,T)
  PART = (RAU(J,K)*PHI(LL,P-1)/SIGMA(J,K))
  PART2 = (RAU(J,K)*PHI(LL,P-1)/SIGMA(J,K))
  DER = DER+PART*PART2
60  CONTINUE

DO 70, T = 1, COUNT (KK,IY)
  P = SEQ (KK,IY,T)
  PART = (-PHI(LL,P)/SIGMA(KK,K))
  PART2 = (-PHI(LL,P)/SIGMA(KK,K))
  DER = DER+PART*PART2
70  CONTINUE
40  CONTINUE

IF (M .EQ. 1) THEN
  DER2 (LL,LLL) = -DER
ELSEIF (M .EQ. 2) THEN
  DER2 (LL+3,LLL+3) = -DER
ENDIF

30  CONTINUE
DO 80, IY = 1, NY
  DO 90, T = 1, COUNT (3,IY)
    P = SEQ (3,IY,T)
    PART = (RAU(3,K)*PHI(LL,P-1)/SIGMA(3,K))
    DER3 = DER3+PART*(-PHI(LL,P)/SIGMA(3,K))
90  CONTINUE

    DO 100, T = 1, COUNT (4,IY)
      P = SEQ (4,IY,T)
      PART = (-PHI(LL,P)/SIGMA(4,K))
      DER3 = DER3+PART*(RAU(4,K)*PHI(LL,P-1)/SIGMA(4,K))
100  CONTINUE
80  CONTINUE

  DER2 (LL,LLL+3) = -DER3
20  CONTINUE
10  CONTINUE

RETURN
END

```

```

C  SUBROUTINE TMULT(MAT2,M2,N2,PROD,M3,N3,II,KK,JJ)
   -----
   REAL      MAT2(M2,N2), PROD(M3,N3)

   DO 7000 I = 1,II,1
     DO 7010 J = 1,JJ,1
       PROD(I,J) = 0.0
       DO 7020 K = 1,KK,1
         PROD(I,J) = PROD(I,J) + MAT2(K,I) * MAT2(K,J)
7020  CONTINUE
7010  CONTINUE
7000  CONTINUE
      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE 2ND DERIVATIVES FOR:
C       RAU-ALPHAJD, ALPHAJW-SIGMA FOR MODEL3.
C -----

C SUBROUTINE M3DER3 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,
C -----
C &              PHI,DER2,NV,K,NRAU)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER2 (NP,NP)
      REAL         MU (2,0:NT)
      REAL         SIGMA (NRAU,NV)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

      OPEN (UNIT=9,FILE='LPT1')

      DO 850, LL = 1, NPARM
        DO 870, M = 1, 2
          IF (M .EQ. 1) THEN
            N = 2
            NN = 1
            J = 3
            KK = 4
          ELSEIF (M .EQ. 2) THEN
            N = 1
            NN = 2
            J = 4
            KK = 3
          ENDIF

C ..... THE VARIABLE DER1 COMPUTES THE 2ND DERIVATIVES FOR RAU-
C       ALPHA, WHILE DER4 COMPUTES THE 2ND DERIVATIVES FOR ALPHA-
C       SIGMA

          DER1 = 0
          DER4 = 0
          DO 10, IY = 1, NY
            DO 530, T = 1, COUNT (M,IY)
              P = SEQ (M,IY,T)
              IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
C &              .NE.-999)) THEN
                PART1 = (CLIMA(IY,P)-MU(M,P))/SIGMA(M,K)
                PART2 = (CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,K)
                MIDDLE = PART1-RAU(M,K)*PART2
C &              DER1 = DER1+MIDDLE*(-PHI(LL,P-1)/SIGMA(M,K))+
                (-PHI(LL,P)/SIGMA(M,K)+RAU(M,K)*PHI(LL,P-1)

```

```

&          /SIGMA(M,K))*PART2
PART1 = ((CLIMA(IY,P)-MU(M,P))/SIGMA(M,K)**2)
PART2 = ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,K)**2)
PART5 = PHI(LL,P)
PART6 = PHI(LL,P-1)
DER4 = DER4+MIDDLE*(PART5/(SIGMA(M,K)**2)-RAU
&          (M,K)*PART6/(SIGMA(M,K)**2))+(-PART1
&          +RAU(M,K)*PART2)*(-PART5/SIGMA
&          (M,K)+RAU(M,K)*PART6/SIGMA(M,K))
      ENDIF
530      CONTINUE
10      CONTINUE
      IF (M .EQ. 1) THEN
        DER2(LL,11) = DER1
        DER2(LL,7) = -DER4
      ELSEIF (M .EQ. 2) THEN
        DER2(LL+3,12) = DER1
        DER2(LL+3,8) = -DER4
      ENDIF

      DER1 = 0
      DER4 = 0
      DO 20, IY = 1, NY
        DO 550, T = 1, COUNT (J,IY)
          P = SEQ (J,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&          .NE.-999)) THEN
            PART1 = (CLIMA(IY,P)-MU(N,P))/SIGMA(J,K)
            PART2 = (CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(J,K)
            MIDDLE = PART1-RAU(J,K)*PART2
            DER1 = DER1+MIDDLE*(-PHI(LL,P-1)/SIGMA(J,K))+
&          (RAU(J,K)*PHI(LL,P-1)/SIGMA(J,K))*PART2
            PART1 = ((CLIMA(IY,P)-MU(N,P))/SIGMA(J,K)**2)
            PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(J,K)**2)
            PART5 = PHI(LL,P)
            PART6 = PHI(LL,P-1)
            DER4=DER4+MIDDLE*(-RAU(J,K)*PART6/(SIGMA(J,K)**2))
&          +(-PART1+RAU(J,K)*PART2)*RAU(J,K)*PART6
&          /SIGMA(J,K)
          ENDIF
550      CONTINUE
20      CONTINUE
      IF (M .EQ. 1) THEN
        DER2(LL,13) = DER1
        DER2(LL,9) = -DER4
      ELSEIF (M .EQ. 2) THEN
        DER2 (LL+3,14) = DER1
        DER2(LL+3,10) = -DER4
      ENDIF

      DER1 = 0
      DER4 = 0
      DO 30, IY = 1, NY
        DO 560, T = 1, COUNT (KK,IY)

```

```

      P = SEQ (KK,IY,T)
      IF (CLIMA(IY,P-1).NE.-999) THEN
        PART2 = (CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(KK,K)
        DER1 = DER1+(-PHI(LL,P)/SIGMA(KK,K))*PART2
      ENDIF
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
        PART1 = (CLIMA(IY,P)-MU(NN,P))/SIGMA(KK,K)
        MIDDLE = PART1-RAU(KK,K)*PART2
        PART1 = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(KK,K)**2)
        PART2 = ((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(KK,K)**2)
        PART5 = PHI(LL,P)
        PART6 = PHI(LL,P-1)
        DER4=DER4+MIDDLE*(PART5/(SIGMA(KK,K)**2))+
&      (-PART1+RAU(KK,K)*PART2)*(-PART5/SIGMA(KK,K))

      ENDIF
560      CONTINUE
30      CONTINUE
      IF (M .EQ. 1) THEN
        DER2(LL,14) = DER1
        DER2(LL,10) = -DER4
      ELSEIF (M .EQ. 2) THEN
        DER2(LL+3,13) = DER1
        DER2(LL+3,9) = -DER4
      ENDIF
870      CONTINUE
850      CONTINUE

      RETURN
      END

```

```

C      SUBROUTINE COPY(MAT1,M1,N1,MAT2,M2,N2,DIM1,DIM2)
      -----

      INTEGER DIM1,DIM2
      REAL MAT1(M1,N1), MAT2(M2,N2)

      DO 10020 I = 1,M2,1
        DO 10030 J = 1,N2,1
          MAT2(I,J) = 0.0
10030      CONTINUE
10020      CONTINUE
      DO 10000 I = 1,DIM1,1
        DO 10010 J = 1,DIM2,1
          MAT2(I,J) = MAT1(I,J)
10010      CONTINUE
10000      CONTINUE
      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE 2ND DERIVATIVES FOR:
C       RAU-SIGMA, RAU-RAU, SIGMA-SIGMA FOR MODEL3
C -----

SUBROUTINE M3DER4 (NY,NT,NP,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,DER2,
C -----
C       & NV,K,NRAU)
C -----

      INTEGER      COUNT (NRAU,NY)
      INTEGER      SEQ (NRAU,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER2 (NP,NP)
      REAL         MU (2,0:NT)
      REAL         SIGMA (NRAU,NV)
      REAL         RAU (NRAU,NV)

      OPEN (UNIT=9,FILE='LPT1')

C ..... THE 2ND DERIVATIVE FOR RAU-RAU, SIGMA-SIGMA AND
C       RAU-SIGMA ARE COMPUTED

      DO 20, IY = 1, NY
        DO 330, T = 1, COUNT (1,IY)
          P = SEQ (1,IY,T)
          IF (CLIMA(IY,P-1).NE.-999) THEN
            PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,K)
            DER2 (11,11) = DER2 (11,11)-(PART2**2)
          ENDIF
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&           .NE.-999)) THEN
            MIDDLE = ((CLIMA(IY,P)-MU(1,P))/SIGMA(1,K)-RAU(1,K)*
&                    ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,K)))
            PART3 = ((CLIMA(IY,P)-MU(1,P))/SIGMA(1,K)**2)
            PART4 = ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,K)**2)
            DER2(7,7) = DER2(7,7)+MIDDLE*(2*PART3/SIGMA(1,K)-2*
&                    RAU(1,K)*PART4/SIGMA(1,K))+((-PART3+PART4
&                    *RAU(1,K))**2)-1/(SIGMA(1,K)**2)
            DER2(11,7) = DER2(11,7)+MIDDLE*(-PART4)+(-PART3+RAU(1,K)
&                    *PART4)*PART2
          ENDIF
330      CONTINUE

        DO 340, T = 1, COUNT (2,IY)
          P = SEQ (2,IY,T)
          IF (CLIMA(IY,P-1).NE.-999) THEN
            PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,K)
            DER2 (12,12) = DER2(12,12)-(PART2**2)
          ENDIF

```

```

      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
        MIDDLE = ((CLIMA(IY,P)-MU(2,P))/SIGMA(2,K)-RAU(2,K)*
&      ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,K)))
        PART3 = ((CLIMA(IY,P)-MU(2,P))/SIGMA(2,K)**2)
        PART4 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,K)**2)
        DER2(8,8)=DER2(8,8)+MIDDLE*(2*PART3/SIGMA(2,K)-2*
&      RAU(2,K)*PART4/SIGMA(2,K))+((-PART3+PART4
&      *RAU(2,K))**2)-1/(SIGMA(2,K)**2)
        DER2(12,8)=DER2(12,8)+MIDDLE*(-PART4)+((-PART3+RAU(2,K)
&      *PART4)*PART2
      ENDIF
340    CONTINUE

    DO 350, T = 1, COUNT (3,IY)
      P = SEQ (3,IY,T)
      IF (CLIMA(IY,P-1).NE.-999) THEN
        PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(3,K)
        DER2 (13,13) = DER2 (13,13)-(PART2**2)
      ENDIF
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
        MIDDLE = ((CLIMA(IY,P)-MU(2,P))/SIGMA(3,K)-RAU(3,K)*
&      ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(3,K)))
        PART3 = ((CLIMA(IY,P)-MU(2,P))/SIGMA(3,K)**2)
        PART4 = ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(3,K)**2)
        DER2(9,9)=DER2(9,9)+MIDDLE*(2*PART3/SIGMA(3,K)-2*
&      RAU(3,K)*PART4/SIGMA(3,K))+((-PART3+PART4
&      *RAU(3,K))**2)-1/(SIGMA(3,K)**2)
        DER2(13,9)=DER2(13,9)+MIDDLE*(-PART4)+((-PART3+RAU(3,K)
&      *PART4)*PART2
      ENDIF
350    CONTINUE

    DO 360, T = 1, COUNT (4,IY)
      P = SEQ (4,IY,T)
      IF (CLIMA(IY,P-1).NE.-999) THEN
        PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(4,K)
        DER2 (14,14) = DER2 (14,14)-(PART2**2)
      ENDIF
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1)
&      .NE.-999)) THEN
        MIDDLE = ((CLIMA(IY,P)-MU(1,P))/SIGMA(4,K)-RAU(4,K)*
&      ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(4,K)))
        PART3 = ((CLIMA(IY,P)-MU(1,P))/SIGMA(4,K)**2)
        PART4 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(4,K)**2)
        DER2(10,10)=DER2(10,10)+MIDDLE*(2*PART3/SIGMA(4,K)-2*
&      RAU(4,K)*PART4/SIGMA(4,K))+((-PART3+PART4
&      *RAU(4,K))**2)-1/(SIGMA(4,K)**2)
        DER2(14,10)=DER2(14,10)+MIDDLE*(-PART4)+((-PART3+RAU(4,K)
&      *PART4)*PART2
      ENDIF
360    CONTINUE
20    CONTINUE
    DER2(7,7) = -DER2(7,7)
    DER2(8,8) = -DER2(8,8)
    DER2(9,9) = -DER2(9,9)
    DER2(10,10) = -DER2(10,10)

```

```

RETURN
END

```

```

C -----
C ..... THIS SUBROUTINE ITERATIVELY ESTIMATES THE MODEL
C PARAMETERS BY THE NEWTON-RAPHSON METHOD (M4).
C -----

SUBROUTINE NEWT4 (ALPHA,PSI,RAU,NP,NT,NY,CLIMA,SEQ,
C -----
& COUNT,DER,DER2,PHI,EPS,NP,NV,K,A,THETA,NRAU,CONVG)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      REAL         A (NP,0:NP)
      REAL         PSI (2,NV,NP)
      REAL         ALPHA (2,NV,NP)
      REAL         PHI (NP,0:NT)
      REAL         DER (NP)
      REAL         DER2 (NP,NP)
      REAL         CLIMA (NY,0:NT)
      REAL         THETA (NP)
      REAL         RAU (NRAU,NV)

15  FORMAT (' THE SUCCESSIVE THETA VALUES FOR VARIABLE: ', I4)
25  FORMAT (' .... DID NOT CONVERGE')
35  FORMAT ('/, ' .... ', I3, ' ITERATION', /)

      OPEN (UNIT=9,FILE='LPT1')

      IC = 0
      WRITE (9,15) K

C ..... TRANSFORMING THE SEPARATE PARAMETER ARRAYS INTO ONE
C VECTOR

      DO 20, J = 1, NP
        THETA (J) = ALPHA (1,K,J)
        THETA (J+3) = ALPHA (2,K,J)
        THETA (J+6) = PSI (1,K,J)
        THETA (J+9) = PSI (2,K,J)
20  CONTINUE
      DO 70, J = 1, NRAU
        THETA (J+12) = RAU (J,K)
70  CONTINUE

C ..... ITERATIVE PARAMETER ESTIMATION

      DO 10, ITER = 1, MAXITER

C ..... VECTOR OF 1ST DERIVATIVES AND MATRIX OF 2ND DERIVATIVES
C IS COMPUTED

      CALL M4DERV (NP,NT,NY,ALPHA,PSI,RAU,CLIMA,SEQ,COUNT,
& DER,DER2,PHI,NP,NV,K,NRAU)

```

```

        DO 40, KK = 1, NP
          DO 50, J = KK, NP
            DER2 (J, KK) = DER2 (KK, J)
50          CONTINUE
40          CONTINUE
          PRINT 35, ITER

C      ..... NEW PARAMETER ESTIMATES ARE COMPUTED

          CALL NEWPARM (NP, DER, DER2, THETA, EPS, IC, A)

C      ..... UPDATE PARAMETER ESTIMATES

          DO 30 J = 1, NPARM
            ALPHA (1, K, J) = THETA (J)
            ALPHA (2, K, J) = THETA (J+3)
            PSI (1, K, J) = THETA (J+6)
            PSI (2, K, J) = THETA (J+9)
30          CONTINUE

          DO 80, J = 1, NRAU
            RAU (J, K) = THETA (J+12)
80          CONTINUE

C      ..... TEST FOR CONVERGENCE

          IF (IC) 10, 10, 60
10          CONTINUE
            WRITE (9, 25)
            CONVG = 0

60          RETURN
          END

C      SUBROUTINE PMAT(MAT, M, N, DIM1, DIM2)
      -----

      REAL MAT(M, N)
      INTEGER DIM1, DIM2

      OPEN (UNIT=12, FILE=' \WATER\DATA\RESI.DAT', STATUS='UNKNOWN')

CC      *** THIS ROUTINE PRINTS OUT A MATRIX OF SIZE M BY N
CC      *** EACH ELEMENT IS PRINT IN A FIELD OF . CHARACTERS WITH
CC      *** TWO DECIMAL PLACES (I.E. NNN NNN.NN)

      DO 50 I = 1, DIM1, 1
        WRITE (12, 510) (MAT(I, J), J = 1, DIM2)
510      FORMAT(' ', 7(F10.4))
50      CONTINUE
      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C AND THE MATRIX OF SECOND DERIVATIVES FOR MODEL4.
C -----

SUBROUTINE M4DERV (NPARM,NY,NT,ALPHA,PSI,RAU,CLIMA,SEQ,COUNT,
C -----
& DER,DER2,PHI,NP,NV,K,NRAU)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      REAL         CLIMA (NY,0:NT)
      REAL         MU (2,0:365)
      REAL         SIGMA (2,0:365)
      REAL         DER (NP)
      REAL         DER2 (NP,NP)
      REAL         PSI (2,NV,NPARM)
      REAL         ALPHA (2,NV,NPARM)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

      DO 10, M = 1, 2
        DO 30, I = 0, NT
          MU (M,I) = 0.0
          SIGMA (M,I) = 0.0
          DO 40, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
            SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
40          CONTINUE
30          CONTINUE
10          CONTINUE

      DO 80, I = 1, NP
        DER (I) = 0.0
        DO 90, J = 1, NP
          DER2 (I,J) = 0.0
90          CONTINUE
80          CONTINUE

      CALL M4DER1(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER,
& NV,K,NRAU)
      CALL M4DER2(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER2,
& NV,K,NRAU)
      CALL M4DER3(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER2,
& NV,K,NRAU)
      CALL M4DER4(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER2,
& NV,K,NRAU)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C FOR MODEL4.
C -----

```

```

C SUBROUTINE M4DER1 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,
C & RAU,PHI,DER,NV,K,NRAU)
C -----

```

```

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER (NP)
      REAL         MU (2,0:365)
      REAL         SIGMA (2,0:365)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

```

```

      DO 850, LL = 1, NPARM
        DO 870, M = 1, 2
          IF (M .EQ. 1) THEN
            N = 2
            NN = 1
            J = 3
            KK = 4
          ELSEIF (M .EQ. 2) THEN
            N = 1
            NN = 2
            J = 4
            KK = 3
          ENDIF

```

```

C ..... THE VARIABLE DER1 COMPUTES THE DERIVATIVE FOR THE MEAN
C FUNCTION, WHILE DER2 AND DER3 COMPUTE THE DERIVATIVE FOR
C THE VARIANCE FUNCTION

```

```

      DER2 = 0
      DER1 = 0
      DER3 = 0
      DO 10, IY = 1, NY
        DO 330, T = 1, COUNT (M,IY)
          P = SEQ (M,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&            MIDDLE = ((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)-RAU(1,K)*
&            ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)))
&            DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(M,P)+RAU(1,K)*
&            PHI(LL,P-1)/SIGMA(M,P-1))
&            PART1 = (-((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)**2)*
&            PHI(LL,P))

```

```

      PART2 = ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)**2)
&          *PHI(LL,P-1)
      DER3 = DER3+MIDDLE*(PART1+RAU(1,K)*PART2)
      ENDIF
      DER2 = DER2+PHI(LL,P)/SIGMA(M,P)
330    CONTINUE

      DO 350, T = 1, COUNT (J,IY)
        P = SEQ (J,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&          MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(N,P)-RAU(1,K)*
&          ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
&          DER1 = DER1+MIDDLE*(RAU(1,K)*PHI(LL,P-1)/SIGMA
&          (NN,P-1))
&          PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)
&          **2)*PHI(LL,P-1)
&          DER3 = DER3+MIDDLE*(RAU(1,K)*PART2)
&          ENDIF
350    CONTINUE

      DO 360, T = 1, COUNT (KK,IY)
        P = SEQ (KK,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&          MIDDLE = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(1,K)
&          *((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(N,P-1)))
&          DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(NN,P))
&          PART1 = (-((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)**2)*
&          PHI(LL,P))
&          DER3 = DER3+MIDDLE*PART1
&          ENDIF
&          DER2 = DER2+PHI(LL,P)/SIGMA(NN,P)
360    CONTINUE
10    CONTINUE

      IF (M .EQ. 1) THEN
        DER (LL) = -DER1
        DER (LL+6) = (-DER3-DER2)
      ELSEIF (M .EQ. 2) THEN
        DER(LL+3) = -DER1
        DER(LL+9) = (-DER3-DER2)
      ENDIF
870    CONTINUE
850    CONTINUE

C      ..... THE DERIVATIVE WITH RESPECT TO THE AUTOCORRELATION
C      COEFFICIENT IS COMPUTED

      DER (NP) = 0
      DO 20, IY = 1, NY
        DO 700, T = 1, COUNT (1,IY)
          P = SEQ (1,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))

```

```

&          THEN
          PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
          PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
          DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
        ENDIF
700    CONTINUE

        DO 701, T = 1, COUNT (2,IY)
          P = SEQ (2,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&          THEN
            PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
            PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
          ENDIF
701    CONTINUE

        DO 702, T = 1, COUNT (3,IY)
          P = SEQ (3,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&          THEN
            PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
            PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
          ENDIF
702    CONTINUE

        DO 703, T = 1, COUNT (4,IY)
          P = SEQ (4,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&          THEN
            PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
            PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
          ENDIF
703    CONTINUE
20    CONTINUE

        RETURN
        END

```

```

C -----
C ..... SUBROUTINE TO COMPUTE THE TRANSPOSE OF A MATRIX
C -----

```

```

C SUBROUTINE TRNSP (PHI,NP,NTT,TRSP,NPARM,NT)
C -----

```

```

      REAL          PHI (NT,NPARM)
      REAL          TRSP (NPARM,NT)

      DO 10, I = 1, NP
        DO 20, J = 1, NTT
          TRSP (I,J) = PHI (J,I)
20      CONTINUE
10     CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE FOLLOWING 2ND DERIVATIVES:
C      ALPHAjD-ALPHAiD, PSIjD-PSIiD, ALPHAjD-PSIiD, ALPHAjW-
C      ALPHAiW, PSIjW-PSIiW, ALPHAjW-PSIiW FOR MODEL 4
C -----

```

```

C SUBROUTINE M4DER2 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,
C & -----
C      RAU,PHI,DER2,NV,K,NRAU)
C -----

```

```

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      INTEGER      T,P
      REAL          CLIMA (NY,0:NT)
      REAL          MIDDLE
      REAL          DER2 (NP,NP)
      REAL          MU (2,0:365)
      REAL          SIGMA (2,0:365)
      REAL          PHI (NPARM,0:NT)
      REAL          RAU (NRAU,NV)

```

```

      DO 850, LL = 1, NPARM
        DO 870, LLL = 1, NPARM
          DO 880, M = 1, 2
            IF (M .EQ. 1) THEN
              N = 2
              NN = 1
              J = 3
              KK = 4
            ELSEIF (M .EQ. 2) THEN
              N = 1
              NN = 2
              J = 4
              KK = 3
            ENDIF

```

```

C ..... THE VARIABLE DER COMPUTES THE 2ND DERIVATIVES FOR
C      ALPHA-ALPHA, DER3 THE DERIVATIVES PSI-PSI AND DER4 THE
C      DERIVATIVES ALPHA-PSI

```

```

      DER = 0
      DER3 = 0
      DER4 = 0
      DO 10, IY = 1, NY
        DO 330, T = 1, COUNT (M,IY)
          P = SEQ (M,IY,T)
          PART = (-PHI(LL,P)/SIGMA(NN,P))+RAU(1,K)*PHI(LL,P-1)
&          /SIGMA(NN,P-1)
          PART2 = (-PHI(LLL,P)/SIGMA(NN,P))+RAU(1,K)*PHI(LLL,
&          P-1)/SIGMA(NN,P-1)
          DER = DER+PART*PART2
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.

```

```

&          -999)) THEN
MIDDLE = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(1,K)*
&          ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
PART1 = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)**2)
PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)**2)
PART3 = PHI(LLL,P)
PART4 = PHI(LLL,P-1)
PART5 = PHI(LL,P)
PART6 = PHI(LL,P-1)
DER3 = DER3+MIDDLE*(2*PART1/SIGMA(NN,P)*PART3*PART5
&          -2*RAU(1,K)*PART2/SIGMA(NN,P-1)*PART4*PART6)+
&          (-PART1*PART3+PART2*RAU(1,K)*PART4)*(-PART1*
&          PART5+PART2*PART6*RAU(1,K))-PART3*PART5/
&          (SIGMA(NN,P)**2)
DER4 = DER4+MIDDLE*(PART3*PART5/(SIGMA(NN,P)**2)-
&          RAU(1,K)*PART4*PART6/(SIGMA(NN,P-1)**2))+(-
&          PART1*PART3+RAU(1,K)*PART2*PART4)*(-PART5/
&          SIGMA(NN,P)+RAU(1,K)*PART6/SIGMA(NN,P-1))
ENDIF
330 CONTINUE

DO 350, T = 1, COUNT (J,IY)
P = SEQ (J,IY,T)
PART = (RAU(1,K)*PHI(LL,P-1)/SIGMA(NN,P-1))
PART2 = (RAU(1,K)*PHI(LLL,P-1)/SIGMA(NN,P-1))
DER = DER+PART*PART2
IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(N,P)-RAU(1,K)*
&          ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)**2)
PART3 = PHI(LLL,P)
PART4 = PHI(LLL,P-1)
PART5 = PHI(LL,P)
PART6 = PHI(LL,P-1)
DER3 = DER3+MIDDLE*(-2*RAU(1,K)*PART2/SIGMA(NN,P-1)*
&          PART4*PART6)+(RAU(1,K)*PART2*PART4)*(RAU(1,K)*
&          PART2*PART6)
DER4 = DER4+MIDDLE*(-RAU(1,K)*PART4*PART6/(SIGMA(NN
&          ,P-1)**2))+RAU(1,K)*PART2*PART4*RAU(1,K)*PART6/
&          SIGMA(NN,P-1)
ENDIF
350 CONTINUE

DO 360, T = 1, COUNT (KK,IY)
P = SEQ (KK,IY,T)
PART = (-PHI(LL,P)/SIGMA(NN,P))
PART2 = (-PHI(LLL,P)/SIGMA(NN,P))
DER = DER+PART*PART2
IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
MIDDLE = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(1,K)*
&          ((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(N,P-1)))
PART1 = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P) ** 2)

```

```

      PART3 = PHI(LLL,P)
      PART4 = PHI(LLL,P-1)
      PART5 = PHI(LL,P)
      PART6 = PHI(LL,P-1)
      DER3 = DER3+MIDDLE*(2*PART1/SIGMA(NN,P))*PART3*PART5
&          +(-PART1*PART3)*(-PART1*PART5)-PART3*PART5/
&          (SIGMA(NN,P)**2)
      DER4 = DER4+MIDDLE*PART3*PART5/(SIGMA(NN,P)**2)+
&          (-PART1*PART3*(-PART5/SIGMA(NN,P)))
      ENDIF
360      CONTINUE
10      CONTINUE

      IF (M .EQ. 1) THEN
        DER2 (LL,LLL) = -DER
        DER2 (LL+6,LLL+6) = -DER3
        DER2 (LL,LLL+6) = -DER4
      ELSEIF (M .EQ. 2) THEN
        DER2 (LL+3,LLL+3) = -DER
        DER2 (LL+9,LLL+9) = -DER3
        DER2 (LL+3,LLL+9) = -DER4
      ENDIF
880      CONTINUE
870      CONTINUE
850      CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS FUNCTION COMPUTES THE GAMMA FUNCTION OF X GIVEN
C BY:
C THE DEFINITE INTEGRAL BETWEEN 0 & INFINITY OF THE
C FUNCTION:
C          Y ** (X-1) * EXP(-Y)
C w.r.t. Y.
C -----

```

```

C FUNCTION GAMMA (ALPH)
C -----
      A = ALPH
      G = 1
4      IF (A .GE. 10) THEN
        GOTO 2
      ELSE
        G = G * A
        A = A + 1
        GOTO 4
      ENDIF
2      T = (1 + (0.0833333 + 0.00347222 - 0.002681327 / A) / A) / A
      GAMMA = EXP(-1 * A + (A - 0.5) * LOG(A) + 0.918939)*T*A/G

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE 2ND DERIVATIVES FOR:
C       RAU-ALPHAjD, RAU-PSIjD, RAU-RAU, RAU-ALPHAjW AND
C       RAU-PSIjW FOR MODEL 4
C -----

C SUBROUTINE M4DER3 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,
C &
C       PHI,DER2,NV,K,NRAU)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER2 (NP,NP)
      REAL         MU (2,0:365)
      REAL         SIGMA (2,0:365)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

      DO 850, LL = 1, NPARM
        DO 870, M = 1, 2
          IF (M .EQ. 1) THEN
            N = 2
            NN = 1
            J = 3
            KK = 4
          ELSEIF (M .EQ. 2) THEN
            N = 1
            NN = 2
            J = 4
            KK = 3
          ENDIF

C ..... THE VARIABLE DER1 COMPUTES THE 2ND DERIVATIVES FOR RAU-
C       ALPHA, WHILE DER3 COMPUTES THE 2ND DERIVATIVES FOR RAU-
C       PSI

          DER1 = 0
          DER3 = 0
          DO 10, IY = 1, NY
            DO 530, T = 1, COUNT (M,IY)
              P = SEQ (M,IY,T)
              IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
C &               -999)) THEN
                PART1 = (CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)
                PART2 = (CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)
                MIDDLE = PART1-RAU(1,K)*PART2
                DER1 = DER1+MIDDLE*(-PHI(LL,P-1)/SIGMA(M,P-1))+
C &               (-PHI(LL,P)/SIGMA(M,P)+RAU(1,K)*PHI(LL,P-1)
C &               /SIGMA(M,P-1))*PART2

```

```

      DER3 = DER3+MIDDLE*(PART2/SIGMA(M,P-1))*(-PHI
&          (LL,P-1))+((PART1/SIGMA(M,P))*(-PHI(LL,P))+RAU
&          (1,K)*(PART2/SIGMA(M,P-1))*PHI(LL,P-1))*PART2
      ENDIF
530      CONTINUE

      DO 550, T = 1, COUNT (J,IY)
        P = SEQ (J,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
          PART1 = (CLIMA(IY,P)-MU(N,P))/SIGMA(N,P)
          PART2 = (CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)
          MIDDLE = PART1-RAU(1,K)*PART2
          DER1 = DER1+MIDDLE*(-PHI(LL,P-1)/SIGMA(NN,P-1))+
&          (RAU(1,K)*PHI(LL,P-1)/SIGMA(NN,P-1))*PART2
          DER3 = DER3+MIDDLE*(-PHI(LL,P-1)/SIGMA(NN,P-1))*
&          PART2+(RAU(1,K)*PHI(LL,P-1)/SIGMA(NN,P-1))*
&          PART2**2
        ENDIF
550      CONTINUE

      DO 560, T = 1, COUNT (KK,IY)
        P = SEQ (KK,IY,T)
        IF (CLIMA(IY,P-1).NE.-999) THEN
          PART2 = (CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(N,P-1)
          DER1 = DER1+(-PHI(LL,P)/SIGMA(NN,P))*PART2
          ENDIF
&        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
          PART3 = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P))
          MIDDLE = (PART3-RAU(1,K)*PART2)
          DER3 = DER3+(-PHI(LL,P)/SIGMA(NN,P))*PART3*PART2
          ENDIF
560      CONTINUE
10      CONTINUE

      IF (M.EQ. 1) THEN
        DER2(LL,NP) = DER1
        DER2(LL+6,NP) = DER3
      ELSEIF (M.EQ. 2) THEN
        DER2 (LL+3,NP) = DER1
        DER2 (LL+9,NP) =DER3
      ENDIF
870      CONTINUE
850      CONTINUE

```

C ..... THE 2ND DERIVATIVE RAU-RAU IS COMPUTED

```

      DER2 (NP,NP) = 0
      DO 20, IY = 1, NY
        DO 330, T = 1, COUNT (1,IY)
          P = SEQ (1,IY,T)
          IF (CLIMA(IY,P-1).NE.-999) THEN
            PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)

```

```

        DER2(NP,NP) = DER2(NP,NP)+PART2**2
        ENDIF
330    CONTINUE

        DO 340, T = 1, COUNT (2,IY)
            P = SEQ (2,IY,T)
            IF (CLIMA(IY,P-1).NE.-999) THEN
                PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
                DER2(NP,NP) = DER2(NP,NP)+PART2**2
            ENDIF
340    CONTINUE

        DO 350, T = 1, COUNT (3,IY)
            P = SEQ (3,IY,T)
            IF (CLIMA(IY,P-1).NE.-999) THEN
                PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
                DER2(NP,NP) = DER2(NP,NP)+PART2**2
            ENDIF
350    CONTINUE

        DO 360, T = 1, COUNT (4,IY)
            P = SEQ (4,IY,T)
            IF (CLIMA(IY,P-1).NE.-999) THEN
                PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
                DER2(NP,NP) = DER2(NP,NP)+PART2**2
            ENDIF
360    CONTINUE
20    CONTINUE
        DER2 (NP,NP) = -DER2 (NP,NP)

        RETURN
        END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES PI=PROBABILITY THAT A WET DAY
C          FOLLOWS A WET DAY OR THE PROBABILITY THAT A WET DAY
C          FOLLOWS A DRY DAY.
C -----

```

```

C SUBROUTINE PIEST (NP,GAM,STATE,K,PHI,PI,NT)
C -----

        INTEGER          STATE
        REAL             LAMBDA
        REAL             GAM (2,NP)
        REAL             PHI (NP,0:NT)
        REAL             PI

        LAMBDA = 0
        DO 10, I = 1, NP
            LAMBDA = LAMBDA + GAM (STATE,I) * PHI (I,K)
10    CONTINUE
        PI = EXP (LAMBDA) / (1 + EXP (LAMBDA))

        RETURN
        END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE 2ND DERIVATIVES FOR:
C     ALPHAjD-ALPHAiW, ALPHAjD-PSIiW, PSIjD-PSIiW AND
C     ALPHAjW-PSIiD FOR MODEL 4.
C -----

```

```

C SUBROUTINE M4DER4 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,
C & -----
C     PHI,DER2,NV,K,NRAU)
C -----

```

```

C     INTEGER      COUNT (4,NY)
C     INTEGER      SEQ (4,NY,NT)
C     INTEGER      T,P
C     REAL         CLIMA (NY,0:NT)
C     REAL         DER2 (NP,NP)
C     REAL         MU (2,0:365)
C     REAL         SIGMA (2,0:365)
C     REAL         PHI (NPARM,0:NT)
C     REAL         RAU (NRAU,NV)

```

```

C     DO 850, LL = 1, NPARM
C       DO 870, LLL = 1, NPARM

```

```

C ..... THE VARIABLE DER COMPUTES THE 2ND DERIVATIVE ALPHAD-
C     ALPHAW, DER3 THE DERIVATIVE ALPHAD-PSIW, DER4 THE
C     DERIVATIVE PSID-PSIW AND DER5 THE DERIVATIVE ALPHAW-
C     PSID

```

```

C     DER = 0
C     DER3 = 0
C     DER4 = 0
C     DER5 = 0
C     DO 10, IY = 1, NY
C       DO 350, T = 1, COUNT (3,IY)
C         P = SEQ (3,IY,T)
C         PART = (RAU(1,K)*PHI(LL,P-1)/SIGMA(1,P-1))
C         DER = DER+PART*(-PHI(LLL,P)/SIGMA(2,P))
C         IF (CLIMA(IY,P).NE.-999) THEN
C           DER3 = DER3+PART*(-PHI(LLL,P)/SIGMA(2,P))*((CLIMA
C             (IY,P)-MU(2,P))/SIGMA(2,P))
C           ENDIF
C           IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
C             -999)) THEN
C             PART2 = ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)**2)
C               *PHI(LL,P-1)
C             PART1 = (-((CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)**2)*
C               PHI(LLL,P)
C             DER4 = DER4+(PART1*RAU(1,K)*PART2)
C             ENDIF
C             IF (CLIMA(IY,P-1).NE.-999) THEN
C               PART = (-PHI(LL,P)/SIGMA(2,P))
C               DER5 = DER5+PART*RAU(1,K)*PHI(LLL,P-1)/SIGMA(1,P-1)*

```

```

&          (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
&          ENDIF
350      CONTINUE

      DO 360, T = 1, COUNT (4,IY)
      P = SEQ (4,IY,T)
      PART = (-PHI(LL,P)/SIGMA(1,P))
      DER = DER+PART*(RAU(1,K)*PHI(LLL,P-1)/SIGMA(2,P-1))
      IF (CLIMA(IY,P-1).NE.-999) THEN
&          DER3 = DER3+PART*(RAU(1,K)*PHI(LLL,P-1)/SIGMA(2,P-1)
&          )*(CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
&          ENDIF
&          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&          PART1 = (-((CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)**2)*
&          PHI(LL,P))
&          PART2 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)**2)*
&          PHI(LLL,P-1)
&          DER4 = DER4+RAU(1,K)*PART2*PART1
&          ENDIF
&          IF (CLIMA(IY,P).NE.-999) THEN
&          PART = (RAU(1,K)*PHI(LL,P-1)/SIGMA(2,P-1))
&          DER5 = DER5+PART*(-PHI(LLL,P)/SIGMA(1,P))*(CLIMA
&          (IY,P)-MU(1,P))/SIGMA(1,P)
&          ENDIF
360      CONTINUE
10      CONTINUE

      DER2 (LL,LLL+3) = -DER
      DER2 (LL,LLL+9) = -DER3
      DER2 (LL+6,LLL+9) = -DER4
      DER2 (LL+3,LLL+6) = -DER5
870      CONTINUE
850      CONTINUE

      RETURN
      END

```

```

C      SUBROUTINE PPMAT(MAT,M,N,DIM1,DIM2)
C      -----

      REAL MAT(M,N)
      INTEGER DIM1,DIM2

      OPEN (UNIT=9,FILE='LPT1')

      CC      *** THIS ROUTINE PRINTS OUT A MATRIX OF SIZE M BY N
      CC      *** EACH ELEMENT IS PRINT IN A FIELD OF . CHARACTERS WITH
      CC      *** TWO DECIMAL PLACES (I.E. NNN NNN.NN)

      WRITE (9,5020)
      DO 5000 I = 1,DIM1,1
      WRITE (9,5010) (MAT(I,J), J = 1,DIM2)
5010      FORMAT(' ',7(F15.6))
5000      CONTINUE
      WRITE (9,5020)
5020      FORMAT(/)
      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE ITERATIVELY ESTIMATES THE MODEL
C PARAMETERS BY THE NEWTON-RAPHSON METHOD FOR M5.
C -----

SUBROUTINE NEWT5 (ALPHA,PSI,RAU,NP,NP,MAXITER,NT,NY,CLIMA,SEQ,
C -----
C & COUNT,DER,DER2,PHI,EPS,NP,NV,K,A,THETA,NRAU,CONVG)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      REAL         A (NP,0:NP)
      REAL         PSI (2,NV,NP,NP)
      REAL         ALPHA (2,NV,NP,NP)
      REAL         PHI (NP,NP,0:NT)
      REAL         DER (NP)
      REAL         DER2 (NP,NP)
      REAL         CLIMA (NY,0:NT)
      REAL         THETA (NP)
      REAL         RAU (NRAU,NV)

15  FORMAT (' THE SUCCESSIVE THETA VALUES FOR VARIABLE: ', I4)
25  FORMAT (' .... DID NOT CONVERGE')
35  FORMAT ('/, ' .... ', I3, ' ITERATION', '/')

      OPEN (UNIT=9,FILE='LPT1')

      IC = 0
      WRITE (9,15) K

C ..... TRANSFORMING THE SEPARATE PARAMETER ARRAYS INTO ONE
C VECTOR

      DO 20, J = 1, NP,NP
          THETA (J) = ALPHA (1,K,J)
          THETA (J+3) = ALPHA (2,K,J)
          THETA (J+6) = PSI (1,K,J)
          THETA (J+9) = PSI (2,K,J)
20  CONTINUE
      DO 70, J = 1, NRAU
          THETA (J+12) = RAU (J,K)
70  CONTINUE

C ..... ITERATIVE PARAMETER ESTIMATION

      DO 10, ITER = 1, MAXITER

C ..... VECTOR OF 1ST DERIVATIVES AND MATRIX OF 2ND DERIVATIVES
C IS COMPUTED

      CALL M5DERV (NP,NP,NT,ALPHA,PSI,RAU,CLIMA,SEQ,COUNT,
&      DER,DER2,PHI,NP,NV,K,NRAU)

```

```

      DO 40, KK = 1, NP
        DO 50, J = KK, NP
          DER2 (J, KK) = DER2 (KK, J)
50      CONTINUE
40      CONTINUE
      PRINT 35, ITER

C      ..... NEW PARAMETER ESTIMATES ARE COMPUTED

      CALL NEWPARM (NP, DER, DER2, THETA, EPS, IC, A)

C      ..... UPDATE PARAMETER ESTIMATES

      DO 30 J = 1, NPARM
        ALPHA (1, K, J) = THETA (J)
        ALPHA (2, K, J) = THETA (J+3)
        PSI (1, K, J) = THETA (J+6)
        PSI (2, K, J) = THETA (J+9)
30      CONTINUE

      DO 80, J = 1, NRAU
        RAU (J, K) = THETA (J+12)
80      CONTINUE

C      ..... TEST FOR CONVERGENCE

      IF (IC) 10, 10, 60
10      CONTINUE
      WRITE (9, 25)
      CONVG = 0

60      RETURN
      END

C      -----
C      ..... FUNCTION OF ONE VARIABLE
C      -----

FUNCTION DIM1 (X)
C      -----

      INTEGER          NPMAX
      PARAMETER        (NPMAX=20)

      COMMON /ONE/ NPP, THET(NPMAX), DERI(NPMAX)
      DIMENSION        XT(NPMAX)

      OPEN (UNIT=9, FILE='LPT1')

      DO 10, J=1, NPP
        XT(J)=THET(J)+X*DERI(J)
10      CONTINUE
      DIM1=FUNC (XT)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C AND THE MATRIX OF SECOND DERIVATIVES FOR MODEL5.
C -----

C SUBROUTINE M5DERV (NPARM,NY,NT,ALPHA,PSI,RAU,CLIMA,SEQ,COUNT,
C & DER,DER2,PHI,NP,NV,K,NRAU)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      REAL         CLIMA (NY,0:NT)
      REAL         MU (2,0:365)
      REAL         SIGMA (2,0:365)
      REAL         DER (NP)
      REAL         DER2 (NP,NP)
      REAL         PSI (2,NV,NPARM)
      REAL         ALPHA (2,NV,NPARM)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

      DO 10, M = 1, 2
        DO 30, I = 0, NT
          MU (M,I) = 0.0
          SIGMA (M,I) = 0.0
          DO 40, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
            SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
40          CONTINUE
30        CONTINUE
10      CONTINUE

      DO 80, I = 1, NP
        DER (I) = 0.0
        DO 90, J = 1, NP
          DER2 (I,J) = 0.0
90        CONTINUE
80      CONTINUE

      CALL M5DER1(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER,
& NV,K,NRAU)
      CALL M5DER2(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER2,
& NV,K,NRAU)
      CALL M5DER3(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER2,
& NV,K,NRAU)
      CALL M5DER4(NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,PHI,DER2,
& NV,K,NRAU)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C FOR MODEL 5.
C -----

```

```

C SUBROUTINE M5DER1 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,
C & RAU,PHI,DER,NV,K,NRAU)
C -----

```

```

      INTEGER      COUNT (NRAU,NY)
      INTEGER      SEQ (NRAU,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER (NP)
      REAL         MU (2,0:365)
      REAL         SIGMA (2,0:365)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

```

```

      DO 850, LL = 1, NPARM
        DO 870, M = 1, 2
          IF (M .EQ. 1) THEN
            N = 2
            NN = 1
            J = 3
            KK = 4
          ELSEIF (M .EQ. 2) THEN
            N = 1
            NN = 2
            J = 4
            KK = 3
          ENDIF

```

```

C ..... THE VARIABLE DER1 COMPUTES THE DERIVATIVE FOR THE MEAN
C FUNCTION, WHILE DER2 AND DER3 COMPUTE THE DERIVATIVE FOR
C THE VARIANCE FUNCTION

```

```

      DER2 = 0
      DER1 = 0
      DER3 = 0
      DO 10, IY = 1, NY
        DO 330, T = 1, COUNT (M,IY)
          P = SEQ (M,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&            MIDDLE = ((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)-RAU(M,K)*
&            ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)))
&            DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(M,P)+RAU(M,K)*
&            PHI(LL,P-1)/SIGMA(M,P-1))
&            PART1 = (-((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)**2)*
&            PHI(LL,P))

```

```

PART2 = ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)**2)
&          *PHI(LL,P-1)
DER3 = DER3+MIDDLE*(PART1+RAU(M,K)*PART2)
ENDIF
DER2 = DER2+PHI(LL,P)/SIGMA(M,P)
330 CONTINUE

DO 350, T = 1, COUNT (J,IY)
P = SEQ (J,IY,T)
IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(N,P)-RAU(J,K)*
&          ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
DER1 = DER1+MIDDLE*(RAU(J,K)*PHI(LL,P-1)/SIGMA
&          (NN,P-1))
PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)
&          **2)*PHI(LL,P-1)
DER3 = DER3+MIDDLE*(RAU(J,K)*PART2)
ENDIF
350 CONTINUE

DO 360, T = 1, COUNT (KK,IY)
P = SEQ (KK,IY,T)
IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
MIDDLE=((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(KK,K)
&          *((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(N,P-1)))
DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(NN,P))
PART1 = (-((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)**2)*
&          PHI(LL,P))
DER3 = DER3+MIDDLE*PART1
ENDIF
DER2 = DER2+PHI(LL,P)/SIGMA(NN,P)
360 CONTINUE
10 CONTINUE

IF (M .EQ. 1) THEN
DER (LL) = -DER1
DER (LL+6) = (-DER3-DER2)
ELSEIF (M .EQ. 2) THEN
DER(LL+3) = -DER1
DER(LL+9) = (-DER3-DER2)
ENDIF

870 CONTINUE
850 CONTINUE

C ..... THE DERIVATIVE WITH RESPECT TO THE AUTOCORRELATION
C COEFFICIENT IS COMPUTED

DO 20, IY = 1, NY
DO 700, T = 1, COUNT (1,IY)
P = SEQ (1,IY,T)
IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.

```

```

&          -999)) THEN
          PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
          PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
          DER(13) = DER(13)+(PART1-RAU(1,K)*PART2)*PART2
        ENDIF
700      CONTINUE

      DO 701, T = 1, COUNT (2,IY)
      P = SEQ (2,IY,T)
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
          PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
          PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
          DER(14) = DER(14)+(PART1-RAU(2,K)*PART2)*PART2
        ENDIF
701      CONTINUE

      DO 702, T = 1, COUNT (3,IY)
      P = SEQ (3,IY,T)
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
          PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
          PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
          DER(15) = DER(15)+(PART1-RAU(3,K)*PART2)*PART2
        ENDIF
702      CONTINUE

      DO 703, T = 1, COUNT (4,IY)
      P = SEQ (4,IY,T)
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
          PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
          PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
          DER(NP) = DER(NP)+(PART1-RAU(4,K)*PART2)*PART2
        ENDIF
703      CONTINUE
20      CONTINUE

      RETURN
      END

```

```

C -----
C ..... SUBROUTINE TO COMPUTE THE TRANSPOSE OF A MATRIX
C -----

```

```

C SUBROUTINE TRANSP (PHI,NPARM,NT,TRSP)
C -----

```

```

      REAL          PHI (NT,NPARM)
      REAL          TRSP (NPARM,NT)

      DO 10, I = 1, NPARM
        DO 20, J = 1, NT
          TRSP (I,J) = PHI (J,I)
20      CONTINUE
10      CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE FOLLOWING 2ND DERIVATIVES:
C     ALPHAjD-ALPHAid, PSIjD-PSIid, ALPHAjD-PSIid, ALPHAjW-
C     ALPHAiw, PSIjW-PSIiw, ALPHAjW-PSIiw FOR MODEL 5.
C -----

```

```

C SUBROUTINE MSDER2 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,
C & -----
C     RAU,PHI,DER2,NV,K,NRAU)
C -----

```

```

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER2 (NP,NP)
      REAL         MU (2,0:365)
      REAL         SIGMA (2,0:365)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

```

```

      DO 850, LL = 1, NPARM
        DO 870, LLL = 1, NPARM
          DO 880, M = 1, 2
            IF (M .EQ. 1) THEN
              N = 2
              NN = 1
              J = 3
              KK = 4
            ELSEIF (M .EQ. 2) THEN
              N = 1
              NN = 2
              J = 4
              KK = 3
            ENDIF

```

```

C ..... THE VARIABLE DER COMPUTES THE 2ND DERIVATIVES FOR
C     ALPHA-ALPHA, DER3 THE DERIVATIVES PSI-PSI AND DER4 THE
C     DERIVATIVES ALPHA-PSI

```

```

      DER = 0
      DER3 = 0
      DER4 = 0
      DO 10, IY = 1, NY
        DO 330, T = 1, COUNT (M,IY)
          P = SEQ (M,IY,T)
          PART = (-PHI(LL,P)/SIGMA(NN,P))+RAU(M,K)*PHI(LL,P-1)
C &          /SIGMA(NN,P-1)
          PART2 = (-PHI(LLL,P)/SIGMA(NN,P))+RAU(M,K)*PHI
C &          (LLL,P-1)/SIGMA(NN,P-1)
          DER = DER+PART*PART2
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))

```

```

&          THEN
&          MIDDLE = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(M,K)*
&                    ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
&          PART1 = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)**2)
&          PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)**2)
&          PART3 = PHI(LLL,P)
&          PART4 = PHI(LLL,P-1)
&          PART5 = PHI(LL,P)
&          PART6 = PHI(LL,P-1)
&          DER3 = DER3+MIDDLE*(2*PART1/SIGMA(NN,P)*PART3*PART5
&                    -2*RAU(M,K)*PART2/SIGMA(NN,P-1)*PART4*PART6)+
&                    (-PART1*PART3+PART2*RAU(M,K)*PART4)*(-PART1*
&                    PART5+PART2*PART6*RAU(M,K))-PART3*PART5/
&                    (SIGMA(NN,P)**2)
&          DER4 = DER4+MIDDLE*(PART3*PART5/(SIGMA(NN,P)**2)-RAU
&                    (M,K)*PART4*PART6/(SIGMA(NN,P-1)**2))+(-PART1
&                    *PART3+RAU(M,K)*PART2*PART4)*(-PART5/SIGMA
&                    (NN,P)+RAU(M,K)*PART6/SIGMA(NN,P-1))
&          ENDIF
330      CONTINUE

DO 350, T = 1, COUNT (J,IY)
    P = SEQ (J,IY,T)
    PART = (RAU(J,K)*PHI(LL,P-1)/SIGMA(NN,P-1))
    PART2 = (RAU(J,K)*PHI(LLL,P-1)/SIGMA(NN,P-1))
    DER = DER+PART*PART2
    IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&        THEN
&        MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(N,P)-RAU(J,K)*
&                    ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
&        PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)**2)
&        PART3 = PHI(LLL,P)
&        PART4 = PHI(LLL,P-1)
&        PART5 = PHI(LL,P)
&        PART6 = PHI(LL,P-1)
&        DER3 = DER3+MIDDLE*(-2*RAU(J,K)*PART2/SIGMA(NN,P-1)*
&                    PART4*PART6)+(RAU(J,K)*PART2*PART4)*(RAU(J,K)
&                    *PART2*PART6)
&        DER4 = DER4+MIDDLE*(-RAU(J,K)*PART4*PART6/(SIGMA
&                    (NN,P-1)**2))+RAU(J,K)*PART2*PART4*RAU(J,K)*
&                    PART6/SIGMA(NN,P-1)
&        ENDIF
350    CONTINUE

DO 360, T = 1, COUNT (KK,IY)
    P = SEQ (KK,IY,T)
    PART = (-PHI(LL,P)/SIGMA(NN,P))
    PART2 = (-PHI(LLL,P)/SIGMA(NN,P))
    DER = DER+PART*PART2
    IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&        THEN
&        MIDDLE = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(KK,K)
&                    *((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(N,P-1)))
&        PART1 = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P) ** 2)

```

```

        PART3 = PHI(LLL,P)
        PART4 = PHI(LLL,P-1)
        PART5 = PHI(LL,P)
        PART6 = PHI(LL,P-1)
        DER3 = DER3+MIDDLE*(2*PART1/SIGMA(NN,P))*PART3*PART5
&          +((-PART1*PART3)*(-PART1*PART5)-PART3*PART5/
&          (SIGMA(NN,P)**2)
        DER4 = DER4+MIDDLE*PART3*PART5/(SIGMA(NN,P)**2)+
&          (-PART1*PART3*(-PART5/SIGMA(NN,P)))
        ENDIF
360      CONTINUE
10      CONTINUE

        IF (M .EQ. 1) THEN
            DER2 (LL,LLL) = -DER
            DER2 (LL+6,LLL+6) = -DER3
            DER2 (LL,LLL+6) = -DER4
        ELSEIF (M .EQ. 2) THEN
            DER2 (LL+3,LLL+3) = -DER
            DER2 (LL+9,LLL+9) = -DER3
            DER2 (LL+3,LLL+9) = -DER4
        ENDIF

880      CONTINUE
870      CONTINUE
850      CONTINUE

        RETURN
        END

```

```

C      SUBROUTINE CNTRAL(MAT,M,N,MATOR,M1,N1,DIM1,DIM2)
      -----
      REAL      MAT(M,N), MATOR(M1,N1)
      INTEGER   DIM1, DIM2
      REAL      AVE(25)

      DO 6000 J = 1,DIM2,1
          AVE(J) = 0.0
          DO 6010 I = 1,DIM1,1
              AVE(J) = AVE(J) + MATOR(I,J)
6010      CONTINUE
          AVE(J) = AVE(J) / FLOAT(DIM1)
6000      CONTINUE

      CC      ***AVE(J) NOW CONTAINS THE AVERAGE OF THE ELEMENTS IN EACH
      CC      ***COLUMN OF THE MATRIX                                     ***

      DO 6020 I = 1,DIM1,1
          DO 6030 J = 1,DIM2,1
              MAT(I,J) =MATOR(I,J) - AVE(J)
6030      CONTINUE
6020      CONTINUE
      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE 2ND DERIVATIVES FOR:
C      RAU-ALPHAjD, RAU-PSIjD, RAU-RAU, RAU-ALPHAjW AND
C      RAU-PSIjW FOR MODEL5
C -----

SUBROUTINE M5DER3 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,
C -----
C      &      PHI,DER2,NV,K,NRAU)
C -----

      INTEGER      COUNT (4,NY)
      INTEGER      SEQ (4,NY,NT)
      INTEGER      T,P
      REAL         CLIMA (NY,0:NT)
      REAL         MIDDLE
      REAL         DER2 (NP,NP)
      REAL         MU (2,0:365)
      REAL         SIGMA (2,0:365)
      REAL         PHI (NPARM,0:NT)
      REAL         RAU (NRAU,NV)

      DO 850, LL = 1, NPARM
        DO 870, M = 1, 2
          IF (M .EQ. 1) THEN
            N = 2
            NN = 1
            J = 3
            KK = 4
          ELSEIF (M .EQ. 2) THEN
            N = 1
            NN = 2
            J = 4
            KK = 3
          ENDIF

          ..... THE VARIABLE DER1 COMPUTES THE 2ND DERIVATIVES FOR RAU-
          ALPHA, WHILE DER3 COMPUTES THE 2ND DERIVATIVES FOR RAU-
          PSI

          DER1 = 0
          DER3 = 0
          DO 10, IY = 1, NY
            DO 530, T = 1, COUNT (M,IY)
              P = SEQ (M,IY,T)
              IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
              &      -999)) THEN
                PART1 = (CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)
                PART2 = (CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)
                MIDDLE = PART1-RAU(M,K)*PART2
                DER1 = DER1+MIDDLE*(-PHI(LL,P-1)/SIGMA(M,P-1))+
                &      (-PHI(LL,P)/SIGMA(M,P)+RAU(M,K)*PHI(LL,P-1)
                &      /SIGMA(M,P-1))*PART2

```



```

&                -999)) THEN
                PART3 = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P))
                DER3 = DER3+(-PHI(LL,P)/SIGMA(NN,P))*PART3*PART2
                ENDIF

560                CONTINUE
30                CONTINUE
                IF (M .EQ. 1) THEN
                    DER2(LL,16) = DER1
                    DER2(LL+6,16) = DER3
                ELSEIF (M .EQ. 2) THEN
                    DER2 (LL+3,15) = DER1
                    DER2 (LL+9,15) =DER3
                ENDIF
870                CONTINUE
850                CONTINUE

C                ..... THE 2ND DERIVATIVE RAU-RAU IS COMPUTED

DO 40, IY = 1, NY
    DO 330, T = 1, COUNT (1,IY)
        P = SEQ (1,IY,T)
        IF (CLIMA(IY,P-1).NE.-999) THEN
            PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
            DER2 (13,13) = DER2 (13,13)-(PART2**2)
        ENDIF
330                CONTINUE

        DO 340, T = 1, COUNT (2,IY)
            P = SEQ (2,IY,T)
            IF (CLIMA(IY,P-1).NE.-999) THEN
                PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
                DER2 (14,14) = DER2(14,14)-(PART2**2)
            ENDIF
340                CONTINUE

            DO 350, T = 1, COUNT (3,IY)
                P = SEQ (3,IY,T)
                IF (CLIMA(IY,P-1).NE.-999) THEN
                    PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
                    DER2 (15,15) = DER2 (15,15)-(PART2**2)
                ENDIF
350                CONTINUE

                DO 360, T = 1, COUNT (4,IY)
                    P = SEQ (4,IY,T)
                    IF (CLIMA(IY,P-1).NE.-999) THEN
                        PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
                        DER2 (NP,NP) = DER2 (NP,NP)-(PART2**2)
                    ENDIF
360                CONTINUE
40                CONTINUE

                RETURN
            END

```

```

-----
..... THIS SUBROUTINE COMPUTES THE 2ND DERIVATIVES FOR:
      ALPHAJD-ALPHAiw, ALPHAJD-PSIiw, PSIJD-PSIiw AND
      ALPHAJW-PSIid FOR MODEL 5
-----

```

```

SUBROUTINE M5DER4 (NY,NT,NP,NPARM,COUNT,SEQ,CLIMA,MU,SIGMA,RAU,
-----
&      PHI,DER2,NV,K,NRAU)
-----

```

```

INTEGER      COUNT (4,NY)
INTEGER      SEQ (4,NY,NT)
INTEGER      T,P
REAL         CLIMA (NY,0:NT)
REAL         DER2 (NP,NP)
REAL         MU (2,0:365)
REAL         SIGMA (2,0:365)
REAL         PHI (NPARM,0:NT)
REAL         RAU (NRAU,NV)

```

```

DO 850, LL = 1, NPARM
  DO 870, LLL = 1, NPARM

```

```

..... THE VARIABLE DER COMPUTES THE 2ND DERIVATIVE ALPHAD-
      ALPHAW, DER3 THE DERIVATIVE ALPHAD-PSIW, DER4 THE
      DERIVATIVE PSID-PSIW AND DER5 THE DERIVATIVE ALPHAW-
      PSID

```

```

      DER = 0
      DER3 = 0
      DER4 = 0
      DER5 = 0
      DO 10, IY = 1, NY
        DO 350, T = 1, COUNT (3,IY)
          P = SEQ (3,IY,T)
          PART = (RAU(3,K)*PHI(LL,P-1)/SIGMA(1,P-1))
          DER = DER+PART*(-PHI(LL,P)/SIGMA(2,P))
          IF (CLIMA(IY,P).NE.-999) THEN
            DER3 = DER3+PART*(-PHI(LL,P)/SIGMA(2,P))*((CLIMA
&              (IY,P)-MU(2,P))/SIGMA(2,P))
            ENDIF
            IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&              -999)) THEN
              PART2 = ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)**2)
&              *PHI(LL,P-1)
              PART1 = (-((CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)**2)*
&              PHI(LL,P)
              DER4 = DER4+(PART1*RAU(3,K)*PART2)
              ENDIF
              IF (CLIMA(IY,P-1).NE.-999) THEN
                PART = (-PHI(LL,P)/SIGMA(2,P))
                DER5 = DER5+PART*RAU(3,K)*PHI(LL,P-1)/SIGMA(1,P-1)*

```

```

&          (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
&          ENDIF
350      CONTINUE

      DO 360, T = 1, COUNT (4,IY)
      P = SEQ (4,IY,T)
      PART = (-PHI(LL,P)/SIGMA(1,P))
      DER = DER+PART*(RAU(4,K)*PHI(LLL,P-1)/SIGMA(2,P-1))
      IF (CLIMA(IY,P-1).NE.-999) THEN
&          DER3 = DER3+PART*(RAU(4,K)*PHI(LLL,P-1)/SIGMA(2,P-1)
&              )*(CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
&          ENDIF
&          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&              -999)) THEN
&              PART1 = (-((CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)**2)*
&                  PHI(LL,P))
&              PART2 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)**2)*
&                  PHI(LLL,P-1)
&              DER4 = DER4+RAU(4,K)*PART2*PART1
&              ENDIF
&              IF (CLIMA(IY,P).NE.-999) THEN
&                  PART = (RAU(4,K)*PHI(LL,P-1)/SIGMA(2,P-1))
&                  DER5 = DER5+PART*(-PHI(LLL,P)/SIGMA(1,P))*(CLIMA
&                      (IY,P)-MU(1,P))/SIGMA(1,P)
&                  ENDIF
360      CONTINUE
10      CONTINUE

      DER2 (LL,LLL+3) = -DER
      DER2 (LL,LLL+9) = -DER3
      DER2 (LL+6,LLL+9) = -DER4
      DER2 (LL+3,LLL+6) = -DER5

870      CONTINUE
850      CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE GENERATES RAINFALL DEPTH ON DAYS
C WHEN RAIN OCCURS
C -----

```

```

C SUBROUTINE DEPTH3 (IDUM9,NP,RAIN,K,AMP,PHASE,GAMM,BI,W)
C -----

```

```

      REAL          AMP (0:NP)
      REAL          PHASE (NP)
      REAL          RAIN

      AM = (AMP(0)+AMP(1)*COS(W*((K-1)-PHASE(1)))+AMP(2)*COS(2*W*
&          ((K-1)-PHASE(2)))) / GAMM
      UNIFOR = URAN9 (IDUM9)
      RAIN = AM * (-1 * LOG(UNIFOR)) ** BI

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE READS IN THE PARAMETER ESTIMATES OF THE
C RAINFALL MODEL AND OF THE CLIMATE MODEL.
C -----

SUBROUTINE DATA (GAM,RAU3,RAU4,RAU5,MU,SIGMA3,SIGMA4,SIGMA5,NP,
C -----
& NV,AMP,PHASE,CV,PHI,CORR,NT,NRAU,NV3,NV4,NV5)
C -----

REAL          GAM (2,NP)
REAL          PSI4 (2,1,3)
REAL          PSI5 (2,4,3)
REAL          ALPHA (2,7,3)
REAL          SIGMA3 (NRAU,NV3)
REAL          SIGMA4 (2,NV4,0:NT)
REAL          SIGMA5 (2,NV5,0:NT)
REAL          MU (2,NV,0:NT)
REAL          PHI (NP,0:NT)
REAL          RAU3 (NRAU,NV3)
REAL          RAU4 (NV4)
REAL          RAU5 (NRAU,NV5)
REAL          AMP (0:NP)
REAL          PHASE (NP)
REAL          CORR (NV,NV)

5  FORMAT (7F10.3)

OPEN (UNIT=12,FILE=' \WATER\DATA\EST.DAT',STATUS='OLD')

READ (12, *) (GAM (1,J), J = 1, NP)
READ (12, *) (GAM (2,J), J = 1, NP)
DO 30, K = 1, NV
  DO 10, M = 1, 2
    READ (12, *) (ALPHA (M,K,J), J = 1, NP)
10  CONTINUE
    IF ((K.EQ.4).OR.(K.EQ.6)) THEN
      IF (K.EQ.4) THEN
        KK = 1
      ELSEIF (K.EQ.6) THEN
        KK = 2
      ENDIF
      READ (12,*) (SIGMA3 (L,KK), L = 1, NRAU)
      READ (12,*) (RAU3 (L,KK), L = 1, NRAU)
    ELSEIF (K .EQ. 5) THEN
      KK = 1
      DO 20, M = 1, 2
        READ (12, *) (PSI4 (M,KK,J), J = 1, NP)
20  CONTINUE
        READ (12,*) RAU4 (KK)
      ELSEIF ((K.EQ.1).OR.(K.EQ.2).OR.(K.EQ.3).OR.(K.EQ.7)) THEN
        IF (K.EQ.1) THEN

```

```

        KK = 1
    ELSEIF (K.EQ.2) THEN
        KK = 2
    ELSEIF (K.EQ.3) THEN
        KK = 3
    ELSEIF (K.EQ.7) THEN
        KK = 4
    ENDIF
    DO 120, M = 1, 2
        READ (12, *) (PSI5 (M, KK, J), J = 1, NP)
120    CONTINUE
        READ (12, *) (RAU5 (I, KK), I = 1, NRAU)
    ENDIF
30    CONTINUE
    READ (12, *) (AMP (I), I = 0, 1)
    READ (12, *) (PHASE (I), I = 1, 1)
    READ (12, *) CV

C      ..... INPUT CORRELATION MATRIX

    DO 80, I = 1, NV
        READ (12, *) (CORR (I, J), J = I, NV)
80    CONTINUE

C      ..... COMPUTE THE MEAN AND STD.DEV. FUNCTION

    DO 40, M = 1, 2
        DO 50, I = 0, NT
            SIGMA4 (M, 1, I) = 0.0
            SIGMA5 (M, 1, I) = 0.0
            SIGMA5 (M, 2, I) = 0.0
            SIGMA5 (M, 3, I) = 0.0
            SIGMA5 (M, 4, I) = 0.0
            DO 70, K = 1, NV
                MU (M, K, I) = 0.0
                DO 60, L = 1, NP
                    MU(M, K, I) = MU(M, K, I) + ALPHA(M, K, L) * PHI(L, I)
                    IF ((K.EQ.1).OR.(K.EQ.2).OR.(K.EQ.3).OR.(K.EQ.7))
&                        THEN
                            IF (K.EQ.1) THEN
                                KK = 1
                            ELSEIF (K.EQ.2) THEN
                                KK = 2
                            ELSEIF (K.EQ.3) THEN
                                KK = 3
                            ELSEIF (K.EQ.7) THEN
                                KK = 4
                            ENDIF
                            SIGMA5(M, KK, I) = SIGMA5(M, KK, I) + PSI5(M, KK, L) *
&                                PHI(L, I)
                        ELSEIF (K.EQ.5) THEN
                            KK = 1
                            SIGMA4(M, KK, I) = SIGMA4(M, KK, I) + PSI4(M, KK, L) *
&                                PHI(L, I)

```

```

        ENDIF
60      CONTINUE
70      CONTINUE
50      CONTINUE
40      CONTINUE

```

```

RETURN
END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE CHOLESKY DECOMPOSITION
C OF A MATRIX
C -----

```

```

C SUBROUTINE CHOLKY (DECOMP,CORR,NV)
C -----

```

```

REAL          CORR (NV,NV)
REAL          DECOMP (NV,NV)

```

```

C ..... COMPUTE CHOLESKY DECOMPOSITION

```

```

      DO 40, I = 1, NV
        DO 50, J = 1, NV
          DECOMP (I,J) = 0
50      CONTINUE
40      CONTINUE
      DECOMP (1,1) = SQRT (CORR (1,1))
      DO 60, J = 2, NV
        DECOMP (J,1) = CORR (1,J) / DECOMP (1,1)
60      CONTINUE
      DO 70, I = 2, NV-1
        TERM = 0
        DO 80, J = 1, I-1
          TERM = TERM + DECOMP (I,J) ** 2
80      CONTINUE
        DECOMP (I,I) = SQRT (CORR (I,I) - TERM)
        DO 90, J = I+1, NV
          TERM = 0
          DO 100, K = 1, I-1
            TERM = TERM + DECOMP (I,K) * DECOMP (J,K)
100     CONTINUE
          DECOMP (J,I) = (CORR (I,J) - TERM) / DECOMP (I,I)
90      CONTINUE
70      CONTINUE
      TERM = 0
      DO 110, J = 1, NV-1
        TERM = TERM + DECOMP (NV,J) ** 2
110     CONTINUE
      DECOMP (NV,NV) = SQRT (CORR (NV,NV) - TERM)

RETURN
END

```

```

C -----
C ..... THIS SUBROUTINE READS IN THE PARAMETER ESTIMATES OF THE
C RAINFALL MODEL AND OF THE CLIMATE MODEL 1.
C -----

SUBROUTINE DATA1 (GAM,MU,SIGMA,NP,NV,AMP,PHASE,CV,PHI,A,B,NT)
C -----

REAL          GAM (2,NP)
REAL          PSI (2,7,3)
REAL          ALPHA (2,7,3)
REAL          SIGMA (2,NV,0:NT)
REAL          MU (2,NV,0:NT)
REAL          PHI (NP,0:NT)
REAL          AMP (0:NP)
REAL          PHASE (NP)
REAL          A (NV,NV)
REAL          B (NV,NV)

5  FORMAT (7F10.3)

OPEN (UNIT=12,FILE=' \WATER\DATA\EST1.DAT',STATUS='OLD')

DO 10, M = 1, 2
  READ (12, *) (GAM (M,J), J = 1, NP)
10 CONTINUE
DO 20, K = 1, NV
  DO 30, M = 1, 2
    READ (12, *) (ALPHA (M,K,J), J = 1, NP)
    READ (12, *) (PSI (M,K,J), J = 1, NP)
30 CONTINUE
20 CONTINUE
READ (12, *) (AMP (I), I = 0, 1)
READ (12, *) (PHASE (I), I = 1, 1)
READ (12, *) CV

C ..... INPUT A & B MATRICES

DO 80, I = 1, NV
  READ (12, *) (A (I,J), J = 1, NV)
80 CONTINUE
DO 180, I = 1, NV
  READ (12, *) (B (I,J), J = 1, NV)
180 CONTINUE

C ..... COMPUTE THE MEAN AND STD.DEV. FUNCTION

DO 40, M = 1, 2
  DO 50, I = 0, NT
    DO 70, K = 1, NV
      MU (M,K,I) = 0.0
      SIGMA (M,K,I) = 0.0
      DO 60, L = 1, NP

```

```

        MU (M,K,I) = MU (M,K,I) + ALPHA (M,K,L) * PHI (L,I)
        SIGMA (M,K,I) = SIGMA (M,K,I) + PSI (M,K,L) *
        PHI (L,I)
&
60      CONTINUE
70      CONTINUE
50      CONTINUE
40      CONTINUE

```

```

RETURN
END

```

```

C      -----
C      ..... A ROUTINE TO GENERATE PSEUDO RANDOM NUMBERS FROM A
C      GAUSSIAN DISTRIBUTION WITH A MEAN OF ZERO AND A
C      STANDARD DEVIATION OF UNITY AS SPECIFIED BY THE USER.
C      THIS ROUTINE REFERENCES UNIF TO GENERATE THE UNIFORMLY
C      DISTRIBUTED RANDOM NUMBERS.
C      -----

```

```

C      SUBROUTINE GRAND2 (NRAND,NV)
C      -----

```

```

REAL          NRAND (1,NV)

COMMON        IDUM1, IDUM2, IDUM3, IDUM4, IDUM5, IDUM6, IDUM7

R1 = URAN1 (IDUM1)
R2 = URAN1 (IDUM1)
T = SQRT(-2*LOG(R1))
NRAND(1,1) = T * SIN (6.283185*R2)
R1 = URAN2 (IDUM2)
R2 = URAN2 (IDUM2)
T = SQRT(-2*LOG(R1))
NRAND(1,2) = T * SIN (6.283185*R2)
R1 = URAN3 (IDUM3)
R2 = URAN3 (IDUM3)
T = SQRT(-2*LOG(R1))
NRAND(1,3) = T * SIN (6.283185*R2)
R1 = URAN4 (IDUM4)
R2 = URAN4 (IDUM4)
T = SQRT(-2*LOG(R1))
NRAND(1,4) = T * SIN (6.283185*R2)
R1 = URAN5 (IDUM5)
R2 = URAN5 (IDUM5)
T = SQRT(-2*LOG(R1))
NRAND(1,5) = T * SIN (6.283185*R2)
R1 = URAN6 (IDUM6)
R2 = URAN6 (IDUM6)
T = SQRT(-2*LOG(R1))
NRAND(1,6) = T * SIN (6.283185*R2)
R1 = URAN7 (IDUM7)
R2 = URAN7 (IDUM7)
T = SQRT(-2*LOG(R1))
NRAND(1,7) = T * SIN (6.283185*R2)

RETURN
END

```

```

C -----
C ..... THIS SUBROUTINE SOLVES A SYSTEM OF EQUATIONS
C -----

SUBROUTINE LINEAR (NPMAX,NP,DER,DER2,THETA)
C -----

REAL          A (13,0:13)
REAL          DER (NPMAX)
REAL          DER2 (NPMAX,NPMAX)
REAL          THETA (NPMAX)

15  FORMAT (' MATRIX IS SINGULAR')
25  FORMAT (' NEW PARAMETER ESTIMATES: ',F10.4)

OPEN (UNIT=9,FILE='LPT1')

C ..... THIS SETS UP THE A MATRIX WHICH IS USED IN SOLVING
C THE SYSTEM OF EQUATIONS

DO 10, I = 1, NP
  A (I,0) = DER (I)
  DO 20, J = 1, NP
    A (I,J) = DER2 (I,J)
20  CONTINUE
10  CONTINUE

C ..... THIS SOLVES THE SYSTEM OF EQUATIONS
C THE DIFFERENCE BETWEEN THE VALUE OF THETA(Q) IN THIS
C ITERATION AND THE PREVIOUS ITERATION IS STORED IN A(Q,0)

DO 30, I1 = 1, NP
  I2 = I1
  T1 = 0
  DO 40, I3 = I1, NP
    IF (ABS (A (I3,I1)) .GT. (ABS (T1))) THEN
      I2 = I3
      T1 = A (I3,I1)
    ENDIF
40  CONTINUE
    IF (T1 .EQ. 0) THEN
      PRINT 15
      STOP
    ENDIF
    IF (I2 .NE. I1) THEN
      DO 50, I0 = 0, NP
        TEMP = A (I1,I0)
        A (I1,I0) = A (I2,I0)
        A (I2,I0) = TEMP
50  CONTINUE
      ENDIF
      T2 = 1 / (A (I1,I1))
      NQ = NP

```

```

        DO 60, I4 = 0, NQ
          A (I1,I4) = A (I1,I4) * T2
60      CONTINUE
        DO 70, I3 = 1, NP
          IF (I1 .NE. I3) THEN
            T2 = A (I3,I1)
            A (I3,0) = A (I3,0)-A(I1,0)* T2
            DO 80, I0 = I1, NP
              A(I3,I0) = A(I3,I0) - A(I1,I0) * T2
80          CONTINUE
            ENDIF
          CONTINUE
70      CONTINUE
30      CONTINUE

C      ..... THIS EXTRACTS THE NEW PARAMETER VALUES

        DO 90, I = 1, NP
          THETA(I) = THETA(I) - A(I,0)
          WRITE (9,25) THETA(I)
90      CONTINUE

        RETURN
        END

C      -----
C      ..... THIS SUBROUTINE COMPUTES TOTAL MEANS AND STD DEVS
C      TO BE USED IN THE COMPUTATION OF CROSS-CORRELATIONS
C      -----

C      SUBROUTINE AVSTD3 (CLIMA,AVEG,DEV,NY,NT,NV)
C      -----

        INTEGER          DENOM (7)
        REAL             CLIMA (NV,NY*NT)
        REAL             AVEG (NV)
        REAL             DEV (NV)

        DO 10, K = 1, NV
          AVEG (K) = 0.0
          DEV (K) = 0.0
          DENOM (K) = 0
10      CONTINUE

        DO 30, K = 1, NV
          DO 20, I = 1, NY*NT
            IF (CLIMA (K,I) .GT. -900) THEN
              AVEG (K) = AVEG (K) + CLIMA (K,I)
              DEV (K) = DEV (K) + (CLIMA (K,I) ** 2)
              DENOM (K) = DENOM (K) + 1
            ENDIF
          CONTINUE
20      CONTINUE
          DEV (K) = SQRT( (DEV(K)-((AVEG(K)**2)/DENOM(K)))/DENOM(K))
          AVEG (K) = AVEG(K)/DENOM(K)
          PRINT *, AVEG (K)
          PRINT *, DEV (K)
30      CONTINUE

        RETURN
        END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE CHOLESKY DECOMPOSITION
C           OF A MATRIX
C -----

C SUBROUTINE CHOLESKY (DECOMP,CORR,NV)
C -----

      REAL          CORR (NV,NV)
      REAL          DECOMP (NV,NV)

5     FORMAT (7F10.3)

C ..... FILL IN SYMMETRICAL PART OF MATRIX

      DO 20, I = 1, NV-1
        DO 30, J = I+1, NV
          CORR (J,I) = CORR (I,J)
30     CONTINUE
20     CONTINUE

C ..... COMPUTE CHOLESKY DECOMPOSITION

      DO 40, I = 1, NV
        DO 50, J = 1, NV
          DECOMP (I,J) = 0
50     CONTINUE
40     CONTINUE
      DECOMP (1,1) = SQRT (CORR (1,1))
      DO 60, J = 2, NV
        DECOMP (J,1) = CORR (1,J) / DECOMP (1,1)
60     CONTINUE
      DO 70, I = 2, NV-1
        TERM = 0
        DO 80, J = 1, I-1
          TERM = TERM + DECOMP (I,J) ** 2
80     CONTINUE
        DECOMP (I,I) = SQRT (CORR (I,I) - TERM)
        DO 90, J = I+1, NV
          TERM = 0
          DO 100, K = 1, I-1
            TERM = TERM + DECOMP (I,K) * DECOMP (J,K)
100    CONTINUE
          DECOMP (J,I) = (CORR (I,J) - TERM) / DECOMP (I,I)
90     CONTINUE
70     CONTINUE
      TERM = 0
      DO 110, J = 1, NV-1
        TERM = TERM + DECOMP (NV,J) ** 2
110    CONTINUE
      DECOMP (NV,NV) = SQRT (CORR (NV,NV) - TERM)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE GENERATES CLIMATE SEQUENCES ACCORDING
C TO THE SPECIFICATIONS OF MODEL5
C -----

C SUBROUTINE MOD5 (RAND,STATE,NV5,NV,SIGMA5,MU,RAUS,K,J,OBSN,
& PSTATE,NT,NRAU)
C -----

      INTEGER          PSTATE, STATE
      REAL             RAUS (NRAU,NV5)
      REAL             RAND (1,NV)
      REAL             OBSN (NV)
      REAL             SIGMA5 (2,NV5,0:NT)
      REAL             MU (2,NV,0:NT)

5  FORMAT (7F10.3)

      IF ((STATE .EQ. 1) .AND. (PSTATE .EQ. 1)) THEN
        JJ = 1
      ELSEIF ((STATE .EQ. 2) .AND. (PSTATE .EQ. 2)) THEN
        JJ = 2
      ELSEIF ((STATE .EQ. 1) .AND. (PSTATE .EQ. 2)) THEN
        JJ = 3
      ELSEIF ((STATE .EQ. 2) .AND. (PSTATE .EQ. 1)) THEN
        JJ = 4
      ENDIF

      IF (J-1 .EQ. 0) THEN
        L = NT
      ELSE
        L = J-1
      ENDIF

      IF (K.EQ.2) THEN
        KK = 1
      ENDIF

      OBSN (K) = SIGMA5(PSTATE,KK,J)*(RAND(1,K)+RAUS(JJ,KK)*(OBSN(K)-
& MU(STATE,K,L))/SIGMA5(STATE,KK,L))+MU(PSTATE,K,J)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE GENERATES CLIMATE SEQUENCES ACCORDING
C TO THE SPECIFICATIONS OF MODEL3
C -----

SUBROUTINE MOD3 (RAND,STATE,NV3,NV,SIGMA3,MU,RAU3,K,J,OBSN,
C -----
C & PSTATE,NT,NRAU)
C -----

      INTEGER          PSTATE, STATE
      REAL             RAU3 (NRAU,NV3)
      REAL             RAND (1,NV)
      REAL             OBSN (NV)
      REAL             SIGMA3 (NRAU,NV3)
      REAL             MU (2,NV,0:NT)

5  FORMAT (7F10.3)

      IF ((STATE .EQ. 1) .AND. (PSTATE .EQ. 1)) THEN
        JJ = 1
      ELSEIF ((STATE .EQ. 2) .AND. (PSTATE .EQ. 2)) THEN
        JJ = 2
      ELSEIF ((STATE .EQ. 1) .AND. (PSTATE .EQ. 2)) THEN
        JJ = 3
      ELSEIF ((STATE .EQ. 2) .AND. (PSTATE .EQ. 1)) THEN
        JJ = 4
      ENDIF

      IF (J-1 .EQ. 0) THEN
        L = NT
      ELSE
        L = J-1
      ENDIF

      IF (K.EQ.1) THEN
        KK = 1
      ELSEIF (K.EQ.3) THEN
        KK = 2
      ELSEIF (K.EQ.4) THEN
        KK = 3
      ELSEIF (K.EQ.7) THEN
        KK = 4
      ENDIF

      OBSN (K) = SIGMA3(JJ,KK)*(RAND(1,K)+RAU3(JJ,KK)*(OBSN(K)-
& MU(STATE,K,L)) / SIGMA3(JJ,KK))+MU(PSTATE,K,J)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE AMPLITUDE & PHASE
C REPRESENTATION
C -----

SUBROUTINE AMPHA (AM,PH,THETA,NPMAX,KMAX,K,PI,NT)
C -----

      REAL          AM(0:KMAX)
      REAL          PH(KMAX)
      REAL          THETA(NPMAX)

45  FORMAT (/, ' AMPLITUDE: ')
55  FORMAT (/, ' PHASE: ')
65  FORMAT (9F8.3)

      OPEN (UNIT=9,FILE='LPT1')

      AM(0) = THETA(1)
      DO 140, I = 1, K
        TA = THETA (2*I)
        TB = THETA(2*I+1)
        AM(I) = SQRT(TA**2 + TB**2)
        IF (TA .LT. 0) THEN
          PH(I) = ATAN(TB / TA) + PI
        ELSEIF (TA .EQ. 0) THEN
          IF (TB .GE. 0) THEN
            PH(I) = 0.5 * PI
          ELSE
            PH(I) = 1.5 * PI
          ENDIF
        ELSEIF (TA .GT. 0) THEN
          IF (TB .GE. 0) THEN
            PH(I) = ATAN (TB/TA)
          ELSE
            PH(I) = ATAN(TB/TA) + 2 * PI
          ENDIF
        ENDIF
        PH(I) = PH(I) * NT / (2*PI*I)
140  CONTINUE

      WRITE (9,45)
      WRITE (9,65) (AM(I),I=0,K)
      WRITE (9,55)
      WRITE (9,65) (PH(I),I=1,K)

      RETURN
      END

```

```

C -----
C THIS SUBROUTINE COMPUTES THE MATRIX OF SIN AND COS TERMS FOR
C THE FOURIER TRANSFORMATION. THIS DIFFERS FROM THE SUBROUTINE
C COSSIN IN THAT HERE THE MATRIX PHI HAS DIMENSION GIVEN BY:
C ---> PHI (NT,NPARM)
C -----

```

```

C SUBROUTINE TRIG (PHI,NPARM,NT)
C -----

```

```

      REAL          PI
      PARAMETER      (PI = 3.14159265)
      REAL          PHI (NT,NPARM)
      REAL          THETA
      REAL          OMEGA
      INTEGER        T

      OMEGA = 2 * PI / NT
      K = (NPARM - 1) / 2
      DO 10, T = 1, NT
        PHI (T,1) = 1
10     CONTINUE
      DO 20, J = 1, K
        J1 = 2 * J
        J2 = J1 + 1
        THETA = OMEGA * J
        A = 2 * COS (THETA)
        PHI (1,J1) = 1
        PHI (2,J1) = A / 2
        PHI (1,J2) = 0
        PHI (2,J2) = SIN (THETA)
        DO 30, T = 3, NT
          PHI (T,J1) = A * PHI (T-1,J1) - PHI (T-2,J1)
          PHI (T,J2) = A * PHI (T-1,J2) - PHI (T-2,J2)
30     CONTINUE
20     CONTINUE
      RETURN
      END

```

```

C -----
C ..... SUBROUTINE TO COMPUTE THE TRANSPOSE OF A MATRIX. THE
C RESULT IS WRITTEN INTO THE SAME MATRIX.
C -----

```

```

C SUBROUTINE GTRANP (DECOMP,NV)
C -----

```

```

      REAL          DECOMP (NV,NV)

      DO 10, I = 1, NV
        DO 20, J = I+1, NV
          TEMP1 = DECOMP (I,J)
          TEMP2 = DECOMP (J,I)
          DECOMP (I,J) = TEMP2
          DECOMP (J,I) = TEMP1
20     CONTINUE
10     CONTINUE

      RETURN
      END

```

```

C -----
C ..... SUBROUTINE TO BRACKET THE MINIMUM
C -----

SUBROUTINE BRACK (A,B,C,FA,FB,FC,DIM1)
-----

PARAMETER      (GLD=1.618034)
PARAMETER      (GLIM=100.)
PARAMETER      (T=1.E-20)

OPEN (UNIT=9,FILE='LPT1')

FA=DIM1(A)
FB=DIM1(B)
WRITE (9,*) ' FA FB', FA, FB
IF (FB.GT.FA) THEN
    DUM=A
    A=B
    B=DUM
    DUM=FB
    FB=FA
    FA=DUM
ENDIF
C=B+GLD*(B-A)
FC=DIM1(C)
WRITE (9,*) ' FC', FC
1 IF (FB.GE.FC) THEN
    R=(B-A)*(FB-FC)
    Q=(B-C)*(FB-FA)
    U=B-((B-C)*Q-(B-A)*R)/(2.*SIGN(MAX(ABS(Q-R),T),Q-R))
    ULIM=B+GLIM*(C-B)
    IF ((B-U)*(U-C).GT.0.) THEN
        FU=DIM1(U)
        IF (FU.LT.FC) THEN
            A=B
            FA=FB
            B=U
            FB=FU
            GOTO 1
        ELSEIF (FU.GT.FB) THEN
            C=U
            FC=FU
            GOTO 1
        ENDIF
        U=C+GLD*(C-B)
        FU=DIM1(U)
    ELSEIF ((C-U)*(U-ULIM).GT.0.) THEN
        FU=DIM1(U)
        IF (FU.LT.FC) THEN
            B=C
            C=U
            U=C+GLD*(C-B)

```

```

        FB=FC
        FC=FU
        FU=DIM1(U)
    ENDIF
ELSEIF ((U-ULIM)*(ULIM-C).GE.0.) THEN
    U=ULIM
    FU=DIM1(U)
ELSE
    U=C+GLD*(C-B)
    FU=DIM1(U)
ENDIF
A=B
B=C
C=U
FA=FB
FB=FC
FC=FU
GOTO 1
ENDIF

RETURN
END

```

```

C -----
C ..... SUBROUTINE TO MULTIPLY TWO MATRICES
C -----

```

```

C SUBROUTINE MULT (FIRST,SECOND,THIRD,ROWX,COLX,ROWA,COLA)
C -----

```

```

        INTEGER      ROWX,COLX,ROWA,COLA,TEST
        REAL          FIRST (ROWX,COLX)
        REAL          SECOND (ROWA,COLA)
        REAL          THIRD (ROWX,COLA)

        TEST=1
        DO 40, KK = 1, ROWA
            DO 50, JJ = 1, COLA
                IF (SECOND (KK,JJ) .EQ. -999.0) THEN
                    TEST=0
                ENDIF
            50 CONTINUE
        40 CONTINUE

        IF (TEST.EQ.1) THEN
            IF (COLX .NE. ROWA) THEN
                PRINT *, 'MATRICES ARE NOT COMPATIBLE'
            ELSE
                DO 10, I = 1, ROWX
                    DO 20, J = 1, COLA
                        THIRD (I,J) = 0
                        DO 30, K = 1, COLX
                            THIRD(I,J) = THIRD(I,J)+FIRST(I,K)*SECOND(K,J)
                        30 CONTINUE
                    20 CONTINUE
                10 CONTINUE
            ENDIF
        ENDIF

        RETURN
        END

```

```

C -----
C ..... FUNCTION TO FIND LOCAL MINIMUM
C -----

C FUNCTION BMIN (AA,BB,C,DIM1,EPS,XMIN)
C -----

      INTEGER          MAXITER
      PARAMETER        (MAXITER=100)
      PARAMETER        (CG=.3819660)
      PARAMETER        (T=1.0E-10)
      REAL              HALF

      A=MIN(AA,C)
      B=MAX(AA,C)
      V=BB
      W=V
      X=V
      E=0.
      FX=DIM1(X)
      FV=FX
      FW=FX

      DO 10, I=1,MAXITER
        HALF=0.5*(A+B)
        TOL=EPS*ABS(X)+T
        T2=2.*TOL
        IF (ABS(X-HALF).LE.(T2-.5*(B-A))) THEN
          GOTO 3
        ENDIF
        IF (ABS(E).GT.TOL) THEN
          R=(X-W)*(FX-FV)
          Q=(X-V)*(FX-FW)
          P=(X-V)*Q-(X-W)*R
          Q=2.*(Q-R)
          IF (Q.GT.0.) THEN
            P=-P
          ELSE
            Q=-Q
          ENDIF
          TEMP=E
          E=D
          IF ((ABS(P).GE.ABS(.5*Q*TEMP)).OR.(P.LE.Q*(A-X)).OR.
            & (P.GE.Q*(B-X))) THEN
            GOTO 1
          ENDIF
          D=P/Q
          U=X+D
          IF ((U-A.LT.T2).OR.(B-U.LT.T2)) THEN
            D=SIGN(TOL,HALF-X)
          ENDIF
          GOTO 2
        ENDIF
      ENDIF

```

```

1      IF (X.GE.HALF) THEN
          E=A-X
        ELSE
          E=B-X
        ENDIF
        D=CG*E
2      IF (ABS(D).GE.TOL) THEN
          U=X+D
        ELSE
          U=X+SIGN(TOL,D)
        ENDIF
        FU=DIM1(U)
        IF (FU.LE.FX) THEN
            IF (U.LT.X) THEN
                B=X
            ELSE
                A=X
            ENDIF
            V=W
            FV=FW
            W=X
            FW=FX
            X=U
            FX=FU
        ELSE
            IF (U.LT.X) THEN
                A=U
            ELSE
                B=U
            ENDIF
            IF ((FU.LE.FW).OR.(W.EQ.X)) THEN
                V=W
                FV=FW
                W=U
                FW=FU
            ELSEIF ((FU.LE.FV).OR.(V.EQ.X).OR.(V.EQ.W)) THEN
                V=U
                FV=FU
            ENDIF
        ENDIF
10     CONTINUE

        PRINT *, 'DID NOT CONVERGE '
3      XMIN=X
        BMIN=FX

        RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C FOR MODEL 3 FOR USAGE IN NUMERICAL RECIPES.
C -----

SUBROUTINE DFUNC (THETA,DER)
C -----
  INTEGER      T,P,NV,NY,NT,NP,NPARM,NRAU
  PARAMETER    (NV=6,NY=7,NT=365,NP=14,NPARM=3,NRAU=4)

  COMMON       K,ICOUNT(NRAU,NY),ISEQ(NRAU,NY,NT),CLIMA(NY,0:NT),
&              ALPHA(2,NV,NPARM),SIGMA(NRAU,NV),PHI(NPARM,0:NT),
&              RAU(NRAU,NV),ISCALE(3,NV)

  DIMENSION    THETA (NP), DER (NP)

  REAL         MU (2,0:NT)
  REAL         MIDDLE

C ..... UPDATE PARAMETER ESTIMATES

  DO 160 J = 1, NPARM
    ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
    ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
160  CONTINUE

  DO 170, J = 1, NRAU
    SIGMA (J,K) = THETA (J+6)/ISCALE(2,K)
    RAU (J,K) = THETA (J+10)/ISCALE(1,K)
170  CONTINUE

  DO 10, M = 1, 2
    DO 20, I = 0, NT
      MU (M,I) = 0.0
      DO 30, L = 1, NPARM
        MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
30    CONTINUE
20    CONTINUE
10    CONTINUE

  DO 40, I = 1, NP
    DER (I) = 0.0
40  CONTINUE

  DO 50, LL = 1, NPARM
    DO 60, M = 1, 2
      IF (M .EQ. 1) THEN
        N = 2
        NN = 1
        J = 3
        KK = 4
      ELSEIF (M .EQ. 2) THEN
        N = 1

```

```

      NN = 2
      J = 4
      KK = 3
    ENDIF

```

```

C      ..... THE VARIABLE DER1 COMPUTES THE DERIVATIVE FOR THE MEAN
C      FUNCTION

      DER1 = 0
      DO 70, IY = 1, NY
        DO 80, T = 1, ICOUNT (M,IY)
          P = ISEQ (M,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999
&          )) THEN
&            MIDDLE = ((CLIMA(IY,P)-MU(M,P))/SIGMA(M,K)-RAU(M,K)*
&            ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,K)))
&            DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(M,K)+RAU(M,K)*
&            PHI(LL,P-1)/SIGMA(M,K))
          ENDIF
80        CONTINUE

        DO 90, T = 1, ICOUNT (J,IY)
          P = ISEQ (J,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999
&          )) THEN
&            MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(J,K)-RAU(J,K)*
&            ((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(J,K)))
&            DER1 = DER1+MIDDLE*(RAU(J,K)*PHI(LL,P-1)/SIGMA
&            (J,K))
          ENDIF
90        CONTINUE

        DO 100, T = 1, ICOUNT (KK,IY)
          P = ISEQ (KK,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999
&          )) THEN
&            MIDDLE=((CLIMA(IY,P)-MU(NN,P))/SIGMA(KK,K)-RAU(KK,K)
&            *((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(KK,K)))
&            DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(KK,K))
          ENDIF
100        CONTINUE
70      CONTINUE

      IF (M .EQ. 1) THEN
        DER (LL) = -DER1
      ELSEIF (M .EQ. 2) THEN
        DER(LL+3) = -DER1
      ENDIF

60    CONTINUE
50    CONTINUE

C      ..... THE DERIVATIVE WITH RESPECT TO THE AUTOCORRELATION

```

C                    COEFFICIENT IS COMPUTED AS WELL AS THE DERIVATIVE  
C                    W.R.T. THE STANDARD DEVIATIONS

```

DO 110, IY = 1, NY
  DO 120, T = 1, ICOUNT (1,IY)
    P = ISEQ (1,IY,T)
    IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999
    &      )) THEN
      PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,K)
      PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,K)
      MIDDLE = PART1-RAU(1,K)*PART2
      DER(11) = DER(11)+MIDDLE*PART2
      PART1 = -((CLIMA(IY,P)-MU(1,P))/SIGMA(1,K)**2)
      PART2 = ((CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,K)**2)
      DER(7) = DER(7)-MIDDLE*(PART1+RAU(1,K)*PART2)-1/SIGMA(1,K)
      ENDIF
120  CONTINUE

    DO 130, T = 1, ICOUNT (2,IY)
      P = ISEQ (2,IY,T)
      IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999
      &      )) THEN
        PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,K)
        PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,K)
        MIDDLE = PART1-RAU(2,K)*PART2
        DER(12) = DER(12)+MIDDLE*PART2
        PART1 = -((CLIMA(IY,P)-MU(2,P))/SIGMA(2,K)**2)
        PART2 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,K)**2)
        DER(8) = DER(8)-MIDDLE*(PART1+RAU(2,K)*PART2)-1/SIGMA(2,K)
        ENDIF
130  CONTINUE

      DO 140, T = 1, ICOUNT (3,IY)
        P = ISEQ (3,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999
        &      )) THEN
          PART1 = (CLIMA(IY,P)-MU(3,P))/SIGMA(3,K)
          PART2 = (CLIMA(IY,P-1)-MU(3,P-1))/SIGMA(3,K)
          MIDDLE = PART1-RAU(3,K)*PART2
          DER(13) = DER(13)+MIDDLE*PART2
          PART1 = -((CLIMA(IY,P)-MU(3,P))/SIGMA(3,K)**2)
          PART2 = ((CLIMA(IY,P-1)-MU(3,P-1))/SIGMA(3,K)**2)
          DER(9) = DER(9)-MIDDLE*(PART1+RAU(3,K)*PART2)-1/SIGMA(3,K)
          ENDIF
140  CONTINUE

        DO 150, T = 1, ICOUNT (4,IY)
          P = ISEQ (4,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999
          &      )) THEN
            PART1 = (CLIMA(IY,P)-MU(4,P))/SIGMA(4,K)
            PART2 = (CLIMA(IY,P-1)-MU(4,P-1))/SIGMA(4,K)
            MIDDLE = PART1-RAU(4,K)*PART2
            DER(14) = DER(14)+MIDDLE*PART2

```

```

        PART1 = -((CLIMA(IY,P)-MU(1,P))/SIGMA(4,K)**2)
        PART2 = ((CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(4,K)**2)
        DER(10)=DER(10)-MIDDLE*(PART1+RAU(4,K)*PART2)-1/SIGMA(4,K)
        ENDIF
150      CONTINUE
110      CONTINUE

      DO 200, I = 1, NP
        DER(I) = -DER(I)
200      CONTINUE

      RETURN
      END

```

```

C      -----
C      ..... SUBROUTINE TO MINIMIZE ALONG A LINE
C      -----

```

```

C      SUBROUTINE MINL (THETA,DER,NP,FMIN)
C      -----

      INTEGER          NPMAX
      PARAMETER        (NPMAX=20)
      PARAMETER        (EPS=1.E-4)

      EXTERNAL DIM1
      DIMENSION        THETA(NP),DER(NP)
      COMMON /ONE/ NPP,THET(NPMAX),DERI(NPMAX)

      OPEN (UNIT=9,FILE='LPT1')

      NPP=NP
      DO 10, J=1,NP
        THET(J)=THETA(J)
        DERI(J)=DER(J)
10      CONTINUE

      A=0.
      B=1.
      C=2.

      CALL BRACK (A,B,C,FA,FB,FC,DIM1)
      FMIN=BMIN (A,B,C,DIM1,EPS,XMIN)
      DO 20, J=1,NP
        DER(J)=XMIN*DER(J)
        THETA(J)=THETA(J)+DER(J)
20      CONTINUE

      RETURN
      END

```

```

C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C FOR MODEL 4 FOR USAGE IN NUMERICAL RECIPES.
C -----
C
C SUBROUTINE DFUNC (THETA,DER)
C -----
C INTEGER T,P,NV,NY,NT,NP,NPARM,NRAU
C PARAMETER (NV=6,NY=7,NT=365,NP=13,NPARM=3,NRAU=1)
C
C COMMON K,ICOUNT(4,NY),ISEQ(4,NY,NT),CLIMA(NY,0:NT),
& ALPHA(2,NV,NPARM),PSI(2,NV,NPARM),PHI(NPARM,0:NT),
& RAU(NRAU,NV),ISCALE(3,NV)
C
C DIMENSION THETA (NP), DER (NP)
C
C REAL MU (2,0:NT)
C REAL SIGMA (2,0:NT)
C REAL MIDDLE
C
C ..... UPDATE PARAMETER ESTIMATES
C
C DO 160 J = 1, NPARM
C   ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
C   ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
C   PSI (1,K,J) = THETA (J+6)/ISCALE(2,K)
C   PSI (2,K,J) = THETA (J+9)/ISCALE(2,K)
160 CONTINUE
C
C DO 170, J = 1, NRAU
C   RAU (J,K) = THETA (J+12)/ISCALE(1,K)
170 CONTINUE
C
C DO 10, M = 1, 2
C   DO 20, I = 0, NT
C     MU (M,I) = 0.0
C     SIGMA (M,I) = 0.0
C     DO 30, L = 1, NPARM
C       MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
C       SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
30 CONTINUE
20 CONTINUE
10 CONTINUE
C
C DO 40, I = 1, NP
C   DER (I) = 0.0
40 CONTINUE
C
C DO 850, LL = 1, NPARM
C   DO 870, M = 1, 2
C     IF (M.EQ. 1) THEN
C       N = 2
C       NN = 1

```

```

      J = 3
      KK = 4
    ELSEIF (M .EQ. 2) THEN
      N = 1
      NN = 2
      J = 4
      KK = 3
    ENDIF

```

```

C      ..... THE VARIABLE DER1 COMPUTES THE DERIVATIVE FOR THE MEAN
C      FUNCTION, WHILE DER2 AND DER3 COMPUTE THE DERIVATIVE FOR
C      THE VARIANCE FUNCTION

```

```

      DER2 = 0
      DER1 = 0
      DER3 = 0
      DO 310, IY = 1, NY
        DO 330, T = 1, ICOUNT (M,IY)
          P = ISEQ (M,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&              -999)) THEN
&              MIDDLE = ((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)-RAU(1,K)*
&              ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)))
&              DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(M,P)+RAU(1,K)*
&              PHI(LL,P-1)/SIGMA(M,P-1))
&              PART1 = (-((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)**2)*
&              PHI(LL,P))
&              PART2 = ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)**2)
&              *PHI(LL,P-1)
&              DER3 = DER3+MIDDLE*(PART1+RAU(1,K)*PART2)
          ENDIF
          DER2 = DER2+PHI(LL,P)/SIGMA(M,P)
330      CONTINUE

```

```

      DO 350, T = 1, ICOUNT (J,IY)
        P = ISEQ (J,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&              -999)) THEN
&              MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(N,P)-RAU(1,K)*
&              ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
&              DER1 = DER1+MIDDLE*(RAU(1,K)*PHI(LL,P-1)/SIGMA
&              (NN,P-1))
&              PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)
&              **2)*PHI(LL,P-1)
&              DER3 = DER3+MIDDLE*(RAU(1,K)*PART2)
          ENDIF
350      CONTINUE

```

```

      DO 360, T = 1, ICOUNT (KK,IY)
        P = ISEQ (KK,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&              -999)) THEN
&              MIDDLE = ((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(1,K)
&              *((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(N,P-1)))

```

```

      DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(NN,P))
      PART1 = (-((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)**2)*
&          PHI(LL,P))
      DER3 = DER3+MIDDLE*PART1
      ENDIF
      DER2 = DER2+PHI(LL,P)/SIGMA(NN,P)
360      CONTINUE
310      CONTINUE

      IF (M .EQ. 1) THEN
        DER (LL) = -DER1
        DER (LL+6) = (-DER3-DER2)
      ELSEIF (M .EQ. 2) THEN
        DER(LL+3) = -DER1
        DER(LL+9) = (-DER3-DER2)
      ENDIF
870      CONTINUE
850      CONTINUE

C      ..... THE DERIVATIVE WITH RESPECT TO THE AUTOCORRELATION
C      COEFFICIENT IS COMPUTED

      DER (NP) = 0
      DO 420, IY = 1, NY
        DO 700, T = 1, ICOUNT (1,IY)
          P = ISEQ (1,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&          THEN
            PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
            PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
          ENDIF
700      CONTINUE

        DO 701, T = 1, ICOUNT (2,IY)
          P = ISEQ (2,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&          THEN
            PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
            PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
          ENDIF
701      CONTINUE

        DO 702, T = 1, ICOUNT (3,IY)
          P = ISEQ (3,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&          THEN
            PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
            PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
          ENDIF
702      CONTINUE

```

```

      DO 703, T = 1, ICOUNT (4,IY)
        P = ISEQ (4,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
          &      THEN
            PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
            PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(1,K)*PART2)*PART2
          ENDIF
703      CONTINUE
420    CONTINUE

      DO 200, I = 1, NP
        DER(I) = -DER(I)
200    CONTINUE

      RETURN
      END

```

```

C -----
C THIS SUBROUTINE COMPUTES THE MATRIX OF SIN AND COS TERMS FOR
C THE FOURIER TRANSFORMATION
C -----

```

```

C SUBROUTINE COSSIN (PHI,NPARM,NT)
C -----

```

```

      REAL          PI
      PARAMETER     (PI = 3.14159265)
      REAL          PHI (NPARM,0:NT)
      REAL          THETA
      REAL          OMEGA
      INTEGER       T

      OMEGA = 2 * PI / NT
      K = (NPARM - 1) / 2
      DO 10, T = 1, NT
        PHI (1,T) = 1
10     CONTINUE
      DO 20, J = 1, K
        J1 = 2 * J
        J2 = J1 + 1
        THETA = OMEGA * J
        A = 2 * COS (THETA)
        PHI (J1,1) = 1
        PHI (J1,2) = A / 2
        PHI (J2,1) = 0
        PHI (J2,2) = SIN (THETA)
        DO 30, T = 3, NT
          PHI (J1,T) = A * PHI (J1,T-1) - PHI (J1,T-2)
          PHI (J2,T) = A * PHI (J2,T-1) - PHI (J2,T-2)
30     CONTINUE
20     CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE COMPUTES THE VECTOR OF FIRST DERIVATIVES
C FOR MODEL 5 FOR USAGE IN NUMERICAL RECIPES.
C -----

SUBROUTINE DFUNC (THETA,DER)
C -----
C
C      INTEGER      T,P,NV,NY,NT,NP,NPARM,NRAU
C      PARAMETER    (NV=6,NY=7,NT=365,NP=16,NPARM=3,NRAU=4)
C
C      COMMON       K,ICOUNT(4,NY),ISEQ(4,NY,NT),CLIMA(NY,0:NT),
&      ALPHA(2,NV,NPARM),PSI(2,NV,NPARM),PHI(NPARM,0:NT),
&      RAU(NRAU,NV),ISCALE(3,NV)
C
C      DIMENSION    THETA (NP), DER (NP)
C
C      REAL         MU (2,0:NT)
C      REAL         SIGMA (2,0:NT)
C      REAL         MIDDLE
C
C ..... UPDATE PARAMETER ESTIMATES
C
C      DO 160 J = 1, NPARM
C        ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
C        ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
C        PSI (1,K,J) = THETA (J+6)/ISCALE(2,K)
C        PSI (2,K,J) = THETA (J+9)/ISCALE(2,K)
160    CONTINUE
C
C      DO 170, J = 1, NRAU
C        RAU (J,K) = THETA (J+12)/ISCALE(1,K)
170    CONTINUE
C
C      DO 10, M = 1, 2
C        DO 20, I = 0, NT
C          MU (M,I) = 0.0
C          SIGMA (M,I) = 0.0
C          DO 30, L = 1, NPARM
C            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
C            SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
30          CONTINUE
20        CONTINUE
10      CONTINUE
C
C      DO 40, I = 1, NP
C        DER (I) = 0.0
40    CONTINUE
C
C      DO 850, LL = 1, NPARM
C        DO 870, M = 1, 2
C          IF (M .EQ. 1) THEN
C            N = 2
C            NN = 1

```

```

      J = 3
      KK = 4
    ELSEIF (M .EQ. 2) THEN
      N = 1
      NN = 2
      J = 4
      KK = 3
    ENDIF

```

```

C      ..... THE VARIABLE DER1 COMPUTES THE DERIVATIVE FOR THE MEAN
C      FUNCTION, WHILE DER2 AND DER3 COMPUTE THE DERIVATIVE FOR
C      THE VARIANCE FUNCTION

```

```

      DER2 = 0
      DER1 = 0
      DER3 = 0
      DO 50, IY = 1, NY
        DO 330, T = 1, ICOUNT (M,IY)
          P = ISEQ (M,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&            MIDDLE = ((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)-RAU(M,K)*
&            ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)))
&            DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(M,P)+RAU(M,K)*
&            PHI(LL,P-1)/SIGMA(M,P-1))
&            PART1 = (-((CLIMA(IY,P)-MU(M,P))/SIGMA(M,P)**2)*
&            PHI(LL,P))
&            PART2 = ((CLIMA(IY,P-1)-MU(M,P-1))/SIGMA(M,P-1)**2)
&            *PHI(LL,P-1)
&            DER3 = DER3+MIDDLE*(PART1+RAU(M,K)*PART2)
&            ENDIF
&            DER2 = DER2+PHI(LL,P)/SIGMA(M,P)
330          CONTINUE

```

```

        DO 350, T = 1, ICOUNT (J,IY)
          P = ISEQ (J,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&            MIDDLE = ((CLIMA(IY,P)-MU(N,P))/SIGMA(N,P)-RAU(J,K)*
&            ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)))
&            DER1 = DER1+MIDDLE*(RAU(J,K)*PHI(LL,P-1)/SIGMA
&            (NN,P-1))
&            PART2 = ((CLIMA(IY,P-1)-MU(NN,P-1))/SIGMA(NN,P-1)
&            **2)*PHI(LL,P-1)
&            DER3 = DER3+MIDDLE*(RAU(J,K)*PART2)
&            ENDIF
350          CONTINUE

```

```

        DO 360, T = 1, ICOUNT (KK,IY)
          P = ISEQ (KK,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
&            MIDDLE=((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)-RAU(KK,K)
&            *((CLIMA(IY,P-1)-MU(N,P-1))/SIGMA(N,P-1)))

```

```

DER1 = DER1+MIDDLE*(-PHI(LL,P)/SIGMA(NN,P))
PART1 = (-((CLIMA(IY,P)-MU(NN,P))/SIGMA(NN,P)**2)*
&      PHI(LL,P))
DER3 = DER3+MIDDLE*PART1
ENDIF
DER2 = DER2+PHI(LL,P)/SIGMA(NN,P)
360      CONTINUE
50      CONTINUE

IF (M.EQ. 1) THEN
    DER (LL) = -DER1
    DER (LL+6) = (-DER3-DER2)
ELSEIF (M.EQ. 2) THEN
    DER(LL+3) = -DER1
    DER(LL+9) = (-DER3-DER2)
ENDIF

870      CONTINUE
850      CONTINUE

C      ..... THE DERIVATIVE WITH RESPECT TO THE AUTOCORRELATION
C      COEFFICIENT IS COMPUTED

DO 60, IY = 1, NY
    DO 700, T = 1, ICOUNT (1,IY)
        P = ISEQ (1,IY,T)
        IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.-999))
&            THEN
                PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
                PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
                DER(13) = DER(13)+(PART1-RAU(1,K)*PART2)*PART2
            ENDIF
700      CONTINUE

        DO 701, T = 1, ICOUNT (2,IY)
            P = ISEQ (2,IY,T)
            IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&                -999)) THEN
                PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
                PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
                DER(14) = DER(14)+(PART1-RAU(2,K)*PART2)*PART2
            ENDIF
701      CONTINUE

            DO 702, T = 1, ICOUNT (3,IY)
                P = ISEQ (3,IY,T)
                IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&                    -999)) THEN
                    PART1 = (CLIMA(IY,P)-MU(2,P))/SIGMA(2,P)
                    PART2 = (CLIMA(IY,P-1)-MU(1,P-1))/SIGMA(1,P-1)
                    DER(15) = DER(15)+(PART1-RAU(3,K)*PART2)*PART2
                ENDIF
702      CONTINUE

```

```

        DO 703, T = 1, ICOUNT (4,IY)
          P = ISEQ (4,IY,T)
          IF ((CLIMA(IY,P).NE.-999).AND.(CLIMA(IY,P-1).NE.
&          -999)) THEN
            PART1 = (CLIMA(IY,P)-MU(1,P))/SIGMA(1,P)
            PART2 = (CLIMA(IY,P-1)-MU(2,P-1))/SIGMA(2,P-1)
            DER(NP) = DER(NP)+(PART1-RAU(4,K)*PART2)*PART2
          ENDIF
703      CONTINUE
60      CONTINUE

        DO 200, I = 1, NP
          DER(I) = -DER(I)
200      CONTINUE

        RETURN
        END

```

```

C      -----
C      ..... SUBROUTINE TO COMPUTE THE INVERSE OF A MATRIX WHEN
C      NOT PUTTING THE SOLUTION INTO THE OLD MATRIX
C      -----

```

```

C      SUBROUTINE INVT (CLAGO,INV,NV)
C      -----

```

```

      REAL          INV (NV,NV)
      REAL          CLAGO (NV,NV)
      REAL          RESULT (7,7)

      DO 50, K = 1, NV
        DO 60, KK = 1, NV
          INV (K,KK) = CLAGO(K,KK)
60      CONTINUE
50      CONTINUE

      DO 10, I = 1, NV
        DIAG = 1 / INV (I,I)
        INV (I,I) = 1
        DO 20, J = 1, NV
          INV (I,J) = INV (I,J) * DIAG
20      CONTINUE
        DO 30, K = 1, NV
          IF (I .NE. K) THEN
            DIAG = INV (K,I)
            INV (K,I) = 0
            DO 40, J = 1, NV
              INV (K,J) = INV (K,J) - INV (I,J) * DIAG
40      CONTINUE
          ENDIF
30      CONTINUE
10      CONTINUE

      RETURN
      END

```

```

C      -----
C      ..... SUBROUTINE TO COMPUTE LOG LIKELIHOOD FUNCTION FOR
C      MODEL 3.
C      -----

C      FUNCTION FUNC (THETA)
C      -----

      INTEGER          T,NV,NY,NT,NP,NPARM,NRAU
      PARAMETER        (NV=6,NY=7,NT=365,NP=14,NPARM=3,NRAU=4)

      COMMON           K,ICOUNT(NRAU,NY),ISEQ(NRAU,NY,NT),CLIMA(NY,0:NT),
&                     ALPHA(2,NV,NPARM),SIGMA(NRAU,NV),
&                     PHI(NPARM,0:NT),RAU(NRAU,NV),ISCALE(3,NV)

      REAL              MU (2,0:NT)
      REAL              LNLIKE,PI
      PARAMETER        (PI=3.141593)

      DIMENSION        THETA(NP)

C      ..... UPDATE PARAMETER ESTIMATES

      DO 160 J = 1, NPARM
        ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
        ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
160    CONTINUE

      DO 170, J = 1, NRAU
        SIGMA (J,K) = THETA (J+6)/ISCALE(2,K)
        RAU (J,K) = THETA (J+10)/ISCALE(1,K)
170    CONTINUE

      DO 10, M = 1, 2
        DO 20, I = 0, NT
          MU (M,I) = 0.0
          DO 30, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
30          CONTINUE
20        CONTINUE
10      CONTINUE

      LNLIKE = 0
      TERM = 0

      DO 40, J = 1, 4
        IF (J .EQ. 1) THEN
          M = 1
          L = 1
        ELSEIF (J .EQ. 2) THEN
          M = 2
          L = 2
        ELSEIF (J .EQ. 3) THEN

```

```

        M = 2
        L = 1
    ELSEIF (J .EQ. 4) THEN
        M = 1
        L = 2
    ENDIF
    DO 50, I = 1, NY
        DO 60, KK = 1, ICOUNT (J,I)
            T = ISEQ (J,I,KK)
            IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
&                THEN
&                RESID = (CLIMA(I,T)-MU(M,T))/SIGMA(J,K)-RAU(J,K)
&                        *((CLIMA(I,T-1)-MU(L,T-1))/SIGMA(J,K))
                LNLIKE = LNLIKE + (RESID)**2
                TERM = TERM + LOG(SIGMA(J,K))
            ENDIF
60        CONTINUE
50    CONTINUE
40    CONTINUE

FUNC = -(-(NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2

RETURN
END

```

```

C -----
C ..... SUBROUTINE TO COMPUTE LOG LIKELIHOOD FUNCTION FOR
C          MODEL 4.
C -----

C
FUNCTION FUNC (THETA)
-----

      INTEGER      T,NV,NY,NT,NP,NPARM,NRAU
      PARAMETER    (NV=6,NY=7,NT=365,NP=13,NPARM=3,NRAU=1)

      COMMON       K,ICOUNT(4,NY),ISEQ(4,NY,NT),CLIMA(NY,0:NT),
&                ALPHA(2,NV,NPARM),PSI(2,NV,NPARM),
&                PHI(NPARM,0:NT),RAU(NRAU,NV),ISCALE(3,NV)

      REAL         MU (2,0:NT)
      REAL         SIGMA (2,0:NT)
      REAL         LNLIKE,PI
      PARAMETER    (PI=3.141593)

      DIMENSION    THETA(NP)

C ..... UPDATE PARAMETER ESTIMATES

      DO 160 J = 1, NPARM
        ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
        ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
        PSI (1,K,J) = THETA (J+6)/ISCALE(2,K)
        PSI (2,K,J) = THETA (J+9)/ISCALE(2,K)
160    CONTINUE

      DO 170, J = 1, NRAU
        RAU (J,K) = THETA (J+12)/ISCALE(1,K)
170    CONTINUE

      DO 10, M = 1, 2
        DO 20, I = 0, NT
          MU (M,I) = 0.0
          SIGMA (M,I) = 0.0
          DO 30, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
            SIGMA (M,I) = SIGMA (M,I) + PSI(M,K,L) * PHI (L,I)
30          CONTINUE
20        CONTINUE
10      CONTINUE

      LNLIKE = 0
      TERM = 0

      DO 40, J = 1, 4
        IF (J .EQ. 1) THEN
          M = 1
          L = 1

```

```

        ELSEIF (J .EQ. 2) THEN
            M = 2
            L = 2
        ELSEIF (J .EQ. 3) THEN
            M = 2
            L = 1
        ELSEIF (J .EQ. 4) THEN
            M = 1
            L = 2
        ENDIF
        DO 50, I = 1, NY
            DO 60, KK = 1, ICOUNT (J,I)
                T = ISEQ (J,I,KK)
                IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
&                THEN
&                    RESID = (CLIMA(I,T)-MU(M,T))/SIGMA(M,T)-RAU(1,K)
&                        *((CLIMA(I,T-1)-MU(L,T-1))/SIGMA(L,T-1))
                    LNLIKE = LNLIKE + (RESID)**2
                    TERM = TERM + LOG(SIGMA(M,T))
                ENDIF
            CONTINUE
60        CONTINUE
50        CONTINUE
40        CONTINUE

        FUNC = -(-(NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2

        RETURN
        END

```

```

C -----
C ..... THIS SUBROUTINE READS IN INITIAL PARAMETER ESTIMATES
C -----

```

```

C SUBROUTINE INT3 (ALPHA,SIGMA,RAU,NPARM,NV,NRAU,ISCALE)
C -----

```

```

        INTEGER          ISCALE (3,NV)
        REAL              ALPHA (2,NV,NPARM)
        REAL              SIGMA (NRAU,NV)
        REAL              RAU (NRAU,NV)

15    FORMAT (' ALPHA PARAMETERS = ')
25    FORMAT (' SIGMA PARAMETERS = ')
35    FORMAT (' RAU PARAMETER = ')

C    OPEN (UNIT=4,FILE='CON')
    OPEN (UNIT=4,FILE='\\WATER\\DATA\\INIT.DAT',STATUS='OLD')

    DO 20, K = 1, NV
        READ (4,*) ISCALE (1,K),ISCALE(2,K), ISCALE(3,K)
        DO 10, M = 1, 2
            READ (4,*) (ALPHA (M,K,I), I = 1, NPARM)
10        CONTINUE
            READ (4,*) (SIGMA (L,K), L = 1, NRAU)
            READ (4,*) (RAU (L,K), L = 1, NRAU)
20    CONTINUE

        RETURN
        END

```

```

C -----
C ..... SUBROUTINE TO COMPUTE LOG LIKELIHOOD FUNCTION FOR
C MODEL 5.
C -----

C FUNCTION FUNC (THETA)
C -----

      INTEGER      T,NV,NY,NT,NP,NPARM,NRAU
      PARAMETER    (NV=6,NY=7,NT=365,NP=16,NPARM=3,NRAU=4)

      COMMON       K,ICOUNT(4,NY),ISEQ(4,NY,NT),CLIMA(NY,0:NT),
&                ALPHA(2,NV,NPARM),PSI(2,NV,NPARM),
&                PHI(NPARM,0:NT),RAU(NRAU,NV),ISCALE(3,NV)

      REAL         MU (2,0:NT)
      REAL         SIGMA (2,0:NT)
      REAL         LNLIKE,PI
      PARAMETER    (PI=3.141593)

      DIMENSION    THETA(NP)

C ..... UPDATE PARAMETER ESTIMATES

      DO 160 J = 1, NPARM
        ALPHA (1,K,J) = THETA (J)/ISCALE(3,K)
        ALPHA (2,K,J) = THETA (J+3)/ISCALE(3,K)
        PSI (1,K,J) = THETA (J+6)/ISCALE(2,K)
        PSI (2,K,J) = THETA (J+9)/ISCALE(2,K)
160    CONTINUE

      DO 170, J = 1, NRAU
        RAU (J,K) = THETA (J+12)/ISCALE(1,K)
170    CONTINUE

      DO 10, M = 1, 2
        DO 20, I = 0, NT
          MU (M,I) = 0.0
          SIGMA (M,I) = 0.0
          DO 30, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
            SIGMA (M,I) = SIGMA (M,I) + PSI(M,K,L) * PHI (L,I)
30          CONTINUE
20        CONTINUE
10      CONTINUE

      LNLIKE = 0
      TERM = 0

      DO 40, J = 1, 4
        IF (J .EQ. 1) THEN
          M = 1
          L = 1

```

```

      ELSEIF (J .EQ. 2) THEN
        M = 2
        L = 2
      ELSEIF (J .EQ. 3) THEN
        M = 2
        L = 1
      ELSEIF (J .EQ. 4) THEN
        M = 1
        L = 2
      ENDIF
      DO 50, I = 1, NY
        DO 60, KK = 1, ICOUNT (J,I)
          T = ISEQ (J,I,KK)
          IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
            & THEN
            RESID = (CLIMA(I,T)-MU(M,T))/SIGMA(M,T)-RAU(J,K)
            & *((CLIMA(I,T-1)-MU(L,T-1))/SIGMA(L,T-1))
            LNLIKE = LNLIKE + (RESID)**2
            TERM = TERM + LOG(SIGMA(M,T))
          ENDIF
        60 CONTINUE
      50 CONTINUE
    40 CONTINUE

    FUNC = -(-(NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2

    RETURN
    END

```

```

C -----
C ..... SUBROUTINE TO MULTIPLY TWO MATRICES
C -----

```

```

SUBROUTINE XNP (FIRST,SECOND,THIRD,ROWX,COLX,ROWA,COLA,RX,
C -----
C      & CX,RA,CA)
C -----

      INTEGER      ROWX,COLX,ROWA,COLA,RX,CX,RA,CA
      REAL          FIRST (RX,CX)
      REAL          SECOND (RA,CA)
      REAL          THIRD (RX,CA)

      IF (COLX .NE. ROWA) THEN
        PRINT *, 'MATRICES ARE NOT COMPATIBLE'
      ELSE
        DO 10, I = 1, ROWX
          DO 20, J = 1, COLA
            THIRD (I,J) = 0
            DO 30, K = 1, COLX
              IF (SECOND (K,J) .NE. -999.0) THEN
                THIRD(I,J) = THIRD(I,J)+FIRST(I,K)*SECOND(K,J)
              ENDIF
            30 CONTINUE
          20 CONTINUE
        10 CONTINUE
      ENDIF

      RETURN
      END

```

```

C      SUBROUTINE INV(MATT,NN,MM)
      -----

      REAL MATT(NN,NN), INVER(25,25)
      REAL MATR1(25,25)

      II = 0
20    II = II + 1

      MATR1(II,II) = 1.0 / MATT(II,II)

      DO 40 J = 1, MM
        DO 30 I = 1, MM

          IF (J .EQ. II .AND. I .EQ. II) THEN
            INVER(I,J) = (-1.0) * MATR1(II,II)

          ELSEIF (J .EQ. II .AND. I .NE. II) THEN
            INVER(I,J) = MATT(I,J) * MATR1(II,II)

          ELSEIF (I .EQ. II .AND. J .NE. II) THEN
            INVER(I,J) = MATT(I,J) * MATR1(II,II)

          ELSE
            INVER(I,J) = MATT(I,J) - ((MATT(I,II) *
&              MATR1(II,J)) * MATR1(II,II))
          ENDIF

30      CONTINUE
CC     PRINT*, (INVER(I,J), I = 1, MM)
40    CONTINUE

      CALL COPY(INVER,25,25,MATT,NN,NN,MM,MM)

      IF (II .LT. MM) GO TO 20

      RETURN
      END

```

```

C      -----
C      ..... SUBROUTINE TO GENERATE A VECTOR ACCORDING TO THE MODEL:-
C              S(t) = ALPHA(i)*PHI(i,t)
C      -----

```

```

C      SUBROUTINE GMEAN (MU,PHI,NT,NPARM,ALPHA,K,NV)
      -----

      REAL          MU (2,NT)
      REAL          PHI (NT,NPARM)
      REAL          ALPHA (NV,2,NPARM)

      DO 10, M = 1, 2
        DO 20, J = 1, NT
          MU (M,J) = 0.0
          DO 30, I = 1, NPARM
            MU (M,J) = MU (M,J) + ALPHA (K,M,I) * PHI (J,I)
30      CONTINUE
20    CONTINUE
10    CONTINUE
      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE GENERATES CLIMATE SEQUENCES ACCORDING
C TO THE SPECIFICATIONS OF MODEL4
C -----

C SUBROUTINE MOD4 (RAND,STATE,NV4,NV,SIGMA4,MU,RAU4,K,J,OBSN,
C & PSTATE,NT)
C -----

      INTEGER      PSTATE, STATE
      REAL          RAU4 (NV4)
      REAL          RAND (1,NV)
      REAL          OBSN (NV)
      REAL          SIGMA4 (2,NV4,0:NT)
      REAL          MU (2,NV,0:NT)

5  FORMAT (7F10.3)

      IF (J-1 .EQ. 0) THEN
        L = NT
      ELSE
        L = J-1
      ENDIF

      IF (K.EQ.5) THEN
        KK = 1
      ELSEIF (K.EQ.6) THEN
        KK = 2
      ENDIF

      OBSN (K) = SIGMA4(PSTATE,KK,J)*(RAND(1,K)+RAU4(KK)*(OBSN(K)-
& MU(STATE,K,L))/SIGMA4(STATE,KK,L))+MU(PSTATE,K,J)

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE GENERATES CLIMATE SEQUENCES ACCORDING
C TO THE SPECIFICATIONS OF MODEL3
C -----

SUBROUTINE MODEL1 (RAND,NV,SIGMA,MU,J,OBSN,PSTATE,NT,A,B,RES)
C -----

      INTEGER          PSTATE
      REAL              RAND (NV,1)
      REAL              SOLN (7,1)
      REAL              RES (NV,1)
      REAL              OBSN (NV,1)
      REAL              SIGMA (2,NV,0:NT)
      REAL              MU (2,NV,0:NT)
      REAL              A (NV,NV)
      REAL              B (NV,NV)

5      FORMAT (7F10.3)

      CALL MULT (B,RAND,SOLN,NV,NV,NV,1)
      CALL MULT (A,OBSN,RES,NV,NV,NV,1)

      DO 10, K = 1, NV
        OBSN (K,1) = RES (K,1) + SOLN (K,1)
10     CONTINUE

      DO 20, K = 1, NV
        RES (K,1) = OBSN (K,1)*SIGMA(PSTATE,K,J)+MU(PSTATE,K,J)
20     CONTINUE

      RETURN
      END

C -----
C ..... SUBROUTINE TO COMPUTE THE INVERSE OF A MATRIX
C -----

SUBROUTINE INVNP (NP,SOLN,NPARM)
C -----

      REAL              SOLN (NPARM,NPARM)

      DO 10, I = 1, NP
        DIAG = 1 / SOLN (I,I)
        SOLN (I,I) = 1
        DO 20, J = 1, NP
          SOLN (I,J) = SOLN (I,J) * DIAG
20     CONTINUE
        DO 30, K = 1, NP
          IF (I .NE. K) THEN
            DIAG = SOLN (K,I)
            SOLN (K,I) = 0
            DO 40, J = 1, NP
              SOLN (K,J) = SOLN (K,J) - SOLN (I,J) * DIAG
40     CONTINUE
          ENDIF
        CONTINUE
30     CONTINUE
10     CONTINUE

      RETURN
      END

```

```

C -----
C ..... THIS SUBROUTINE READS IN INITIAL PARAMETER ESTIMATES,
C THE CONVERGENCE CRITERION AND THE MAXIMUM NUMBER OF
C ITERATIONS TO BE PERFORMED
C -----

C SUBROUTINE INITIAL (EPS,MAXITER,ALPHA,PSI,RAU,NPARM,NV,NRAU)
C -----

REAL          ALPHA (2,NV,NPARM)
REAL          PSI (2,NV,NPARM)
REAL          RAU (NRAU,NV)

5  FORMAT (' EPS, MAXITER = ')
15  FORMAT (' ALPHA PARAMETERS = ')
25  FORMAT (' PSI PARAMETERS = ')
35  FORMAT (' RAU PARAMETER = ')

C  OPEN (UNIT=4,FILE='CON')
  OPEN (UNIT=4,FILE='\\WATER\\DATA\\INIT.DAT',STATUS='OLD')

  READ (4,*) EPS, MAXITER
  DO 20, K = 1, NV
    DO 10, M = 1, 2
      READ (4,*) (ALPHA (M,K,I), I = 1, NPARM)
      READ (4,*) (PSI (M,K,I), I = 1, NPARM)
10   CONTINUE
      READ (4,*) (RAU (L,K), L = 1, NRAU)
20  CONTINUE

  RETURN
END

```

```

C -----
C ..... THIS SUBROUTINE READS IN INITIAL PARAMETER ESTIMATES,
C THE CONVERGENCE CRITERION AND THE MAXIMUM NUMBER OF
C ITERATIONS TO BE PERFORMED
C -----

SUBROUTINE INT4 (ALPHA,PSI,RAU,NPARM,NV,NRAU,ISCALE)
C -----

      INTEGER          ISCALE (3,NV)
      REAL             ALPHA (2,NV,NPARM)
      REAL             PSI (2,NV,NPARM)
      REAL             RAU (NRAU,NV)

15  FORMAT (' ALPHA PARAMETERS = ')
25  FORMAT (' PSI PARAMETERS = ')
35  FORMAT (' RAU PARAMETER = ')

C  OPEN (UNIT=4,FILE='CON')
  OPEN (UNIT=4,FILE='\\WATER\\DATA\\INIT.DAT',STATUS='OLD')

  DO 20, K = 1, NV
    READ (4,*) ISCALE (1,K), ISCALE (2,K), ISCALE(3,K)
    DO 10, M = 1, 2
      READ (4,*) (ALPHA (M,K,I), I = 1, NPARM)
      READ (4,*) (PSI (M,K,I), I = 1, NPARM)
10  CONTINUE
    READ (4,*) (RAU (L,K), L = 1, NRAU)
20  CONTINUE

  RETURN
  END

```

```

C -----
C . . . . . THIS SUBROUTINE GENERATES RANDOM NORMAL NUMBERS WITH
C           MEAN ZERO AND STD. DEV. OF 1 AND THEN IS TRANSFORMED
C           INTO A RANDOM NUMBER WITH STD. DEV. OF SIGMA.
C -----

SUBROUTINE GAUSS (DECOMP,RAND)
C -----

    INTEGER          NV
    PARAMETER        (NV=7)
    INTEGER          ROWX,ROWA,COLX,COLA
    REAL             DECOMP (NV,NV)
    REAL             NRAND (1,NV)
    REAL             RAND (1,NV)

5   FORMAT (7F10.3)

C   . . . . . GENERATE RANDOM NORMAL (0,1) NUMBER

    CALL GRAND2 (NRAND,NV)

C   . . . . . GENERATE RANDOM NORMAL (0,S) NUMBER

    ROWX = 1
    ROWA = NV
    COLX = NV
    COLA = NV
    CALL MULT (NRAND,DECOMP,RAND,ROWX,COLX,ROWA,COLA)

    RETURN
    END

```

```

C -----
C ..... SUBROUTINE TO GENERATE A VECTOR ACCORDING TO THE MODEL:-
C          S(t) = ALPHA(i)*PHI(i,t) &
C          S(t) = PSI(i)*PHI(i,t)
C -----

C SUBROUTINE GAVSTD (MU,PHI,NT,NPARM,ALPHA,NV,PSI,SIGMA)
C -----

REAL          MU (2,NV,NT)
REAL          SIGMA (2,NV,NT)
REAL          PHI (NT,NPARM)
REAL          ALPHA (NV,2,NPARM)
REAL          PSI (NV,2,NPARM)

DO 10, M = 1, 2
  DO 20, J = 1, NT
    DO 40, K = 1, NV
      MU (M,K,J) = 0.0
      SIGMA (M,K,J) = 0.0
      DO 30, I = 1, NPARM
        MU (M,K,J) = MU(M,K,J)+ALPHA(K,M,I)*PHI(J,I)
        SIGMA(M,K,J) = SIGMA(M,K,J)+PSI(K,M,I)*PHI(J,I)
30      CONTINUE
40    CONTINUE
20  CONTINUE
10  CONTINUE

RETURN
END

```

```

C -----
C ..... SUBROUTINE TO COMPUTE RESIDUAL SERIES FOR MODEL 4
C -----

SUBROUTINE M4RES (RAU,ALPHA,PSI,PHI,COUNT,SEQ,CLIMA,
C -----
& NT,NY,NPARM,NV,K,NRAU,NP,CONVG)
C -----

INTEGER          COUNT (4,NY)
INTEGER          SEQ (4,NY,NT)
INTEGER          T,CONVG
REAL             AKAIKE,LNLIKE,PI
PARAMETER        (PI=3.141593)
REAL             CLIMA (NY,0:NT)
REAL             MU (2,0:365)
REAL             SIGMA (2,0:365)
REAL             RESID (7,12,365)
REAL             PSI (2,NV,NPARM)
REAL             ALPHA (2,NV,NPARM)
REAL             PHI (NPARM,0:NT)
REAL             RAU (NRAU,NV)

5  FORMAT (7F10.4)
15 FORMAT (' AKAIKE" S CRITERION FOR VARIABLE:', I4, ' IS:', F30.4)

OPEN (UNIT=14,FILE='\\WATER\\DATA\\RESI4.DAT',STATUS='UNKNOWN')
OPEN (UNIT=9,FILE='LPT1')

IF (CONVG.EQ.0) THEN
    GOTO 250
ENDIF

DO 10, M = 1, 2
    DO 20, I = 0, NT
        MU (M,I) = 0.0
        SIGMA (M,I) = 0.0
        DO 30, L = 1, NPARM
            MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
            SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
30        CONTINUE
20        CONTINUE
10        CONTINUE

DO 80, I = 1, NY
    DO 90, J = 1, NT
        RESID (K,I,J) = -999.00
90        CONTINUE
80        CONTINUE
LNLIKE = 0
TERM = 0

DO 40, J = 1, 4

```

```

      IF (J .EQ. 1) THEN
        M = 1
        L = 1
      ELSEIF (J .EQ. 2) THEN
        M = 2
        L = 2
      ELSEIF (J .EQ. 3) THEN
        M = 2
        L = 1
      ELSEIF (J .EQ. 4) THEN
        M = 1
        L = 2
      ENDIF
      DD 50, I = 1, NY
      DO 60, KK = 1, COUNT (J,I)
        T = SEQ (J,I,KK)
        IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
          & THEN
          RESID(K,I,T) = (CLIMA(I,T)-MU(M,T))/SIGMA(M,T)-RAU(1,K)
          & *((CLIMA(I,T-1)-MU(L,T-1))/SIGMA(L,T-1))
          LNLIKE = LNLIKE + (RESID(K,I,T))**2
          ENDIF
          TERM = TERM + LOG(SIGMA(M,T))
60      CONTINUE
50      CONTINUE
40      CONTINUE

      LNLIKE = -((NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2
      AKAIKE = -2*LNLIKE+2*NP

      WRITE (9,15) K, AKAIKE

250     IF (K .EQ. 7) THEN
          DO 100, I = 1, NY
            DO 70, T = 1, NT
              WRITE (14,5) (RESID (K,I,T), K = 1, NV)
70          CONTINUE
100         CONTINUE
          ENDIF

      RETURN
      END

```

```

C -----
C ..... SUBROUTINE TO COMPUTE RESIDUAL SERIES FOR MODEL 5
C -----

SUBROUTINE M5RES (RAU,ALPHA,PSI,PHI,COUNT,SEQ,CLIMA,
C -----
& NT,NY,NPARM,NV,K,NRAU,NP,CONVG)
C -----

INTEGER          COUNT (4,NY)
INTEGER          SEQ (4,NY,NT)
INTEGER          T,CONVG
REAL             LNLIKE,AKAIKE,PI
PARAMETER        (PI=3.141593)
REAL             CLIMA (NY,0:NT)
REAL             MU (2,0:365)
REAL             SIGMA (2,0:365)
REAL             RESID (7,12,365)
REAL             PSI (2,NV,NPARM)
REAL             ALPHA (2,NV,NPARM)
REAL             PHI (NPARM,0:NT)
REAL             RAU (NRAU,NV)

5  FORMAT (7F10.4)
15 FORMAT (' AKAIKE"S CRITERION FOR VARIABLE:', 14, ' IS:', F10.4)

C  OPEN (UNIT=14,FILE='A:RESIDU.DAT',STATUS='UNKNOWN')
  OPEN (UNIT=14,FILE='\\WATER\\DATA\\RESI5.DAT',STATUS='UNKNOWN')
  OPEN (UNIT=9,FILE='LPT1')

  IF (CONVG.EQ.0) THEN
    GOTO 250
  ENDIF

  DO 10, M = 1, 2
    DO 20, I = 0, NT
      MU (M,I) = 0.0
      SIGMA (M,I) = 0.0
      DO 30, L = 1, NPARM
        MU (M,I) = MU (M,I) + ALPHA (M,K,L) * PHI (L,I)
        SIGMA (M,I) = SIGMA (M,I) + PSI (M,K,L) * PHI (L,I)
30      CONTINUE
20      CONTINUE
10      CONTINUE

  DO 80, I = 1, NY
    DO 90, J = 1, NT
      RESID (K,I,J) = -999.00
90      CONTINUE
80      CONTINUE
      LNLIKE = 0
      TERM = 0

```

```

DO 40, J = 1, 4
  IF (J .EQ. 1) THEN
    M = 1
    L = 1
  ELSEIF (J .EQ. 2) THEN
    M = 2
    L = 2
  ELSEIF (J .EQ. 3) THEN
    M = 2
    L = 1
  ELSEIF (J .EQ. 4) THEN
    M = 1
    L = 2
  ENDIF

  DO 50, I = 1, NY
    DO 60, KK = 1, COUNT (J,I)
      T = SEQ (J,I,KK)
      IF ((CLIMA(I,T).NE.-999).AND.(CLIMA(I,T-1).NE.-999))
        & THEN
        & RESID(K,I,T) = (CLIMA(I,T)-MU(M,T))/SIGMA(M,T)-
        & RAU(J,K)*((CLIMA(I,T-1)-MU(L,T-1))/
        & SIGMA(L,T-1))
        LNLIKE = LNLIKE + (RESID(K,I,T))**2
        ENDIF
      TERM = TERM + LOG(SIGMA(M,T))
60    CONTINUE
50    CONTINUE
40    CONTINUE
    LNLIKE = -((NY*NT)/2)*LOG(2*PI)-TERM-LNLIKE/2
    AKAIKE = -2*LNLIKE+2*NP
    WRITE (9,15) K, AKAIKE

250  IF (K .EQ. 7) THEN
    DO 100, I = 1, NY
      DO 70, T = 1, NT
        WRITE (14,5) (RESID (K,I,T), K = 1, NV)
70      CONTINUE
100    CONTINUE
    ENDIF

    RETURN
  END

```

```

C -----
C ..... SUBROUTINE TO MINIMIZE A FUNCTION
C -----

SUBROUTINE POLRIB (THETA,NP,TOL,ITER,FMIN)
C -----

INTEGER          NPMAX,MAXITER
PARAMETER        (NPMAX=20)
PARAMETER        (MAXITER=200)
PARAMETER        (EPS=1.E-10)
REAL             NUM

DIMENSION        THETA(NP),GRAD(NPMAX),DIR(NPMAX),DER(NPMAX)

OPEN (UNIT=9,FILE='LPT1')

FTHETA=FUNC(THETA)
WRITE (9,*) 'FTHETA', FTHETA
CALL DFUNC(THETA,DER)
DO 10, J=1,NP
    GRAD(J)=-DER(J)
    DIR(J)=GRAD(J)
    DER(J)=DIR(J)
10  CONTINUE

DO 20, I=1,MAXITER
    ITER=I
    CALL MINL (THETA,DER,NP,FMIN)
    WRITE (9,*) 'FMIN', FMIN
    IF (2.*ABS(FMIN-FTHETA).LE.TOL*(ABS(FMIN)+ABS(FTHETA)+EPS))
        &      RETURN
    FTHETA=FUNC(THETA)
    CALL DFUNC(THETA,DER)
    DENOM=0.
    NUM=0.
    DO 40, J=1,NP
        DENOM=DENOM+GRAD(J)**2
        NUM=NUM+(DER(J)+GRAD(J))*DER(J)
40  CONTINUE
    IF (DENOM.EQ.0.) RETURN
    GAMMA=NUM/DENOM
    DO 50, J=1,NP
        GRAD(J)=-DER(J)
        DIR(J)=GRAD(J)+GAMMA*DIR(J)
        DER(J)=DIR(J)
50  CONTINUE
20  CONTINUE

PRINT *, 'DID NOT CONVERGE'

RETURN
END

```

```

C -----
C FUNCTION TO GENERATE A UNIFORM NUMBER
C -----

C FUNCTION URAN1 (SEED)
C -----

    DIMENSION R(97)
    PARAMETER (M1=259200,IA1=7141,IC1=54773,RM1=3.8580247E-6)
    PARAMETER (M2=134456,IA2=8121,IC2=28411,RM2=7.4373773E-6)
    PARAMETER (M3=243000,IA3=4561,IC3=51349)

    DATA INIT /0/

    IF (SEED.LT.0.OR.INIT.EQ.0) THEN
        INIT=1
        IX1=MOD(IC1-IDUM,M1)
        IX1=MOD(IA1*IX1+IC1,M1)
        IX2=MOD(IX1,M2)
        IX1=MOD(IA1*IX1+IC1,M1)
        IX3=MOD(IX1,M3)
        DO 10 J = 1,97
            IX1=MOD(IA1*IX1+IC1,M1)
            IX2=MOD(IA2*IX2+IC2,M2)
            R(J)=(FLOAT(IX1)+FLOAT(IX2)*RM2)*RM1
10        CONTINUE
        SEED=1
    ENDIF
    IX1=MOD(IA1*IX1+IC1,M1)
    IX2=MOD(IA2*IX2+IC2,M2)
    IX3=MOD(IA3*IX3+IC3,M3)
    J=1+(97*IX3)/M3
    IF (J.GT.97.OR.J.LT.1) PAUSE
    URAN1=R(J)
    R(J)=(FLOAT(IX1)+FLOAT(IX2)*RM2)*RM1

    RETURN
    END

```

```

C -----
C ..... THIS SUBROUTINE READS IN INITIAL PARAMETER ESTIMATES,
C          THE CONVERGENCE CRITERION AND THE MAXIMUM NUMBER OF
C          ITERATIONS TO BE PERFORMED
C -----

C SUBROUTINE INTAL3 (EPS,MAXITER,ALPHA,SIGMA,RAU,NPARM,NV,NRAU)
C -----

      REAL          ALPHA (2,NV,NPARM)
      REAL          SIGMA (NRAU,NV)
      REAL          RAU (NRAU,NV)

      5  FORMAT ( ' EPS, MAXITER = ' )
     15  FORMAT ( ' ALPHA PARAMETERS = ' )
     25  FORMAT ( ' SIGMA PARAMETERS = ' )
     35  FORMAT ( ' RAU PARAMETER = ' )

C  OPEN (UNIT=4,FILE='CON')
  OPEN (UNIT=4,FILE='\\WATER\\DATA\\INIT.DAT',STATUS='OLD')

  READ (4,*) EPS, MAXITER
  DO 20, K = 1, NV
    DO 10, M = 1, 2
      READ (4,*) (ALPHA (M,K,I), I = 1, NPARM)
     10  CONTINUE
      READ (4,*) (SIGMA (L,K), L = 1, NRAU)
      READ (4,*) (RAU (L,K), L = 1, NRAU)
     20  CONTINUE

  RETURN
  END

```